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ECO2N V2.0: A TOUGH2 fluid property module for modeling CO₂-H₂O-NaCl systems to elevated temperatures of up to 300°C

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Abstract

We have improved ECO2N, the TOUGH2 fluid property module of the CO₂-H₂O-NaCl system. The major enhancements include: (i) the upper temperature limit is increased from 110 to about 300°C; (ii) the thermophysical properties of the CO₂-rich phase are more accurately calculated as a non-ideal mixture of CO₂ and H₂O; (iii) the approach to calculate the specific enthalpy of dissolved CO₂ has been improved to make the code more robust in modeling phase transitions under non-isothermal conditions; and (iv) more sophisticated models for effective heat conductivity of formations saturated with supercritical CO₂ have been provided. The new module includes a comprehensive description of the thermodynamic and thermophysical properties of H₂O-NaCl-CO₂ mixtures, that reproduces fluid properties largely within experimental error for the temperature, pressure and salinity conditions 10°C < T < 300°C, P < 600 bar, and salinity up to halite saturation. This includes density, viscosity, and specific enthalpy of fluid phases as functions of temperature, pressure, and composition, as well as partitioning of mass components H₂O, NaCl and CO₂ among the different phases. ECO2N with the TOUGH2 reservoir simulator can be applied to a wide range of problems in geologic sequestration of CO₂ in saline aquifers, and in enhanced geothermal reservoirs. ECO2N can describe both sub- and supercritical states of CO₂, but applications that involve subcritical conditions are limited to systems in which there is no change of phase between liquid and gaseous CO₂, and in which no mixtures of liquid and gaseous CO₂ occur. © 2016 Society of Chemical Industry and John Wiley & Sons, Ltd.

Introduction

TOUGH2/ECO2N¹⁻³ is a widely used numerical simulator for investigating problems involved in CO₂ geological sequestration. However, the previous version (V1.0) of ECO2N is only applicable to systems with temperature up to 110°C.⁴ As a result, many problems that involve higher temperature could not be simulated by the code. Motivated by the potential for enhanced

geothermal systems (EGS) operating with CO₂, and CO₂ geologic storage at temperatures above 100°C, Spycher and Pruess⁵ developed new mutual solubility correlations for higher temperatures. ECO2N V2.0 is an enhanced version of ECO2N V1.0 that inherits all the capabilities of ECO2N V1.0 and expands the applicable temperature range up to about 300°C by incorporating the newly developed mutual solubility correlations for higher temperature.⁵ The fluid property module can be used to model non-isothermal multiphase flow in the system H₂O-NaCl-CO₂. TOUGH2/ECO2N V2.0 represents fluids as consisting of two phases: a water-rich aqueous phase, hereafter referred to as aqueous or ‘liquid’; and a CO₂-rich phase, hereafter referred to as ‘gas’. In addition, solid salt may also be present. The only chemical reactions modeled by ECO2N V2.0 are equilibrium phase partitioning of water and carbon dioxide between the liquid and gas phases, and precipitation and dissolution of solid salt. The partitioning of H₂O and CO₂ between liquid and gas phases is modeled as a function of temperature, pressure, and salinity, using the recently developed correlations of Spycher and Pruess.^{5, 6} Dissolution and precipitation of salt are treated by means of local equilibrium solubility. Associated changes in fluid porosity and permeability may also be modeled. All phases – gas, liquid, solid – may appear or disappear in any grid block during the course of a simulation. Thermodynamic conditions covered include a temperature range from about 10°C to 300°C (approximately), pressures up to 600 bar, and salinity up to NaCl (halite) saturation. Note that ECO2N can describe both subcritical and supercritical states of CO₂, but applications that involve subcritical conditions are limited to systems in which there is no change of phase between liquid and gaseous CO₂, and in which no mixtures of liquid and gaseous CO₂ occur. For those cases, a user may use the fluid property module ECO2M⁷ instead.

ECO2N V2.0 uses the same framework for describing the thermophysical status of H₂O-NaCl-CO₂ systems as its preceding version (Table 1). This paper will only describe the new enhancements which are not available in V1.0.

Table 1. Summary of ECO2N V2.0 thermophysical capabilities.

