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Author

Shackley, M. Steven

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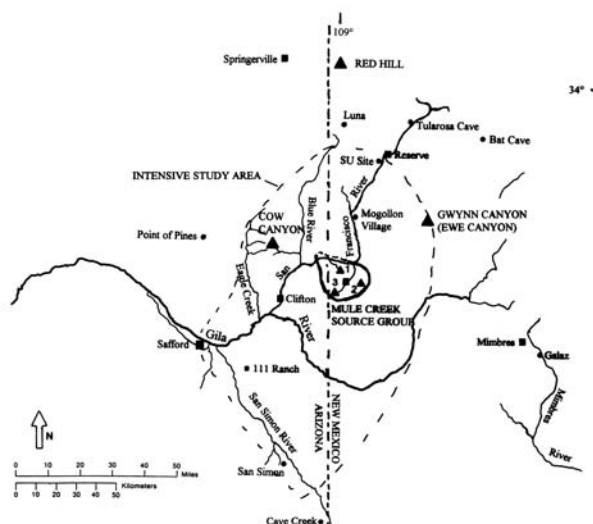
GEOARCHAEOLOGICAL XRF LAB

ARCHAEOLOGICAL X-RAY FLUORESCENCE SPECTROMETRY LABORATORY

8100 WYOMING BLVD., SUITE M4-158

ALBUQUERQUE, NM 87113 USA

SOURCE PROVENANCE OF OBSIDIAN ARTIFACTS FROM PITHOUSE THROUGH CLASSIC PERIOD SITES ALONG EAGLE CREEK, APACHE NATIONAL FOREST, EASTERN ARIZONA



by

M. Steven Shackley Ph.D., Director
Geoarchaeological XRF Laboratory

Report Prepared for

David Lewandowski
Department of Anthropology
Northern Arizona University
Flagstaff, Arizona

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INTRODUCTION

The analysis here of 47 obsidian artifacts from Pithouse through Late Classic Period sites in the Eagle Creek basin indicates a dominance of locally procured Cow Canyon (n=46) obsidian for tool production. Two of the artifacts were produced from two of the localities in the Mule Creek source area in western New Mexico, or from secondary deposits along the Gila River south of these sites.

LABORATORY SAMPLING, ANALYSIS AND INSTRUMENTATION

All archaeological samples are analyzed whole. The results presented here are quantitative in that they are derived from "filtered" intensity values ratioed to the appropriate x-ray continuum regions through a least squares fitting formula rather than plotting the proportions of the net intensities in a ternary system (McCarthy and Schamber 1981; Schamber 1977). Or more essentially, these data through the analysis of international rock standards, allow for inter-instrument comparison with a predictable degree of certainty (Hampel 1984; Shackley 2011).

All analyses for this study were conducted on a ThermoScientific *Quant'X* EDXRF spectrometer, located in the Archaeological XRF Laboratory, Albuquerque, New Mexico. It is equipped with a thermoelectrically Peltier cooled solid-state Si(Li) X-ray detector, with a 50 kV, 50 W, ultra-high-flux end window bremsstrahlung, Rh target X-ray tube and a 76 μm (3 mil) beryllium (Be) window (air cooled), that runs on a power supply operating 4-50 kV/0.02-1.0 mA at 0.02 increments. The spectrometer is equipped with a 200 l min^{-1} Edwards vacuum pump, allowing for the analysis of lower-atomic-weight elements between sodium (Na) and titanium (Ti). Data acquisition is accomplished with a pulse processor and an analogue-to-digital converter. Elemental composition is identified with digital filter background removal, least squares empirical peak deconvolution, gross peak intensities and net peak intensities above background.

