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Net fluxes of broadband shortwave and photosynthetically active radiation complement NDVI and near infrared reflectance of vegetation to explain gross photosynthesis variability across ecosystems and climate

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1 Net Fluxes of Broadband Shortwave and Photosynthetically
2 Active Radiation Complement NDVI and Near Infrared
3 Reflectance of Vegetation to explain Gross Photosynthesis
4 Variability Across Ecosystems and Climate

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28 **Abstract**

29 A grand challenge in global change research is understanding how the interaction of
30 vegetation with the environment influences ecosystem gross primary productivity (GPP)
31 through carbon assimilation. An evolving goal is to continuously predict GPP variability
32 everywhere by finding a robust scaling relationship between flux tower GPP and satellite
33 spectral reflectance. The footprint mismatch between the pixel size of many early satellite
34 measurements and eddy flux measurements is a major hindrance in such an endeavor. By
35 using a large set of growing season data covering 100 site-years in North and Central
36 America, we explored the potential of transforming incident and reflected shortwave (R_g)
37 and photosynthetically active radiation (PAR) measurements into a broadband
38 normalized difference vegetation index (NDVI) and near-infrared (NIR) reflectance of
39 vegetation (NIR_v) which simultaneously explains the GPP variability. We found that the
40 broadband NDVI and NIR_v derived from R_g and PAR measurements at the daily time
41 scale were highly correlated with Planet Fusion, Landsat-8/9, and Sentinel-2 narrowband
42 NDVI and NIR_v across a wide range of climate and ecological gradients. The differences
43 between satellite and broadband NDVI and NIR_v were found to be significantly
44 associated with soil background variations, phenological stages, water stress and signal
45 saturation of broadband NIR reflectance at high biomass. The seasonal variability of
46 broadband NDVI and NIR_v remarkably captured the seasonality of vegetation phenology,
47 evaporative fraction, GPP and rainfall in different ecosystems. Although a saturation of
48 GPP at high NDVI was evident, a linear relationship between broadband NIR_v times
49 incident PAR versus GPP indicated the strength of NIR_v-based approach to capture the
50 hidden light use efficiency impacts on GPP. We conclude that the inexpensive

51 measurement of R_g and PAR components can provide highly reliable information on
52 NDVI, NIR_v, and GPP uninterruptedly thereby augmenting the proximal sensing
53 capability of the flux tower sites without the need for additional spectrometer
54 measurements. The proposed in-situ vegetation indices make a stronger case on the use of
55 radiation signals for handshaking between ecosystem-scale measurements and remote
56 sensing observables relevant to carbon uptake.

57 Keywords: Spectral reflectance, broadband, vegetation index, NIR_v, gross primary
58 productivity, photosynthetically active radiation, ecosystem, climate

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71 **1. Introduction**

72 Vegetation is an integral component of the biosphere influencing the variability of
73 energy, water, and carbon dioxide fluxes (Ryu et al., 2012; Hoek van Dijke et al., 2020;
74 Camps-Valls et al., 2021; Brown et al., 2017). Systematic information of biophysical
75 metrics that describe vegetation vigor, phenological development, and biomass
76 production are required to enhance our understanding of the flux variabilities in the
77 climate system, for ecosystem monitoring and agricultural management practices (Brown
78 et al., 2017; Richardson et al., 2010; Zhang et al., 2004, Sellers et al., 1997; Foley et al.,
79 2011; Godfray et al., 2010). Consequently, leaf area index (LAI) and fraction of absorbed
80 photosynthetically active radiation (FAPAR) are identified as two of the essential climate
81 variables by the Global Climate Observing System (GCOS).

82 For the large-scale monitoring of vegetation development through remote sensing
83 satellites, FAPAR and LAI are not available as direct measurements, and they need to be
84 retrieved through complex radiative transfer models. However, there are two more
85 biophysical metrics namely NDVI (Normalized Difference Vegetation Index) and GCC
86 (Green Chromatic Coordinate) that are closely related to vegetation growth and
87 development yet have a proximity with both FAPAR and LAI (Seyednasrollah et al.,
88 2019; Gitelson et al., 2019; Richardson et al., 2007, 2013; Hao et al., 2012). In this
89 context, NDVI can be directly obtained from the amount of reflectance in red and near-
90 infrared (NIR, hereafter) regions of the electromagnetic spectrum. Similarly, GCC can
91 also be directly calculated from the amount of reflectance in red, green, and blue regions
92 (Richardson et al., 2007; Brown et al., 2017).

93 Using the theory of strong absorption of photosynthetically active radiation (PAR) in the
94 red region and dissipation of energy through reflection of the NIR radiation, a vast body
95 of literature explored the potential of satellite derived NDVI to understand the variability
96 in gross primary productivity (GPP, hereafter) with vegetation growth and development
97 (Ustin and Middleton, 2021; Liu et al., 2020, Magney et al., 2019, Huang et al., 2019;
98 Prabhakara et al., 2015; Mutanga et al., 2023; Tesfaye and Awoke, 2021; Mutanga and
99 Skidmore, 2004). An asymptotic pattern in NDVI at maximum vegetation growth became
100 evident from all these studies, and NDVI yielded poor GPP estimates in evergreen
101 vegetation (Pierrat et al., 2022) or during the peak seasons when vegetation reaches
102 maturity (Mutanga et al., 2023; Mutanga and Skidmore, 2004). This saturation is
103 attributed to the imbalance due to the insensitivity of chlorophyll absorbing red light at
104 dense canopy cover (Kumar et al., 2001, Mutanga et al., 2023) versus a simultaneous rise
105 in the NIR reflectance, apparently leading to negligible changes in NDVI. Studies
106 showed that NIR reflectance scales with leaf nitrogen (Ollinger et al., 2011), and
107 therefore NIR reflectance can be used as an index of photosynthetic capacity (Field and
108 Mooney, 1986).

109 Following the analogy of linearity between GPP versus the product of absorbed PAR and
110 light use efficiency, NIR reflectance of vegetation (NIR_v) (product of NDVI and NIR
111 reflectance) is explored to understand the magnitude and variability of GPP at hourly-to-
112 daily and from ecosystem to global scale (Badgley et al., 2017; Baldocchi et al., 2020).
113 The philosophy of linking GPP with NIR_v is based on the fact that increasing biomass
114 (leaf layer) in the canopy results in multiple scattering, which leads to significant changes
115 in NIR reflectance in moderate-to-high vegetation density (LAI from 2 to 6) (Sellers et

116 al., 1997). Some of the more recent studies demonstrated a much tighter coupling when
117 GPP is linked with the product of NIRv and incident PAR (Dechant et al., 2022). All
118 these studies generated encyclopedic understanding on the pros and cons of NDVI and
119 NIRv towards explaining the GPP variability across and within ecosystems. However,
120 there are some open challenges. Firstly, how to bridge the scale gap in linking global
121 remote sensing reflectance with eddy covariance GPP? Secondly, how to obtain the best
122 and consistent NDVI and NIRv information at the same scale of flux tower GPP
123 measurements? Thirdly, how to extrapolate the findings of a handful of ecosystems to
124 global scale?

125 Over the last decade, there has been great evolution on in-situ monitoring of phenology
126 through the PhenoCam network (<http://phenocam.sr.unh.edu/>) (Richardson et al., 2013;
127 Filippa et al., 2018; Petach et al., 2014; Browning et al., 2017; Burke et al., 2021; Zhou et
128 al., 2020; Tian et al., 2021). The broad objective is to develop deep insights into the
129 temporal variation in phenology across (within) different (same) ecosystems, and how
130 this variability is driven by environmental factors such as radiation, temperature, and
131 precipitation. PhenoCam provides data at an intermediate scale between ground
132 observations and satellite remote sensing. This camera-based monitoring of vegetation
133 phenology is standardized with consumer-grade digital cameras (e.g., Sonnentag et al.,
134 2012) which records a three-layer image (red, green, and blue: RGB) and a NIR
135 monochrome image. Broadband NDVI can be calculated from these paired images,
136 however, a correction is needed if the exposure between the two images is different and
137 there is a need for empirical adjustments to make the camera NDVI match with satellite
138 NDVI (Petach et al., 2014). Most of the cameras have a nearly horizontal field of view

139 with about a quarter of the image sky. This can lead to earlier green up and saturation as
140 compared to PAR and shortwave radiation sensors with vertical fields of view that better
141 match satellite imagery. Another disadvantage of camera-based approach is the
142 pronounced variability in normalized channel brightness resulting from changes in
143 quality and quantity of incident solar radiation (Richardson et al., 2007, Liu et al., 2022).
144 Like the dedicated NDVI sensors, the PhenoCam network is a relatively new invention
145 whereas many measurements of PAR and shortwave radiation extend much further back
146 in time.

147 One of the emerging utilities of FLUXNET are continuous observations of NDVI and
148 NIRv for assessing the contribution of vegetation seasonality on energy, water, and
149 carbon fluxes at the corresponding flux tower footprint (Hoek van Dijke et al., 2020).
150 Despite satellite NDVI providing global coverage of vegetation vigor, current NDVI
151 products suffer from trade-off between high (low)-spatial and low(high)-temporal
152 resolution. While coarse spatial resolution (250 m) satellite observations are available as
153 continuous time series (e.g., MODIS and VIIRS), finer spatial resolution vegetation
154 information (10 – 30 m) is available only as a discrete time series (e.g., Sentinel-2 and
155 Landsat8/9). Contamination of satellite observations due to the cloud interference brings
156 hindrance while diagnosing the seasonal variation of vegetation attributes. Therefore,
157 NDVI and NIRv at the ‘eyes and ears’ of the flux towers and at the temporal resolution of
158 flux measurements is a critical requirement, especially since it is nearly impossible to
159 measure LAI daily and without destruction. This could complement operational remote
160 sensing data and document considerable diversity in plant development and seasonality in

161 greenness over the representative vegetation in which the flux towers operate. This will
162 simultaneously complement and magnify the legacy research of PhenoCam.

163 Many AmeriFlux sites are equipped with Decagon/METER (SRS-Ni, SRS-Nr) or
164 Apogee (S2-111-SS, S2-112-SS, S2-411-SS, S2-412-SS) NDVI sensors on flux towers to
165 capture the rapid change in ecosystem greenness and temporal variability of NDVI.
166 However, these sensors have only been available since 2015, and the Decagon/METER
167 model has already gone out of production. Both sensor models had early issues with
168 stability as well (Anderson et al., 2016). The Apogee sensors have a similar cost to other
169 research-grade radiation sensors (~\$600 US). Thus, a major challenge concerns how to
170 observe temporally continuous NDVI and NIRv at the flux tower sites accurately,
171 inexpensively, and over the entire data record.

172 Based on measurements of incident (i) and reflected (r) components of PAR (symbolized
173 as Q in the equations and figures) (Q^i , Q^r) in conjunction with incident (i) and reflected (r)
174 shortwave radiation (R_g) (R_g^i , R_g^r), Huemmrich et al. (1999) and Wilson and Meyers
175 (2007) showed the possibility of estimating a robust broadband NDVI, but with limited
176 evaluation with respect to spatially coarse satellite NDVI over a restricted number of
177 sites. The approach has great potential as measurements of Q contain information on the
178 visible waveband and greenness, and net flux of $R_g - Q$ measurements inform us about
179 infrared reflectance (shortwave minus visible). While Rocha and Shaver (2009) evaluated
180 broadband NDVI and enhanced vegetation index (EVI) at a burnt and an unburnt site in
181 the high latitude, Rocha et al. (2021) assessed the effects of solar position on the
182 relationship between ecosystem function and NDVI derived from $R_g - Q$ measurements.
183 However, it remains unclear how well a broadband NDVI and NIRv perform in a range

184 of ecosystems and how they relate to carbon uptake when plants are exposed to large
185 fluctuations of covarying biometeorological limits. Therefore, the current work seeks to
186 address the following science questions and objectives: *(SQ1) Can a broadband NDVI
187 and NIRv retrieved from the flux tower shortwave and PAR measurements explain
188 satellite NDVI and NIRv variability across ecosystems and climate of varying energy-
189 water availability limits? (SQ2) How effectively does the broadband NDVI capture the
190 phenological changes and function of vegetation on the land surface? (SQ3) Can we use
191 broadband NDVI and NIRv as a robust modulator of GPP across a wide range of
192 energy-water availability? (SQ4) How does the background soil exposure variations,
193 phenology, radiation components, and water stress impact the estimation of broadband
194 NDVI?*

195 We present a cross-site synthesis of PAR and R_g observations to derive a broadband
196 NDVI and NIRv (hereafter denoted as $NDVI_{bb}$ and $NIRv_{bb}$) across a diverse range of
197 ecosystems, water- and energy-limited conditions. $NDVI_{bb}$ and $NIRv_{bb}$ patterns and their
198 seasonal variability were compared with satellite NDVI, NIRv, and GCC at multiple
199 spatial scales from 3-30 m spatial resolution using PLANET Fusion (continuous time
200 series) and Harmonized Landsat Sentinel (HLS) (discrete time series) datasets. We
201 analyzed the relationship between $NDVI_{bb}$, $NIRv_{bb}$, and PAR variability with GPP to fill
202 a critical scale gap between flux footprint, ecosystem, and satellite remote sensing. As a
203 final test, we assessed the role soil background variations in different phenological stages,
204 radiation components, and water stress variations on the performance of $NDVI_{bb}$.

205 **2. Materials and methods**

206 **2.1. Estimating NDVI_{bb} and NIR_{vbb}**

207 Beer's law already provides the theoretical link between incident, transmitted, and
208 absorbed PAR versus leaf area. We hypothesize that from the net fluxes of PAR and R_g
209 component measurements we can directly estimate NDVI_{bb} and NIR_{vbb}, which
210 simultaneously explains GPP variability (Baldocchi et al., 2020; Wilson and Meyers,
211 2007; Huemmrich et al., 1999). For estimating NDVI_{bb} and NIR_{vbb}, estimation of visible
212 and near-infrared reflectance ($\rho_{\text{vis,bb}}$, $\rho_{\text{nir,bb}}$) in the broad visible (0.4 – 0.7 μm) and near-
213 infrared to shortwave infrared spectrum (0.7 – 3 μm) is needed. The derivation of
214 NDVI_{bb} is based on the theory of satellite narrowband NDVI. Vegetation shows strong
215 absorption (85 – 90%) and low reflectance and transmittance (5 – 10%) in the visible
216 wavelength. However, they show substantially higher reflectance, transmittance, and low
217 absorption in the NIR radiation wavelength (Wilson and Meyers, 2007; Campbell and
218 Norman, 1998). **Figure 1a** shows the conceptual diagram for estimating $\rho_{\text{vis,bb}}$ and $\rho_{\text{nir,bb}}$
219 from PAR and R_g measurements.

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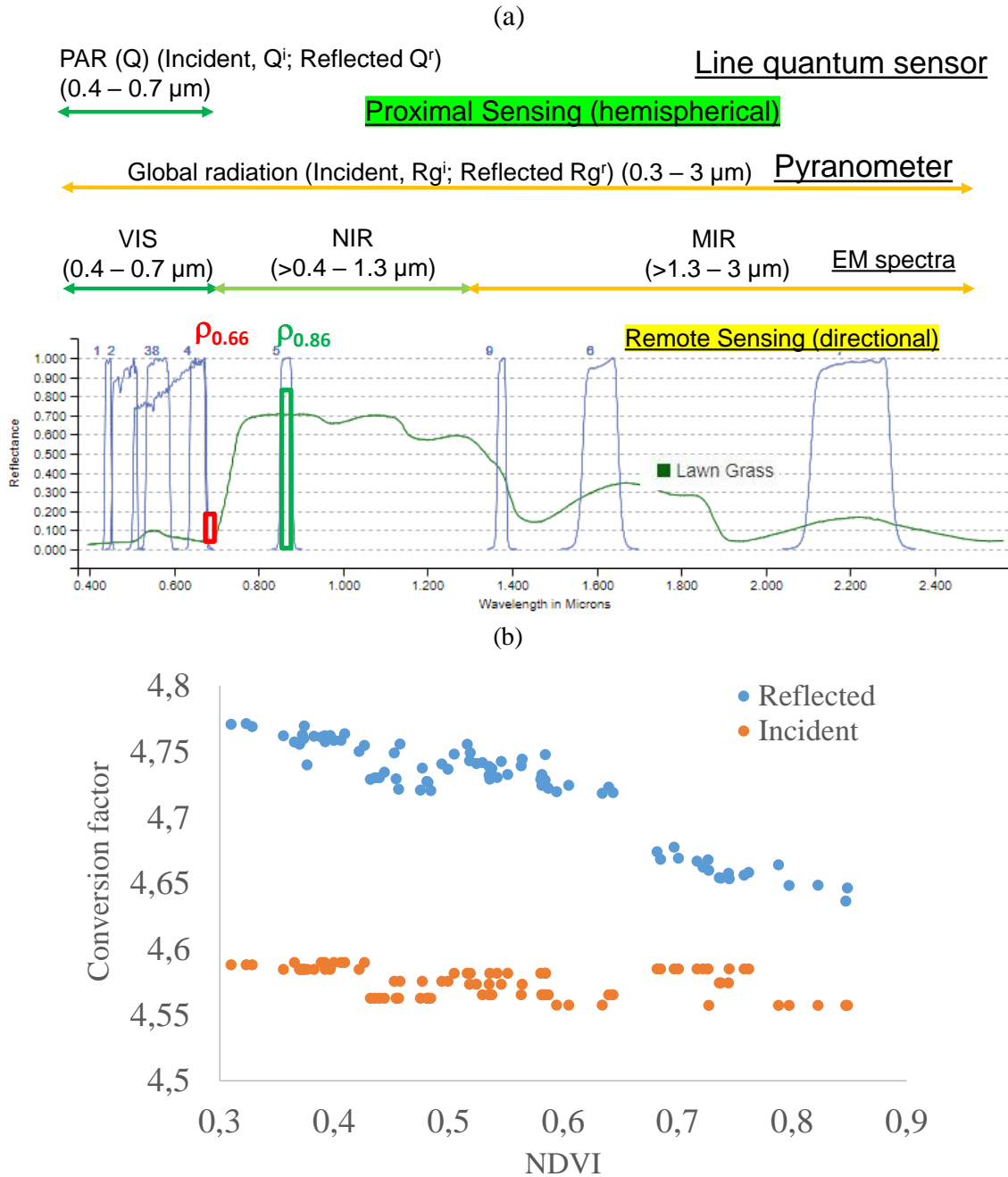


Figure 1: (a) Conceptual diagram showing the hypothesis for estimating broadband spectral reflectance from the measurements of hemispherical broadband radiation components in PAR and total shortwave spectral region. It also shows an example of the narrowband spectral reflectances that we obtain in red and near infrared spectral region from operational remote sensing satellite Landsat-9 (Source: <https://landsat.usgs.gov/spectral-characteristics-viewer>). VIS signified visible, NIR signifies near-infrared, MIR signifies mid-wave infrared. (b) Figure showing the scaling factor for converting PAR (both incident and reflected) from $\mu\text{mols}/\text{m}^2/\text{s}$ to W/m^2 for a range of NDVI as an example over rice crop in California.

226 A pyranometer measures energy flux density in its native spectral range (0.3 – 3 μm) and
 227 the quantum sensor measures photon flux density in its native range (0.4 – 0.7 μm) (**Fig.**
 228 **1**). To produce a broadband NDVI that deduces the reflectance of energy, we need to do a
 229 transformation, starting with the principle of Planck’s law ($E = h\nu$, $h = \text{Planck’s constant}$
 230 and $\nu = \text{frequency of radiation}$) and information on the incoming solar spectrum and
 231 reflected spectrum. Therefore, we used the incident and reflected PAR measurements in
 232 conjunction with incident and reflected R_g measurements to segregate $\rho_{\text{vis,bb}}$ and $\rho_{\text{nir,bb}}$.
 233 The measurements of hemispherical broadband PAR and R_g components act as the
 234 proximal sensing data source to retrieve equivalent estimates of narrowband directional
 235 reflectances in red and NIR regions as obtained from remote sensing satellites (**Fig. 1**).
 236 The central wavelength of narrowband red and NIR directional reflectance of operational
 237 remote sensing satellite is around 0.66 ($\rho_{0.66}$) and 0.86 ($\rho_{0.86}$) μm . We hypothesize that
 238 separation of $\rho_{\text{vis,bb}}$ and $\rho_{\text{nir,bb}}$ from proximal sensing of broadband hemispherical PAR
 239 and R_g components are approximately equivalent to $\rho_{0.66}$ and $\rho_{0.86}$. Therefore,
 240 $\text{NDVI}_{\text{bb}} = (\rho_{\text{nir,bb}} - \rho_{\text{vis,bb}})/(\rho_{\text{nir,bb}} + \rho_{\text{vis,bb}})$
 241 NDVI_{bb} is considered as approximately equivalent proxy for the standard NDVI [$\text{NDVI} =$
 242 $(\rho_{0.66} - \rho_{0.86})/(\rho_{0.66} + \rho_{0.86})$]. The implication of approximating $\rho_{\text{vis,bb}} \approx \rho_{0.66}$ and $\rho_{\text{nir,bb}} \approx$
 243 $\rho_{0.86}$ in different ecosystems are described in detail in section 3.4.
 244 At first, $\rho_{\text{vis,bb}}$ was approximated from reflected (r) and incident (i) components of PAR
 245 ($\mu\text{mol}/\text{m}^2/\text{s}$) measurements (symbolized as Q) as follows:

$$\rho_{\text{vis,bb}} = Q^r/Q^i \tag{1}$$

246 For estimating $\rho_{\text{nir,bb}}$, the shortwave radiation components (R_g^i and R_g^f) were partitioned
247 into downward broadband visible ($\text{VIS}_{i,\text{bb}}$) and near-infrared ($\text{NIR}_{i,\text{bb}}$) components
248 following Weiss and Norman (1985) and Wilson and Meyers (1999), however with little
249 modification.

$$\text{VIS}_{i,\text{bb}} = k_{\text{vis}} R_g^i \quad (2)$$

250 Here k_{vis} is the ratio of Q^i and R_g^i , where Q^i in $\mu\text{mols}/\text{m}^2/\text{s}$ was converted to W/m^2 as
251 ($Q^i/4.5946$). This conversion factor is based on the energy of photons of visible light in
252 $0.4 - 0.7 \mu\text{m}$ region of the electromagnetic spectrum. Considering green wavelength
253 ($0.55 \mu\text{m}$) as the central average wavelength in the entire visible band ($0.4 - 0.7 \mu\text{m}$), we
254 can apply Planck's law as $E = hc/\lambda = Nhc/\lambda$. Where, h = Planck's constant (6.626×10^{-34}
255 Js), c = speed of light ($3 \times 10^8 \text{ m/s}$), N = Avogadro number ($6.022 \times 10^{23} \text{ mol}^{-1}$). By
256 putting the central wavelength of Q^i ($\lambda = 0.55 \mu\text{m}$), we can derive the scaling factor (i.e.,
257 $4.5946 \mu\text{mols}/\text{joules}$) to convert Q^i from $\mu\text{mols}/\text{m}^2/\text{s}$ to W/m^2 . Maximum plant
258 photosynthesis occurs in the blue ($0.44 \mu\text{m}$) and red light ($0.66 \mu\text{m}$) (Liu and Van Iersel,
259 2021). Putting these values in the conversion equation will make the conversion factor
260 $3.6757 \mu\text{mols}/\text{joules}$ and $5.5135 \mu\text{mols}/\text{joules}$ for blue and red bands, respectively.
261 However, these are the maximum and minimum conversion factors. Averaging these
262 three conversion factors from blue, green, and red leads to the mean value of 4.5946
263 $\mu\text{mols}/\text{joules}$. Alternatively, by applying Planck's law at every $0.01 \mu\text{m}$ interval within
264 the visible spectrum, followed by integration over the entire visible band also results in
265 the same value. Thus, for every datapoint k_{vis} varies instead of assuming a constant
266 (Weiss and Norman, 1985; Wilson and Meyers, 1999).

$$\mathbf{NIR}_{i,bb} = k_{nir}R_g^i \quad (3)$$

267 Given there is no universal consensus on converting the energy of photons from μmols to
268 watts beyond the visible region, we cannot apply the same factor to estimate k_{nir} from the
269 reflected components of PAR and R_g for the NIR region. The conversion factor of 4.5946
270 $\mu\text{mols/joules}$ is applicable for Q^i as we have Planck's law and a known solar spectrum.
271 However, the reflected light is filtered and the filtering changes with season, leaf area,
272 soil etc. Until now, there is no report on a scaling factor for converting reflected PAR
273 from the molar to energy unit. Deriving such a scaling factor needs hyperspectral data of
274 reflected PAR and shortwave radiation spectra. The UC Berkeley Biomet lab had
275 collected reflected PAR and shortwave radiation spectra over rice, and we have computed
276 this scaling factor for the reflected PAR for different classes of NDVI (**Fig. 1b**). While
277 the conversion factor for the incoming PAR changes marginally with NDVI (4.56 – 4.60
278 $\mu\text{mols/joules}$), the conversion factor for the reflected PAR changes with season from 4.78
279 to 4.64 $\mu\text{mols/joules}$ (**Fig. 1b**). The conversion factor of 4.72 $\mu\text{mols/joules}$ is the average
280 value for the reflected PAR as derived from the available observations in rice. Computing
281 this value over other vegetation types is not within the scope of this study since it needs
282 hyperspectral measurements.