Components: #1: water, #2: NaCl, #3: CO ₂
Parameter choices
(NK, NEQ, NPH, NB) ^a

= (3, 4, 3, 6) water, NaCl, CO₂, non-isothermal (default)

= (3, 3, 3, 6) water, NaCl, CO₂, isothermal

molecular diffusion can be modeled by setting NB = 8

Primary variables

single fluid phase (only aqueous, or only gas) \underline{h} (P, Xsm, X3, T)

P – pressure (Pa)

Xsm – NaCl salt mass fraction Xs (on the basis of a two-component, CO₂-free water-salt system), or solid NaCl saturation Ss+10

X3 – CO₂ (true) mass fraction in the aqueous phase, or in the gas phase,

in the three-component system water-salt-CO₂

T – temperature (°C)

two fluid phases (aqueous and gas) \underline{h} (P, Xsm, Sg+10, T)

P – gas phase pressure (Pa)

Xsm – NaCl salt mass fraction Xs (on the basis of a two-component, CO₂-free water-salt system), or solid saturation Ss+10

Sg – gas phase saturation

T – temperature (°C)

- a NK – maximum number of mass components; NEQ – maximum number of equations per grid block; NPH – maximum number of active phases (including solid salt); NB – number of secondary parameters other than component mass fractions.
- b When discussing fluid phase conditions, we refer to the potentially mobile (aqueous and gas) phases only; in all cases solid salt may precipitate or dissolve, adding another active phase to the system.

Code enhancements

Extended CO₂-H₂O solubility model

The upper temperature limit of the H₂O-CO₂ mutual solubility model has been extended from 110°C to about 300°C. In particular, partitioning among co-existing aqueous and gas phases is calculated based on the correlations developed by Spycher and Pruess⁶ for the low temperature range (<99°C) and Spycher and Pruess⁵ for the high temperature range (109°C to ~300°C). At temperatures between 99°C and 109°C, a cubic function is applied to interpolate both the equilibrium mass fraction of CO₂ in the aqueous phase and the equilibrium mass fraction of H₂O in the gas (CO₂-rich) phase. This approach guarantees a smooth transition between the low temperature and the high temperature ranges such that both the solubility function and its first derivative are continuous (Fig. 1).

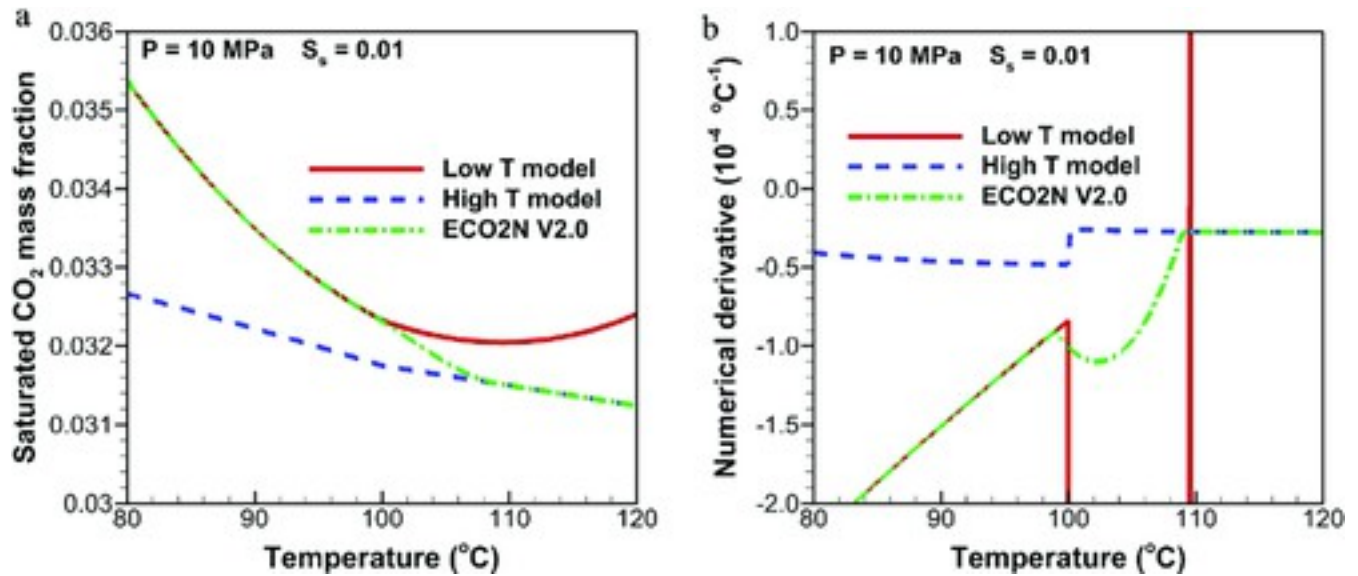


Figure 1

[Open in figure viewer](#)[PowerPoint](#)

Transition between low temperature model (<99°C) and high temperature model (>109°C). (a) Computed dissolved CO₂ mass fraction (at saturation) as a function of temperature; (b) the numerical derivative of the dissolved CO₂ mass fraction with respect to temperature ($\Delta T = 1E-8^{\circ}C^{-1}$). ‘Low T model’ indicates the mutual solubility model developed by Spycher and Pruess,⁶ whereas ‘High T model’ indicates the mutual solubility model for higher temperatures by Spycher and Pruess.⁵ ‘ECO2N V2.0’ indicates the combined model implemented in ECO2N V2.0.