The analysis for mid Zb condition elements Ti-Nb, Pb, Th, the x-ray tube is operated at 30 kV, using a 0.05 mm (medium) Pd primary beam filter in an air path at 200 seconds livetime to generate x-ray intensity Ka-line data for elements titanium (Ti), manganese (Mn), iron (as Fe_2O_3^T), cobalt (Co), nickel (Ni), copper, (Cu), zinc, (Zn), gallium (Ga), rubidium (Rb), strontium (Sr), yttrium (Y), zirconium (Zr), niobium (Nb), lead (Pb), and thorium (Th). Not all these elements are reported since their values in many volcanic rocks are very low. Trace element intensities were converted to concentration estimates by employing a least-squares calibration line ratioed to the Compton scatter established for each element from the analysis of international rock standards certified by the National Institute of Standards and Technology (NIST), the US. Geological Survey (USGS), Canadian Centre for Mineral and Energy Technology, and the Centre de Recherches Pétrographiques et Géochimiques in France (Govindaraju 1994). Line fitting is linear (XML) for all elements but Fe where a derivative fitting is used to improve the fit for iron and thus for all the other elements. When barium (Ba) is analyzed in the High Zb condition, the Rh tube is operated at 50 kV and up to 1.0 mA, ratioed to the bremsstrahlung region (see Davis 2011; Shackley 2011). Further details concerning the petrological choice of these elements in Southwest obsidians is available in Shackley (1988, 1995, 2005; also Mahood and Stimac 1991; and Hughes and Smith 1993). Nineteen specific pressed powder standards are used for the best fit regression calibration for elements Ti-Nb, Pb, Th, and Ba, include G-2 (basalt), AGV-2 (andesite), GSP-2 (granodiorite), SY-2 (syenite), BHVO-2 (hawaiite), STM-1 (syenite), QLO-1 (quartz latite), RGM-1 (obsidian), W-2 (diabase), BIR-1 (basalt), SDC-1 (mica schist), TLM-1 (tonalite), SCO-1 (shale), NOD-A-1 and NOD-P-1 (manganese) all US Geological Survey standards, NIST-278 (obsidian), U.S. National Institute of Standards and Technology, BE-N (basalt) from the Centre de Recherches Pétrographiques et

Géochimiques in France, and JR-1 and JR-2 (obsidian) from the Geological Survey of Japan (Govindaraju 1994).

The data from the WinTrace software were translated directly into Excel for Windows software for manipulation and on into SPSS for Windows for statistical analyses. In order to evaluate these quantitative determinations, machine data were compared to measurements of known standards during each run. RGM-1 a USGS obsidian standard is analyzed during each sample run for obsidian artifacts to check machine calibration (Table 1). Source assignments were made by reference to Shackley (1995, 1998a, 2005; see Table 1 and Figure 1) and source standard data at this lab.

DISCUSSION

All but two of the artifacts were produced from one of the closely related chemical groups known as Cow Canyon (Shackley 2005). The other two artifacts were produced from two of the chemical groups at Mule Creek in western New Mexico; Antelope Creek and Mule Mountains. One of these artifacts is the basal portion of a straight based side-notched point typical of Late Classic points at Mule Creek and Point of Pines (Shackley 2010, 2012; see Figure 2 here). A short discussion of the Cow Canyon geology and archaeology follows (see also Shackley 1990, 2005). While the nodule sizes are small at Cow Canyon, it has been a preferred raw material since Paleoindian times given its clarity, great flaking character, and availability over a large area.

The Cow Canyon Source

It appears that the secondary depositional extent of this source is much greater than originally mapped (Shackley 1988). This Tertiary source is eroding east into the Blue River, south into the San Francisco River and west into the Gila River as originally noted, but also erodes in much higher density west into Eagle Creek west of the Blue River, and on south into

the Gila River and up to 20 km south into the San Simon River Valley. A number of “pockets” of Cow Canyon glass have been located in the San Simon River Valley in Pliocene/Pleistocene lacustrine sediments of the 111 Ranch Formation the result of considerably higher sedimentation rates in these periods than currently (Figure 3). The density of nodules at the upper end of Eagle Creek approximately 15 km west of the primary contexts at Cow Canyon is up to 1 per 10m², and less than 100 times lower in the Gila River (Shackley 1992a). Again, these nodules are mixed with the Mule Creek marekanites in the alluvium. Nodules up to 5 cm in diameter have been recovered in the 111 Ranch Formation, as large as those at the primary source to the north, suggesting that the sediment load during the Plio-Pleistocene was very great.

Recent unpublished analyses by the Berkeley Archaeological XRF Lab of over 300 obsidian samples from Late Middle Archaic to Early Agricultural contexts at McEuen Cave (AZ W:13:6 ASM) in the Gila Mountains about 30 km north of the Gila River at Geronimo, indicates that Cow Canyon obsidian comprises over 80% of the assemblage. Most, if not all of the raw nodules were procured from the Gila River alluvium in the Safford and San Carlos Valleys.

Cow Canyon obsidian was also the only obsidian source known from the Murray Springs Clovis site in southern Arizona. Two of the fluted points were produced from this glass, likely from nodules around 5 cm in diameter, and probably through bipolar reduction. This is the earliest indication of both obsidian procurement and bipolar technology in the southern Southwest (Shackley huckell and haynes).