283 In the present case, we estimate k_{nir} as $(1 - k_{vis})$. This gives us the advantage that no
284 additional uncertainty is introduced due to the conversion from molar unit to energy unit
285 for Q^r . This is also another reason that we did not adopt the approach of Huemmrich et al.
286 (1999) to bypass any uncertainty for converting the reflected component of PAR from
287 molar to energy unit.

288 From the partitioning of $VIS_{i,bb}$, the reflected upward broadband visible component
289 ($VIS_{r,bb}$) is estimated as follows.

$$VIS_{r,bb} = \rho_{vis,bb} VIS_{i,bb} \quad (4)$$

290 For estimating the upward reflected broadband near-infrared ($NIR_{r,bb}$), we used the R_g^r
291 (W/m^2) measurements as follows.

$$NIR_{r,bb} = R_g^r - VIS_{r,bb} \quad (5)$$

292 From the information of $NIR_{r,bb}$ and $NIR_{i,bb}$, we can now estimate $\rho_{nir,bb}$ as follows.

$$\rho_{nir,bb} = NIR_{r,bb}/NIR_{i,bb} \quad (6)$$

293 From eqs. (1) and (6), VI_{bb} is computed as follows.

$$NDVI_{bb} = (\rho_{nir,bb} - \rho_{vis,bb})/(\rho_{nir,bb} + \rho_{vis,bb}) \quad (7)$$

294 Reflected near-infrared radiation from the vegetation, $NIR_{v,bb}$, was calculated in terms of
295 a renormalized $NDVI_{bb}$ times broadband NIR reflectance ($\rho_{nir,bb}$) ($NIR_{v,bb} =$
296 $NDVI_{bb} * \rho_{nir,bb}$) (Baldocchi et al., 2020; Badgley et al., 2019).

297 The daytime PAR and R_g components measured between 10:00 to 14:00 h were used for
298 computing $\rho_{vis,bb}$ and $\rho_{nir,bb}$. The purpose of selecting this time slot is, all the operational
299 remote sensing satellites have equatorial crossing time either around 10:00 – 11:00 h
300 (Terra platform) or around 13:00 – 14:00 h (Aqua platform) (Wilson and Meyers, 2007).
301 Thus, the comparison between satellite versus broadband NDVI will be coherent in this
302 way. The daily values of $\rho_{vis,bb}$ and $\rho_{nir,bb}$ was estimated by averaging their 30-min values
303 from 10:00 to 14:00 h, followed by the calculation of $NDVI_{bb}$ and $NIR_{v,bb}$.

304 It is important to remember that satellite NDVI accounts for the signals of the entire field
305 of view of the sensors. Based on the spatial resolution of the sensors, there is a possibility
306 of inclusion of background soil reflectance in satellite NDVI due to different soil
307 reflectance factors in red and NIR wave bands. Such possibility also exists while deriving
308 $NDVI_{bb}$ from PAR and R_g measurements at the flux tower sites. For instance, in the
309 deciduous vegetation and annual crops, early and late in the growing season when leaf
310 area is small, the soil background can be seen by the sensors. The background reflectance
311 can substantially influence the spectral reflectance (both for satellite and proximal) from
312 the closed canopy due to multiple scattering in the NIR and SWIR bands. The extent of
313 such background effects will be different in two methods of estimating NDVI. We
314 anticipate that the impact of variations in soil background will be consequently reflected
315 in their comparison.

316 **2.2. Sites and data**

317 The site locations, biome, vegetation type, climate and associated information are listed
318 in **Table 1**. The analysis was carried out for cropland (CRO), grassland (GRA), woody
319 savanna (WSA), open shrubland (OSH), forest (FOR), and nontidal wetlands (WET).
320 These are AmeriFlux (11 sites) and NEON (14 sites) sites with publicly available data
321 accessible through the respective AmeriFlux web pages. Seven out of eleven AmeriFlux
322 site are from University of California, Berkeley, Biometeorology lab
323 (<https://nature.berkeley.edu/biometlab/sites.php>) and the sites characteristics are
324 published by the group (Baldocchi et al., 2020; Eichelmann et al., 2018; Hemes et al.,
325 2019; Ma et al., 2016). The remaining four AmeriFlux sites are maintained by University
326 of Nebraska (US-Ne3), University of Illinois (US-UiA, US-UiB), and United States

327 Department of Agriculture (US-Wkg), respectively. The description of the NEON sites is
 328 available in the NEON web page ([https://www.neonscience.org/field-sites/explore-field-](https://www.neonscience.org/field-sites/explore-field-sites)
 329 sites) and also in the site information of AmeriFlux
 330 (<https://ameriflux.lbl.gov/sites>). Croplands were a mix of rainfed (Ne3, UiA, UiB, xSL)
 331 and irrigated sites (Bi1, Bi2). While Bi2 received subsurface flooding irrigation, Bi1
 332 received subsurface ditch irrigation (Bi2 receives a single irrigation in July-August, Bi1
 333 receives 2 irrigations around May-June and August-September). Irrespective of single
 334 and multiple cropping systems, majority of the sites are covered with temporary crops
 335 followed by harvest and a bare soil period (Ne3, UiA, UiB, xSL). The time period of data
 336 availability for the individual sites are also mentioned in Table 1.

337 **Table 1.** Sites characteristics where both incident and reflected photosynthetically active
 338 radiation measurements are available (Superscripts, ^P = PLANET fusion; ^{HLS} =
 339 Harmonized Landsat and Sentinel); Planet fusion: 01/2018 – 12/2021; Landsat: **01/2014 –**
 340 **12/2021**; Sentinel-2: **01/2016 – 12/2021**)

Biome	Site	Vegetation type	Latitude, Longitude	P (mm)	Climate type	Time period	Reference
CRO	Bi1 ^{P, HLS}	Alfalfa	38.0992, -121.4993	338	Csa	2016 – 2021	Wang et al. (2023)
	Bi2 ^{P, HLS}	Corn	38.1091, -121.5351	338	Csa	2017 – 2021	Baldocchi et al. (2020)
	Ne3 ^{HLS}	Corn-soybean	41.1797, -96.4397	783	Dfa	2003 – 2021	Suyker et al. (2005)
	UiA ^{HLS}	Switchgrass	40.0646, -88.1961	1051	Dfa	2015	Blackely et al. (2022)
	UiB ^{HLS}	Miscanthus	40.0628, -88.1984	1051	Dfa	2014 – 2016	Blackely et al. (2022)
	xSL ^{HLS}	Mixed	40.4619, -103.0293	432	Bsk	2017 – 2021	Metzger et al. (2019)
GRA	xAE ^{HLS}	Herbaceous	35.4106, -99.0588	780	Cfa	2017 – 2021	Metzger et al. (2019)
	xCP ^{HLS}	Herbaceous	40.8155, -104.7456	320	Bsk	2017 – 2021	Metzger et al. (2019)

	xKA ^{HLS}	Herbaceous	39.1104, - 96.6129	850	Cfa	2017-- 2021	Metzger et al. (2019)
	xKZ ^{HLS}	Herbaceous	39.1008, - 96.5631	870	Cfa	2017-- 2021	Metzger et al. (2019)
	Var ^{P,HLS}	Herbaceous	38.4133, - 120.9508	559	Csa	2000 – 2021	Baldocchi et al. (2020)
	Wkg ^{HLS}	Herbaceous	31.7365, - 109.9419	407	Bsk	2004 – 2021	Scott et al. (2010)
WSA	Ton ^{P,HLS}	Herbaceous, understory	38.4309, - 120.9660	559	Csa	2001 – 2021	Baldocchi et al. (2020)
	xSJ ^{HLS}	Herbaceous, understory	37.1088, - 119.7323	540	Csa	2018 – 2021	Metzger et al. (2019)
OSH	xJR ^{HLS}	woody (evergreen or deciduous)	32.5907, - 106.8425	270	Bsk	2017 – 2021	Metzger et al. (2019)
	xNQ ^{HLS}	woody (evergreen or deciduous)	40.1776, - 112.4524	288	Dfb	2017-- 2021	Metzger et al. (2019)
	xSR ^{HLS}	woody (evergreen or deciduous)	31.9107, - 110.8355	346	Bsk	2017 – 2021	Metzger et al. (2019)
FOR	xAB ^{HLS}	ENF	45.7624, - 122.3303	2450	Csb	2017 – 2021	Metzger et al. (2019)
	xBL ^{HLS}	DBF	39.0603, - 78.0716	983	Cfa	2017 – 2021	Metzger et al. (2019)
	xDL ^{HLS}	MF	32.5417, - 87.8039	1372	Cfa	2017 – 2021	Metzger et al. (2019)
	xHa ^{HLS}	DBF	42.5369, - 72.1727	1071	Dfb	2017 – 2021	Metzger et al. (2019)
	xJE ^{HLS}	ENF	31.1948, - 84.4686	1307	Cfa	2017-- 2021	Metzger et al. (2019)
WET	Myb ^{P,HLS}	herbaceous, woody	38.0499, - 121.7650	338	Csa	2010 – 2021	Arias-Ortiz et al. (2021)
	TW1 ^{P,HLS}	herbaceous	38.1074, - 121.6469	421	Csa	2012 – 2020	Baldocchi et al. (2020)
	TW4 ^{P,HLS}	herbaceous	38.1027, - 121.6413	421	Csa	2013 – 2021	Eichelmann et al. (2018)

341 P: Annual precipitation (mm)

342 Bsk: Steppe: warm winter; Cfa: Humid Subtropical: mild with no dry season, hot summer; Csa:
343 Mediterranean: mild with dry, hot summer; Csb: Mediterranean: mild with dry, warm summer; Dfa: Humid
344 Continental: humid with severe winter, no dry season, hot summer; Dfb: Warm Summer Continental:
345 significant precipitation in all seasons.

346 CRO: cropland; GRA: Grassland; WSA: Woody savanna; OSH: Open shrubland; FOR: Forest; WET:
347 Wetland; ENF: Evergreen needleleaf forest; DBF: Deciduous broadleaf forest; MF: Mixed forest

348 **2.3. Measurements: Radiation and Energy Flux Density and Biophysical**
349 **Conditions**

350 The incident and reflected PAR (Q^i and Q^r) measurements were made with upward and
351 downward facing quantum sensors (Kipp & Zonen, PAR-Lite or PQS1) at each tower.

352 The shortwave radiation components were measured by pyranometers, one facing upward
353 for measuring the incident component (R_g^i) and the other looking downward for
354 measuring the reflected component (R_g^r) (Kipp & Zonen, CNR1 or Hukseflux NR01). A
355 suite of meteorological variables was measured in conjunction with the mass and energy
356 flux measurements. Air temperature and relative humidity were measured once every 10
357 seconds (0.1Hz) with Vaisala HMP45 RH/Temp sensors, with fan-aspirated solar shields
358 to represent ambient air and prevent solar heating. These data were then stored as a 30-
359 min average.

360 Fluxes were calculated from high-frequency (20 Hz) continuous recordings of
361 temperature, water vapor, and CO₂ concentrations, along with three-dimensional
362 measurements of wind velocities using infrared gas analyzers and a 3-D sonic
363 anemometer mounted on a scaffold or a tower structure at each site. High-frequency data
364 were integrated to 30-min intervals, and half-hourly fluxes were calculated from the
365 covariance between fluctuations in the vertical wind velocity and concentrations of
366 greenhouse gasses and energy. Common across sites are flux corrections and quality
367 control, which include high-frequency data despiking, 2-D coordinate rotations, sensor
368 separation distance, density corrections, and site-specific friction velocity (u^*) filtering
369 (Leuning, 2007, Wang et al., 2023).

370 Net CO₂ exchange was partitioned into ecosystem respiration and gross photosynthesis
371 (GPP) (symbolized as A_G in figures and equations) by training an Artificial Neural
372 Network on nighttime CO₂ fluxes (Biomet lab sites) or by applying partitioning
373 algorithms based on the short-term temperature sensitivity of respiration and nighttime
374 CO₂ fluxes to extrapolate respiration from nighttime to daytime periods and thus predict
375 ecosystem respiration at all times (Reichstein et al. 2005). Regardless of the partitioning
376 method, A_G was estimated as the sum of measured net CO₂ exchange and estimated
377 ecosystem respiration. A_G and surface energy balance fluxes measured between 10:00 to
378 14:00 h were averaged from their 30-min values to support the analysis of $NDVI_{bb}$ and
379 NIR_{vbb} .

380 Continuous measurements of leaf area index (LAI) were available from the University of
381 California, Berkeley, Biometeorology lab for the Tonzi Ranch site. Three identical
382 consumer grade point-and-shoot digital cameras (PowerShot A570IS, Canon, Japan)
383 were used to quantify LAI continuously and details are available in Ryu et al. (2012). The
384 cameras were leveled at 1 m height with the lens pointed towards the zenith. They were
385 approximately 50 m apart and set to a maximum wide angle (focal length of 5.8 mm),
386 automatic exposure, aperture priority mode and minimum aperture (F/2.6). These settings
387 yielded a view zenith angle from 0 to 32° diagonally. The Canon Hack Development Kit
388 (CHDK) (CHDK Project, <http://chdk.wikia.com>) was installed on the flash memory cards
389 of the cameras to extend their capabilities, including digital repeat photography through a
390 simple script written in uBasic (Sonnentag et al., 2012). The cameras were turned on and
391 off with data loggers (CR200, CR10X, Campbell Scientific Inc., USA).

392 **2.4. Remote sensing data**

393 Three different satellite datasets, namely Planet Fusion and Harmonized Landsat
394 Sentinel-2 (HLS) are used with spatial resolution varying between 3 and 30 m,
395 respectively. Planet Fusion data was available from 2017 to 2021 as a daily continuous
396 time series and HLS data was available from 01/2014 (Landsat) and 01/2016 (Sentinel-2)
397 as a daily discrete time series at 3 - 5 days interval. Although HLS data is continuously
398 generated, the Ameriflux database was updated until 12/2021 at the time of start of this
399 analysis, and therefore the present analysis is restricted up to year 2021.

400 Planet Labs' Planet Fusion data set comes from a constellation of more than 100
401 CubeSats in low earth orbit. This provides high resolution (3m x 3m pixels) and high
402 frequency revisits (<1day) but adds the complications of integrating data from many
403 sensors, cross-calibration, and atmospheric contamination. The Planet Fusion processing
404 combines this high resolution CubeSat data with MODIS/VIIRS, Landsat-8, and
405 Sentinel-2 imagery to create a daily, gap filled product that is radiometrically accurate,
406 and free of clouds and shadows. The technical specification can be found in
407 https://assets.planet.com/docs/Planet_fusion_specification_March_2021.pdf. The Planet
408 Fusion data was only available for seven UC Berkeley Biomet lab sites (**Table 1**) (US-
409 Bi1, US-Bi2, US-Var, US-Ton, US-Myb, US-TW1, US-TW4). Therefore, comparison
410 and evaluation of VI_{bb} and $NIR_{V_{bb}}$ at 3 m spatial resolution was restricted to seven sites.
411 For analyzing and comparing $NDVI_{bb}$ with Planet Fusion at 3 m spatial resolution, we
412 extracted the radiation & associated meteorological variables, surface energy balance and
413 carbon fluxes, and ancillary hydrological variables (soil moisture and precipitation) of

414 seven Biomet sites of California corresponding to the Planet Fusion data availability
415 period.

416 HLS data was available across all the 25 sites. The HLS products take input data from the
417 joint National Aeronautics and Space Administration-United States Geological Survey
418 (NASA-USGS) Landsat-8 and Landsat-9 (L8/9, hereafter) and the European Space
419 Agency (ESA) Sentinel-2A and Sentinel-2B (S2, hereafter) satellites to generate a
420 harmonized, analysis-ready surface reflectance data product with observations every two
421 to three days (<https://www.earthdata.nasa.gov/esds/harmonized-landsat-sentinel-2>)
422 (Claverie et al., 2018). HLS data for all the sites were acquired for the central pixel of the
423 tower sites through NASA AppEEARS (The Application for Extracting and Exploring
424 Analysis Ready Samples, <https://appears.earthdatacloud.nasa.gov/>). AppEEARS enables
425 users to acquire datasets using coordinate, temporal, and band/layer information. For
426 analyzing and comparing $NDVI_{bb}$ with HLS data at 30 m spatial resolution, we extracted
427 the radiation & associated meteorological variables, surface energy balance and carbon
428 fluxes, and ancillary hydrological variables (soil moisture and precipitation) of all the 25
429 sites corresponding to the data availability time-period of L8/9 (01/2014 – 12/2021) and
430 S2 (01/2016 – 12/2021), respectively.

431 While we computed NDVI and NIR_v from red and near-infrared surface reflectance as
432 described in Baldocchi et al. (2020) and Badgley et al. (2019), we compared $NDVI_{bb}$ and
433 $NIR_{v_{bb}}$ against daily NDVI and NIR_v from Planet Fusion and HLS (both L8/9 and S2)
434 covering the time frame of both the datasets. The entire analysis is performed across 25
435 flux tower sites of Ameriflux that covers a broad spectrum of ecosystems and energy-
436 water availability limits. For Planet Fusion, we used 5x5 pixel average values of

437 calculated NDVI and NIR_v for all daily scenes around each flux tower sites (seven sites)
438 of Biomet lab. For HLS, we acquired data over the central pixel of the flux tower location
439 and conducted the subsequent evaluation.

440 **2.5. Evaluation method**

441 To evaluate NDVI_{bb} and NIR_{vbb} with respect to the satellite vegetation indices across
442 different ecosystems and climate (SQ1), we used coefficient of determination (R^2), bias,
443 root mean squared difference (RMSD), normalized root mean squared difference
444 (nRMSD, in percent), and systematic root mean squared difference (sRMSD, in percent)
445 as statistical metrics (**Fig. 3 – 8**; section 3.1). To assess the efficacy of NDVI_{bb} in
446 capturing the phenological changes and vegetation function (SQ2), we computed the
447 mean seasonal variation of NDVI_{bb} and satellite NDVI in terms of the daily mean values
448 normalized by their annual mean (Baldocchi et al., 2021) and examined their responses to
449 green chromatic coordinate (GCC) and evaporative fraction (ratio of latent heat flux and
450 net available energy) (F_E). While GCC is used as a phenological metric to assess the
451 responses of NDVIs with the progression from low vegetation cover (or senescent
452 vegetation) until the peak vegetation, F_E indicates the biophysical response of vegetation
453 at different developmental stages (**Fig. 9 – 10**; section 3.2). To understand the
454 explanatory potential of NDVI_{bb} and NIR_{vbb} to the GPP variability (SQ3), we followed a
455 two-step procedure. We first verified the intraseasonal variability (coefficient of
456 variation, CV) in NDVI_{bb} and NIR_{vbb} by comparing them with the intraseasonal
457 variability of GPP and corresponding precipitation variability for the growing season.
458 Growing seasons includes all the days of spring, summer, and early autumn (i.e., periods
459 April to middle October). We subsequently used both these indices in conjunction with

460 EC GPP measurements to examine their relationship for range of energy-water
461 availability limits (**Fig. 11 – 12**; section 3.3). To understand the effects of soil
462 background variations and phenological progression on $NDVI_{bb}$ retrieval and its
463 consequent impact on $NDVI_{bb}$ -NDVI difference (SQ4), we also adopted a two-step
464 analysis. At the first step, we examined $NDVI_{bb}$ -NDVI difference (δ_{NDVI}) with respect to
465 GCC from senescent vegetation or bare soil to the peak vegetation stage. In this analysis,
466 we used Soil Adjusted and Atmospherically Resistant Vegetation Index (SARVI) (Qi et
467 al., 1994; Kaufman et al., 1992) as a third variable to simultaneously understand the
468 impacts of soil background variations on δ_{NDVI} during different phenological stages (**Fig.**
469 **13**; section 3.4). At the second step, we analyzed the effects of individual radiation
470 components on the performance of broadband hemispherical reflectances with respect to
471 satellite directional reflectances (**Fig. 14 – 16**; section 3.4) across ecosystems. Same
472 analysis is also performed for a range of energy-water availability limits and described in
473 Appendix A4 (**Fig. A3 – A5**).

474 **3. Results and discussion**

475 To answer the four science questions, we organized the results and discussion into four
476 sub-sections (3.1 to 3.4). The sequence of results and the corresponding figure numbers
477 are in the order of the following progression.

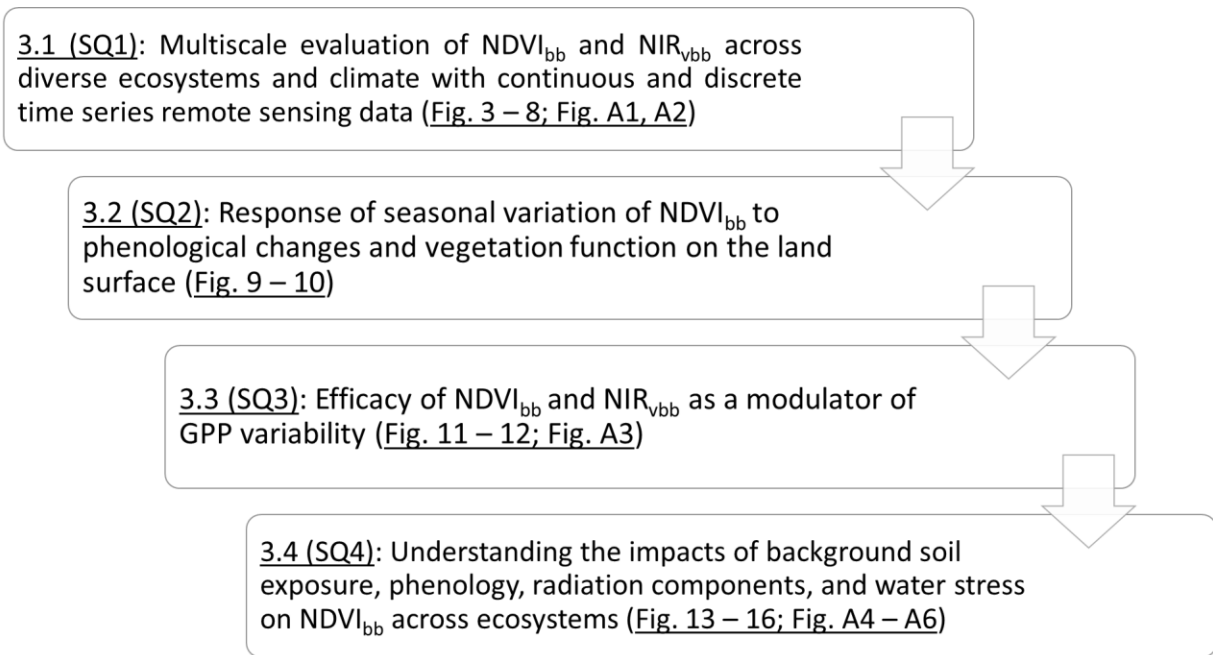


Figure 2: An illustrative diagram showing the sequence of results corresponding to the science questions (SQs) and the respective figure numbers associated with the description of results falling under individual science question.

478 **3.1. Multiscale evaluation of broadband NDVI and NIRv across different**
 479 **ecosystems and climate (SQ1)**

480 *Planet Fusion evaluation:* The scatterplots of NDVI_{bb} versus Planet Fusion NDVI (**Fig.**

481 **3-6**) revealed a robust, stable, and linear relationship at all the four different ecosystems

482 across seven EC flux tower sites of UC Berkeley Biomet lab in California. When all the

483 data points of corn and alfalfa were combined, the goodness of fit of linear regression

484 revealed NDVI_{bb} to explain 86% of the variation of Planet Fusion NDVI ($R^2 = 0.86$) at

485 the croplands (CRO) with bias, RMSD and sRMSD of 0.02, 0.08 and 52% for a wide

486 range of available energy-water limit (represented through evaporative fraction, F_E). In

487 CRO, the unexplained variation in NDVI_{bb} at a given NDVI was larger for $NDVI > 0.7$,

488 which also corresponded to high water and available energy limits (**Fig. 3a**). Besides,

489 some unexplained variation in NDVI_{bb} at a given NDVI was also evident in CRO for low

490 NDVI (NDVI: 0.35 – 0.40) and low F_E (F_E : 0.3 – 0.4), which ultimately led to relatively
 491 less dense scatters around low NDVI region.