[Caption](#)

Non-ideal gas (CO₂-rich) phase properties

Unlike ECO2N V1.0, which approximates the gas phase properties with the properties of pure CO₂, V2.0 calculates the actual properties of non-ideally mixed gas phase of CO₂ and H₂O. In the gas phase, the CO₂ behaves either as a liquid, gas, or as a supercritical fluid while the water could be considered as water vapor. However, its properties tend to deviate from ‘vapor-like’ and approach ‘liquid-like’ values as the gas phase pressure increases.⁸ At elevated pressures, the H₂O partial pressure in the gas phase can be well above the saturation pressure of pure H₂O, $P_{sat}(T)$. To properly model the effects of H₂O on the properties of the CO₂-rich phase, two new approaches have been implemented in ECO2N V2.0: (i) The use of simple, smooth mixing functions of pure component properties (default option, IE(16) = 0), and (ii) the direct use of the cubic EOS implemented for solubility calculations (IE(16) = 2). For IE(16) = 1, gas phase properties are calculated as in ECO2N V1.0. The following is a brief description of the first new approach (default). The second new approach is discussed in detail by Spycher and Pruess^{5, 8} and will not be repeated here.

The gas phase density ρ_{gas} is calculated as a sum of the partial densities, ρ_{CO_2} and $\rho_{\text{H}_2\text{O}}$:

$$\rho_{\text{gas}} = \rho_{\text{CO}_2} + \rho_{\text{H}_2\text{O}} \quad (1a)$$

The partial densities are calculated as follows:

$$\rho_{\text{CO}_2} = (1 - y_{\text{H}_2\text{O}}) \rho_a (P, T)$$

$$\rho_{\text{H}_2\text{O}} = \rho_{\text{sv}} (P_v, T) + X_L^{1.8} y_{\text{H}_2\text{O}} \rho_{\text{sl}} (P, T) \quad (1b)$$

where $y_{\text{H}_2\text{O}}$ is the mole fraction of H_2O in the gas phase and X_L is taken as zero if the actual partial pressure of water ($P_{\text{H}_2\text{O}}$) is equal to, or less than, the saturation pressure of pure water (P_{sat}^0) at the prevailing temperature, or as $(1 - P_{\text{sat}}^0/P_{\text{H}_2\text{O}})$ if $P_{\text{H}_2\text{O}} > P_{\text{sat}}^0$ (X_L could be viewed as a factor proportional to the fraction of liquid-like H_2O within the total H_2O component). ρ_a and ρ_{sl} are the densities of pure CO_2 and H_2O liquid, respectively, at the prevailing temperature and pressure, whereas ρ_{sv} is the density of pure H_2O vapor at the prevailing temperature but corresponding to $P_v (= \min(P_{\text{H}_2\text{O}}, P_{\text{sat}}^0))$. The calculated densities compare well with the experimental data reported in the literature for various composition, pressure, and temperature (Fig. 2).

