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¹ Predicting aerosol reactivity across scales: from the

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laboratory to the atmosphere

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6ABSTRACT

7To fully utilize the results of laboratory-based studies of the chemistry of model atmospheric 8aerosol reactions, it is important to understand how to relate them to the conditions found in 9nature. In this study, we have taken a validated reaction-diffusion mechanism for oxidation of 10C₃₀H₆₂ aerosol by OH under flow tube conditions and examined its predictions for another 11experimental regime (continuous flow stirred tank reactor) and for the atmosphere, spanning 12alkane aerosol viscosities from liquid to semi-solid. The results show that under OH-13concentration-limited and aerosol-mixing-limited conditions, it should be possible to select 14laboratory experimental conditions where many aspects of the particle phase and volatile product 15chemistry under atmospheric conditions can be revealed. If the OH collision and organic 16diffusion rates are comparable, however, reactivity is highly sensitive to the details of both OH 17concentration and internal mixing. The characteristics of the transition between limiting 18conditions provide key insights into which parts of the reaction mechanism dominate in the 19various kinetic regimes.

20Introduction 21

Laboratory studies are central for interpreting field observations and parameterizing 23models to understand and ultimately predict chemical transformations in the atmosphere.¹ Gas-24phase rate coefficients, yields and formation mechanisms of secondary organic aerosol, and 25multiphase transformation mechanisms and rates provide quantitative parameters needed to 26improve the fidelity of regional and climate models. A key challenge, noted by many, is 27developing reliable methods to mimic or reliably extrapolate the chemistry conducted in the 28laboratory-scale reactors to the real atmosphere.²⁻⁷ Central to this challenge is the difficulty of 29matching laboratory and atmospheric concentrations and timescales that often differ by orders of 30magnitude.

To determine the oxidation lifetime of many trace gas species, rate coefficients can be 32easily quantified under laboratory conditions and directly applied to atmospheric conditions. 33Here the experimental challenge centers on understanding complex reaction networks and 34accurately providing oxidation conditions and trace gas levels to reproduce the competition 35between bimolecular and unimolecular pathways in polluted and pristine regions of the 36atmosphere. This is nicely illustrated by the recent discovery of the importance of unimolecular 37autoxidation pathways^{8, 9} of gas-phase peroxy radicals (RO₂) when bimolecular reactions of NO 38are suppressed (e.g. a source of aerosol in the pristine regions of the atmosphere¹⁰⁻¹²). Another 39example is the formation and chain propagation of aerosol-phase alkoxy radicals when reaction 40conditions are designed to favor $RO_2 + NO$ or SO_2 (i.e., the urban atmosphere) rather than typical 41laboratory oxidation conditions that naturally favor $RO_2 + RO_2$ pathways.^{13, 14}

42 It is not so straightforward, however, to examine the coupling of gas-aerosol interactions 43 with oxidative ageing of aerosols. Direct extrapolation of laboratory observations to the 44atmosphere becomes more tenuous due to the emergence of additional timescales (i.e., particle-45phase diffusion, uptake, evaporation) within the aerosol particles that can govern the kinetics. 46For example, heterogeneous oxidation lifetimes (involving OH or O_3) and oxidized product 47 distributions of organic species in aerosol are not determined solely by intrinsic reactivity but are 48often complex, time-dependent functions of internal aerosol mixing timescales.¹⁵⁻²² This is 49because when internal mixing is slow (e.g., a semisolid), reactions are localized at the aerosol 50surface, while rapid mixing in liquids results in uniform composition changes and scaling of 51 overall reactivity with particle volume. Whether a reaction is localized to a surface or scales 52 with volume does not depend on aerosol viscosity alone, but is an emergent property¹⁸ that 53depends upon oxidant concentration in the gas phase and particle diameter. What is most 54 important is the time between oxidant collisions with the particle surface relative to the mixing 55time within the aerosol. Thus, in the laboratory, reactions using semisolid aerosol at high [OH] 56 and short reaction times may not replicate the heterogeneous chemistry occurring over the long 57time, low oxidant field of the atmosphere. A semisolid aerosol in the laboratory may in fact 58appear well-mixed in the atmosphere since aerosol constituents have ample time to sample the 59entire particle volume between reactive events. New approaches are needed to better understand 60 when the results of existing laboratory experiments can be extrapolated to other conditions, and

61to guide the design of experiments that more faithfully replicate the coupling of diffusive and 62reactive timescales of the atmosphere.