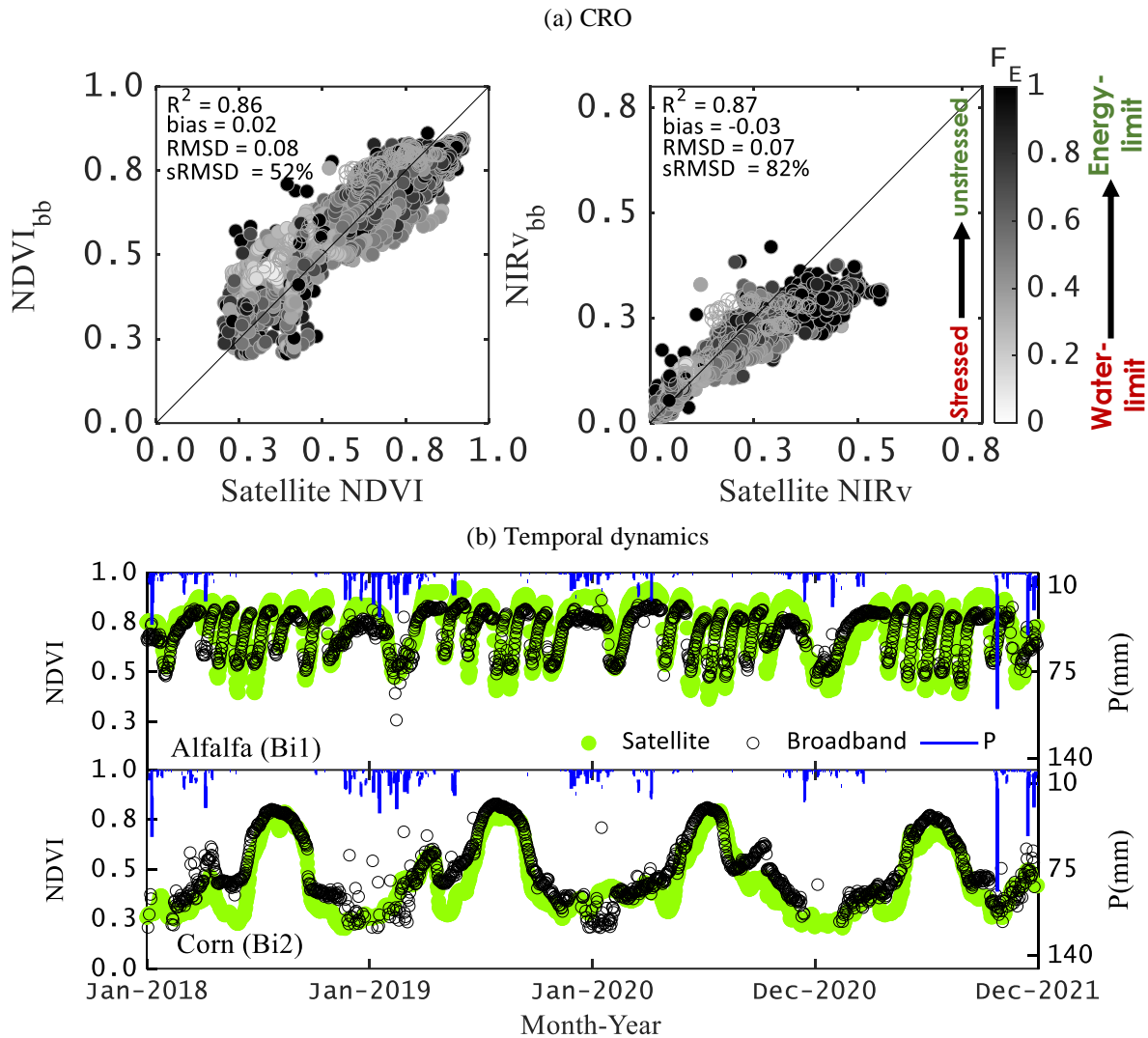


Figure 3: (a) Plots of $NDVI_{bb}$ and $NIRv_{bb}$ versus Planet Fusion $NDVI$ and $NIRv$ (3 m spatial resolution) in the Californian cropland ecosystems for $NDVI > 0.25$ for a range of evaporative fraction (F_E) representing stressed to unstressed conditions. F_E is an indicator of water availability and denotes the ratio of evaporation (latent heat flux) to equilibrium evaporation. (b) Temporal dynamics of $NDVI_{bb}$ (black dots) and $NDVI$ (green dots) along with daily precipitation (P) (blue stairs) shows close correspondence between them in Alfalfa [Bi1] and Corn [Bi2].

492 Same analysis by combining data of grassland (GRA) and woody savanna (WSA) sites
 493 revealed low range of NDVI in both the datasets with low mean 0.44 – 0.48 and median
 494 0.43 – 0.50 as compared to CRO (mean 0.57 – 0.61 and median 0.58 – 0.63). Due to the

495 strong water limitations, these ecosystems very rarely form a closed canopy cover,
 496 ultimately leading to low mean NDVI. The goodness of fit of linear regression revealed
 497 $NDVI_{bb}$ to explain 47% of the variation of Planet Fusion NDVI ($R^2 = 0.47$) (**Fig. 4a**),
 498 which is substantially lower as compared to CRO. Consequently, bias and RMSD was
 499 also higher (0.04 and 0.11) than the croplands (**Fig. 4a**). Some exceptionally high
 500 magnitude of $NDVI_{bb}$ was evident at low satellite NDVI (0.3 – 0.5), which corresponded
 501 to unstressed conditions in GRA-WSA (**Fig. 4a**).

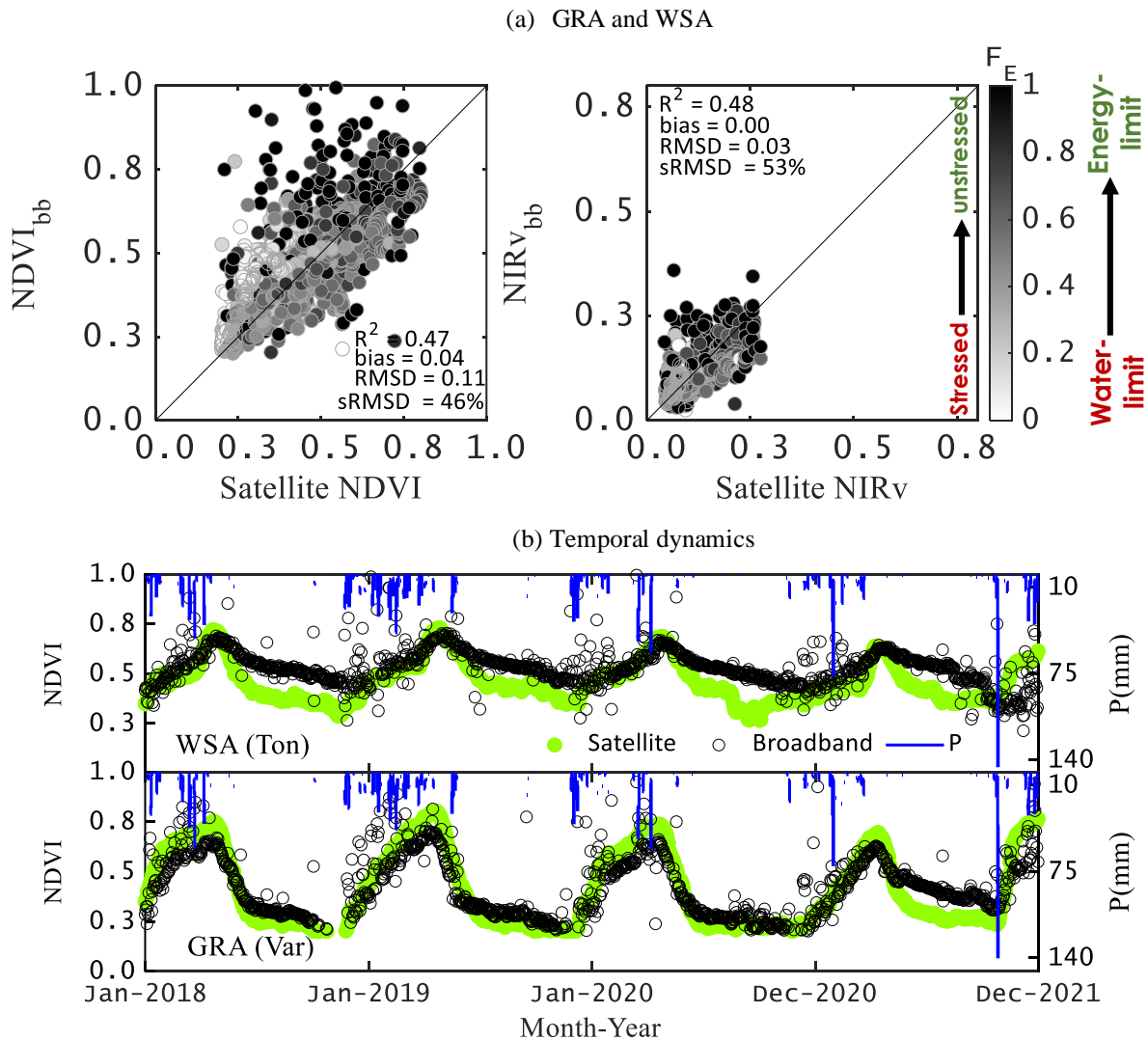
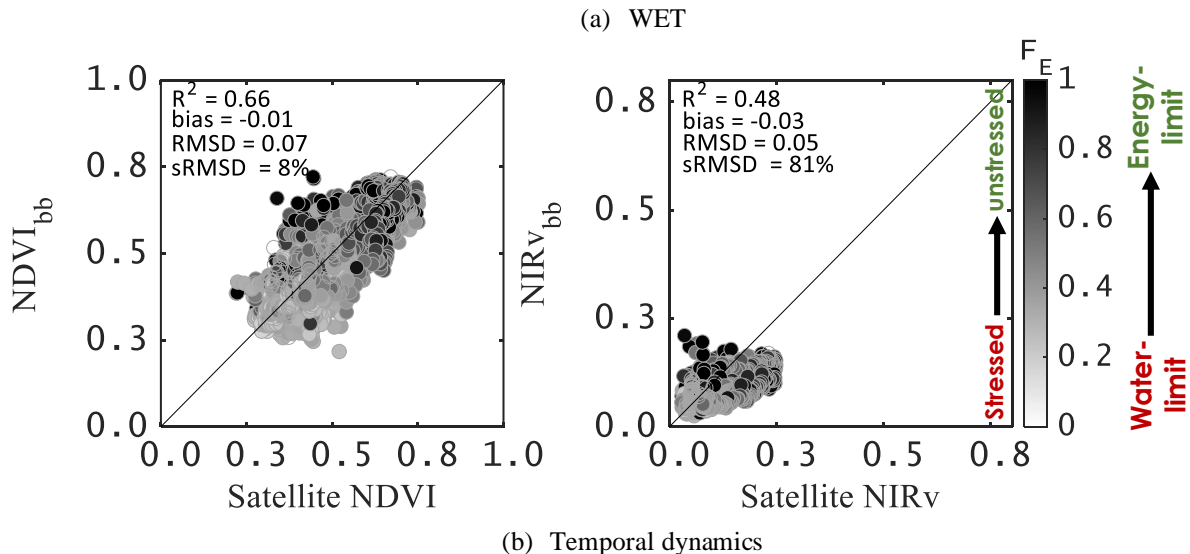


Figure 4: (a) Plots of $NDVI_{bb}$ and $NIRv_{bb}$ versus Planet Fusion NDVI and $NIRv$ (3 m spatial resolution) in water-limited Californian grassland (GRA) and woody savanna (WSA) ecosystems for $NDVI > 0.25$ for a range of

evaporative fraction (F_E) representing stressed to unstressed conditions. F_E is an indicator of water availability and denotes the ratio of evaporation (latent heat flux) to equilibrium evaporation. (b) Temporal dynamics of $NDVI_{bb}$ (black dots) and $NDVI$ (green dots) along with daily precipitation (P) (blue stairs) shows close correspondence between them especially in the grassland [Var] and partly in woody savanna [Ton].

502 In comparison to CRO and GRA-WSA, scatterplot of $NDVI_{bb}$ versus Planet Fusion
 503 $NDVI$ at the wetland (WET) sites revealed relatively lesser spread (**Fig. 5a**) with a mean
 504 and median of 0.51 – 0.52, respectively. $NDVI_{bb}$ explains 66% of the variations in Planet
 505 Fusion $NDVI$ ($R^2 = 0.66$), with a relatively low bias (-0.01), RMSD (0.06) and sRMSD
 506 (17%) for the entire range of available energy and limits. The seasonal dynamics of daily
 507 $NDVI_{bb}$ at the representative sites (**Fig. 3b, 4b, 5b**) revealed a close resemblance with
 508 satellite $NDVI$ for the respective tower pixel at almost all the sites, except at the WSA
 509 (Tonzi ranch) (**Fig. 4b**). Despite the rising and falling behavior of $NDVI_{bb}$ was well
 510 coordinated with the satellite $NDVI$ at WSA, substantial differences between $NDVI_{bb}$ and
 511 Planet Fusion $NDVI$ is also evident as the magnitude $NDVI$ declined with the
 512 progression of summer (**Fig. 4b**). This implies that $NDVI_{bb}$ could not efficiently capture
 513 the very low $NDVI$ magnitude of the open canopy architecture at woody savanna during
 514 the water stressed summer months.



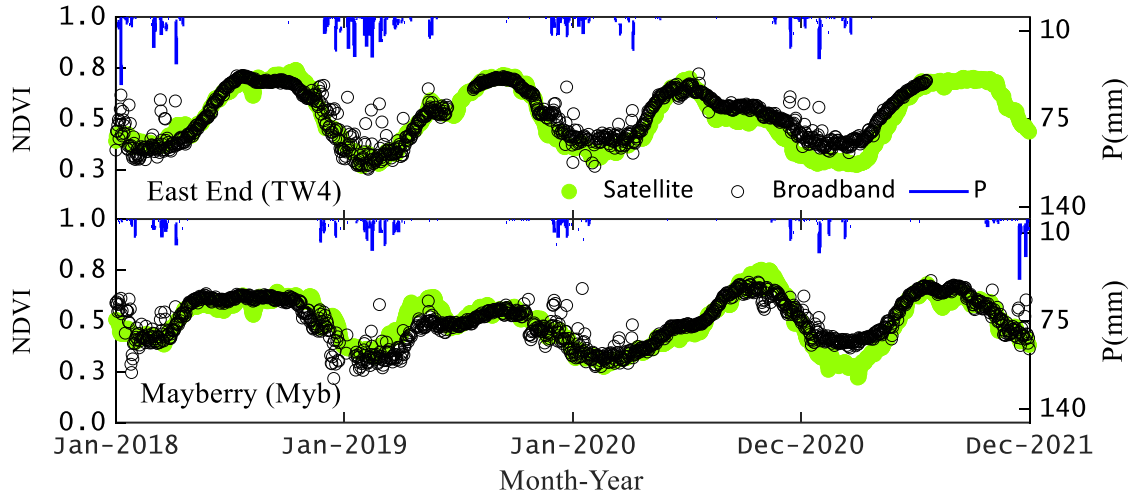


Figure 5: (a) Plots of $NDVI_{bb}$ and $NIR_{v_{bb}}$ versus Planet Fusion $NDVI$ and NIR_v (3 m spatial resolution) in Californian wetland ecosystems (non-tidal) for $NDVI > 0.25$ for a range of evaporative fraction (F_E) representing stressed to unstressed conditions. F_E is an indicator of water availability and denotes the ratio of evaporation (latent heat flux) to equilibrium evaporation. (b) Temporal dynamics of $NDVI_{bb}$ (black dots) and $NDVI$ (green dots) along with daily precipitation (P) (blue stairs) shows close correspondence between them in both East End and Mayberry.

515 Comparison of $NIR_{v_{bb}}$ versus satellite NIR_v revealed $NIR_{v_{bb}}$ to explain 48 – 87%

516 variation in satellite NIR_v (R^2 varying from 0.48 – 0.87; mean R^2 : 0.79) (**Fig. 3a, 4a, 5a,**

517 **6**), RMSD (varying from 0.02 – 0.09; mean RMSD: 0.05), and sRMSD (varying from 35

518 – 92%; mean sRMSD: 70%) for a broad range of F_E . A distinct saturation in $NIR_{v_{bb}}$

519 signal around $NIR_{v_{bb}} > 0.3$ and asymptotic behavior in $NIR_{v_{bb}}$ was evident in CRO and

520 WET with increasing satellite NIR_v . While this saturation of $NIR_{v_{bb}}$ corresponded to

521 high F_E (> 0.7) (**Fig. 3a**) at the CRO sites, the saturation of $NIR_{v_{bb}}$ corresponded to both

522 high and low F_E at the WET sites (**Fig. 5a**). Nevertheless, by combining data of all the

523 sites, the overall performance of $NDVI_{bb}$ and $NIR_{v_{bb}}$ appeared to be stable (**Fig. 6**). The

524 range of $NDVI_{bb}$ and $NIR_{v_{bb}}$ obtained from the net fluxes of PAR and R_G are comparable

525 with the magnitude and dynamics of satellite $NDVI$ and NIR_v within and across different

526 ecosystems falling under the same climatic setting.

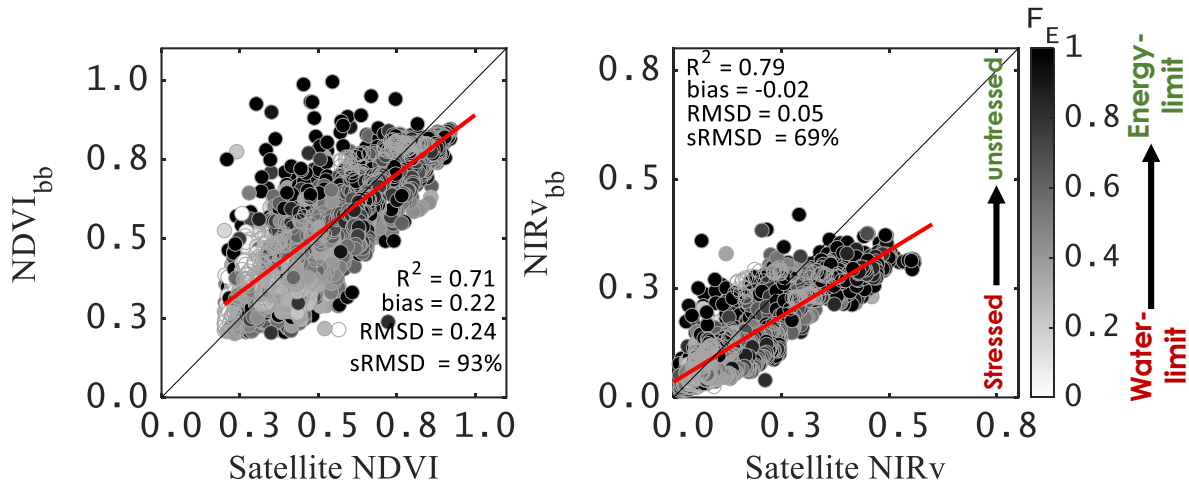


Figure 6: Pooled evaluation plots of $NDVI_{bb}$ and $NIRV_{bb}$ versus Planet Fusion $NDVI$ and $NIRV$ (3 m spatial resolution) by combining all the seven sites of Californian ecosystems for a range of evaporative fraction (F_E) representing stressed to unstressed conditions. F_E is an indicator of water availability and denotes the ratio of evaporation (latent heat flux) to equilibrium evaporation.

527 Landsat & Sentinel evaluation: Analysis of $NDVI_{bb}$ and $NIRV_{bb}$ derived from Q^i and Q^r

528 measurements in all 25 sites with respect to HLS $NDVI$ and $NIRV$ provided another

529 assessment of $NDVI_{bb}$ and $NIRV_{bb}$ in energy- and water-limited environments across

530 diverse ecosystems. Our analysis revealed that $NDVI_{bb}$ consistently captured the

531 variations in $NDVI$ when compared with both L8/9 and S2 sensors (**Fig. 7; Fig. 8**). While

532 the mean and median $NDVI_{bb}$ at the energy-limited ecosystems was found to be 0.56

533 (0.52 for HLS) and 0.59 (0.53 for HLS), these metrics were 0.41 (0.34 for HLS) and 0.38

534 (0.31 for HLS) in the water-limited ecosystems. Four distinct features are notable from

535 this analysis. Firstly, the scatterplots of $NDVI_{bb}$ versus satellite $NDVI$ showed significant

536 spread in the energy-limited ecosystems for both L8/9 and S2 across the entire range of

537 F_E (**Fig. 7a, c**). Secondly, the statistical metrics of $NDVI_{bb}$ with respect to $NDVI$ were

538 very similar across the sensors, with higher coefficient of determination for S2 ($R^2 =$

539 0.59) as compared to L8/9 ($R^2 = 0.43$) and lower systematic difference in L8/9 as

540 compared to S2 (**Fig. 7a, c**).

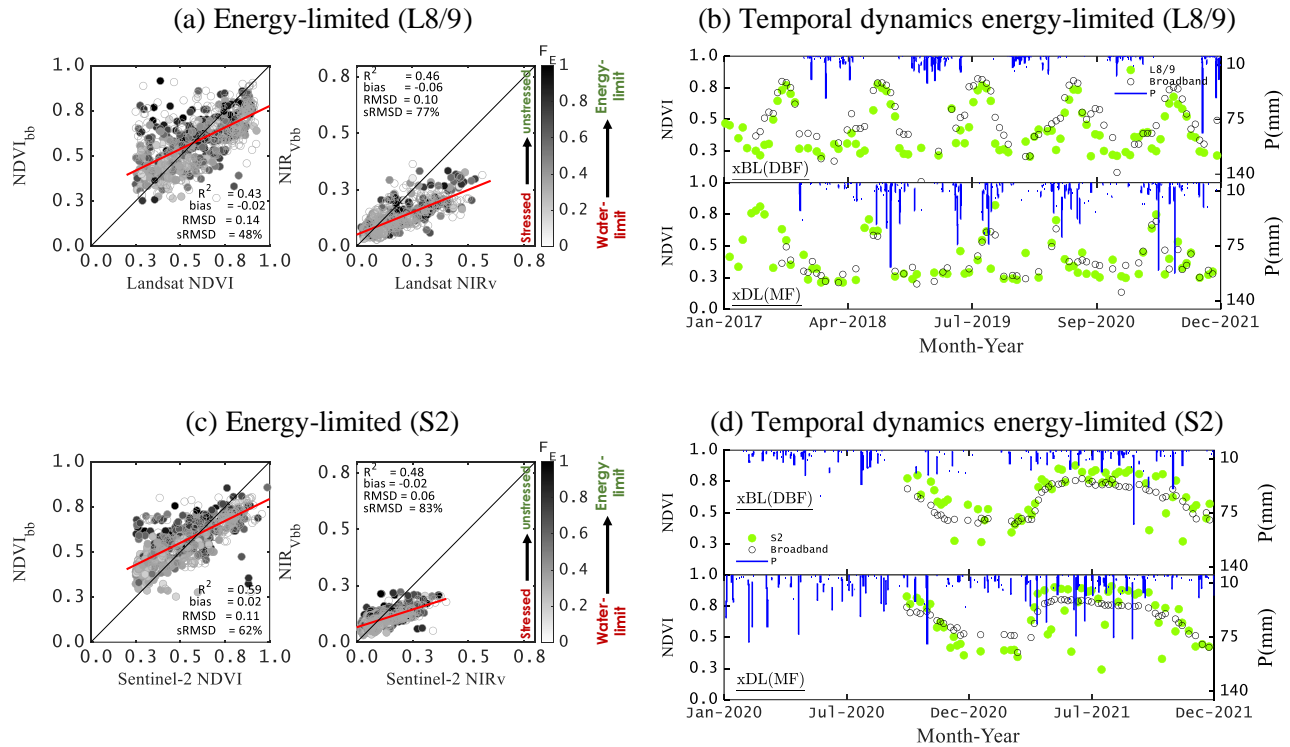


Figure 7: (a, c) Plots of NDVI_{bb} and NIR_V_{bb} versus Landsat and Sentinel-2 NDVI and NIR_v (30 m spatial resolution) in energy-limited ecosystems of Biomet and NEON sites. Color shading is done by evaporative fraction (F_E) showing stressed to unstressed conditions which corresponds to water and energy limits within the energy-limited environment. (b, d) Illustrative examples of temporal dynamics of NDVI_{bb} (black dots) and NDVI (green dots) along with daily precipitation (P) (blue stairs) showing close correspondence in the seasonal and interannual variability of NDVI_{bb} and NDVI at the NEON sites Blandly Experimental Farm (xBL) and Dead Lake (xDL).

541 Thirdly, the agreement between NDVI_{bb} versus satellite NDVI and NIR_V_{bb} versus
542 satellite NIR_v was much stronger (with less systematic difference) in the water-limited
543 ecosystems ($R^2 = 0.59 - 0.66$ and $0.69 - 0.74$; sRMSD: 33 - 42% and 49 - 54%) as
544 compared to energy-limited ecosystems ($R^2 = 0.44 - 0.49$ and $0.41 - 0.46$; sRMSD: 48 -
545 57% and 76 - 77%). Fourthly, a marked asymptotic pattern and saturation in NIR_V_{bb} was
546 evident with increasing satellite NIR_v (>0.3) (**Fig. 7a, c; Fig. 8a, c**) in both the climatic
547 limits. This resulted in large differences and high RMSD between NIR_V_{bb} versus satellite
548 NIR_v across the entire range of F_E . The seasonal dynamics of NDVI_{bb} at the
549 representative cropland and grassland sites in the water-limited ecosystems revealed a
550 close resemblance with satellite NDVI for the respective tower pixel at each site (**Fig. 8b**,

551 d). A detailed description of ecosystem wise analysis by combining data of both L8/9 and
 552 S2 is given in **Appendix A1, Fig. A1, Table A1** (for L8/9) and **Table A2** (for S2),
 553 respectively.