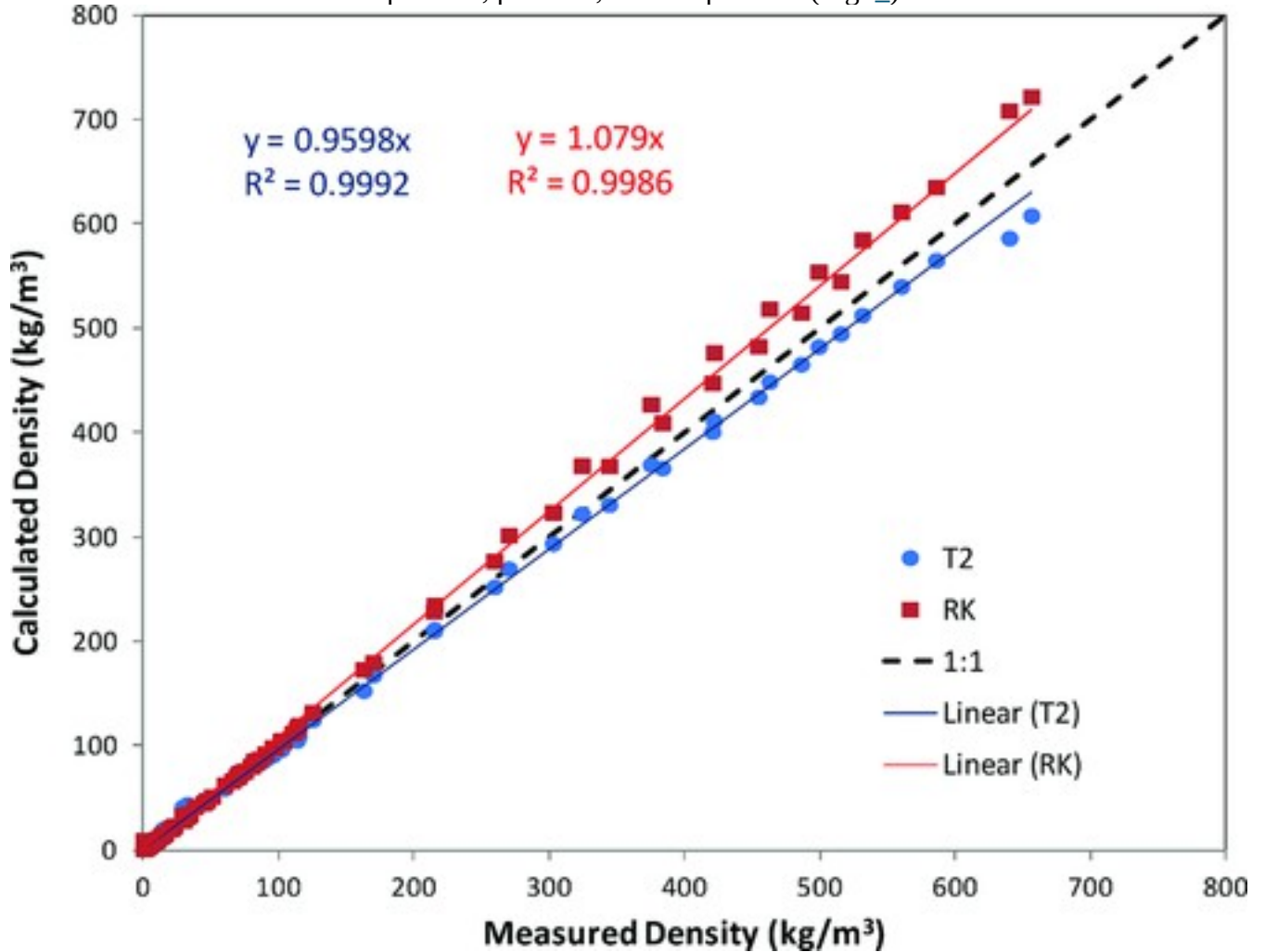


Figure 2

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Comparison of computed densities of the gas phase against the experimental data reported in the literature. 9-13 T2 indicates the default model (Eqn (1)) while RK indicates the alternative model (IE(16) = 2).

[Caption](#)

The gas phase specific enthalpy is also calculated as a sum of contributions from both components plus an empirical mixing heat term:

$$h_{\text{gas}} = y_{\text{H}_2\text{O}} h_{\text{H}_2\text{O}} + (1 - y_{\text{H}_2\text{O}}) h_{\text{CO}_2} + y_{\text{H}_2\text{O}} \left(\frac{P}{\rho_{\text{gas}}} - \frac{P_{\text{H}_2\text{O}}}{\rho_{\text{H}_2\text{O}}} \right) \quad (2)$$

where h_{CO_2} is the specific enthalpy of the CO_2 component, and $h_{\text{H}_2\text{O}}$ is the specific enthalpy of the H_2O component, which is calculated as:

$$h_{\text{H}_2\text{O}} = (1 - X_L) u_{\text{sv}} + X_L u_{\text{sl}} + \frac{P_{\text{H}_2\text{O}}}{\rho_{\text{H}_2\text{O}}} \quad (3)$$

where u_{sv} and u_{sl} are the specific enthalpies of water vapor and liquid water, respectively. The calculated enthalpy values compare well with the experimental data reported in the literature for various compositions, pressures, and temperatures (Fig. 3). Note that the same reference state (i.e., the internal energy of saturated liquid water equals zero at the triple point of pure water) is used in ECO2N V2.0 for both components. As a result, the enthalpy of the CO_2 component in V2.0 is smaller than that in V1.0 by a constant (302 192 J/kg) (unless IE(16) = 1 – when the ECO2N V1.0 formulation for water/ CO_2 properties is used the reference state is unchanged from V1.0).