63 In the present work, we report results of a comprehensive computational study that 64 provides a means to connect laboratory data to atmospheric aerosol chemistry using a single 65reaction-diffusion mechanism that has been validated for liquid and semisolid alkane aerosol 66 oxidation.^{22, 23} It builds on an initial investigation in which we predicted the compositions of 200 67nm diameter aerosol particles with a range of viscosities intermediate between liquid and solid **68** for an OH partial pressure typical of laboratory flow tube conditions ([OH] = 5.04×10^{10} 69molec./cm³).¹⁹ We introduced a reaction-diffusion index (I_{RD}) to provide a quantitative link 70between particle viscosity and reactivity, enabling identification of the specific parts of the free 71 radical chain reaction mechanism that dominate aerosol composition when the relative extent of 72mixing is varied. Here, we use the same reaction scheme to predict both aerosol and volatile 73product compositions as well as aerosol size and mass under OH densities typical of flow tube 74(FT, $[OH] = 5.04 \times 10^{10}$ molec./cm³) and continuous flow stirred tank²⁴ (CFSTR, $[OH] = 5.04 \times 10^{10}$ molec./cm³) 7510^8 molec./cm³) reactors and the atmosphere ([OH] = 5.04×10^6 molec./cm³). The results are 76 analyzed in terms of I_{RD} , allowing identification of regimes where atmospheric reactivity of 77alkanes is readily predictable from laboratory experiments, and where it is not. The simulations 78 provide insights to how product distributions, regimes, and factors that control the kinetics are 79connected.

80Simulations 81

The free radical oxidation scheme used in this work was developed and validated for 83reaction between OH and $C_{30}H_{62}$ in liquid (squalene) and semisolid (triacontane) under laboratory 84flow tube conditions.^{22, 23} The chain reaction is launched by H abstraction from the alkane, 85forming alkyl radicals that rapidly condense with available O_2 to form peroxy radicals. O_2 is 86assumed to be instantly replenished from the surrounding gas due to its fast diffusion through 87liquids and solid organics.^{25, 26} The peroxy radicals disproportionate and fragment through 88subsequent chain reactions, forming a variety of condensed and gas - phase products. A summary 89of the reaction scheme is presented in the Supporting Information (SI), Section 1. Since known 90or calculated rate constants are used for the elementary reaction steps involved, rather than 91adjustable parameters, and the mechanism spans initial abstraction to fully oxidized CO_2 , the 92simulations are predictive for a broad range of reaction conditions even outside of those used to 93establish the reaction scheme.

In the description we have developed for the free radical reactions involved in alkane 950xidation,²³ the individual product molecules formed by free radical functionalization and 96fragmentation reactions are not explicitly included. Rather, molecules are represented as a set of 97functionalities and carbon backbone lengths to allow the essential characteristics of the chemical 98system to be captured, while minimizing the extreme complexity of tracking all the kinds of 99molecules that can be generated by a free radical chain spanning H abstraction from pure 100hydrocarbon to formation and evaporation of CO₂. The initial $C_{30}H_{62}$ molecule is taken to be a 101straight chain with primary and secondary alkyl groups for all aerosol viscosities, thus neglecting 102the branching that adds tertiary carbon reactions²³ and reduces viscosity.²⁷ The free radical 103groups generated during the reaction are alkyl, peroxy, alkoxy, acyl, acyl peroxy and acyloxy, 104located at primary or secondary positions. The radicals react further to form ketone, aldehyde, 105carboxylic acid and alcohol functionalities, accompanied by C_{30} chain fragmentation. Stochastic 106reaction-diffusion simulation methods²³ have been used to perform simulations using this scheme 107to generate spatially resolved concentration *vs* time curves that include the effect of variable 108volume on instantaneous rates. The software package is Kinetiscope.²⁸ The 200-nm-diameter 109aerosol is represented by an array of 200 13.17 nm × 13.17 nm × 0.5 nm compartments located 110along the particle radius when full reaction-diffusion simulations are performed, or by a single 111compartment when the diffusion coefficient is fast enough for the particle to be well-mixed at all 112times.