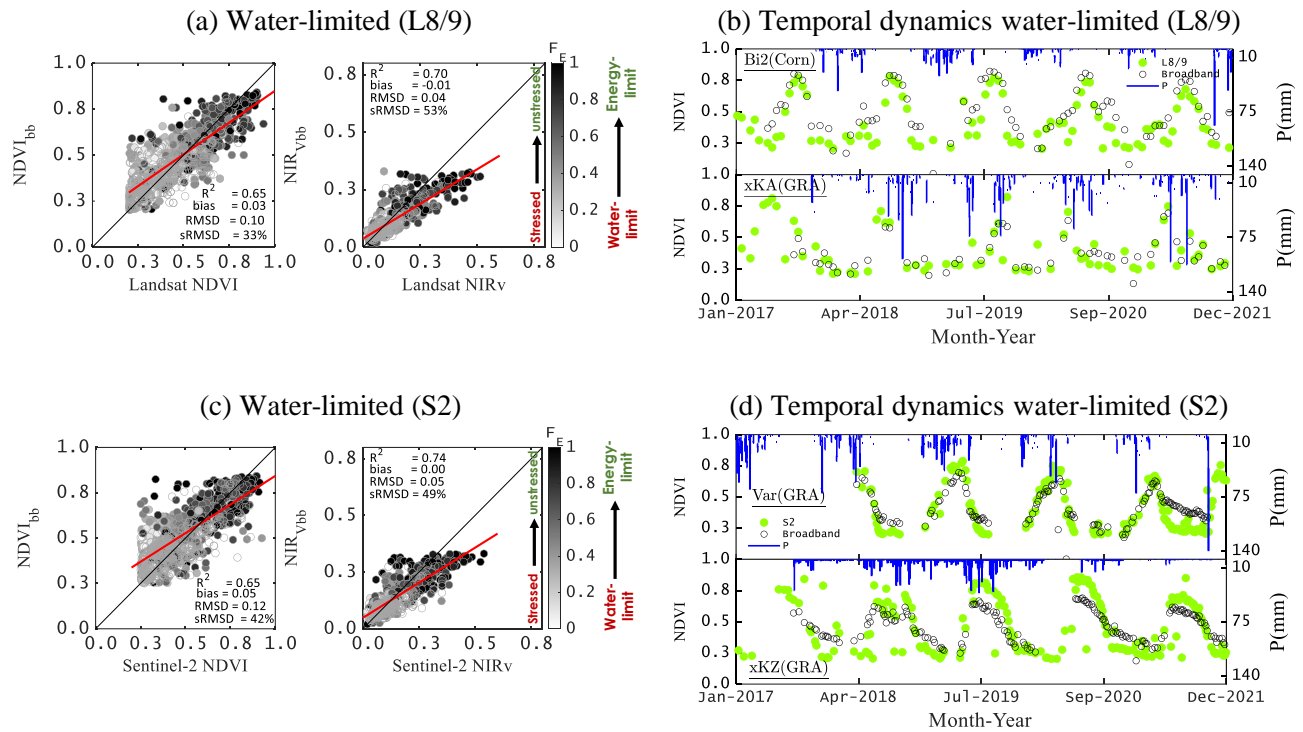


Figure 8: (a, c) Plots of NDVI_{bb} and NIR_{vbb} versus Landsat and Sentinel-2 NDVI and NIR_v (30 m spatial resolution) in water-limited ecosystems of Biomet and NEON sites. Color shading is done by evaporative fraction (F_E) showing stressed to unstressed conditions which corresponds to water and energy limits within the water-limited environment. (b, d) Illustrative examples of temporal dynamics of NDVI_{bb} (black dots) and NDVI (green dots) along with daily precipitation (P) (blue stairs) showing close correspondence in the seasonal and interannual variability of NDVI_{bb} and NDVI over Bouldin corn (Bi2), Vaira ranch (Var) and two grasslands sites of NEON Konza Prairie Biological Station (xKA and xKZ).

554 To understand the explanatory potential of NDVI_{bb} in tracking the variation in LAI, we
 555 also analyzed NDVI_{bb} with respect to continuous LAI observations at the Tonzi ranch site
 556 (detailed explanations are in **Appendix A2; Fig. A2**).

557 It is further important to emphasize that in the comparisons between broadband versus
 558 satellite NDVI, we do not expect to see an ideal 1:1 relationship. Discrepancies between
 559 these two indices could arise, (i) due to the differences in bandwidths for the bands used

560 in satellite and broadband vegetation indices and (ii) due to comparing broadband
561 hemispherical reflectance derived through proximal sensing versus directional
562 narrowband reflectance from remote sensing. The effects of these two important aspects
563 are demonstrated and discussed in detail in section 3.4.

564 **3.2. Mean seasonal variability of broadband NDVI with phenology and** 565 **vegetation function (SQ2)**

566 This section examines the mean seasonal variability of NDVI_{bb} along with satellite
567 NDVI, a phenological metric namely Green Chromatic Coordinate (GCC), and their
568 response to water stress (evaporative fraction, F_E). (**Fig. 9, 10**). This analysis is based on
569 the continuous time series information of NDVI_{bb} , Planet Fusion NDVI, and UC
570 Berkeley Biomet lab flux tower datasets. The reasons to use Planet Fusion data are that
571 firstly they are finely resolved in time to allow for filtering day-to-day variability and
572 secondly, they span over a period of four years to allow investigating the mean seasonal
573 variations. **Figure 9** shows the synthesis of the mean seasonal variation of NDVI_{bb} ,
574 NDVI, GCC, and F_E for the two cropland ecosystems (alfalfa and corn) in California. The
575 seasonal variation is expressed in terms of the daily mean values normalized by their
576 annual mean following Baldocchi et al. (2021).

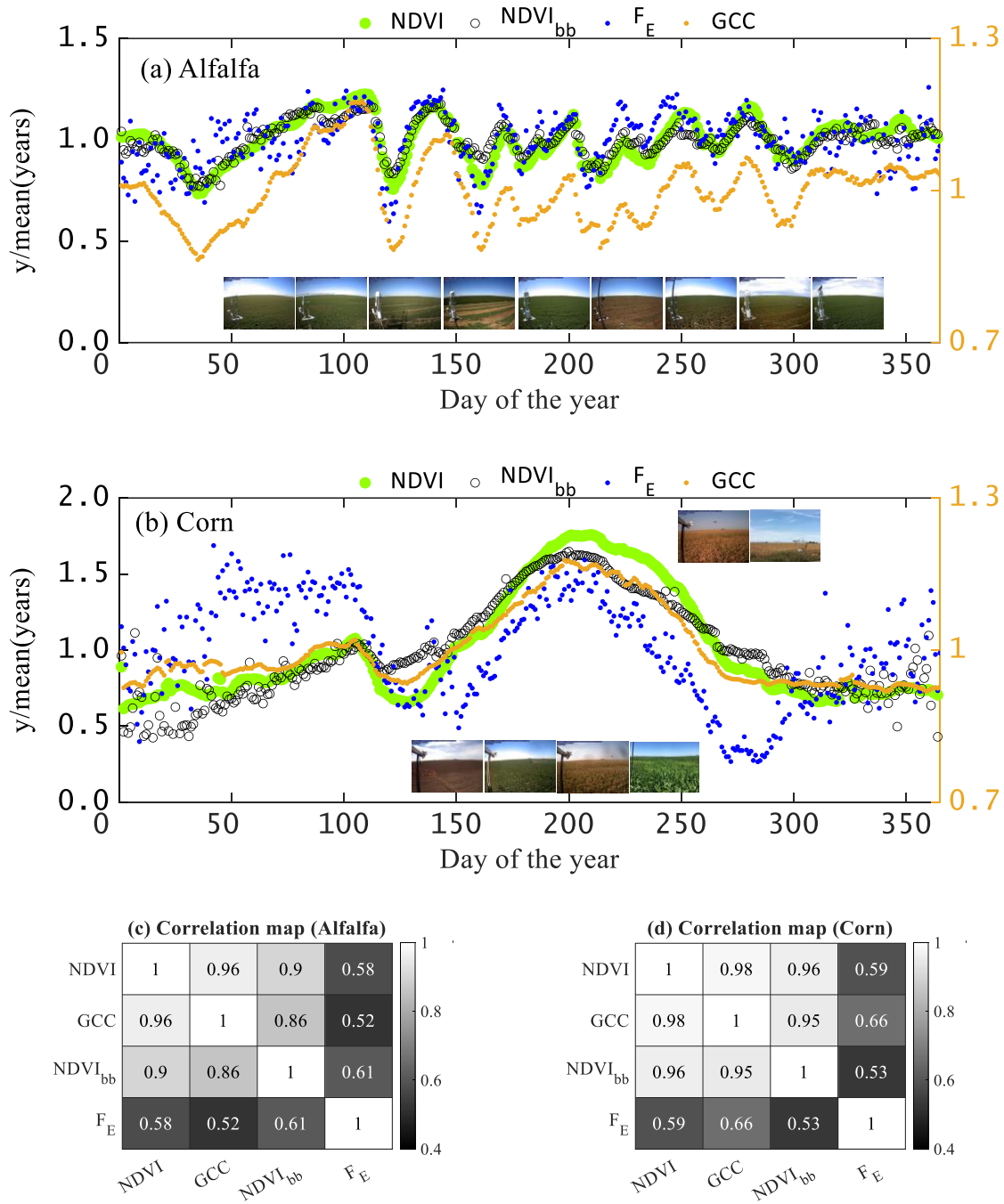
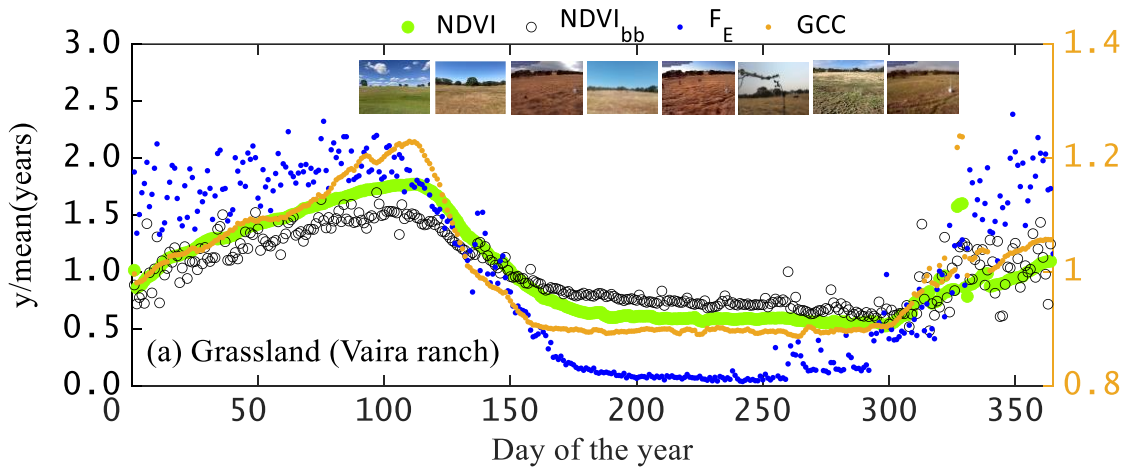


Figure 9: (a-b) Daily variation in NDVI_{bb}, Planet Fusion NDVI, Green Chromatic Coordinate (GCC) (secondary y-axis), and evaporative fraction (F_E) over agricultural ecosystems (alfalfa and corn) in California. Here we plot daily values, averaged over 4 years, normalized by the annual mean for that variable. (c-d) Correlation map showing the strength of seasonal relationship between individual variables. For corn, the correlation map is applicable for the growing season from March to September.

577 Two things became evident in this analysis. (i) The coordination of NDVI_{bb} and satellite

578 NDVI was found to be remarkably high with GCC at alfalfa throughout the entire year

579 (Fig. 9a, c). The coordination of the two NDVIs with GCC was also substantially strong
 580 in corn from the start of the growing season (day of the year 120), green-up phase (day of
 581 the year 150 – 180), peak growth phase (day of the year 180 – 250) and until the end of
 582 the growing season (day of the year 250 – 300) (Fig. 9b, d). (ii) The response of both the
 583 NDVIs and GCC is also highly correlated with F_E at the alfalfa site (Fig. 9a, c) ($r = 0.52$
 584 – 0.61), and their responses to F_E were also very robust in corn during the annual growth
 585 cycle that spans from day of the year 120 to 300 (Fig. 9d) ($r = 0.53 – 0.66$). This
 586 indicates substantial controls of water availability on the growth dynamics of both alfalfa
 587 and corn. Peak daily $NDVI_{bb}$ (and $NDVI$) and GCC coincided when F_E is greater than
 588 their mean annual values (Fig. 9b).



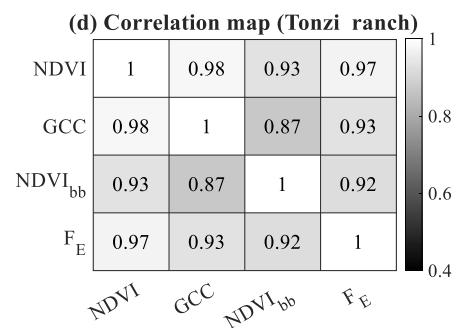
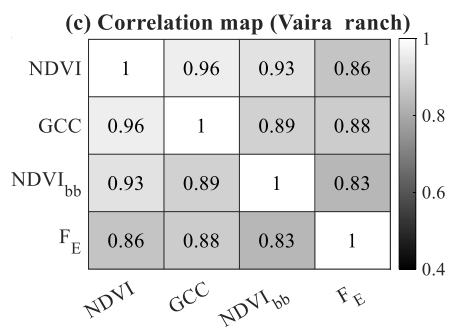
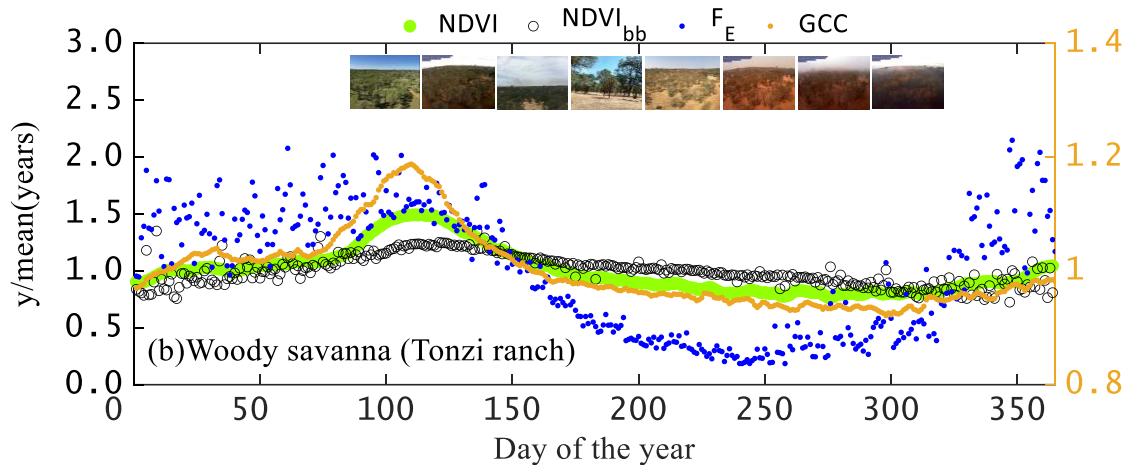


Figure 10: (a-b) Daily variation in NDVI_{bb}, Planet Fusion NDVI, Green Chromatic Coordinate (GCC) (secondary y-axis), and evaporative fraction (F_E) over grassland (Vaira ranch) and woody savanna (Tonzi ranch) ecosystems in California. Here we plot daily values, averaged over 4 years, normalized by the annual mean for that variable. (c-d) Correlation map showing the strength of seasonal relationship between individual variables during the growing season from March to October.

589 **Figure 10** shows the synthesis of the mean seasonal variation of NDVI_{bb}, NDVI, GCC,
 590 and F_E at GRA and WSA ecosystems in California. Here also, some distinctive behavior
 591 of NDVI_{bb}, NDVI and GCC and their response to water stress variations was noted.
 592 Firstly, the overall coordination of GCC and NDVIs with F_E was high in both the
 593 ecosystems ($r = 0.83 - 0.88$ and $r = 0.93 - 0.97$) and the coordination strength of NDVI
 594 versus GCC was equally high ($r = 0.83 - 0.96$ and $r = 0.87 - 0.98$). Secondly, the
 595 coordination strength between the two NDVIs versus water availability in GRA and
 596 WSA is substantially higher as compared to the croplands (**Fig. 10c, d**). In both the
 597 ecosystems, peak daily NDVI_{bb} (and NDVI) and GCC coincided when F_E is greater than

598 their mean annual values (**Fig. 10a, b**). The maximum NDVI and GCC was found in
599 early spring (day of year 120) during the unstressed conditions. After that, the two
600 NDVIs started declining with F_E and it reached the minimum during the middle of the
601 summer, between days 170 and 250. Interestingly, despite the declining pattern of GCC
602 was very similar to NDVI in WSA, it remained invariant in GRA between days 150 and
603 300. This is the period when the soil remains nearly bone dry due to prolong absence of
604 precipitation. The coalition of high soil water stress and atmospheric aridity (as defined
605 by vapor pressure deficit) in association with high net available energy triggers the
606 stomatal closure and consequently the photosynthetic activity is at the minimum level.
607 It is further important to emphasize that croplands receive subsurface irrigation at a depth
608 of 2 m and the irrigation frequency is very low. While corn receives a single irrigation
609 (around day of the year 230), alfalfa receives maximum 2 irrigation (around day of the
610 year 160 and 180; 240 and 260). The atmospheric humidity over cropland is higher (as
611 compared to GRA and WSA) due to being situated close to the delta shores of
612 Sacramento and due to moisture advection (Wang et al., 2023). All these factors lead to
613 an increased evaporative fraction, and vegetation seasonality responds significantly to F_E
614 dynamics. On the contrary, being situated at the valley, GRA and WSA sites face dual
615 challenge due to high soil and atmospheric water stress. The different NDVI, $NDVI_{bb}$ and
616 GCC profiles for these two ecosystems indicate that the vegetation seasonality in GRA
617 and WSA has a stronger coupling to the seasonality in water stress as compared to the
618 croplands, despite being situation in the same Mediterranean climate.

619 **3.3. Efficacy of NDVI_{bb} and NIR_{vbb} to explain GPP variability (SQ3)**

620 This analysis is carried out into two halves. In the first step, we examined whether the
621 intraseasonal variability (expressed as ‘coefficient of variation’, cv) of NDVI_{bb} and
622 NIR_{vbb} can explain the intraseasonal variability of GPP (symbolized as A_G). In the
623 second step, we tested the robustness and feasibility of using daily NDVI_{bb} and NIR_{vbb} as
624 a robust predictor of daily GPP. Here also, we used the continuous time series EC tower
625 GPP record of seven Biomet lab sites and the Planet Fusion data. The reasons to use
626 PLANET Fusion data is the same as mentioned in section 3.2.

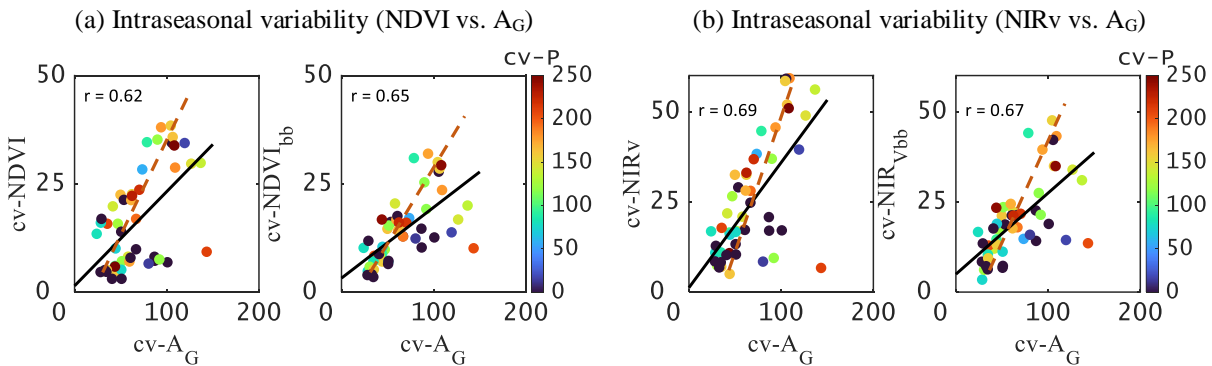
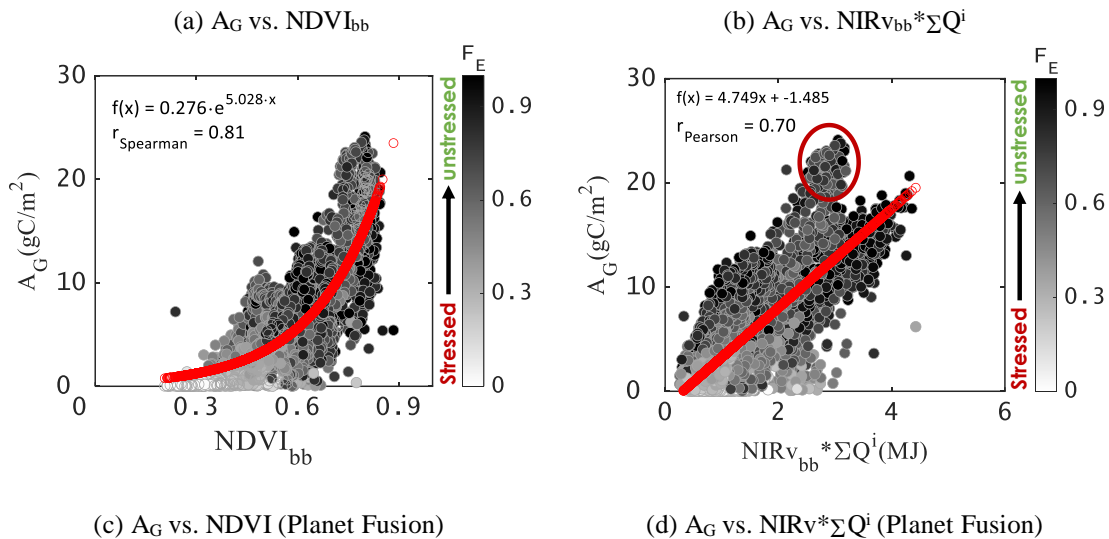
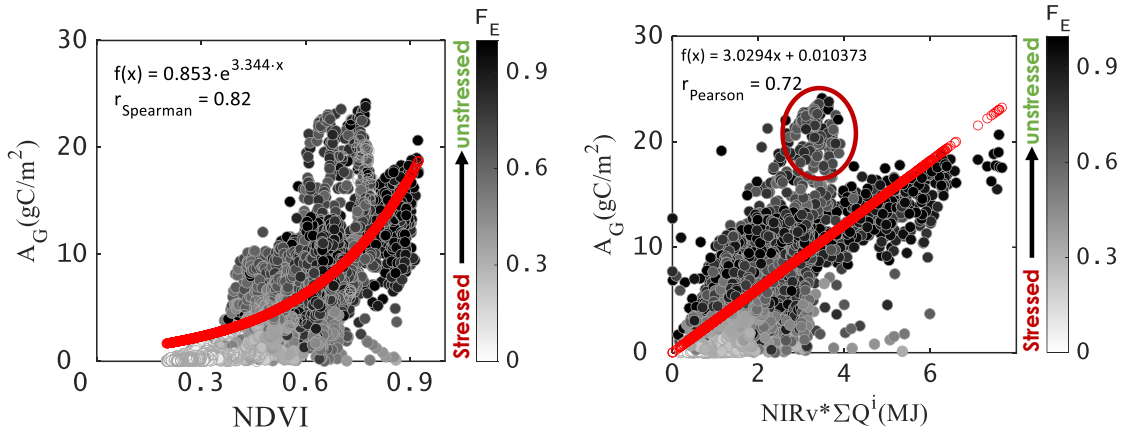


Figure 11: (a-b) Plots of intraseasonal variability (expressed as ‘coefficient of variation’, cv) in NDVI (NDVI_{bb}) and NIR_v (NIR_{vbb}) versus intraseasonal variability in gross photosynthesis (A_G) by combining all the site data in different ecosystems of California. Color shading is by precipitation (P) variability. This also shows the steeper slope of cv-NDVI (and cv-NDVI_{bb}) vs. cv-A_G and cv-NIR_v (and cv-NIR_{vbb}) vs. cv-A_G relationship with increasing precipitation variability.

627 Combining data of all the seven Biomet sites showed a relatively stronger relationship
628 between NIR_v (NIR_{vbb}) variability with GPP variability ($r = 0.67 - 0.69$) as compared to
629 NDVI (NDVI_{bb}) ($r = 0.62 - 0.65$) (**Fig. 11a, b**). The intraseasonal variability of the two
630 NDVIs versus GPP and two NIR_v versus GPP relationship was also found to be strongly
631 associated with the rainfall variability during the growing season in CRO, GRA, and
632 WSA (Appendix A3, **Fig. A3**). Previous studies also showed the tendency of the water-
633 limited ecosystems towards higher interannual variability in vegetation productivity
634 (Ritter et al., 2020). The high water use efficiency of cropland, grassland, and savanna

635 plays a major role. During the pluvial years or wet seasons, water infiltration into deep
 636 soil layers compensates for the preceding water deficit (Ritter et al., 2020), increasing the
 637 soil water content available for transpiration and biomass production for the following
 638 months. This leads to an increased productivity during high rainfall years relative to their
 639 reduction during the dry years (Ritter et al., 2020). Similar mechanism is also reflected in
 640 these two vegetation indices through efficient vegetation greening as a result of optimum
 641 vegetation productivity. This led to steeper slopes between the coefficient of variation of
 642 these two vegetation indices (both broadband and satellite) versus the coefficient of
 643 variation of GPP ($cv-A_G$) with increasing precipitation variability when all the data were
 644 pooled together (**Fig. 11**).