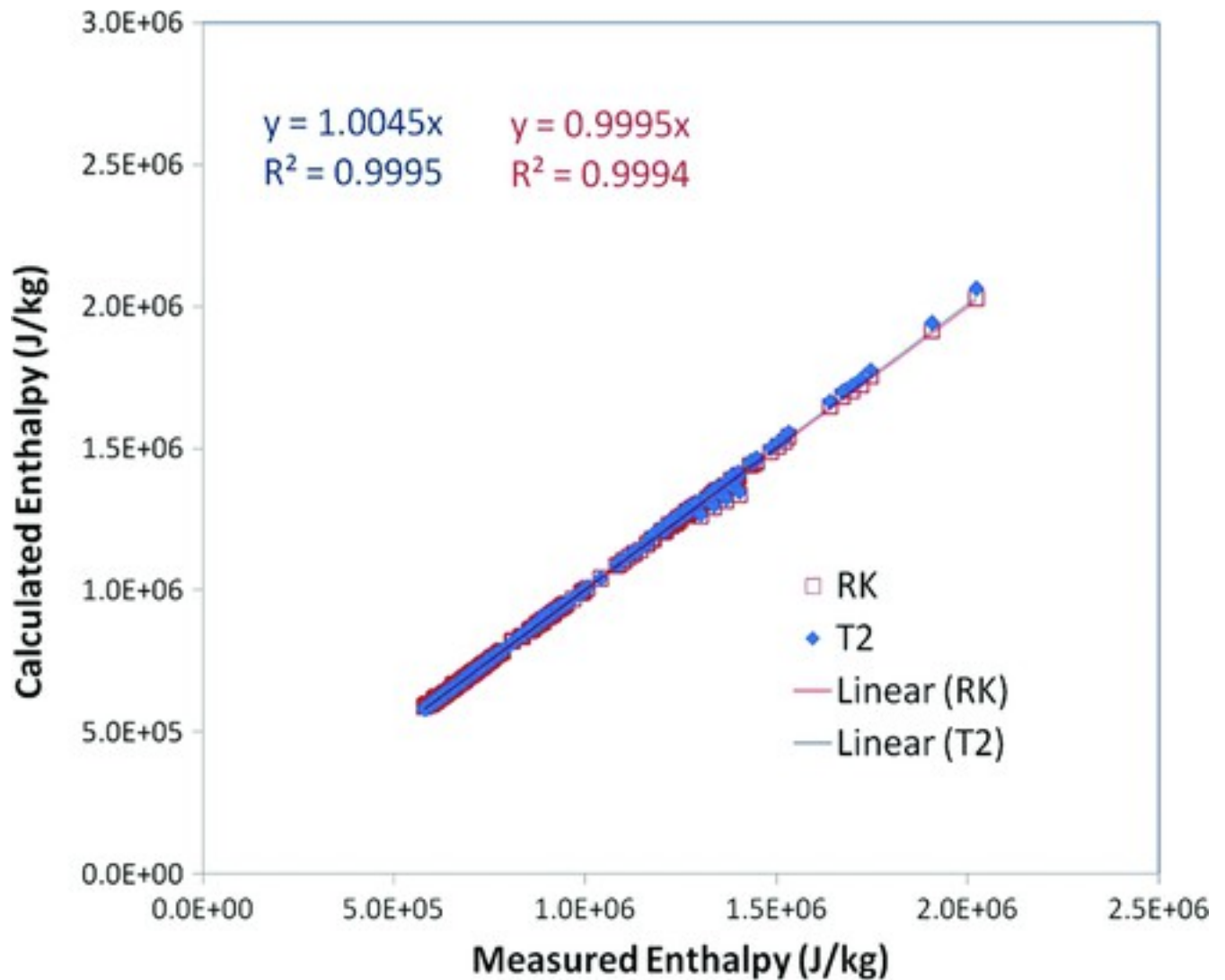


Figure 3

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Comparison of computed specific enthalpy of the gas phase against the experimental data reported in the literature.^{11, 14, 15} T2 indicates the default model while RK indicates the alternative model (IE(16) = 2).