113 In our initial study assuming FT conditions,¹⁹ the $C_{30}H_{62}$ diffusion coefficients were 114varied from 1 x 10⁻¹¹ (liquid) to 8.39 × 10⁻¹⁹ cm²/s (semi-solid using the diffusion coefficient of 115triacontane derived from viscosity measurements²⁹) in increments of factors of 10, and all 116product species in the simulation were assumed to diffuse with the same coefficient as the initial 117alkane throughout the system, neglecting possible decreases in viscosity as the particle 118composition changes. This range is typical of self-diffusion in semisolid α -pinene SOA.³⁰ In this 119paper, we use this scheme to predict both gas and condensed phase products for the same 120viscosity range over a broad OH concentration range, representing FT, CFSTR and atmospheric 121conditions. Table 1 lists the OH densities studied, with reaction times selected to reach a constant 122total OH exposure of 3.024 × 10¹² molecules cm⁻³ s. The largest diffusion coefficients used for 123fully spatially resolved FT, CFSTR, and atmospheric reaction-diffusion simulations are 10⁻¹¹, 10⁻¹² 124¹³, and 10⁻¹⁵ cm²/s, respectively. Above these values, the aerosol is well-mixed and the 125calculations become very expensive. Instead, we can represent the system accurately by single 126compartment simulations with instantaneous mixing as shown in Houle et al¹⁸ and Wiegel et

127al.^{22, 23}

128

129Table 1. Simulation conditions for $OH + C_{30}H_{62}$

Conditions	OH	Reaction time	Self-diffusion	OH Exposure
	concentration	(s)	coefficients	(molec. cm ⁻³
	molec/cm ³		(cm^2/s)	s)
Laboratory flow	$5.04 \times 10^{10(31, 32)}$	60	instantaneous,	3.024×10^{12}
tube (FT) ^a			$10^{-11} - 8.39 \times 10^{-19}$	
Continuous flow	$5.04 \times 10^{8(24,13,14)}$	6000	instantaneous,	
stirred tank			$10^{-13} - 8.39 \times 10^{-19}$	3.024×10^{12}
(CFSTR)				
Atmosphere,	$5 \times 10^{6} {}^{(33,34)}$	604800 (1 week)	instantaneous,	3.024×10^{12}
daytime			$10^{-15} - 8.39 \times 10^{-19}$	

130 a. Simulation results for FT conditions are taken from Houle et al.¹⁹

131Results 132

133 The simulations provide complete data sets for volume, mass, carbon chain lengths and 134functionality concentrations in the particle and gas phases as a function of time. Spatial 135distributions of all particle-phase species are generated when a multicompartment scheme is 136used. These results are analyzed to compare the outcomes of the oxidation reaction over a broad 137range of conditions, providing new information on how reactivity varies with OH concentration 138and aerosol viscosity. Details of the simulated data are presented in the SI, Section 2.

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140I. Trends in consumption of C₃₀H₆₂ by OH reactions

141 A striking result of the simulations is the emergence of distinctive patterns in several 142particle properties as viscosity and OH concentration are varied. For example, nearly identical 143C₃₀H₆₂ decay curves (Figure S1) and spatial distributions (Figure S2) as a function of time are 144found for FT, CFSTR and atmospheric OH densities when diffusion coefficients are offset by 145about 2 orders of magnitude. These self-similarity trends are surprising but can be rationalized 146by considering the relative rates of reaction and internal mixing via self-diffusion, which controls 147the concentration of abstractable hydrogen near the aerosol surface where most reactions with 1480H take place.^{18, 35} As described previously,¹⁹ a reaction-diffusion index, *I_{RD}* can be defined as:

149
$$I_{RD} = \frac{k_{rxn} [OH(g)] d_p^2}{4 \pi^2 D} (1)$$

150where k_{rxn} is the rate constant for reaction between OH and species in the particle, d_p is the **151**particle diameter, [OH(g)] is the concentration of OH in the gas phase and *D* is the self-diffusion

152coefficient of organic molecules in the particle. When important characteristics of the reactions

153 such as initial uptake

154



155 156 157**Figure 1**. Initial uptake coefficients for $C_{30}H_{62}$ reacting with all three OH concentrations as a 158function of I_{RD} . The IRD ranges corresponding to the limiting kinetic regimes are labeled.