(e) A_G vs. $NDVI_{bb}^2 * \Sigma Q^i$

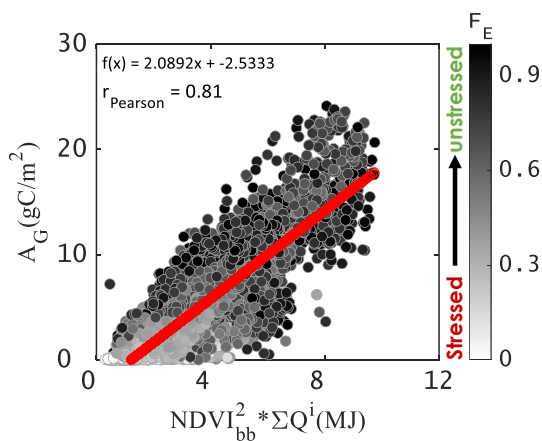


Figure 12: (a, b) Site-level relationships of A_G versus $NDVI_{bb}$ and A_G versus $NIRv_{bb} * \Sigma Q^i$ at seven eddy covariance towers of UC Berkeley Biomet sites that includes 2 crop sites (one C4, one C3), one grassland site, one woody savanna and 3 wetland sites. (c, d) Similar plot is shown by plotting A_G with Planet Fusion NDVI and $NIRv_{bb} * \Sigma Q^i$. Here ΣQ^i and A_G are the daily integrated Q^i (MJ) and A_G (gC/m^2) obtained by summing up the half-hourly observation. Data points inside the red circle showed saturation in A_G with increasing $NIRv_{bb} * \Sigma Q^i$. These data points belong to corn crop and could presumably be associated with the diffuse component of Q^i . (e) Scatter plot of A_G versus $NDVI_{bb}^2 * \Sigma Q^i$.

645 By combining the Planet Fusion and EC data of all the seven Biomet lab sites, we found a
 646 distinct exponential pattern between flux tower GPP and $NDVI_{bb}$ (**Fig. 12a**) with
 647 Spearman's correlation (r_{Spearman}) of 0.79. This further confirms that NDVI saturates at
 648 high biomass and this saturation is mainly attributed to the insensitivity of chlorophyll
 649 absorbing red light at 100% vegetation cover (Sellers et al., 1985; Kumar et al., 2001).
 650 Any addition of vegetation does not impact further changes since the amount of red light
 651 that can be absorbed by leaves reaches a peak, whereas NIR reflectance will increase
 652 because an addition of leaves results in multiple scattering (Tesfaye and Awoke, 2020;
 653 Kumar et al., 2001). The imbalance between red and high NIR reflectance results in a
 654 marginal change in the NDVI ratio and yields saturation at high biomass. Given the

655 product of NIR_v and incident PAR (Q^i) is considered as a proxy for GPP at different
656 spatial scales (Dechant et al., 2022), we further evaluated the relationships between GPP
657 versus the product of NIR_{v_{bb}} and daily integrated PAR ($NIR_{v_{bb}} * \Sigma Q^i$). We found a strong
658 and significant correlation between A_G versus $NIR_{v_{bb}} * \Sigma Q^i$ ($r = 0.68$) (**Fig. 12b**). Despite
659 substantial linearity between A_G versus $NIR_{v_{bb}} * \Sigma Q^i$ relationship, a small portion of data
660 points (inside red circle) showed saturation in A_G with increasing $NIR_{v_{bb}} * \Sigma Q^i$ (**Fig. 12b**).
661 These data points belong to corn and could presumably be associated with the diffuse
662 component of Q^i . The scatterplots of flux tower A_G versus Planet Fusion NDVI and the
663 product of $NIR_v * \Sigma Q^i$ also showed the same exponential and linear pattern and very
664 similar correlation ($r_{\text{Spearman}} = 0.81$ and $r_{\text{Pearson}} = 0.70$) (**Fig. 12c, d**). These results
665 corroborate with the findings of Pierrat et al. (2022), Gamon et al. (1995), and Liu et al.
666 (2021) which showed that NDVI is insensitive to maximum carbon uptake in evergreen
667 trees and reported the similar pattern of saturation as we found in **Fig. 12a**. However,
668 where canopy structure, PAR, and carbon uptake are in synchrony, NDVI was found to
669 be significantly correlated with gross photosynthesis (Gamon et al., 1995, Liu et al.,
670 2021). The $NIR_{v_{bb}} * \Sigma Q^i$ approach constituted a nonlinear stretch of $NDVI_{bb}$ by
671 multiplying $NDVI_{bb}$ with the NIR reflectance, thereby increasing the sensitivity of
672 $NIR_{v_{bb}} * \Sigma Q^i$ for high vegetation carbon uptake and green biomass. $NIR_{v_{bb}}$ implicitly
673 assumes a linear relationship between $NDVI_{bb}$ and fractional absorbed PAR, and this
674 fraction is 100% at maximum $NDVI_{bb}$. Therefore, multiplying $NIR_{v_{bb}}$ by ΣQ^i gives a
675 close estimate of absorbed PAR, and we see a good relationship with gross primary
676 productivity. Although the scatterplot of A_G versus Planet Fusion $NIR_v * \Sigma Q^i$ shows a
677 tendency to saturate at high A_G (**Fig. 12d**), overall $NIR_{v_{bb}} * \Sigma Q^i$ approach reflects much

678 better fidelity to capture the variability in carbon fluxes (Baldocchi et al., 2020; Dechant
679 et al., 2022).

680 The relationship between A_G versus $NIR_{vbb} * \Sigma Q^i$ has a physical basis and it is analogous
681 to the classic light use efficiency (LUE) approach of Monteith (1972), Gitelson and
682 Gamon (2015). According to Monteith (1972), Gitelson and Gamon (2015), $GPP =$
683 $LUE * \Sigma APAR$, where APAR is the absorbed PAR. From this analogy, NIRv seems to
684 carry the dual information of absorbed PAR and LUE. While high (low) GPP is the
685 consequence of high (low) absorbed PAR, NIRv is the consequence of multiple reflection
686 in the near infrared reflectance which increases with vegetation layer. Therefore,
687 $NIR_{vbb} * \Sigma Q^i$ has a clear upper and lower bound to explain the GPP variability for a wide
688 range of vegetation and radiation conditions. For example, high GPP during the peak
689 developmental phase of vegetation is due to high absorbed PAR, which apparently leads
690 to high NIRv. On the other hand, during the early growth phase and maturity, we see
691 increasing and declining GPP with increasing and decreasing absorbed PAR and NIRv.
692 Thus, $NIR_{vbb} * \Sigma Q^i$ is able to separate green from dead vegetation. This is the reason why
693 we found a remarkably good relationship when we plotted daily A_G with $NIR_{vbb} * \Sigma Q^i$
694 (**Fig. 12b, d**). One other aspect worth highlighting is that corn is C4 crop and has a
695 complex canopy structure. The fact that our broadband NIRv can capture this so well,
696 further shows the promise of this analysis.

697 Interestingly, by simply taking the square of NDVI and by multiplying $NDVI^2$ with ΣQ^i ,
698 we obtained even better correlation ($r_{Pearson} = 0.81$) between A_G versus $NDVI^2 * \Sigma Q^i$ (**Fig.**
699 **12e**) as compared to A_G versus $NIR_{vbb} * \Sigma Q^i$. The square of NDVI gives an almost

700 equivalent result as fractional absorbed PAR (Carlson and Ripley, 1997). Then
701 multiplying NDVI^2 with $\sum Q^i$, we get an estimate of absorbed PAR, which is why **Fig.**
702 **12e** showed strong correlation between A_G and $\text{NDVI}^2 \sum Q^i$.

703 It is worth mentioning that in the higher latitude sites, NDVI describes GPP during
704 vegetation green-up when the energy from PAR is generally high (Zhang et al., 2020;
705 Descals et al., 2022). However, NDVI provides little information about GPP in the
706 autumn where photosynthesis is driven by the seasonally decreasing PAR. Therefore,
707 including PAR with NIR_{Vbb} could have added advantage in describing the day-to-day
708 variability in GPP during the periods of varying cloudiness where NDVI remains almost
709 invariant (Zhang et al., 2020; Descals et al., 2022).

710 **3.4. Impacts of background soil exposure, phenology, radiation** 711 **components, and water stress on broadband NDVI (SQ4)**

712 This analysis is categorized into two parts. First, we examined the consequence of
713 background soil exposure and phenology on the estimation of NDVI_{bb} . We carried a
714 residual error analysis across different ecosystems where the differences between
715 broadband and satellite NDVI ($\delta_{\text{NDVI}} = \text{NDVI}_{\text{bb}} - \text{NDVI}$) were assessed with respect to
716 GCC for a range of Soil Adjusted and Atmospherically Resistant Vegetation Index
717 (SARVI) (**Fig. 13** below). At the second step, we compared the broadband VIS and NIR
718 hemispherical reflectances with satellite narrowband directional reflectance for a large
719 range of incident and reflected radiation components (**Fig. 14 - 16**).

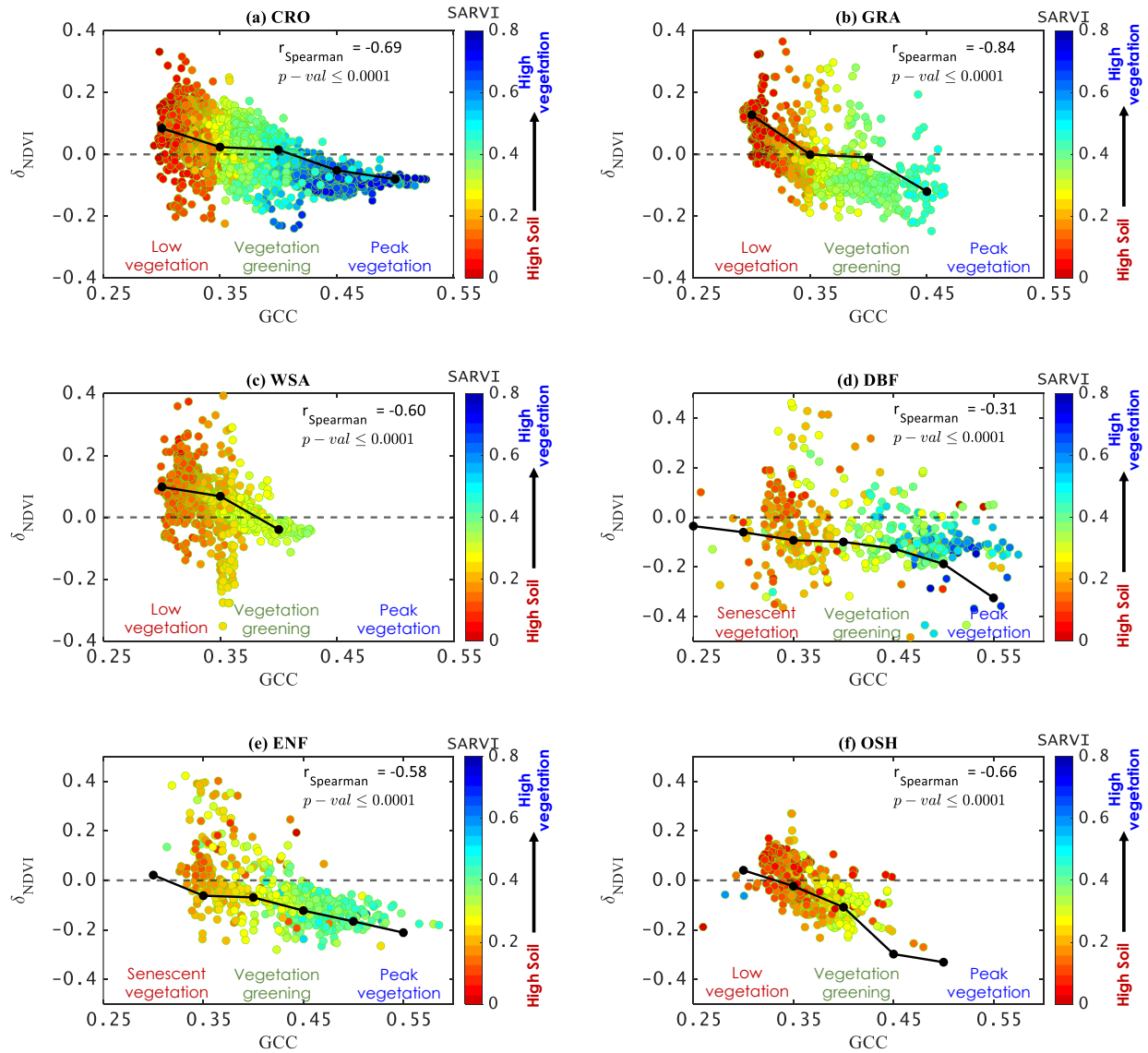


Figure 13: Plots of residual difference between $NDVI_{bb}$ and satellite NDVI ($\delta_{NDVI} = NDVI_{bb} - NDVI$) versus Green Chromatic Coordinate (GCC) for a wide range of soil background conditions across diverse ecosystems. Color shading is by Soil Adjusted and Atmospherically Resistant Vegetation index (SARVI), which serves as an indicator of soil-canopy background. The black line indicates the average bias for each bin. This clearly indicates a consistent positive difference between $NDVI_{bb}$ and satellite NDVI during low vegetation or during vegetation senescence, which also coincides with low SARVI. The black line shows the mean bias pattern for different classes of GCC.

720 The scatterplots of mean δ_{NDVI} clustered for different classes of GCC (from
 721 senescence/low vegetation to peak vegetation) showed significant relationships between
 722 mean δ_{NDVI} and GCC for varying background from high soil cover to high canopy cover
 723 across different ecosystems ($r = 0.31 - 0.84$) (**Fig. 13**). A consistent positive bias in

724 NDVI_{bb} ($\delta_{\text{NDVI}} > 0$) is evident when the vegetation cover is low (low GCC) or during the
725 senescent phase of vegetation (datapoints in red color cluster). The green reflectance
726 contributes very little during senescence and red reflectance has a greater dominance
727 among the three primary band reflectances, ultimately leading to low GCC
728 ($0.30 < \text{GCC} < 0.35$). Low leaf area during the senescent phase in forests or due to grazing
729 in the grasslands leads to greater exposure of soil background at the field-of-view of the
730 sensors, ultimately reading to high NIR reflectance (Huete et al., 1988; Qi et al., 1994).
731 Spectral reflectance of the canopies is mixed with background reflectance due to multiple
732 scattering in the broad NIR band. Such high NIR reflectance apparently leads to an
733 overestimation of NDVI_{bb} under low vegetation cover. These positive biases in NDVI_{bb}
734 also corresponded to low SARVI (0 – 0.2) (**Fig. 13**), indicating soil background to be
735 exerting considerable influence on the canopy spectra and the calculated NDVI_{bb} (Huete
736 et al., 1988; Qi et al., 1994). This dual assessment of δ_{NDVI} with respect to phenology and
737 soil background variations authenticates the sensitivity of NDVI_{bb} to first-order soil
738 exposure effects.

739 In the estimation of NDVI_{bb}, the consistency of broadband hemispherical VIS and NIR
740 reflectances ($\rho_{\text{vis,bb}}$, $\rho_{\text{nir,bb}}$) play a crucial role where proximal sensing of R_g (R_g^i , R_g^r) and
741 PAR components (Q^i , Q^r) are used. Therefore, to further understand the effects of
742 background on this overestimation of NDVI_{bb}, we compared the performance of $\rho_{\text{vis,bb}}$
743 and $\rho_{\text{nir,bb}}$ with respect to satellite narrowband directional reflectances under varying PAR
744 (both Q^i and Q^r) (**Fig. 14 - 16** below). The effects of individual R_g components (R_g^i , R_g^r)
745 on this comparison is very similar to what is seen in **Fig. 14 – 16** and they are not shown
746 for brevity.

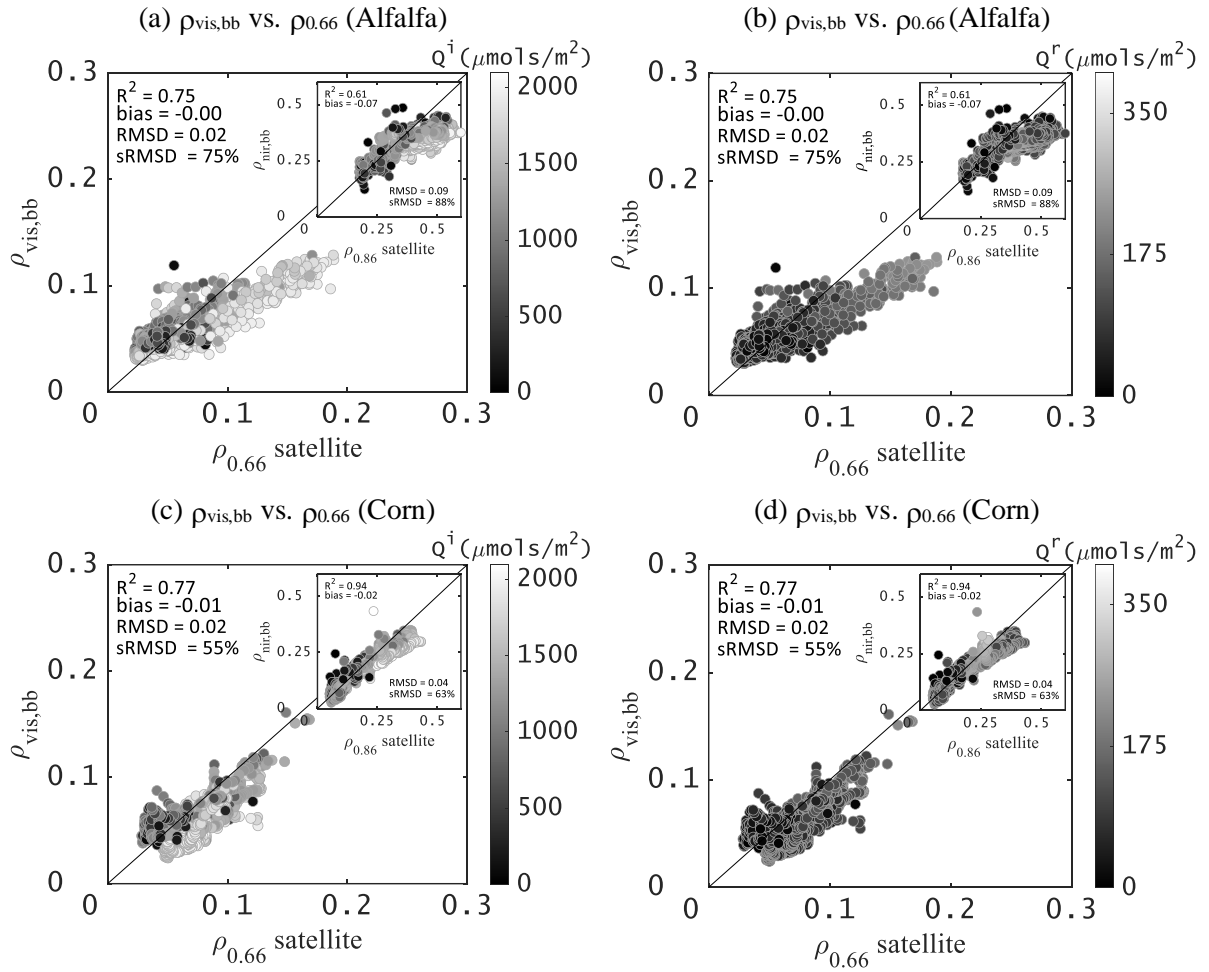


Figure 14: Illustrative examples of comparison between $\rho_{vis,bb}$ versus $\rho_{0.66}$ for a range of incident and reflected PAR (Q^i and Q^r) over alfalfa (a, b) and corn (c, d) in California. This clearly shows a tendency of systematic underestimation of $\rho_{vis,bb}$ with respect to $\rho_{0.66}$, and the underestimation increases at high Q^i and Q^r . Figures in the inset shows a similar comparison between $\rho_{nir,bb}$ versus $\rho_{0.86}$ for a range Q^i and Q^r . This analysis was performed with Planet Fusion data.

747 The spectral reflectance comparison revealed $\rho_{vis,bb} < \rho_{0.66}$ for the majority of the data
 748 points and their differences were magnified with increasing Q^i ($Q^i > 1500 \mu\text{mol}$). The
 749 effects of high Q^i and Q^r on $\rho_{nir,bb}$ was also visible and $\rho_{nir,bb} > \rho_{0.86}$ at low satellite $\rho_{0.86}$
 750 (inset of Fig. 14 – 16). This implies that in the estimation of $NDVI_{bb}$ [$NDVI_{bb} = (\rho_{nir,bb} -$
 751 $\rho_{vis,bb}) / (\rho_{nir,bb} + \rho_{vis,bb})$], there is a consistent overestimation of the numerator ($\rho_{nir,bb} -$
 752 $\rho_{vis,bb} > \rho_{0.86} - \rho_{0.66}$) with increasing Q^i under low fractional vegetation cover. This

753 ultimately led to greater difference between NDVI_{bb} and satellite NDVI ($\delta_{\text{NDVI}} > 0$) at high
 754 Q^i , Q^r , and at high soil background.

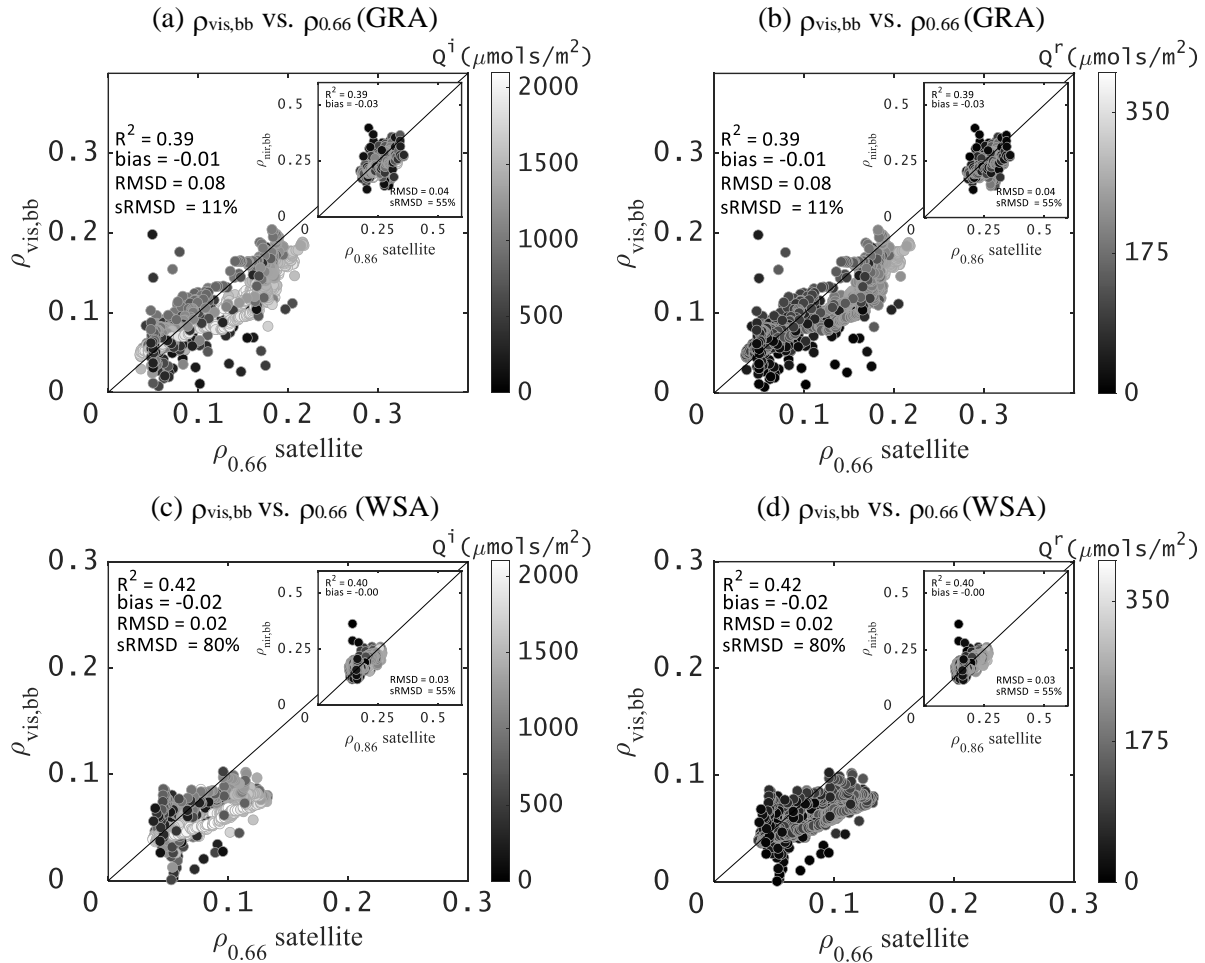
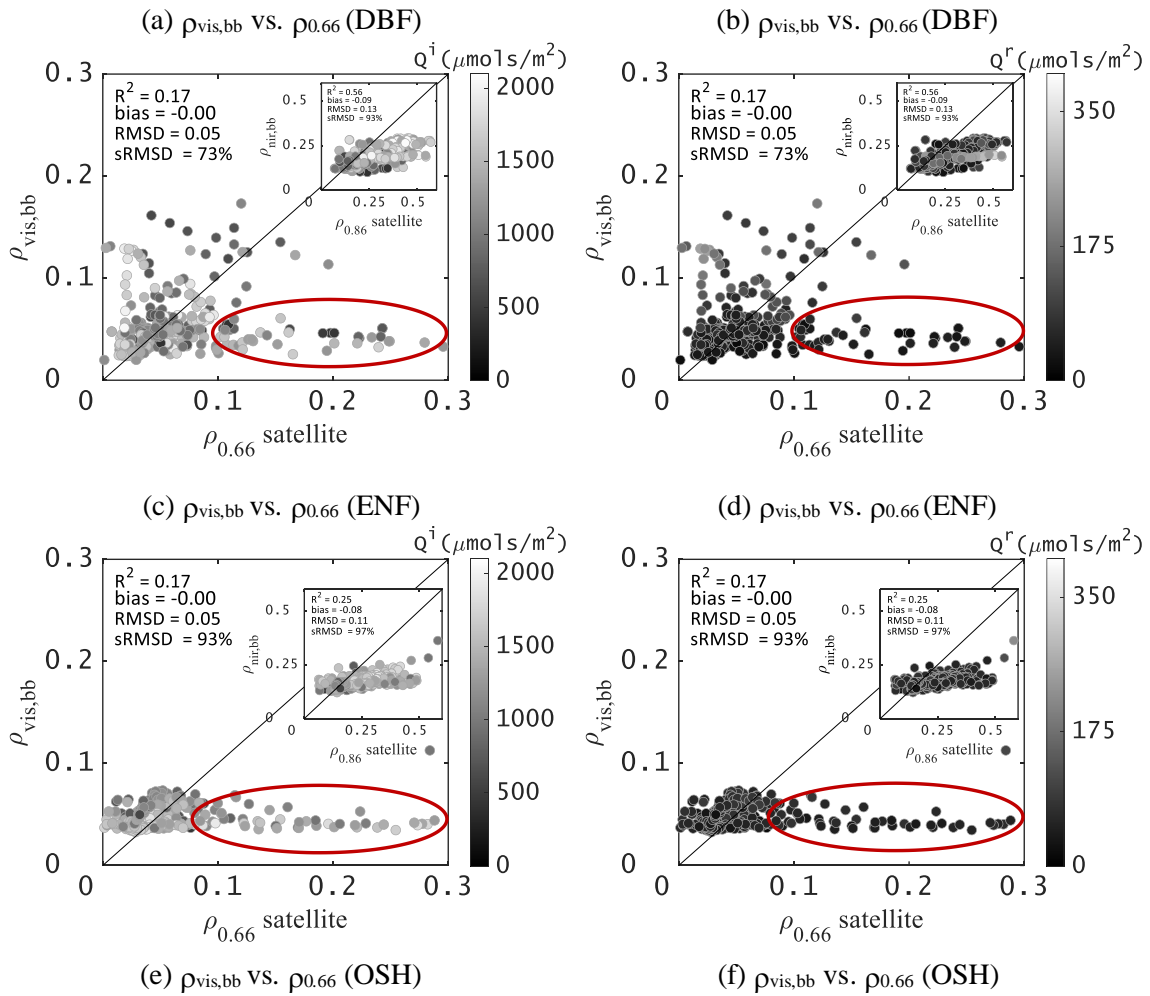


Figure 15: Illustrative examples of comparison between $\rho_{\text{vis,bb}}$ versus $\rho_{0.66}$ for a range of incident and reflected PAR (Q^i and Q^r) over grassland (Vaira ranch) (a, b) and woody savanna (Tonzi ranch) (c, d) in California. This clearly shows a tendency of systematic underestimation of $\rho_{\text{vis,bb}}$ with respect to $\rho_{0.66}$, and the underestimation increases at high Q^i and Q^r . Figures in the inset shows a similar comparison between $\rho_{\text{mir,bb}}$ versus $\rho_{0.86}$ for a range Q^i and Q^r . This analysis was performed with Planet Fusion data.