The “primary source” is located in Apache National Forest, central Greenlee County, Arizona. This is a relatively small primary source along and east of Arizona Hwy 191 located in a Tertiary rhyolite body. The nodules are found within an eroded rhyolite/ash unit that appears to be a remnant dome structure. Perlite or vitrophyre was not evident. The nodules are common in a rhyolite regolith on top of the dome as well as the rhyolite/ash alluvium at the base of the

slopes. Nodules up to 5 cm in diameter are available, but most of the nodules are near or less than 4 cm. The density of nodules is fairly high in places, up to 5 per m². Cortex is mainly a thin gray-black. The interior glass is aphyric and the color is as variable as the mid-Tertiary marekanite sources. The most common color/opacity is a near transparent brown-green sometimes with thin banding. A few specimens exhibit a nearly opaque aphyric gray-green banded megascopic character, and very few have mahogany blotches.

The nodules are eroding and funneled into the Cow Canyon Creek/Turkey Creek system possibly reaching the Blue River 15 km east. Reduced nodules and flakes occur everywhere on the regolith and in the alluvium. The pattern of reduction is similar to Vulture with sporadic bipolar reduction throughout the source area. These specific areas may reach 10-20 per m². Overall the rejected core/flake density is less than 1 per 5m². There are no published sources on this locality other than the county geology map (Wilson and Moore 1958), and Shackley (1988, 1995, 2005).

As evident in this assemblage, Cow Canyon as a source is slightly bimodal in Sr (see Figure 1). Given the Tertiary age of the source and corresponding extensive secondary deposition of the marekanites, no are has been found that correlates with those groups. They are certainly closely magmatically related, and isn't really relevant archaeologically.

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Table 1. Elemental concentrations and source assignments for the archaeological specimens, and analyses of USGS RGM-1 obsidian standard. All measurements in parts per million (ppm).

Sample	Ti	Mn	Fe	Rb	Sr	Y	Zr	N b	Pb	Th	Source
Pithouse Period											
1	141 0	44 0	1021 1	13 7	13 3	20	12 3	13	20	15	Cow Canyon
2	138 8	43 0	1015 5	13 9	14 0	19	12 5	13	20	10	Cow Canyon
3	137 3	46 0	9851	14 2	13 5	18	12 7	16	20	18	Cow Canyon
4	122 6	45 5	9262	13 2	11 2	21	12 6	18	16	11	Cow Canyon
5	122 2	41 6	9356	13 2	13 1	21	11 6	15	18	16	Cow Canyon
6	141 9	53 0	1051 3	14 5	12 0	26	13 3	21	19	13	Cow Canyon
7	132 7	40 5	9349	13 3	13 4	21	12 4	13	22	21	Cow Canyon
8	132 2	54 3	1017 8	14 6	12 0	27	13 3	21	24	15	Cow Canyon
9	140 9	59 1	1056 1	15 4	12 3	28	13 7	21	21	11	Cow Canyon
10	128 0	39 1	9398	13 3	13 3	17	11 7	13	20	14	Cow Canyon
11	141 2	53 0	1044 4	14 3	11 9	22	13 7	20	19	15	Cow Canyon
12	128 0	42 0	9515	14 0	13 5	21	12 1	16	19	11	Cow Canyon
13	121 9	42 5	9089	13 2	12 9	17	11 6	16	20	14	Cow Canyon
14	138 3	41 4	9569	12 8	13 0	20	12 0	14	20	11	Cow Canyon
15	141 7	43 0	9807	13 8	13 4	17	12 0	17	20	18	Cow Canyon
16	138 5	48 1	1003 1	13 9	14 0	20	12 7	16	19	11	Cow Canyon
17	146 1	42 6	9985	13 9	13 3	16	12 2	11	22	14	Cow Canyon
18	138 5	43 3	9743	13 8	13 4	18	12 0	18	21	14	Cow Canyon
19	143 7	48 8	1034 1	14 8	14 0	21	12 3	14	24	15	Cow Canyon
20	147 2	47 5	1035 8	13 6	13 2	18	11 6	11	21	18	Cow Canyon
21	131 8	45 1	9788	13 8	13 6	19	11 6	14	22	17	Cow Canyon
22	163 8	55 9	1079 4	14 9	14 6	22	12 7	15	20	12	Cow Canyon
23	168 3	49 0	1161 5	14 8	14 1	19	12 1	15	24	15	Cow Canyon
24	139 5	47 4	9135	14 5	10 6	19	91	11	20	14	Cow Canyon