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160coefficients and ratios of products as a function of OH concentration and *D* are plotted as a 161function of I_{RD} , trends are captured in a way that allows these characteristics to be compared self-162consistently. As shown in Figure 1, initial uptake coefficients calculated from the decay curves in 163Figure S1 are nearly coincident, indicating that at the very early stages of oxidation, the kinetic 164factors that control $C_{30}H_{62}$ consumption are essentially unchanged over a 10⁴ variation in OH 165concentration. Table S1 lists the values for I_{RD} for each combination of diffusion coefficient and 166OH concentration used in Figure 1. 167 The definition of I_{RD} is closely related to the definition of the uptake coefficient, γ , 168through the reacto-diffusion length, L.¹⁸ If the aerosol is well-mixed, 2*L* is equal to the diameter 169of the particle. If it is not, $d_p > 2L$, and the more general expression for γ is

170
$$\gamma = \frac{2k_{rxn}\rho_0 N_A}{3cMd_p^2} \left[d_p^3 - (d_p - 2L)^3 \right] (2)$$

171where ρ_0 is the aerosol density, *c* is the velocity of the OH molecules, *M* is the molecular weight 172of the aerosol material, and N_A is Avogadro's number. Specifically,¹⁸

173
$$L = \left(\frac{2k_{rxn}}{I_{RD}\pi^3 c\sigma}\right)^{1/2} (3)$$

174where σ is the probability that a colliding OH will stick to the aerosol surface long enough to 175react.

As I_{RD} increases due to decreasing *D* (increasing viscosity) and/or increasing OH gas 177concentration, both *L* and γ decrease. This trend is evident in Figure 1, where all three γ vs I_{RD} 178datasets fall on a single curve, with 3 distinct regimes. When $I_{RD} < 10^{-2}$, γ is essentially constant at 179a value similar to that for OH reaction with liquid squalane under FT conditions, ³⁶ and reflects 180reactivity controlled only by OH reactions at the aerosol surface. When $I_{RD} > 10^2$, γ is similarly 181constant, but with a value lower than reported in previous work.^{13, 37} However, a quantitative 182comparison of the predicted and experimental values for γ is complicated by the fact that the 183decay curves used to calculate γ are not exponential and the experimental particle geometry is 184not well-defined. The OH concentration- and organic-mixing-limited cases are separated by a 185transition region where γ decreases rapidly.

186

187II. Trends in physical properties

188 Although the uptake coefficients follow a regular pattern, other characteristics do not. 189Aerosol final diameters and masses are compared in Figure 2, and masses as a function of time 190for the 3 OH densities are shown in Figure S3. The variation in final diameter with I_{RD} is not 191monotonic: it passes through a minimum around I_{RD} =10-100, which coincides with the uptake 192coefficient transition regime shown in Figure 1. The final particle mass, on the other hand, 193decreases as I_{RD} increases. Under well-mixed conditions when $I_{RD} < 10^{-2}$, there is a net mass



194

195Figure 2. Aerosol diameter and mass characteristics after OH exposure. The mass is relative to **196**the initial mass of unreacted $C_{30}H_{62}$.

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198increase relative to the pure alkane due to oxygen incorporation. The mass increase competes 199with mass loss due to fragmentation and volatilization as the diffusion coefficient decreases and 200the chemistry becomes self-diffusion limited. The extent of mass loss increases with decreasing 201gas - phase OH concentration, and is most pronounced under atmospheric conditions, where OH 202concentration is lowest. This result indicates that there are significant differences in reactivity as 203OH concentration and particle viscosity are varied even though the reaction mechanism is the 204same.



205

206Figure 3. Total functional groups present (a) in the aerosol particle and (b) in the gas phase **207**(including CO₂) after OH exposure, normalized to the total amount of $C_{30}H_{62}$ reacted, for the **208**three OH densities.