755 The overestimation tendency of NDVI_{bb} apparently diminished with vegetation greening
 756 and it showed underestimation under dense vegetation cover ($\delta_{\text{NDVI}} < 0$). However,
 757 exceptions were also found in the deciduous broadleaf forest (DBF) and evergreen
 758 needleleaf forest (ENF), where the underestimation tendency of NDVI_{bb} was visible
 759 across all the clusters of GCC (**Fig. 13d, e**). In the NDVI_{bb} retrieval, the broadband NIR

760 reflectance covers up to 3.0 μm , which also accounts the reflectance signals from the
 761 shortwave infrared domain. Studies reported that in the coniferous needles, the surface
 762 reflectance in the shortwave infrared region is low (Pierrat et al., 2022; Roberts et al.,
 763 2004). This consequently leads to an underestimation of $\rho_{\text{nir,bb}}$ and NDVI_{bb} at the ENF.



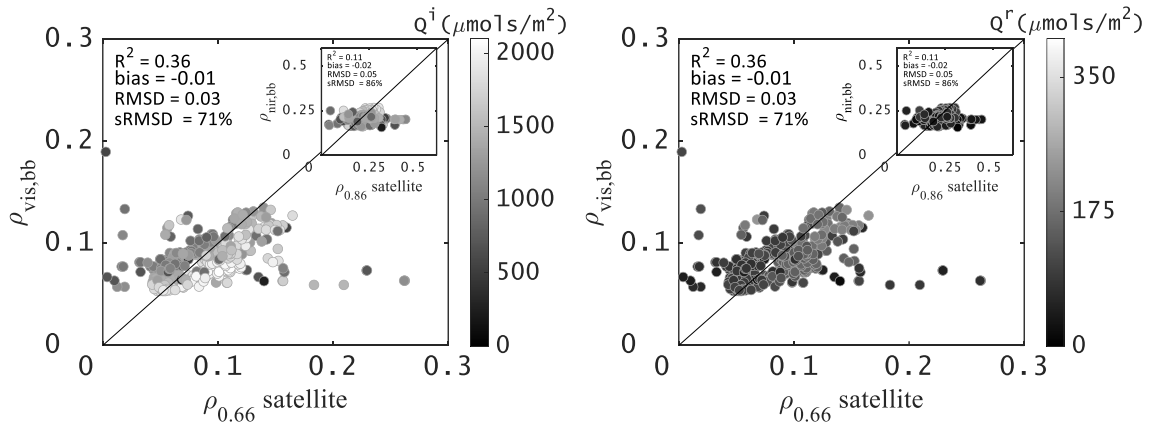


Figure 16: Illustrative examples of comparison between $\rho_{vis,bb}$ versus $\rho_{0.66}$ for a range of incident and reflected PAR (Q^i and Q^r) over deciduous broadleaf forests (DBF) (a, b), evergreen needleleaf forests (ENF) (c, d), and open shrubland (OSH) (e, f). This clearly shows a tendency of systematic underestimation of $\rho_{vis,bb}$ with respect to $\rho_{0.66}$, and the underestimation increases at high Q^i and Q^r . Figures in the inset shows a similar comparison between $\rho_{nir,bb}$ versus $\rho_{0.86}$ for a range Q^i and Q^r . This analysis was performed with HLS data since no Planet Fusion data was available for these ecosystems.

764 **Figure 13 (a-f)** also showed underestimation of $NDVI_{bb}$ at low GCC ($GCC < 0.35$)
 765 corresponding to high background soil exposure (red data cluster at $\delta_{NDVI} < 0$). These
 766 datapoints could be associated with the low magnitude of Q^i and the details are revealed
 767 in **Fig. 14 – 16**. In all the ecosystems, there were data clusters with $\rho_{vis,bb} > \rho_{0.66}$ and these
 768 datapoints are associated with low Q^i (Q^i : 0 - 650 μmols) (**Fig. 14 - 16**). In the croplands,
 769 $\rho_{vis,bb}$ could also be affected due to irrigation, and $\rho_{vis,bb}$ might pick up the signal of wet
 770 soil, ultimately leading to $\rho_{vis,bb} > \rho_{0.66}$ (Ma et al., 2019). The effects of low Q^i and low Q^r
 771 were also evident in $\rho_{nir,bb}$ to some extent (**inset of Fig. 14 – 16**). These conditions led to
 772 an underestimation of the numerator in eq. (7) ($\rho_{nir,bb} - \rho_{vis,bb} < \rho_{0.86} - \rho_{0.66}$), ultimately
 773 leading to an underestimation of $NDVI_{bb}$ as compared to satellite $NDVI$.
 774 These detailed analysis (**Fig. 14 – 16**) additionally helped understanding the reasons for
 775 underestimation of $NDVI_{bb}$ (**Fig. 3 – 6, Fig. 13**) and saturation of NIR_{vbb} under high
 776 biomass when compared with the satellites (**Fig. 3 – 6**). A closer look at these figures

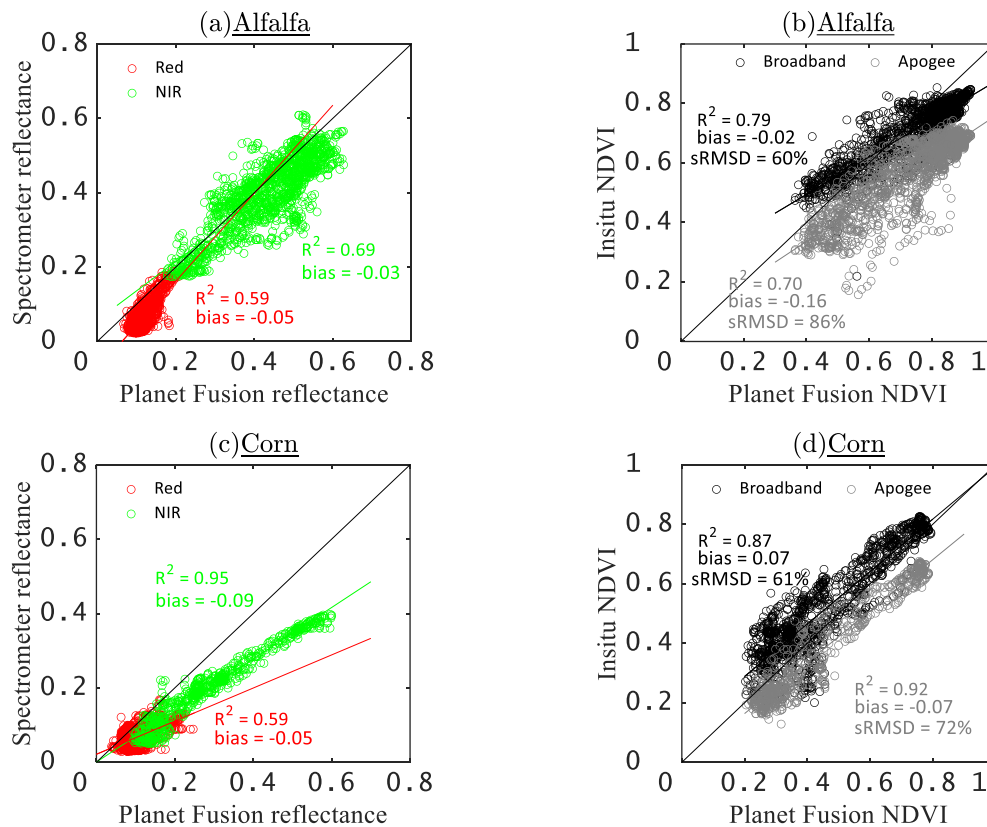
777 (inset of Fig. 14 - 16) revealed that almost 75-85% of the $\rho_{\text{nir,bb}}$ signal tends to become
778 invariant with increasing $\rho_{0.86}$ beyond 0.30 μm at the cropland (alfalfa, corn) and forests
779 (DBF, ENF) and beyond 0.25 μm at the OSH. The wavebands have very different
780 bandwidths. The broadband NIR reflectance has differential sensitivity to increasing
781 biomass as compared to $\rho_{0.86}$, which ultimately led to saturation in NIR_{bb} compared with
782 satellite NIR_{v} at high F_{E} . The consequence of the water stress on $\rho_{\text{vis,bb}}$ and $\rho_{\text{nir,bb}}$
783 estimation is described in Appendix A4 (Fig. A4-A6).

784 4. Broader implications

785 Overall, our analysis shows that continuous and combined measurement of Q and R_{g}
786 components serves a robust proximal sensing capability for diagnosing the seasonal
787 variability in NDVI across ecological and climatic gradients. The NDVI_{bb} versus satellite
788 NDVI relationship was highly significant when compared with satellite sensors at
789 different spatial resolutions (Fig. 3 - 8), across a broad spectrum of managed and
790 unmanaged ecological settings, crop management regimes (e.g., irrigated vs. rainfed) that
791 experience dynamic water stress, productivity variability, and physiological variations.
792 With the availability of PAR and R_{g} components worldwide from different FLUXNET
793 sub-networks, a global comparison with satellite NDVI and other vegetation indices is
794 foreseen in the future. Due to the nature of the broadband reflectance retrieval from the
795 proximal sensing of hemispherical radiation components, the spectral differences
796 between broadband versus narrowband reflectances at the selected band regions are
797 obvious. Therefore, we do not anticipate a perfect one to one relation between NDVI_{bb}
798 versus satellite NDVI unless we have hyperspectral tower-based remote sensing, or
799 custom built sensors of Q and R_{g} components at the narrowband wavelengths. In fact,

800 comparing different NDVI from different satellite sensors showed substantial differences
 801 across different ecosystems (Fan and Liu, 2016; Huang et al., 2021).

802 The advantage of $NDVI_{bb}$ in comparison to spectrometer based NDVI became further
 803 evident when we compared $NDVI_{bb}$ and Planet Fusion NDVI with Apogee spectrometer
 804 NDVI at some flux tower sites covering four representative land cover types (alfalfa,
 805 corn, pasture, herbaceous vegetation) in California (**Fig. 17b, d, f, h**). The consistent
 806 negative bias in Apogee NDVI (-0.07 to -0.26) is mainly attributed to weak relationship
 807 in Apogee versus Planet Fusion red and NIR spectral reflectance (**Fig. 17a, c, e, g**),
 808 confirming that exploring the radiation components to estimate in-situ vegetation
 809 attributes is credible.



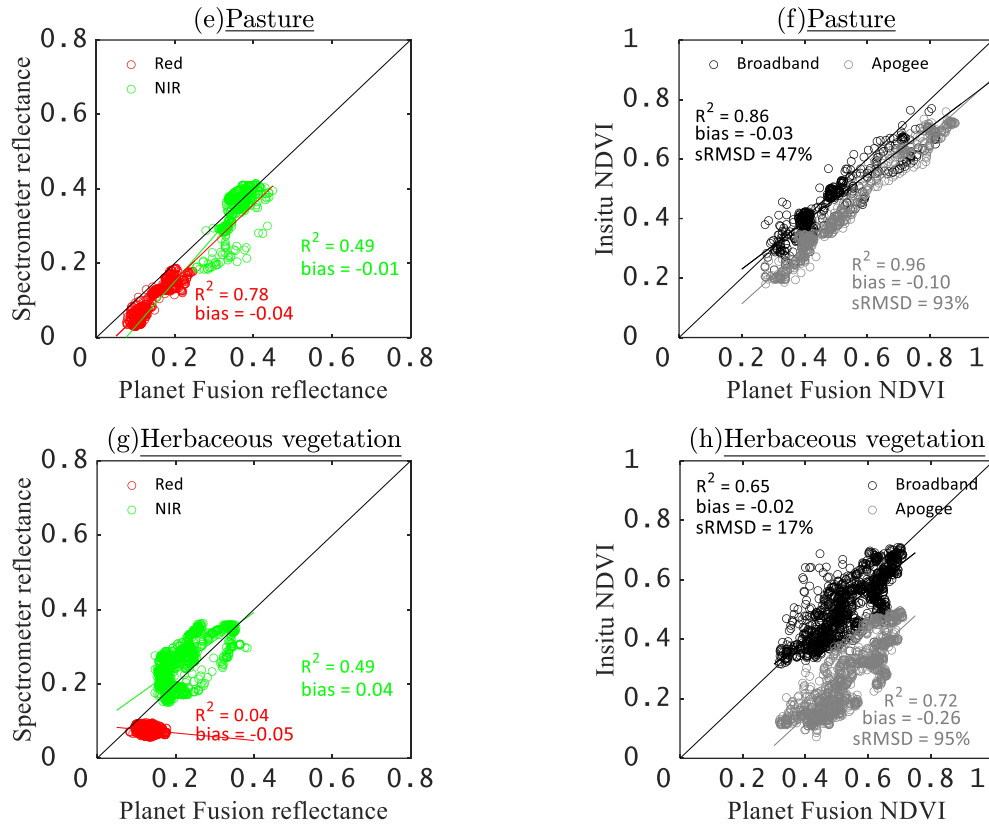


Figure 17: (a, c, e, g) Illustrative examples of comparison between Apogee spectrometer versus Planet Fusion reflectances in red and NIR wavelengths in four representative ecosystems in California. (b, d, f, h) Comparison between Apogee spectrometer versus Planet Fusion and broadband NDVI in four representative ecosystems in California. This clearly shows a tendency of systematic underestimation of Apogee NDVI with respect to satellite, which is attributed to the disagreement in spectral reflectances between Apogee spectrometer and Planet Fusion.

810 Relatively greater disagreement of $NDVI_{bb}$ and $NIR_{V_{bb}}$ with respect to HLS data is due to
 811 the relatively coarser spatial resolution of 30 x 30 m NDVI used in their comparison.
 812 Significant variability in greenness and fractional vegetation cover can be present at the
 813 sub-pixel scale depending on the ecosystem types (Turner et al, 2002). If HLS versus
 814 $NDVI_{bb}$ and $NIR_{V_{bb}}$ agrees on a seasonal scale, we can assume that the greenness and
 815 vegetation fraction surrounding the tower is representative of the land cover within the 30
 816 x 30 m S2 and L8/9 pixel containing the tower location. For several sites that showed
 817 moderate agreement between the tower $NDVI_{bb}$ and HLS NDVI, the variability in

818 fraction ground cover within the 30 x 30 m S2 and L8/9 pixel might be responsible for
819 such a behavior. This indicates the challenges and intricacies associated with respect to
820 directly comparing flux tower $NDVI_{bb}$ with satellite NDVI at a coarser spatial scale in the
821 presence of profound spatial variability in vegetation cover. The extent to which the
822 variability in fractional vegetation cover within one HLS pixel could impact such
823 comparison, could only be estimated upon having coincident Planet Fusion (3 m) and
824 HLS (30 m) data across all the sites. In the present study, only a small subset of sites
825 (seven biomet sites) had both Planet fusion and HLS datasets. This led us examining the
826 effects of the variability of vegetation fraction on $NDVI_{bb}$ versus satellite NDVI
827 evaluation at four different ecosystems (**Appendix A5; Fig. A7**). The statistical
828 comparison clearly showed the effects due to the variability in vegetation fraction when
829 NDVI average from 10 x 10 pixels of Planet Fusion was used for the evaluation of
830 $NDVI_{bb}$. Although a detailed spatial variability analysis could shed greater insight, such
831 analysis is beyond the scope of the present study. Nevertheless, $NDVI_{bb}$ is a valid proxy
832 of satellite NDVI for a wide range of conditions tested.

833 We found that $NIR_{v_{bb}} * \sum Q^i$ is a robust structural proxy for GPP by combining 28 site-
834 years of data (seven sites and four years for each site). $NIR_{v_{bb}} * \sum Q^i$ tends to have higher
835 signal quality (Baldocchi et al., 2020) as compared to $NDVI_{bb}$, and $NDVI_{bb}$ is known to
836 become invariant at high GPP (Baldocchi et al., 2020, Dechant et al., 2022). The
837 correlative relationship between GPP and NDVI emerges because green leaves do
838 photosynthesis, and there is a seasonality in greenness and photosynthesis. Therefore, one
839 should be careful to use such correlative relationships to upscale GPP from the
840 information of NDVI. A linear relationship between GPP and $NIR_{v_{bb}} * \sum Q^i$ was also

841 reported for croplands (Dechant et al., 2020, 2022; Liu et al., 2020; Wu et al., 2019),
842 which could further be exploited to understand GPP variability at different spatio-
843 temporal scales. Our results confirm and considerably extend previous findings and
844 demonstrated that the linearity between GPP and $\text{NIR}_{\text{vbb}} * \Sigma Q^i$ also holds for a range of
845 ecosystems that experience variable water stress. The significant outcome of this analysis
846 is that from the measurements of four radiation components, we are able to detect the
847 most critical vegetation variables that have a direct link with ecosystem carbon
848 assimilation across a range of climatic gradients. Our results also substantiate the findings
849 of Pierrat et al. (2022) who showed that in the boreal ecosystems where seasonal
850 downregulation of photosynthesis occurs without significant changes in canopy structure
851 or chlorophyll content, NDVI scales poorly with carbon assimilation. We believe that
852 more work is needed to develop a robust scaling function for GPP versus $\text{NIR}_{\text{vbb}} * \Sigma Q^i$
853 relationship across a wide spectrum of ecological gradients. Such studies should use high
854 spatial resolution satellite data, standardized PAR and R_g sensors and calibration
855 methods. Nevertheless, our tower-based broadband NDVI and NIR_{v} is promising enough
856 to be treated as highly valuable and critical vegetation attributes relevant to flux
857 measurement footprints for ecosystem modeling.

858 Despite its own limitations, the present study could be seen promising enough that
859 highlights the utility of shortwave and photosynthetically active radiation measurements
860 to augment the proximal sensing capability at the flux tower sites. The in-situ broadband
861 NDVI derived through transforming these radiation signals could make a stronger case
862 for how these data could be used for handshaking between ecosystem-scale
863 measurements and remote sensing for scaling and/or understanding satellite observables.

864 **5. Summary and conclusion**

865 We conclude that the net fluxes of broadband shortwave radiation components in
866 conjunction with the components of photosynthetically active radiation offer a novel
867 proximal sensing perspective to directly retrieve a robust broadband NDVI and NIRv
868 relevant to explain ecosystem productivity for a wide spectrum of ecosystems and
869 climatic gradients. This novel perspective is obtained through a simplified method which
870 neither needs explicit radiative transfer for solving canopy reflectance, nor does it need
871 any additional spectrometer measurements. Our analysis revealed that the discrepancies
872 between the broadband NDVI and operational satellite-based NDVI products are due to
873 the differences in hemispherical versus directional reflectance, differential sensitivity of
874 broad visible and near infrared reflectance to background soil exposure, water stress and
875 biomass accumulation and resultant saturation of the hemispherical reflectance signals at
876 high biomass.

877 These critical insights and multiscale comparison with satellite products are highly
878 significant to monitoring the intraseasonal and interannual variability of NDVI directly at
879 the flux tower sites and relevant to validating operational NDVI products from the Earth
880 observation mission. Statistical analysis over a range of ecosystems and climatic limits
881 demonstrates the potential of the broadband NDVI and NIRv as a valid alternative to
882 study the effects of vegetation seasonality on energy-water-carbon flux interactions and
883 their interannual variability worldwide. This novel approach can be implemented across
884 all the flux tower sites of AmeriFlux and Fluxnet subnetworks to generate insightful
885 vegetation dynamics information for the ecosystem modeling community and
886 complementing the PhenoCam observations. As more flux sites are equipped with the

887 necessary radiometric instrumentation, i.e., quantum sensors and pyranometers, we
888 expect the available ground-based data to increase dramatically. This will provide the
889 community with a critical tool to link flux tower measurements with satellite-borne
890 observations.

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1211 **Figure captions:**

Figure 1: (a) Conceptual diagram showing the hypothesis for estimating broadband spectral reflectance from the measurements of hemispherical broadband radiation components in PAR and total shortwave spectral region. It also shows an example of the narrowband spectral reflectances that we obtain in red and near infrared spectral region from operational remote sensing satellite Landsat-9 (Source: <https://landsat.usgs.gov/spectral-characteristics-viewer>). VIS signified visible, NIR signifies near-infrared, MIR signifies mid-wave infrared. (b) Figure showing the scaling factor for converting PAR (both incident and reflected) from $\mu\text{mol}/\text{m}^2/\text{s}$ to W/m^2 for a range of NDVI as an example over rice crop in California.

Figure 2: An illustrative diagram showing the sequence of results corresponding to the science questions (SQs) and the respective figure numbers associated with the description of results falling under individual science question.

Figure 3: (a) Plots of NDVI_{bb} and NIR_{vbb} versus Planet Fusion NDVI and NIR_{v} (3 m spatial resolution) in the Californian cropland ecosystems for $\text{NDVI} > 0.25$ for a range of evaporative fraction (F_{E}) representing stressed to unstressed conditions. F_{E} is an indicator of water availability and denotes the ratio of evaporation (latent heat flux) to equilibrium evaporation. (b) Temporal dynamics of NDVI_{bb} (black dots) and NDVI (green dots) along with daily precipitation (P) (blue stairs) shows close correspondence between them in Alfalfa [Bi1] and Corn [Bi2].

Figure 4: (a) Plots of NDVI_{bb} and NIR_{vbb} versus Planet Fusion NDVI and NIR_{v} (3 m spatial resolution) in water-limited Californian grassland (GRA) and woody savanna (WSA) ecosystems for $\text{NDVI} > 0.25$ for a range of evaporative fraction (F_{E}) representing stressed to unstressed conditions. F_{E} is an indicator of water availability and denotes the ratio of evaporation (latent heat flux) to equilibrium evaporation. (b) Temporal dynamics of NDVI_{bb} (black dots) and NDVI (green dots) along with daily precipitation (P) (blue stairs) shows close correspondence between them especially in the grassland [Var] and partly in woody savanna [Ton].

Figure 5: (a) Plots of NDVI_{bb} and NIR_{vbb} versus Planet Fusion NDVI and NIR_{v} (3 m spatial resolution) in Californian wetland ecosystems (non-tidal) for $\text{NDVI} > 0.25$ for a range of evaporative fraction (F_{E}) representing stressed to unstressed conditions. F_{E} is an indicator of water availability and denotes the ratio of evaporation (latent heat flux) to equilibrium evaporation. (b) Temporal dynamics of NDVI_{bb} (black dots) and NDVI (green dots) along with daily precipitation (P) (blue stairs) shows close correspondence between them in both East End and Mayberry.