26	131 3	47 4	9381	13 5	11 3	21	12 7	21	19	8	Cow Canyon
27	131 1	42 2	9469	13 2	13 0	20	11 4	15	19	19	Cow Canyon
28	135 4	51 7	1014 2	14 4	12 0	24	13 5	23	23	20	Cow Canyon
29	108 5	41 2	8071	16 3	13	23	10 0	30	19	28	Mule Mtns/Mule Creek
30	134 3	43 7	9636	13 7	13 5	18	11 5	14	19	18	Cow Canyon
31	171 8	42 4	1148 1	14 0	13 2	17	12 1	12	19	15	Cow Canyon
Classic Period											
32	129 7	44 1	9587	13 6	13 8	24	11 9	13	16	11	Cow Canyon
33	107 3	36 9	1044 8	24 2	24	42	11 2	21	28	28	Antelope Cr/Mule Creek
34	129 7	44 8	9362	13 5	13 1	21	11 6	14	20	18	Cow Canyon
36	134 7	47 0	9888	14 4	13 8	18	11 7	14	18	16	Cow Canyon
37	133 1	44 0	9897	13 9	13 6	19	11 6	14	19	10	Cow Canyon
39	146 7	43 8	1041 3	14 1	13 7	16	11 7	11	19	14	Cow Canyon
40	130 2	53 6	1027 4	14 5	11 9	28	13 0	17	20	10	Cow Canyon
41	139 7	44 5	9951	13 8	13 4	16	11 6	13	20	12	Cow Canyon
42	135 8	43 1	9474	13 7	13 1	19	12 4	8	18	7	Cow Canyon
43	133 7	44 0	9547	13 9	13 2	20	11 8	11	23	14	Cow Canyon
44	124 8	51 1	9526	14 0	11 2	21	12 4	16	17	16	Cow Canyon
45	138 5	48 0	9939	14 5	13 8	20	12 4	15	22	11	Cow Canyon
Sample	Ti	Mn	Fe	Rb	Sr	Y	Zr	N b	Pb	Th	Source
46	133 1	57 3	1012 2	14 6	11 5	22	13 0	17	21	14	Cow Canyon
47	129 7	45 2	9647	14 3	13 9	16	12 4	15	19	12	Cow Canyon
48	132 0	43 4	9703	13 1	13 5	18	12 0	14	17	10	Cow Canyon
49	126 7	53 3	9806	14 0	11 7	23	13 0	14	17	9	Cow Canyon
50	135 9	52 7	1033 4	15 2	12 3	23	13 1	21	22	13	Cow Canyon
RGM1-S4	161 9	27 7	1329 7	14 7	10 8	24	21 6	8	21	12	standard
RGM1-S4	139 2	44 8	9588	13 8	13 3	16	11 6	12	21	9	standard
RGM1-S4	160 3	29 4	1334 1	14 6	10 9	23	21 0	5	17	15	standard