209 210III Trends in particle and gas - phase oxidation products

The simulations track the functionalities formed during oxidation in both the condensed 212and gas phases. Figure 3 shows the total accumulated oxygen-containing functionalities in the 213particle and gas phases per $C_{30}H_{62}$ reacted, which reflects the extent of reaction over the range of 214OH densities and viscosities. Figure 3a shows that functionality accumulation in the particle is 215similar for all OH densities when mixing is fast, increases significantly in the I_{RD} transition 216region, then decreases as the reaction becomes mixing limited. Figure 3b shows that this trend is 217accompanied by very low fragmentation and volatilization to form oxidized gas products when 218mixing is fast transitioning to abundant gas product formation as I_{RD} increases into the mixing-219limited regime.

Plots of the total amounts of each stable functionality as a function of time – ketones, 221alcohols, aldehydes, carboxylic acids and CO_2 – normalized to the instantaneous amount of 222unreacted $C_{30}H_{62}$ are presented in Figures S4-S12, organized according to I_{RD} . The data show 223that while the functionality compositions of the aerosol particles are similar for each I_{RD} across 224the full range of OH densities, those of the volatilized gases are not. Figures 4 and 5 summarize 225these results for the transition regime, showing the final functionality distributions 226corresponding to the totals in Figure 3 for particles and gas phase, respectively.



229Figure 4. Composition of aerosol particles after exposure to OH in the transition kinetic regime230between [OH] limited and self-diffusion limited.



233

234**Figure 5.** Composition of the gas phase surrounding the aerosol particles after exposure to OH in 235the transition kinetic regime between [OH] limited and self-diffusion limited. (a) gas products 236that originate from minority functionality components in the aerosol; (b) gas products that 237originate from majority functionality components in the aerosol.

The simulation results in Figure 4 show that there is a dramatic change in composition as 239a function of I_{RD} in the transition regime, with FT, CFSTR and atmospheric OH densities having 240the same chemistry although the extents of oxidation are quite different (Figure 3). Ketones 241become the dominant functional group, accompanied by a modest increase in carboxylic acids. 242The ketones are relatively unreactive to OH and accumulate as the reaction becomes self-243diffusion-limited, unlike the alcohols which undergo a significant decrease in concentration. The 244aldehydes, present in small amounts, also decrease with increasing I_{RD} . Carboxylic acids, 245although minor products, accumulate as the reaction becomes mixing limited.

The gaseous product distribution is qualitatively different from the particle phase 247composition, as shown in Figure 5. Although ketones and alcohols are the major particle phase 248components of the aerosol, aldehydes and CO₂ are the major volatile products. The OH 249concentration conditions affect the gas composition for two of the products: FT conditions, 250which have the highest OH-aerosol collision frequency, favor CO₂ volatilization, while 251atmospheric conditions favor ketone volatilization.

Inspection of the molecular chain lengths (carbon numbers) at I_{RD} =10 for the three OH 253regimes (Figure S13) shows that although the aerosol particle functionality distributions are very 254similar for the three OH concentration cases in the transition regime (Figure 4), the carbon 255numbers are not. As the OH concentration decreases from FT to atmosphere, the average carbon 256number at the end of OH exposure increases. Figure 2 shows that the particle mass decreases, 257however, with mass loss being greatest for the lowest OH concentration when I_{RD} =10. These 258results together with the data in Figures 3a, 5 and S7 show that under atmospheric 259concentrations, volatilization of oxidized functionalities is favored relative to FT conditions.

260Discussion

Using a detailed reaction-diffusion mechanism developed and validated for the reaction 262of OH + semisolid and liquid $C_{30}H_{62}$ under FT conditions,^{19, 23, 29} we have performed a 263computational study to predict oxidative aging chemistry for the much lower OH partial 264pressures typical of laboratory CFSTR and atmospheric conditions over a broad range of alkane 265viscosities. Strong patterns in chemical and physical properties are evident as OH concentration 266and alkane self-diffusion coefficients are varied over ranges of 4 and 7 orders of magnitude, 267respectively. These patterns are rationalized in terms of the reaction-diffusion index I_{RD} , which is 268a measure of the balance between the availability of OH from the gas phase for H abstraction 269reactions and the extent of mixing by self-diffusion in the aerosol, for each combination of OH 270concentration and particle viscosity.