Figure 6: Pooled evaluation plots of NDVI_{bb} and NIR_{vbb} versus Planet Fusion NDVI and NIR_{v} (3 m spatial resolution) by combining all the seven sites of Californian ecosystems for a range of evaporative fraction (F_{E}) representing stressed to unstressed conditions. F_{E} is an indicator of water availability and denotes the ratio of evaporation (latent heat flux) to equilibrium evaporation.

Figure 7: (a, c) Plots of NDVI_{bb} and NIR_{vbb} versus Landsat and Sentinel-2 NDVI and NIR_{v} (30 m spatial resolution) in energy-limited ecosystems of Biomet and NEON sites. Color shading is done by evaporative fraction (F_{E}) showing stressed to unstressed conditions which corresponds to water and energy limits within the energy-limited environment. (b, d) Illustrative examples of temporal dynamics of NDVI_{bb} (black dots) and NDVI (green dots) along with daily precipitation (P) (blue stairs) showing close correspondence in the seasonal and interannual variability of NDVI_{bb} and NDVI at the NEON sites Blandy Experimental Farm (xBL) and Dead Lake (xDL).

Figure 8: (a, c) Plots of $NDVI_{bb}$ and NIR_{vbb} versus Landsat and Sentinel-2 $NDVI$ and NIR_v (30 m spatial resolution) in water-limited ecosystems of Biomet and NEON sites. Color shading is done by evaporative fraction (F_E) showing stressed to unstressed conditions which corresponds to water and energy limits within the water-limited environment. (b, d) Illustrative examples of temporal dynamics of $NDVI_{bb}$ (black dots) and $NDVI$ (green dots) along with daily precipitation (P) (blue stairs) showing close correspondence in the seasonal and interannual variability of $NDVI_{bb}$ and $NDVI$ over Bouldin corn (Bi2), Vaira ranch (Var) and two grasslands sites of NEON Konza Prairie Biological Station (xKA and xKZ).

Figure 9: (a-b) Daily variation in $NDVI_{bb}$, Planet Fusion $NDVI$, Green Chromatic Coordinate (GCC) (secondary y-axis), and evaporative fraction (F_E) over agricultural ecosystems (alfalfa and corn) in California. Here we plot daily values, averaged over 4 years, normalized by the annual mean for that variable. (c-d) Correlation map showing the strength of seasonal relationship between individual variables. For corn, the correlation map is applicable for the growing season from March to September.

Figure 10: (a-b) Daily variation in $NDVI_{bb}$, Planet Fusion $NDVI$, Green Chromatic Coordinate (GCC) (secondary y-axis), and evaporative fraction (F_E) over grassland (Vaira ranch) and woody savanna (Tonzi ranch) ecosystems in California. Here we plot daily values, averaged over 4 years, normalized by the annual mean for that variable. (c-d) Correlation map showing the strength of seasonal relationship between individual variables during the growing season from March to October.

Figure 11: (a-b) Plots of intraseasonal variability (expressed as ‘coefficient of variation’, cv) in $NDVI$ ($NDVI_{bb}$) and NIR_v (NIR_{vbb}) versus intraseasonal variability in gross photosynthesis (A_G) by combining all the site data in different ecosystems of California. Color shading is by precipitation (P) variability. This also shows the steeper slope of cv - $NDVI$ (and cv - $NDVI_{bb}$) vs. cv - A_G and cv - NIR_v (and cv - NIR_{vbb}) vs. cv - A_G relationship with increasing precipitation variability.

Figure 12: (a, b) Site-level relationships of A_G versus $NDVI_{bb}$ and A_G versus $NIR_{vbb} * \Sigma Q^i$ at seven eddy covariance towers of UC Berkeley Biomet sites that includes 2 crop sites (one C4, one C3), one grassland site, one woody savanna and 3 wetland sites. (c, d) Similar plot is shown by plotting A_G with Planet Fusion $NDVI$ and $NIR_{vbb} * \Sigma Q^i$. Here ΣQ^i and A_G are the daily integrated Q^i (MJ) and A_G (gC/m^2) obtained by summing up the half-hourly observation. Data points inside the red circle showed saturation in A_G with increasing $NIR_{vbb} * \Sigma Q^i$. These data points belong to corn crop and could presumably be associated with the diffuse component of Q^i . (e) Scatter plot of A_G versus $NDVI_{bb}^{2*} \Sigma Q^i$.

Figure 13: Plots of residual difference between $NDVI_{bb}$ and satellite $NDVI$ ($\delta_{NDVI} = NDVI_{bb} - NDVI$) versus Green Chromatic Coordinate (GCC) for a wide range of soil background conditions across diverse ecosystems. Color shading is by Soil Adjusted and Atmospherically Resistant Vegetation index (SARVI), which serves as an indicator of soil-canopy background. The black line indicates the average bias for each bin. This clearly indicates a consistent positive difference between $NDVI_{bb}$ and satellite $NDVI$ during low vegetation or during vegetation senescence, which also coincides with low SARVI. The black line shows the mean bias pattern for different classes of GCC.

Figure 14: Illustrative examples of comparison between $\rho_{vis,bb}$ versus $\rho_{0.66}$ for a range of incident and reflected PAR (Q^i and Q^r) over alfalfa (a, b) and corn (c, d) in California. This clearly shows a tendency of systematic underestimation of $\rho_{vis,bb}$ with respect to $\rho_{0.66}$, and the underestimation increases at high Q^i and Q^r . Figures in the inset shows a similar comparison between $\rho_{nir,bb}$ versus $\rho_{0.86}$ for a range Q^i and Q^r . This analysis was performed with Planet Fusion data.

Figure 15: Illustrative examples of comparison between $\rho_{\text{vis,bb}}$ versus $\rho_{0.66}$ for a range of incident and reflected PAR (Q^i and Q^r) over grassland (Vaira ranch) (a, b) and woody savanna (Tonzi ranch) (c, d) in California. This clearly shows a tendency of systematic underestimation of $\rho_{\text{vis,bb}}$ with respect to $\rho_{0.66}$, and the underestimation increases at high Q^i and Q^r . Figures in the inset shows a similar comparison between $\rho_{\text{nir,bb}}$ versus $\rho_{0.86}$ for a range Q^i and Q^r . This analysis was performed with Planet Fusion data.

Figure 16: Illustrative examples of comparison between $\rho_{\text{vis,bb}}$ versus $\rho_{0.66}$ for a range of incident and reflected PAR (Q^i and Q^r) over deciduous broadleaf forests (DBF) (a, b), evergreen needleleaf forests (ENF) (c, d), and open shrubland (OSH) (e, f). This clearly shows a tendency of systematic underestimation of $\rho_{\text{vis,bb}}$ with respect to $\rho_{0.66}$, and the underestimation increases at high Q^i and Q^r . Figures in the inset shows a similar comparison between $\rho_{\text{nir,bb}}$ versus $\rho_{0.86}$ for a range Q^i and Q^r . This analysis was performed with HLS data since no Planet Fusion data was available for these ecosystems.

Figure 17: (a, c, e, g) Illustrative examples of comparison between Apogee spectrometer versus Planet Fusion spectral reflectances in red and NIR wavelengths in four representative ecosystems in California. (b, d, f, h) Comparison between Apogee spectrometer versus Planet Fusion and broadband NDVI in four representative ecosystems in California. This clearly shows a tendency of systematic underestimation of Apogee NDVI with respect to satellite, which is attributed to the disagreement in spectral reflectances between Apogee spectrometer and Planet Fusion.

Figure A1: Boxplot of statistical error metric of NDVI_{bb} and NIR_{vbb} with respect to satellite NDVI and NIRv by combining data of both L8/9 and S2 of HLS by combining data of different sites falling in different ecosystem categories. Here, nRMSD is the normalized root mean squared deviation. This is computed by normalizing RMSD with the range (maximum - minimum) of satellite NDVI and NIRv.

Figure A2: (a) Time series of daily broadband vegetation index (NDVI_{bb}), LAI and canopy gap fraction (P_{gap}) at the oak grass savanna (Tonzi ranch). The data points inside the red rectangular box represent the periods when NDVI_{bb} and LAI showed maximum divergence, which coincided with larger gap fraction during the winter season. (b) Scatterplot of NDVI_{bb} versus LAI by averaging daily data for four years, showing significantly high correlation during the growing season.

Figure A3: (a-f) Plots of intraseasonal variability (expressed as ‘coefficient of variation’, cv) in satellite (Planet Fusion) NDVI (NDVI_{bb}) and NIRv (NIR_{vbb}) versus intraseasonal variability in GPP (A_G) in different ecosystems of California. Color shading is by precipitation (P) variability. This also shows the steeper slope of cv-NDVI (and cv- NDVI_{bb}) vs. cv- A_G and cv-NIRv (and cv- NIR_{vbb}) vs. cv- A_G relationship with increasing precipitation variability.

Figure A4: Illustrative examples of comparison between $\rho_{\text{vis,bb}}$ versus $\rho_{0.66}$ and $\rho_{\text{nir,bb}}$ versus $\rho_{0.86}$ for a range of water stress (evaporative fraction, F_E) over alfalfa and corn in California. This clearly shows a tendency of systematic underestimation of $\rho_{\text{vis,bb}}$ with respect to $\rho_{0.66}$, and the underestimation increases with elevated water stress. This analysis was performed with Planet Fusion data.

Figure A5: Illustrative examples of comparison between $\rho_{\text{vis,bb}}$ versus $\rho_{0.66}$ and $\rho_{\text{nir,bb}}$ versus $\rho_{0.86}$ for a range of water stress (evaporative fraction, F_E) over grassland (Vaira ranch) and woody savanna (Tonzi ranch) in California. This clearly shows a tendency of systematic underestimation of $\rho_{\text{vis,bb}}$ with respect to $\rho_{0.66}$, and the underestimation increases with elevated water stress. This analysis was performed with Planet Fusion data.

Figure A6: Illustrative examples of comparison between $\rho_{vis,bb}$ versus $\rho_{0.66}$ and $\rho_{nir,bb}$ versus $\rho_{0.86}$ for a range of water stress (evaporative fraction, F_E) over deciduous broadleaf forest (DBF) and evergreen needleleaf forest (ENF). This clearly shows a tendency of systematic underestimation of $\rho_{vis,bb}$ with respect to $\rho_{0.66}$, in majority of the datapoints. This also shows a clear saturation of $\rho_{nir,bb}$ signal at high $\rho_{0.86}$. This analysis was performed with HLS data since no Planet Fusion data was available for these ecosystems.

Figure A7: Illustrative examples of the effects of variability (coefficient of variation, cv in percent) in fractional vegetation cover (f_c) on broadband versus satellite NDVI comparison at four different ecosystems. We took the cluster average of 10 x 10 pixel NDVI from Planet Fusion surrounding the tower sites. To understand the impact of f_c variability we took the standard deviation of f_c of 10 x 10 pixel and normalized with mean f_c of the same pixels.

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1228 **Appendix**

1229 **A1. Assessment of NDVI_{bb} and NIR_{vbb} in different ecosystems**

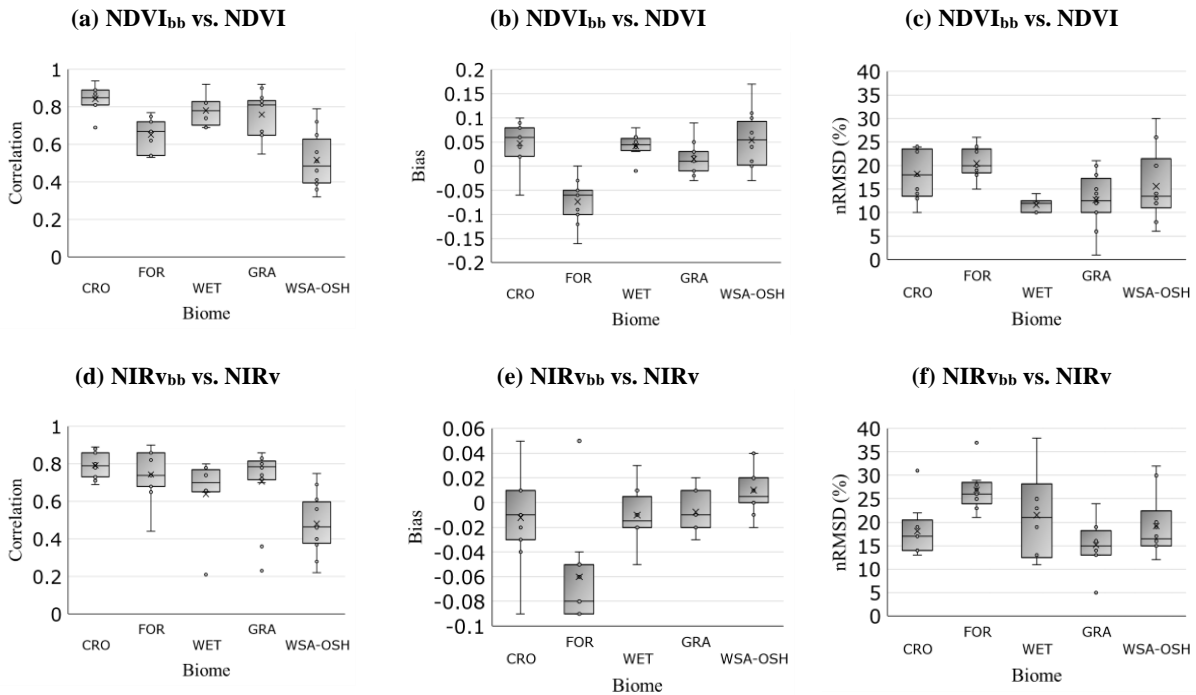


Figure A1: Boxplot of statistical error metric of NDVI_{bb} and NIR_{vbb} with respect to satellite NDVI and NIRv by combining data of both L8/9 and S2 of HLS by combining data of different sites falling in different ecosystem categories. Here, nRMSD is the normalized root mean squared deviation. This is computed by normalizing RMSD with the range (maximum - minimum) of satellite NDVI and NIRv.

1230 Ecosystem wise analysis by combining data of both L8/9 and S2 revealed (**Fig. A1**)
 1231 significantly high correlation between NDVI_{bb} versus satellite NDVI in cropland (0.81 -
 1232 0.85), grassland (0.82 - 0.87), and wetlands ($r = 0.68 - 0.78$) (**Fig. A1a**). However, a
 1233 relatively degraded, yet significant relationship was noted in forest ($r = 0.52 - 0.59$) and
 1234 woody savanna-shrubland (0.48 - 0.64). The normalized RMSD (nRMSD) showed higher
 1235 percentage difference in cropland and forest (18 - 22%) as compared to the other
 1236 ecosystems (**Fig. A1b**). Boxplots also revealed high systematic negative mean bias in
 1237 forest and systematic positive mean bias in cropland, grassland, and woody savanna-
 1238 shrubland (**Fig. A1c**). Although no systematic difference in error metrics was identified
 1239 with respect to L8/9 and S2 sensors, boxplot indicated relatively higher errors in NIR_{vbb}

1240 as compared to $NDVI_{bb}$. Error statistics of individual sites with both L8/9 and S2 HLS
1241 data are listed in **Table A1** (for L8/9) and **Table A2** (for S2).

1242 **A2. Comparing $NDVI_{bb}$ with leaf area index (LAI)**

1243 To understand whether $NDVI_{bb}$ is able to capture the variation in leaf area index (LAI),
1244 we also analyzed $NDVI_{bb}$ with respect to LAI observations at the Tonzi ranch site (oak
1245 grass savanna). Significantly high correlation ($r = 0.82$) was found during the active
1246 growing season between day of the year (DOY) 130 to 280 and for the canopy gap
1247 fraction (P_{gap}) of 0.4 to 0.6 (**Fig. A2**). $NDVI_{bb}$ and LAI started to diverge with increasing
1248 P_{gap} beyond 0.65, which corresponds to the period around DOY 300 onwards (autumn),
1249 and this divergence remained until the warm spring around DOY 100. This is the time
1250 when grasses at the understory start greening up and $NDVI_{bb}$ can be affected due to the
1251 background effects. Differences in the nature of LAI measurement (upward looking
1252 digital camera) versus the estimation of broadband NDVI (downward looking radiation
1253 sensors) could be responsible for this divergence. The divergence between LAI and
1254 NDVI in autumn is also because leaf browning results in lower NDVI while leaves can
1255 still remain in the canopy affecting transmittance and LAI. This is one of the advantages
1256 of using NDVI over direct measurements of LAI from transmittance, as NDVI is related
1257 to the amount of green leaves and transmittance measures total LAI.

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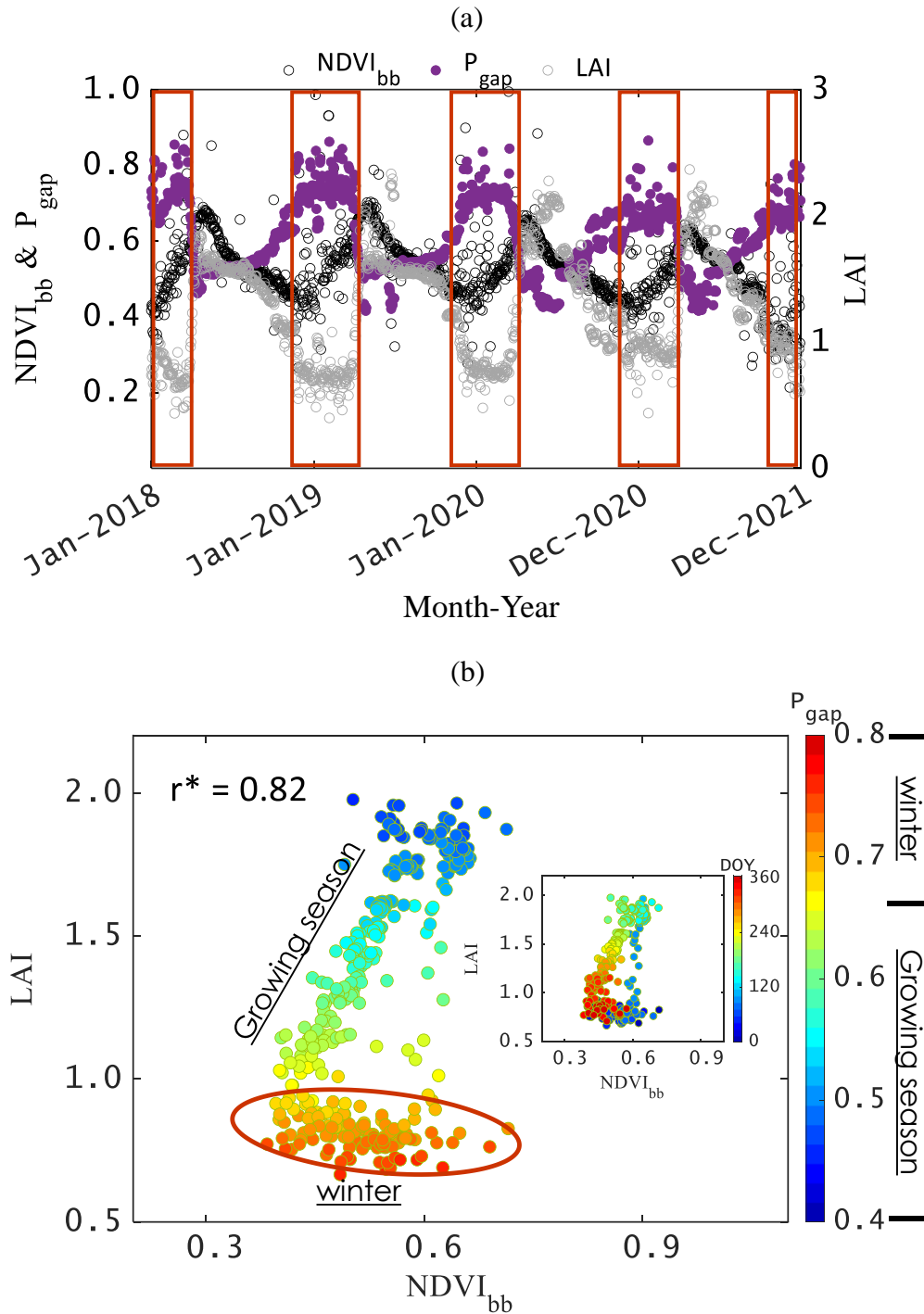


Figure A2: (a) Time series of daily broadband vegetation index (NDVI_{bb}), LAI and canopy gap fraction (P_{gap}) at the oak grass savanna (Tonzi ranch). The data points inside the red rectangular box represent the periods when NDVI_{bb} and LAI showed maximum divergence, which coincided with larger gap fraction during the winter season. (b) Scatterplot of NDVI_{bb} versus LAI by averaging daily data for four years, showing significantly high correlation during the growing season.

1262 **A3. NDVI_{bb} and NIR_{vbb} variability versus GPP variability across**
 1263 **ecosystems (SQ3)**

1264 This analysis is based on the continuous time series EC tower A_G record of seven Biomet
 1265 lab sites and the Planet Fusion data and is linked with section 3.3. The intraseasonal
 1266 variability of both NDVI_{bb} and NIR_{vbb} was significantly correlated with the intraseasonal
 1267 variability of A_G across different ecosystems ($r = 0.51 - 0.65$ and $r = 0.49 - 0.74$). Similar
 1268 pattern was also noted in satellite NDVI and NIR_v versus A_G for all the ecosystems ($r =$
 1269 $0.40 - 0.79$ and $r = 0.52 - 0.84$) (**Fig. A3**).

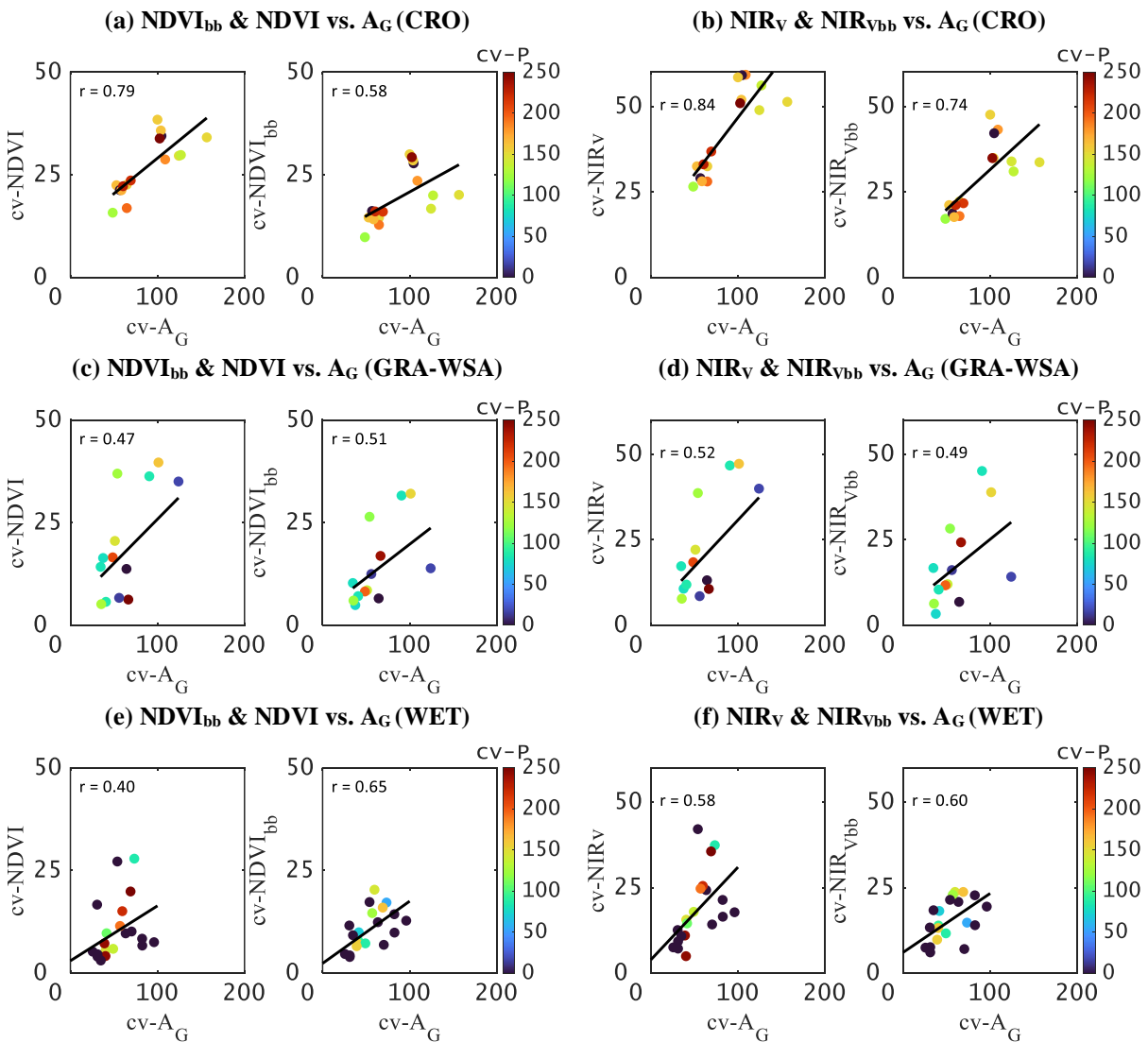


Figure A3: (a-f) Plots of intraseasonal variability (expressed as ‘coefficient of variation’, cv) in satellite (Planet Fusion) NDVI (NDVI_{bb}) and NIR_V (NIR_{Vbb}) versus intraseasonal variability in GPP (A_G) in different ecosystems of California. Color shading is by precipitation (P) variability. This also shows the steeper slope of cv-NDVI (and cv-NDVI_{bb}) vs. cv-A_G and cv-NIR_V (and cv-NIR_{Vbb}) vs. cv-A_G relationship with increasing precipitation variability.