[Caption](#)

The viscosity of the gas phase is calculated based on the fluidity method proposed by Davidson.¹⁶

Non-iterative calculation of specific enthalpy of dissolved CO₂ under single-phase aqueous conditions

In the previous version of ECO2N (V1.0), the specific enthalpy of dissolved CO₂ for the entire range of CO₂ mass fraction (the 3rd primary variable, X3) can be summarized (the subscript g or l was dropped for simplicity) as:

$$h_{\text{CO}_2,\text{aq}}(T, P, X_s)$$

$$= \begin{cases} h_a(T, P) + h_{\text{dis}}(T, X_s) & X_3 \geq X_{\text{CO}_2, \text{eq}} \\ h_a(T, P_{X_3}) + h_{\text{dis}}(T, X_s) & X_3 < X_{\text{CO}_2, \text{eq}} \end{cases} \quad (4)$$

where h_{dis} is the heat of dissolution of CO_2 . $X_{\text{CO}_2,\text{eq}}$ is the equilibrium mass fraction of CO_2 in the aqueous phase at given P , T , and X_s above which free gas occurs. The specific enthalpy of CO_2 , $h_a(T, P)$, is calculated as a function of temperature and pressure through bivariate interpolation from a tabulation of Altunin's¹⁷ correlation. P_{X_3} is the saturation pressure corresponding to the given CO_2 mass fraction in the aqueous phase at given T , which is obtained by inversion of the solubility correlation using an iterative procedure. However, this approach was found often to suffer convergence problems for non-isothermal simulations, especially during appearance or disappearance of a free CO_2 phase. The reason is that the specific enthalpy of dissolved CO_2 calculated using Eqn 4 does not have a continuous first derivative at the phase change line, which could lead to an inaccurate Jacobian matrix and cause convergence problems when the system is close to a phase change. Figure 4 shows an example contour map of the specific enthalpy of dissolved CO_2 (excluding h_{dis} for simplicity) at given temperature (40°C) and salinity (0.01). Above the phase change line, the specific enthalpy only depends on the pressure. Below the phase change line, it only depends on the mass fraction X_3 . Although $h_{\text{CO}_2,\text{aq}}$ is continuous across the phase change line, the partial derivative of $h_{\text{CO}_2,\text{aq}}$ with respect to X_3 is non-zero when approaching the line from below but is zero when approaching the line from above. Similar discontinuities can be found in the partial derivatives with respect to other variables (e.g. P , T , or X_s).