The results in Figures 4, and S4-S6 suggest that laboratory measurements of alkane 272aerosol particle compositions during reactions with OH under both FT and CFSTR conditions 273may be predictive of compositions under atmospheric conditions when the chemistry is strongly 274limited by arrival of OH at the aerosol surface ($I_{RD} < 10$). When $I_{RD} > 10$, the aerosol 275compositions are qualitatively similar, but not identical (Figures 4 and S7-S8). These trends 276continue into the higher I_{RD} range. The results in Figures 5 and S4-S12 show clearly that 277laboratory measurements will only be qualitatively predictive for the limiting cases and not 278predictive at all in the transition regime. Experimental tests of these findings will be very 279valuable for model validation and improvement, particularly measurements of particle size in the 280transition regime, and of the ratio of ketones to alcohols in the particle phase across the I_{RD} range.

To understand these trends, we examine the free radical oxidation chemistry involved in 282the three kinetic regimes, which is sketched in Figure 6. When the oxidation kinetics are in the 283OH concentration-limited range, the aerosol is relatively C-H rich. Abstraction events generate 284peroxy radicals, which form ketones and alcohols as the primary oxidation products for the OH 285exposure times studied. This simple picture only applies when internal mixing is fast and 286oxidized moieties are dilute. As the rates of OH reaction and self-diffusion become competitive,

287the *sec*-peroxy radical concentration increases from near-zero to significant amounts and the 288reaction becomes confined to the near-surface region of the aerosol, as shown in Figures S2, S14 289and S15.



290

292Figure 6. Abstraction, propagation, and product formation steps for the heterogeneous reaction 293of OH with alkanes. The blue arrows mark steps that are second order in peroxy radical 294concentration. Acyloxy and alkoxy radicals undergo unimolecular decomposition reactions and 295generate additional peroxy radicals. Under reaction-limited conditions (dashed line box), ketone 296and alcohol formation dominate. Under mixing limited conditions, the entire free radical chain is 297active.

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The other radicals, which are more reactive, are present in negligible quantities. Under 300these conditions, disproportionation of peroxy radicals to form alkoxy radicals and O_2 , which is 301quite slow when the aerosol is well-mixed,¹⁸ becomes kinetically significant. Alkoxy radicals

302fragment to form aldehydes and primary alkyl radicals that react and dissociate to form mixture 303of alkyl, peroxy, acyl, acyloxy and acylperoxy radicals. This intensely reacting environment 304rapidly consumes available C-H and O-H bonds, and the steady state population of peroxy 305radicals decreases to near-zero. These conditions favor acyl radical formation and fragmentation 306to form CO₂.

307 The chemical trends for alkane oxidation shown in Figure 6 connect the reaction 308mechanism to the product functionality distributions shown in Figures S4-S12. When $I_{RD} < 1$, 309particle phase ketones and alcohols dominate for all OH densities, with very little fragmentation 310to form volatiles, indicating that the chemistry does not go deeply into the free radical chain. **311**This I_{RD} range corresponds to $D > 10^{-12}$ cm²/s for FT conditions, $D > 10^{-14}$ cm²/s for CFSTR 312 conditions, and $D > 10^{-16} \text{ cm}^2/\text{s}$ for atmospheric conditions for this particle size (200 nm initially). 313This suggests that FT measurements with liquid phase aerosol can be informative for 314atmospheric conditions even when the target atmospheric material is quite viscous. When I_{RD} > 31510^2 , the product distribution is also simple, dominated by ketones in the particle phase and CO₂ 316in the gas phase, consistent with formation of terminal products. The self-diffusion coefficient 317 range for this regime is in the highly viscous fluid to semisolid range: $D > 10^{-14}$ cm²/s for FT **318** conditions, $D > 10^{-16}$ cm²/s for CFSTR conditions, and $D > 10^{-18}$ cm²/s for atmospheric conditions 319 for 200 nm particles. Given the challenges of generating well-controlled solid aerosol 320nanoparticles, laboratory investigations of the oxidation of highly viscous liquids may be a 321 reasonable proxy to learn about reactivity in the atmosphere. The correlation by I_{RD} allows 322development of similar useful experimental strategies for other particle diameters.