1270 **A4. Effects energy-water-limitations on broadband spectral reflectance**
 1271 **across different ecosystems?**

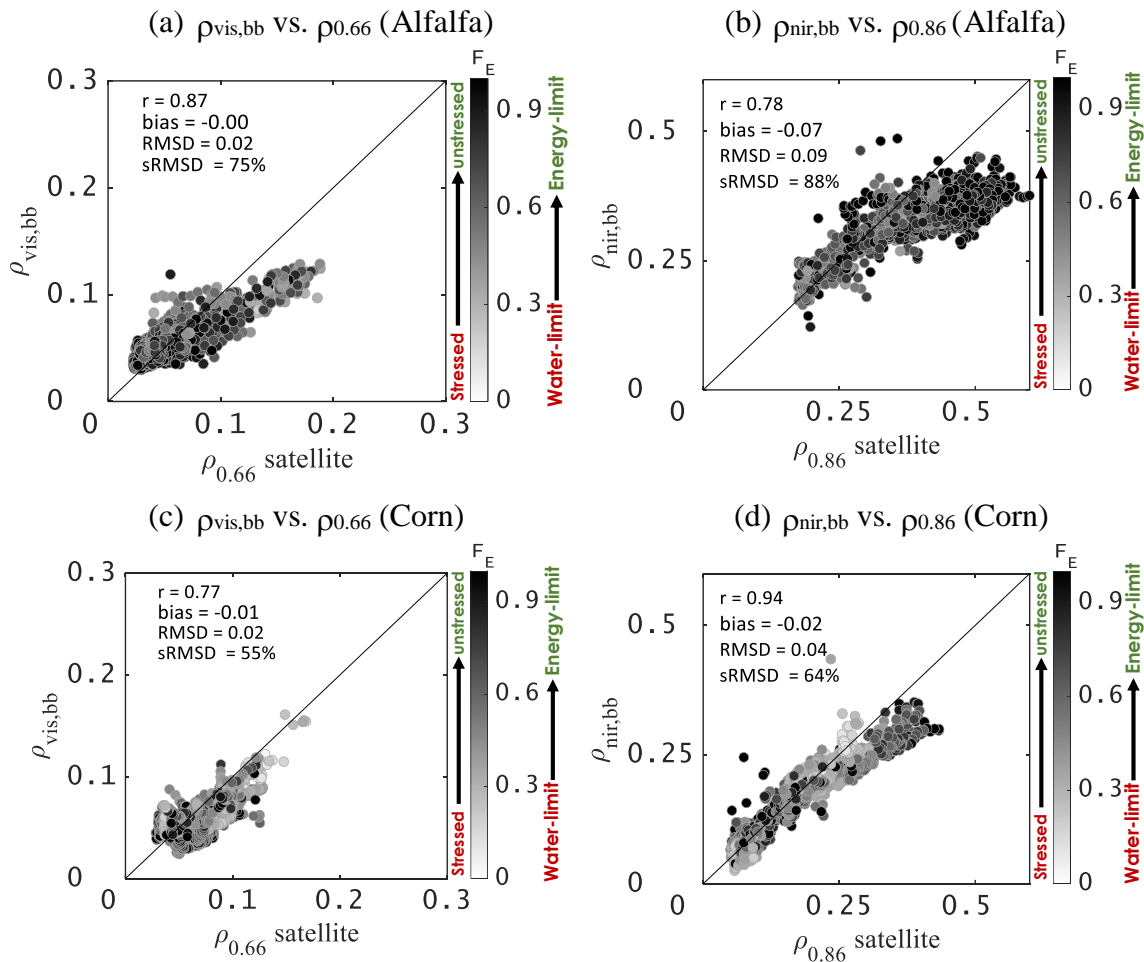


Figure A4: Illustrative examples of comparison between $\rho_{vis,bb}$ versus $\rho_{0.66}$ and $\rho_{nir,bb}$ versus $\rho_{0.86}$ for a range of water stress (evaporative fraction, F_E) over alfalfa and corn in California. This clearly shows a tendency of systematic underestimation of $\rho_{vis,bb}$ with respect to $\rho_{0.66}$, and the underestimation increases with elevated water stress. This analysis was performed with Planet Fusion data.

1272 An analysis of $\rho_{vis,bb}$ and $\rho_{nir,bb}$ with respect to satellite narrowband red and NIR

1273 reflectance ($\rho_{0.66}$ and $\rho_{0.86}$) showed $\rho_{vis,bb}$ was systematically less as compared to $\rho_{0.66}$

1274 ($\rho_{vis,bb} < \rho_{0.66}$) for the majority of the data points (**Fig. A3 – A5**), and their differences were
 1275 magnified with increasing water stress ($F_E < 0.3$).

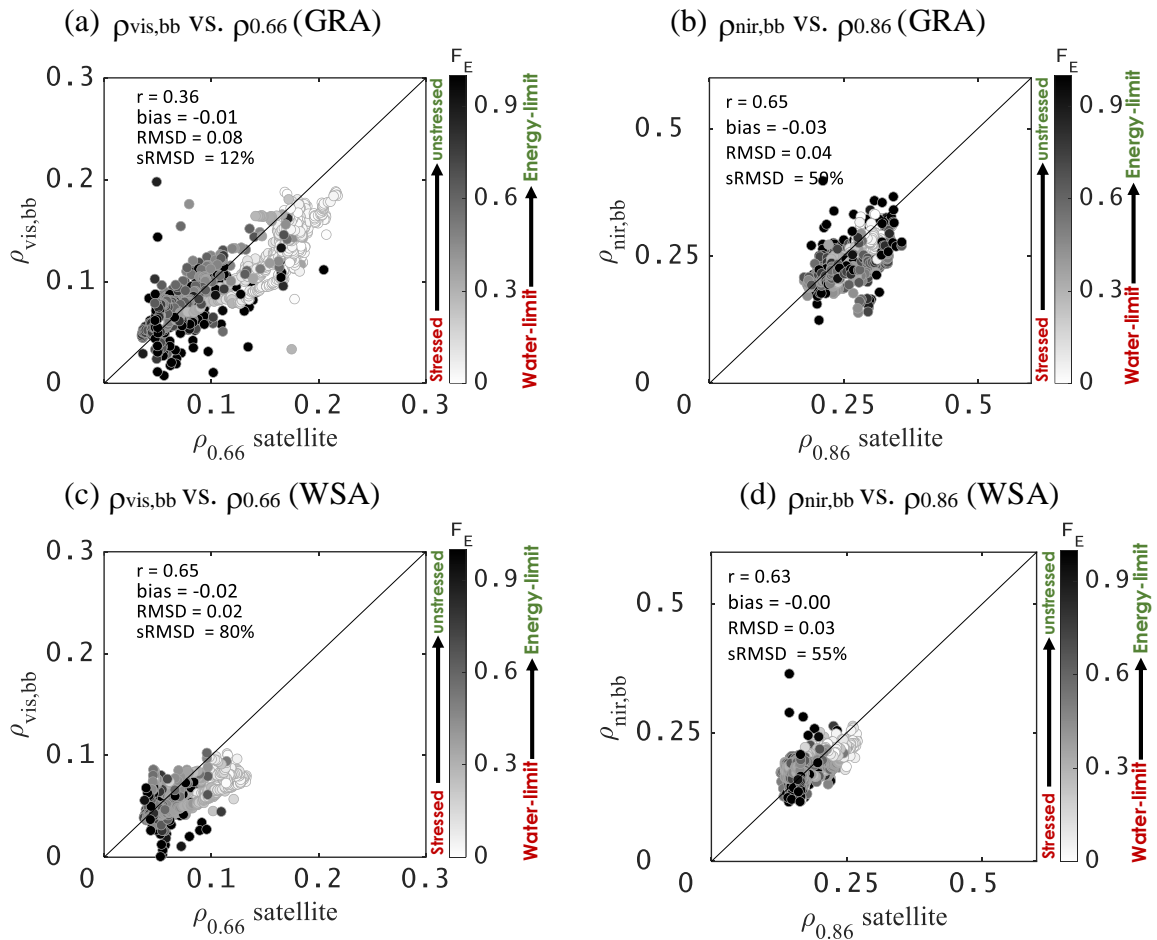


Figure A5: Illustrative examples of comparison between $\rho_{vis,bb}$ versus $\rho_{0.66}$ and $\rho_{nir,bb}$ versus $\rho_{0.86}$ for a range of water stress (evaporative fraction, F_E) over grassland (Vaira ranch) and woody savanna (Tonzi ranch) in California. This clearly shows a tendency of systematic underestimation of $\rho_{vis,bb}$ with respect to $\rho_{0.66}$, and the underestimation increases with elevated water stress. This analysis was performed with Planet Fusion data.

1276 The effects of water limitations on $\rho_{nir,bb}$ was also evident at low $\rho_{0.86}$ and $\rho_{nir,bb} > \rho_{0.86}$ at
 1277 high water stress. This implies substantial overestimation of the numerator in $NDVI_{bb}$ as
 1278 compared to satellite $NDVI$ ($\rho_{nir,bb} - \rho_{vis,bb} > \rho_{0.86} - \rho_{0.66}$) with increasing water limitation,
 1279 ultimately leading to a large positive difference between them when water stress
 1280 progresses.

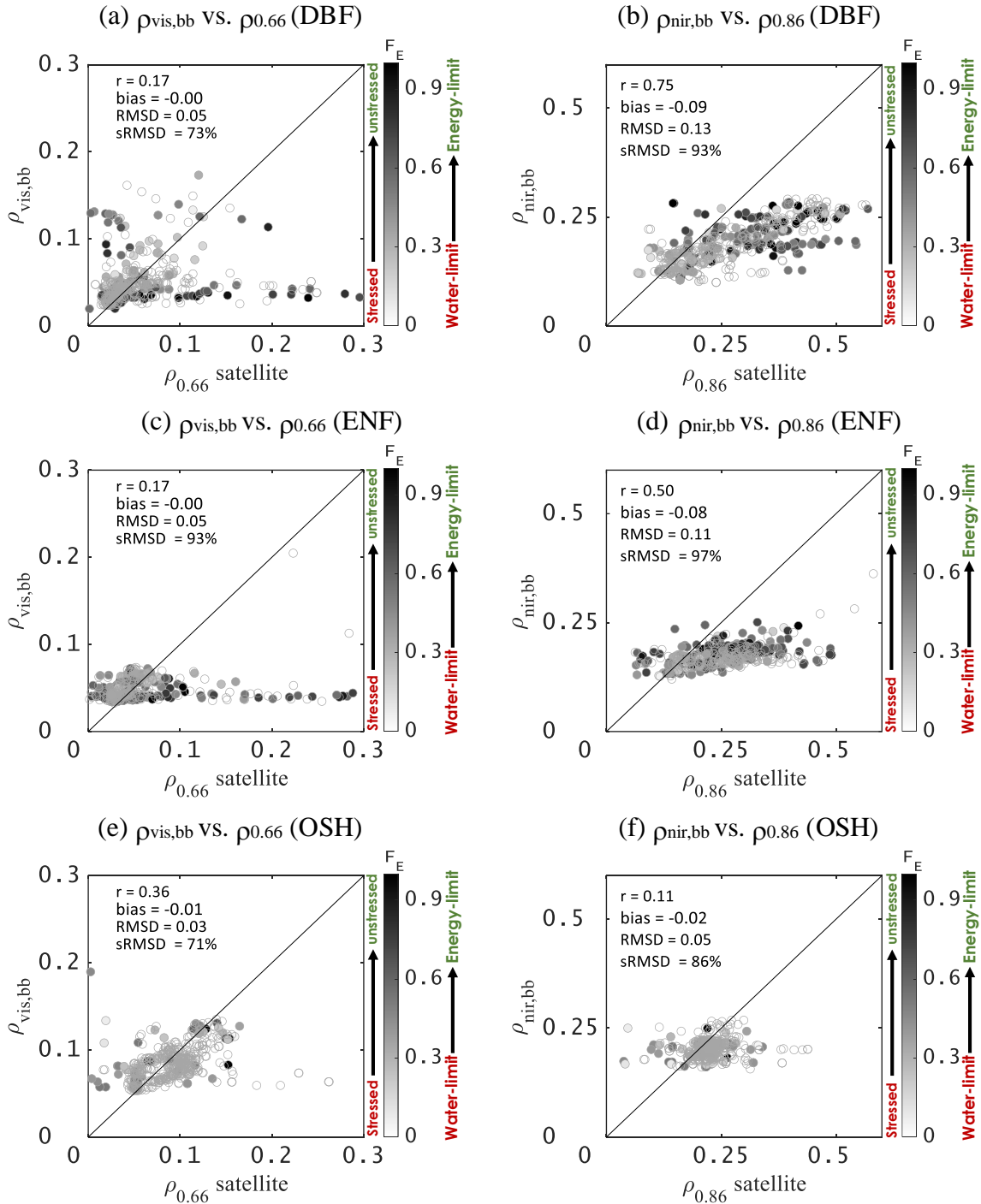


Figure A6: Illustrative examples of comparison between $\rho_{\text{vis,bb}}$ versus $\rho_{0.66}$ and $\rho_{\text{nir,bb}}$ versus $\rho_{0.86}$ for a range of water stress (evaporative fraction, F_E) over deciduous broadleaf forest (DBF) and evergreen needleleaf forest (ENF). This clearly shows a tendency of systematic underestimation of $\rho_{\text{vis,bb}}$ with respect to $\rho_{0.66}$, in majority of the datapoints. This also shows a clear saturation of $\rho_{\text{nir,bb}}$ signal at high $\rho_{0.86}$. This analysis was performed with HLS data since no Planet Fusion data was available for these ecosystems.

1281 **A5. Effects of variability of vegetation fraction on $NDVI_{bb}$ versus**
 1282 **satellite NDVI relationship**

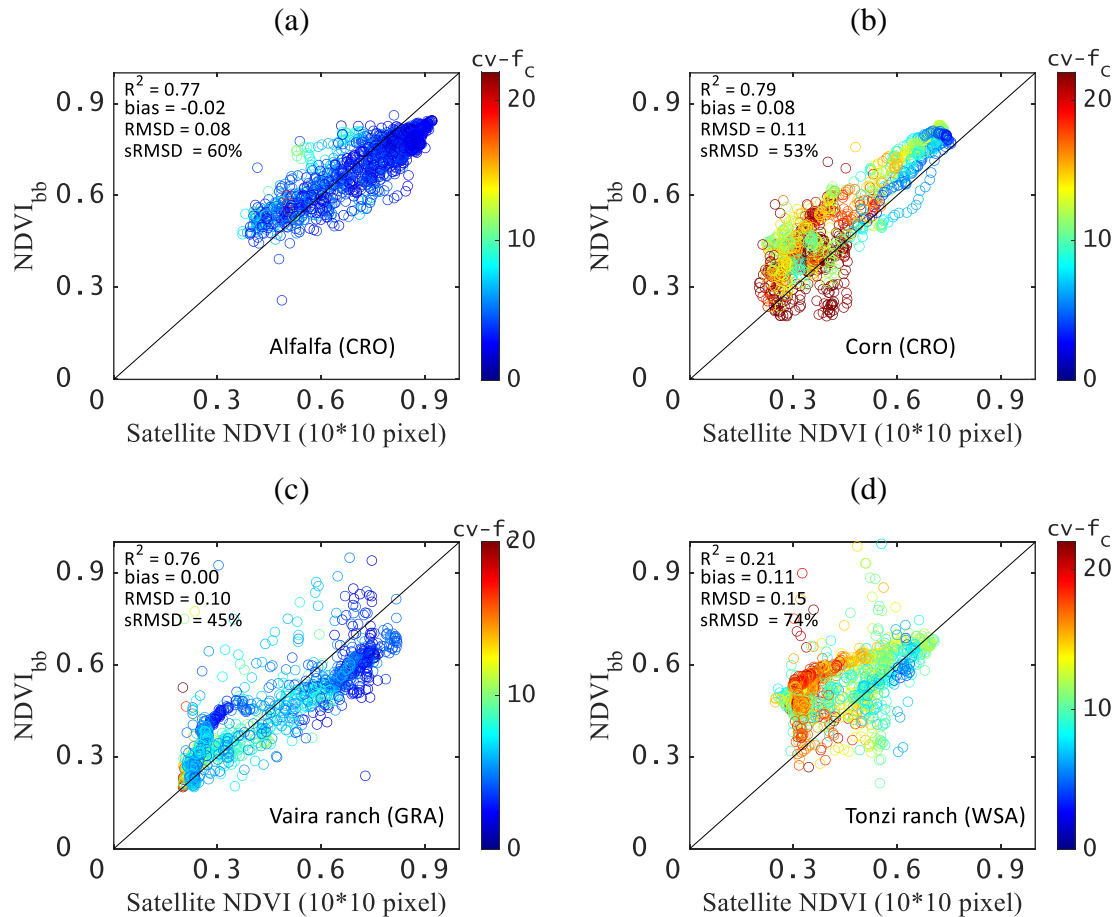


Figure A7: Illustrative examples of the effects of variability (coefficient of variation, cv in percent) in fractional vegetation cover (f_c) on broadband versus satellite NDVI comparison at four different ecosystems. We took the cluster average of 10×10 pixel NDVI from Planet Fusion surrounding the tower sites. To understand the impact of f_c variability we took the standard deviation of f_c of 10×10 pixel and normalized with mean f_c of the same pixels.

1283 A simple analysis is conducted to examine the effects of coarser spatial resolution on the
 1284 agreement between $NDVI_{bb}$ versus satellite NDVI. This analysis reveals the reasons for
 1285 relatively higher errors in $NDVI_{bb}$ when it was compared with 30 m spatial resolution
 1286 L8/9 and S2 NDVI from HLS datasets (as compared to Planet Fusion). This analysis is
 1287 only possible over a small subset of sites in California where coincident data from both
 1288 Planet Fusion and HLS is available. Therefore, investigation is made at four different
 1289 ecosystems using data of UC Berkeley Biomet sites.

1290 **Figure A7** shows the effects of variability in fractional vegetation cover (f_c) on the
1291 comparison between $NDVI_{bb}$ versus satellite NDVI. To understand the impact of f_c
1292 variability, we estimated the coefficient of variation of f_c ($cv-f_c$, in percent.). We took the
1293 standard deviation of Planet Fusion (3 m spatial resolution) f_c over 10 x 10 pixel cutouts
1294 surrounding the towers and normalized it with mean f_c of the same numbers of pixels.
1295 Taking the cluster average of 10 x 10 pixel NDVI from Planet Fusion surrounding the
1296 tower sites, we found that the difference between $NDVI_{bb}$ versus satellite NDVI to be
1297 sensitive to high $cv-f_c$ ($\geq 20\%$). This showed overestimation of $NDVI_{bb}$ for low values of
1298 satellite NDVI at GRA and WSA (Fig. A6c, d). Overall, high systematic RMSD was
1299 found in both these ecosystems (45% in GRA and 74% in WSA) as compared to the
1300 errors obtained from NDVI averaging over 3x3 pixels (40% in GRA and 61% in WSA).
1301 The variability of f_c was very little over alfalfa and no significant difference due to pixel
1302 averaging was found at this site. For corn, the difference was found only at the start of the
1303 growing season when $cv-f_c$ was high ($\geq 20\%$) and the positive difference between $NDVI_{bb}$
1304 versus satellite NDVI was narrowed down with full vegetation growth.

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Table A1: Site wise error statistics of NDVI_{bb} and NIR_{vbb} (parenthesis) with respect to L8/L9 Harmonized Landsat and Sentinel (HLS) data

Ecosystem	Site	r	bias	RMSD	nRMSD (%)
CRO	US-Bi1	0.81 (0.78)	0.02 (-0.03)	0.10 (0.08)	14 (17)
	US-Bi2	0.89 (0.86)	0.10 (0.01)	0.13 (0.04)	23 (14)
	US-Ne3	0.69 (0.69)	0.08 (0.05)	0.16 (0.10)	23 (19)
	US-UiA	0.94 (0.89)	-0.06 (-0.09)	0.12 (0.14)	18 (31)
	US-UiB	0.82 (0.80)	0.08 (-0.04)	0.17 (0.11)	24 (22)
	US-xSL	0.90 (0.73)	0.06 (0.00)	0.07 (0.02)	24 (14)
GRA	US-xAE	0.65 (0.72)	0.01 (-0.01)	0.08 (0.02)	12 (16)
	US-xCP	0.65 (0.70)	0.05 (0.01)	0.08 (0.02)	14 (13)
	US-xKA	0.81 (0.84)	0.01 (-0.02)	0.09 (0.04)	15 (14)
	US-xKZ	0.83 (0.80)	0.01 (-0.03)	0.13 (0.08)	20 (19)
	US-Var	0.92 (0.86)	-0.01 (-0.02)	0.08 (0.04)	10 (13)
	US-Wkg	0.67 (0.36)	-0.03 (-0.02)	0.07 (0.03)	13 (24)
WSA	US-Ton	0.65 (0.56)	0.10 (0.02)	0.13 (0.03)	13 (15)
	US-xSJ	0.51 (0.61)	-0.03 (-0.02)	0.10 (0.03)	13 (16)
OSH	US-xJR	0.72 (0.69)	0.07 (0.02)	0.08 (0.02)	20 (20)
	US-xNQ	0.36 (0.47)	0.04 (0.00)	0.07 (0.01)	8 (17)
	US-xSR	0.39 (0.28)	0.00 (-0.01)	0.06 (0.02)	14 (15)
FOR	US-xAB	0.62 (0.65)	-0.09 (-0.09)	0.14 (0.11)	24 (37)
	US-xBL	0.72 (0.90)	-0.10 (-0.10)	0.15 (0.12)	25 (27)
	US-xDL	0.72 (0.82)	-0.06 (-0.09)	0.13 (0.12)	18 (29)
	US-xHa	0.54 (0.74)	-0.05 (-0.09)	0.18 (0.13)	20 (26)
	US-xJE	0.54 (0.68)	-0.12 (-0.08)	0.17 (0.10)	20 (26)
WET	US-Myb	0.69 (0.66)	0.05 (0.01)	0.09 (0.03)	12 (13)
	US-TW1	0.74 (0.74)	-0.01 (-0.05)	0.07 (0.06)	10 (25)
	US-TW4	0.69 (0.21)	0.04 (-0.02)	0.10 (0.09)	14 (38)

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Table A2: Site wise error statistics of NDVI_{bb} and NIR_{vbb} with respect to S2 Harmonized Landsat and Sentinel (HLS) data

Ecosystem	Site	r	bias	RMSD	nRMSD
CRO	US-Bi1	0.85 (0.79)	0.02 (-0.02)	0.09 (0.07)	10 (13)
	US-Bi2	0.81 (0.88)	0.09 (0.01)	0.15 (0.04)	15 (14)
	US-Ne3	-	-	-	-
	US-UiA	-	-	-	-
	US-UiB	-	-	-	-
	US-xSL	0.87 (0.71)	0.04 (-0.01)	0.07 (0.04)	13 (19)
GRA	US-xAE	0.90 (0.81)	0.02 (0.01)	0.05 (0.02)	10 (16)
	US-xCP	0.65 (0.74)	0.09 (0.02)	0.11 (0.02)	1 (5)
	US-xKA	0.82 (0.78)	0.03 (-0.01)	0.09 (0.03)	12 (14)
	US-xKZ	0.81 (0.83)	-0.01 (-0.02)	0.14 (0.07)	21 (19)
	US-Var	0.85 (0.79)	0.03 (0.01)	0.11 (0.04)	18 (13)
	US-Wkg	0.55 (0.23)	-0.02 (-0.01)	0.07 (0.03)	6 (16)
WSA	US-Ton	0.46 (0.37)	0.17 (0.04)	0.19 (0.05)	30 (30)
	US-xSJ	0.56 (0.46)	-0.01 (-0.01)	0.11 (0.03)	12 (12)
OSH	US-xJR	0.79 (0.75)	0.11 (0.04)	0.12 (0.04)	26 (32)
	US-xNQ	0.32 (0.40)	0.07 (0.01)	0.09 (0.01)	6 (19)
	US-xSR	0.41 (0.22)	0.01 (-0.01)	0.06 (0.02)	14 (16)
FOR	US-xAB	-	-	-	-
	US-xBL	0.77 (0.87)	-0.05 (-0.05)	0.12 (0.07)	19 (21)
	US-xDL	0.75 (0.86)	0.01 (-0.04)	0.13 (0.08)	19 (23)
	US-xHa	0.53 (0.44)	-0.16 (-0.06)	0.26 (0.12)	26 (28)
	US-xJE	0.67 (0.74)	-0.03 (-0.05)	0.11 (0.07)	15 (25)
WET	US-Myb	0.83 (0.65)	0.08 (0.03)	0.10 (0.04)	14 (23)
	US-TW1	0.82 (0.78)	0.03 (-0.02)	0.10 (0.04)	12 (19)
	US-TW4	0.92 (0.80)	0.06 (-0.01)	0.08 (-0.03)	10 (11)

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