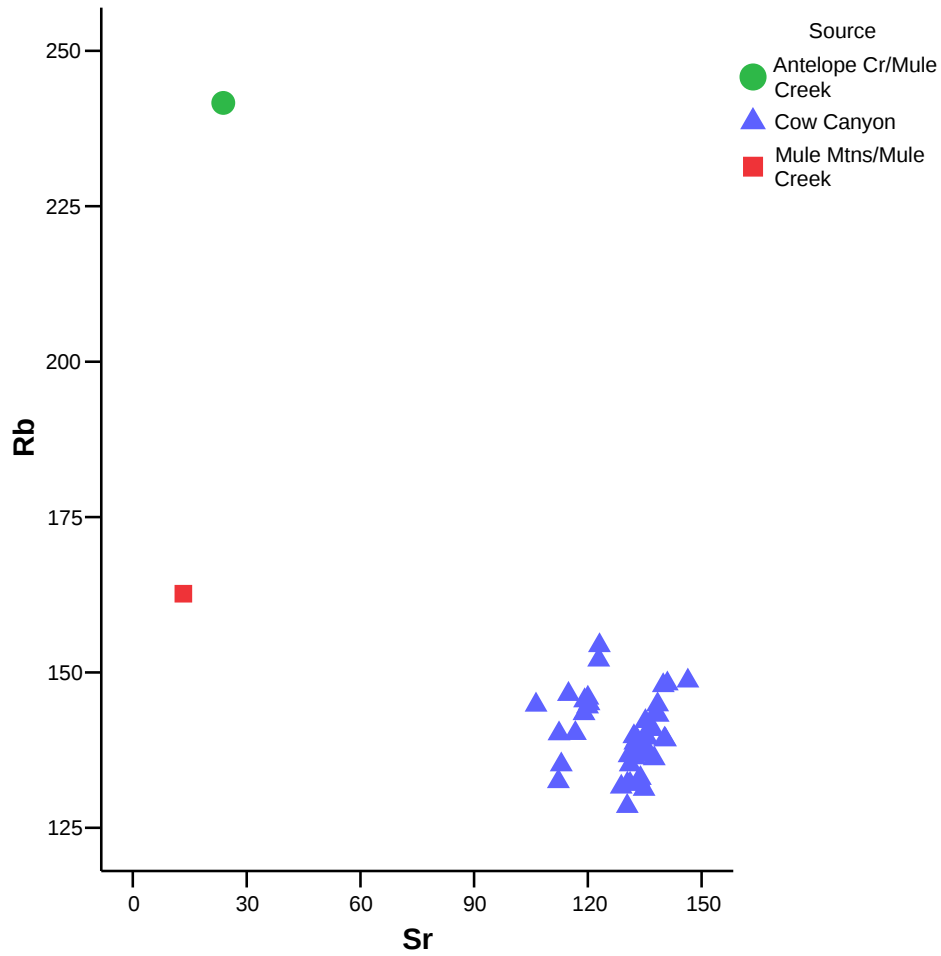


Figure 1. Sr versus Rb bivariate plot of the elemental concentrations for the archaeological specimens. Many of the Eagle Creek artifacts are more similar to the higher Sr group.



Figure 2. Projectile point forms from Point of Pines (from Shackley 2012).

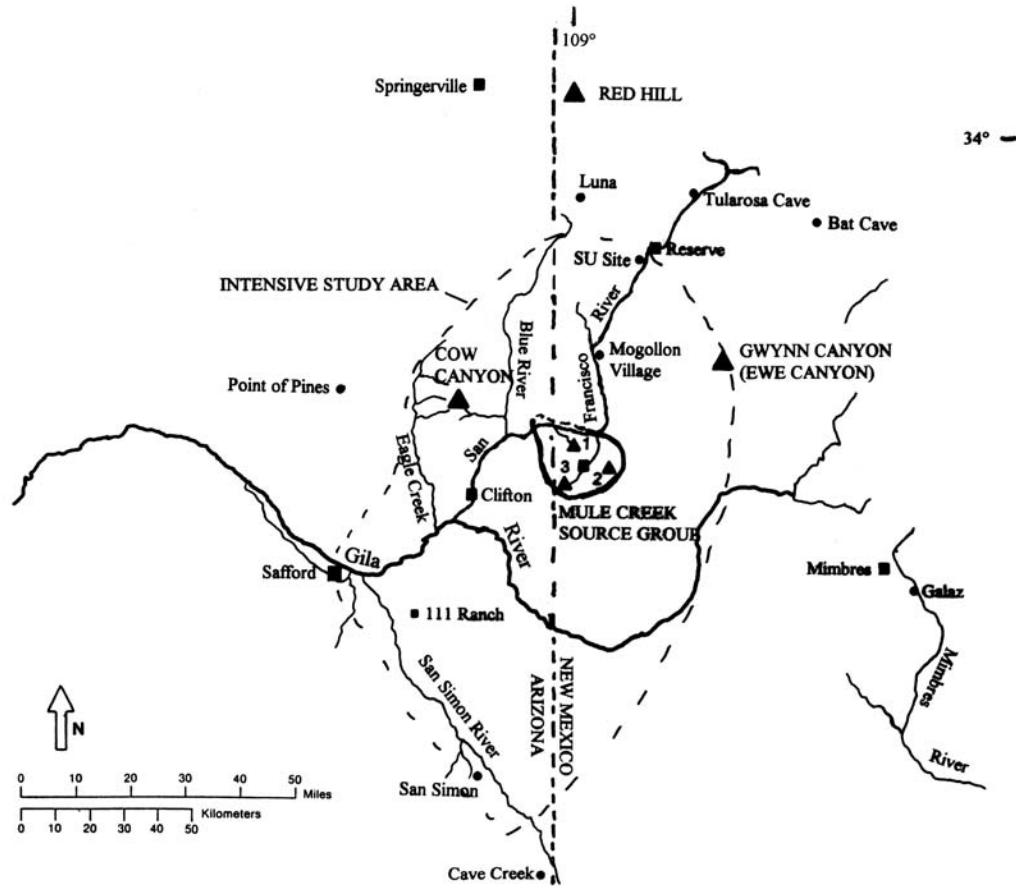


Figure 3. Location of Mogollon-Datil Volcanic Province archaeological obsidian sources and Tertiary Period Cow Canyon as well as major drainage basins (from Shackley 1998b, 2005).