323 While the initial uptake coefficients and condensed phase functionality distributions are 324independent of OH concentration in the transition region, the gas product distribution, extent of 325 fragmentation, particle mass and particle diameter are not, and it is not possible to extrapolate 326between the OH concentration ranges without use of explicit reaction-diffusion models. **327**Moreover, some features appear only in this range. For example, as shown in Figure 3, there is a 328dramatic increase in formation of oxygen-containing functionalities in the particle and mass loss 329relative to the OH concentration and self-diffusion limited regimes, as well as an increase in 330peroxy radical densities. Examination of Figures S4-S12 reveals that the primary chemical 331signature of the transition region is the OH concentration-dependence of the relative amounts of 332gaseous products, specifically CO₂ relative to aldehydes. In addition, Figure S15 shows that 333peroxy radicals have a lower steady state concentration above $I_{RD} = 1$ for CFSTR conditions 334 relative to FT and atmosphere. This pattern reveals that the delicate balance between mixing and 335reaction mainly affects the formation and subsequent reactions of alkoxy radicals. Since 336aldehydes are formed via H abstraction from alcohols and subsequent fragmentation to form 337aldehydes and new peroxy radicals, this suggests that the peroxy radical disproportionation 338chemistry dominates under CFSTR conditions, reducing the peroxy radical population relative to 339 higher and lower OH densities and generating volatile aldehydes. FT and atmospheric OH 340densities, on the other hand, favor acyl radical formation and fragmentation. Although OH 341concentration is lowest under atmospheric conditions, the extent of erosion of the aerosol is 342 greatest in the transition region, leading to a change of about 20% in particle diameter. This 343 indicates that fragmentation to form volatiles is favored when free radical chain initiation events 344are relatively rare, consistent with carbon numbers in Figure S13.

Previous investigations have also sought to develop predictive correlations for 346heterogeneous aerosol reactions using reaction- and diffusion-limited kinetic regimes.¹⁵ In 347particular, low diffusion rates in semisolids have been proposed to lead to extended lifetimes for 348organics in the troposphere.³⁸⁻⁴⁰ It has been suggested that aerosol becomes more viscous during 349aging, and hence less prone to degradation.⁴¹ The broad trends investigated here for an alkane 350model system indicate that extrapolations across a range of [OH] and viscosity conditions may 351not be possible in any simple way depending on the details of the chemistry, reacting 352environment and particle sizes involved.

The combinations of chemical and physical factors that control lifetimes at various stages 354of aging remain to be fully identified, but are likely to involve the following elements. Important 355reactions introduced by the presence of alkenes include oligomerization and efficient free radical 356cycling via hydroxyalkyl chemistry.⁴² While the presence of NO₂ can be neglected if the aerosol 357is relatively dry, heterogeneous NO reactions can accelerate the oxidation process. NO promotes 358free radical cycling and fragmentation by increasing the extent of formation of alkoxy radicals 359even under well-mixed conditions.¹³ Uptake of gaseous water becomes important as the particle 360becomes more hydrophilic. Particle-phase water adds acid-base chemical pathways that can 361promote fragmentation⁴³ and can decrease viscosity as can chain fragmentation under dry 362conditions. Inclusion of all of these elements, which are very important to be able to describe 363aerosol aging correctly, can be done in a straightforward way by modifying the basic mechanistic 364framework used here, if sufficient experimental data are available to validate the calculations. 365We can anticipate some of the effects. As viscosities decrease, they will cause a decrease in the 366relevant I_{RD} (Eq 1), and may cause the system to enter a kinetic regime that is more gas-collision367controlled. However, the effect of such a change on particle composition and properties would 368likely not be pronounced unless the change is from the mixing-limited to the transition regime or 369from the transition regime to the OH limited regime. While we compare 200 nm particles in all 370cases, the I_{RD} analysis allows experimental data obtained on one size of particle to be used to 371predict oxidation on a different size if there are no changes in the chemistry or viscosity. 372Extensions of the present work to include additional reactions and composition-dependent 373viscosities will be valuable for development of new generalizations of chemical transformations 374for atmospheric chemistry models.

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377ASSOCIATED CONTENT

378Supporting Information. Summary of the OH oxidation mechanism. Detailed plots of simulation379results including spatial maps and compositions as a function of time. This material is available380free of charge via the Internet at http://pubs.acs.org.

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555TOC graphic