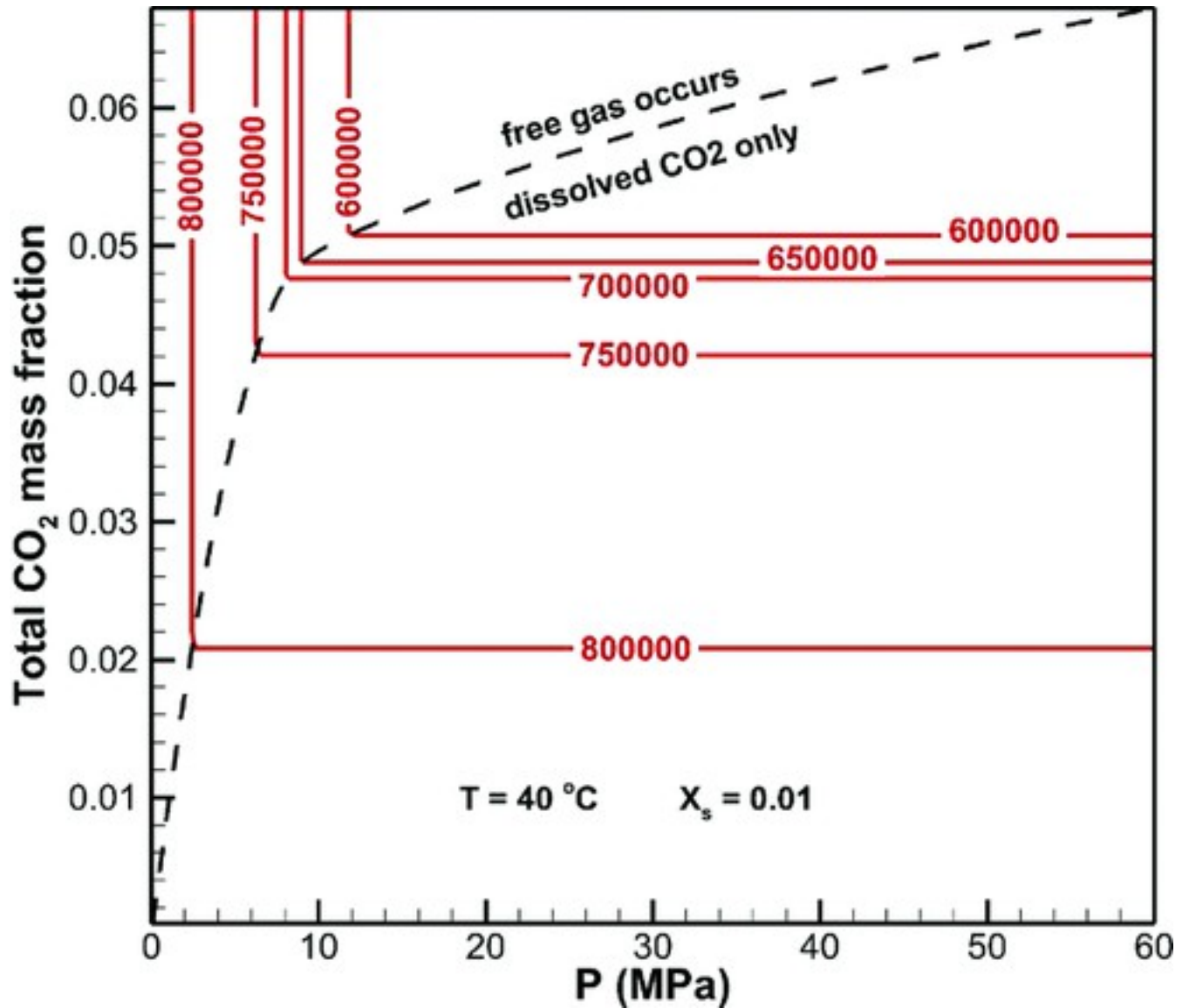


Figure 4

[Open in figure viewerPowerPoint](#)

Contours of the specific enthalpy, $h_{\text{CO}_2,\text{aq}} - h_{\text{dis}}$ (J/kg) of dissolved CO₂ (red lines) as a function of total CO₂ mass fraction X_3 and pressure at a given temperature (40°C) and salt mass fraction (0.01) as calculated in ECO2N V1.0. The phase change line (black dashed line) is calculated using the correlations developed by Spycher and Pruess.⁶ The specific enthalpy is continuous at the phase change line, but its partial derivative with respect to either pressure or CO₂ mass fraction is not.

[Caption](#)

In ECO2N V2.0, we use a modified approach to calculate the enthalpy of dissolved CO₂ so that its derivatives with respect to either primary variable are continuous across the phase change line. We evaluate the specific enthalpy of the dissolved CO₂ for single-phase aqueous conditions as a non-linearly scaled value of its counterpart under two-phase conditions.

$$h_{\text{CO}_2,\text{aq}}(T, P, X_s)$$

$$= \begin{cases} h_a(T, P) + h_{\text{dis}}(T, X_s) & X_3 \geq X_{\text{CO}_2,\text{eq}} \\ h_a(T, P) f(X_3) + h_{\text{dis}}(T, X_s) & X_3 < X_{\text{CO}_2,\text{eq}} \end{cases} \quad (5)$$

where the scaling function $f(X_3)$ is defined as follows:

$$f(X_3) = 2 - \cos \left[3\pi (X_{\text{CO}_2,\text{eq}} - X_3) \right] \quad (6)$$

As a result, ECO2N V2.0 does not suffer the convergence problems caused by the discontinuous first derivatives of the specific enthalpy across the CO₂ saturation line and appears to be more robust than V1.0 especially for non-isothermal applications, even though the calculated enthalpies of dissolved CO₂ are practically identical in V2.0 and V1.0 (Fig. 5). The iterative calculation to obtain P_{x_3} is no longer needed in V2.0 because $f(X_3)$ is a direct function of the third primary variable X_3 .

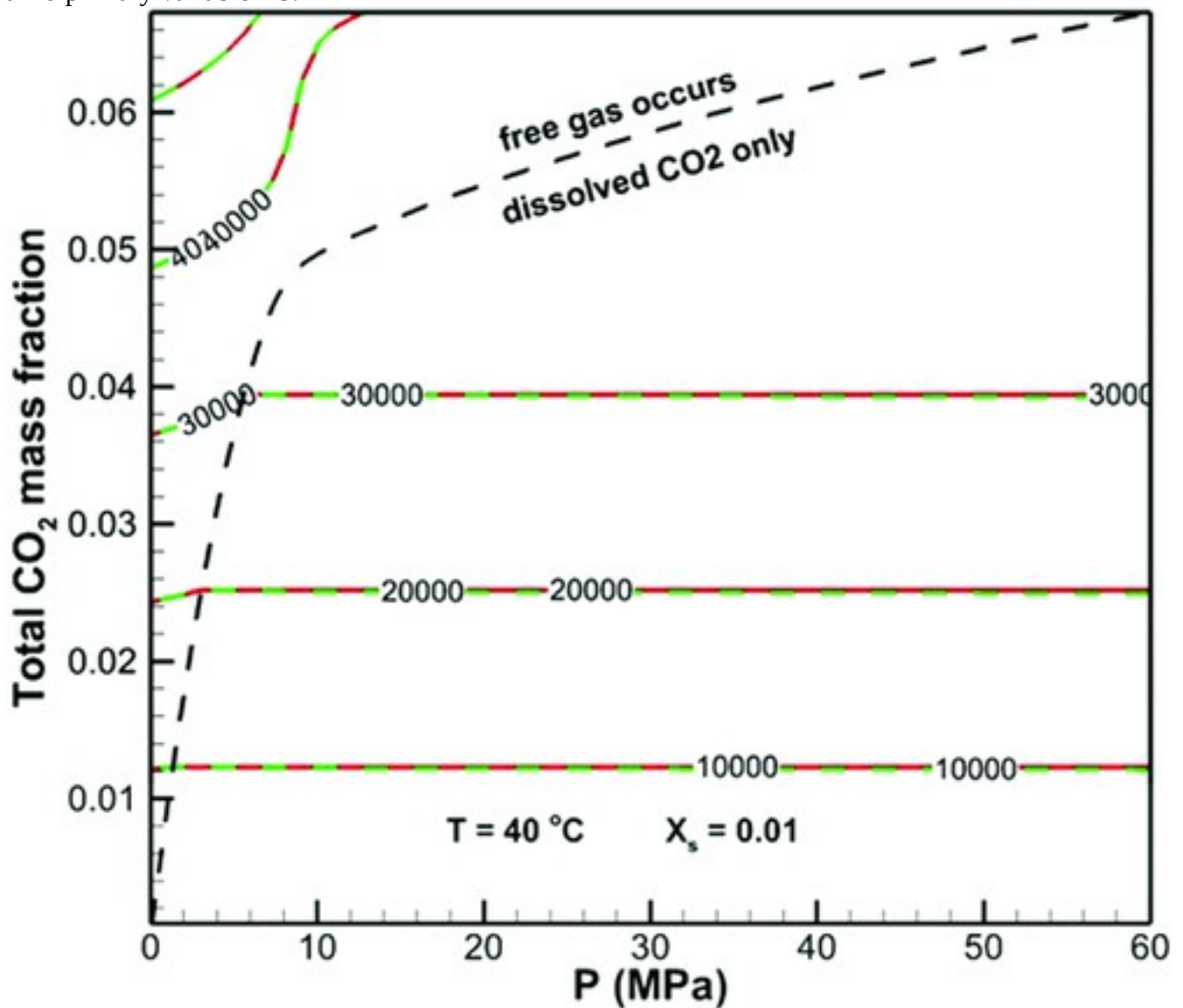


Figure 5

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Contours of the scaled specific enthalpy of dissolved CO₂, $\min(X_3, X_{CO_2,eq})(h_{CO_2,aq} - h_{dis})$, (J/kg) as a function of total CO₂ mass fraction X₃ and pressure at a given temperature (40°C) and salt mass fraction (0.01). V1.0 is the red lines whereas V2.0 is the green dashed lines. The phase change line (black dashed line) is calculated using the correlations developed by Spycher and Pruess.[6](#)

[Caption](#)

Sophisticated models for effective heat conductivity

Two more sophisticated models for effective heat conductivity of formations containing CO₂ have been provided (optional, require some modifications to the TOUGH2 core code) for further improvement of the simulation of CO₂-H₂O-NaCl systems. These are useful because the thermal conductivity of CO₂ varies greatly depending on its occurrence as a gas, a liquid, or a supercritical fluid (Fig. [6](#)). The first new model employs expressions for CO₂[18](#) and brine[19](#), [20](#) thermal conductivity as a function of pressure and temperature, then uses the effective medium theory of Zimmerman[21](#) to combine them with the user-specified rock thermal conductivity to get an effective thermal conductivity for the grid block. The second new model retains the original TOUGH feature of the user specifying the effective thermal conductivity for liquid (aqueous) and gas (CO₂)-saturated conditions, but modifies the CO₂-saturated thermal conductivity in proportion to CO₂ density. The details of models are described in the users' guide.[22](#)

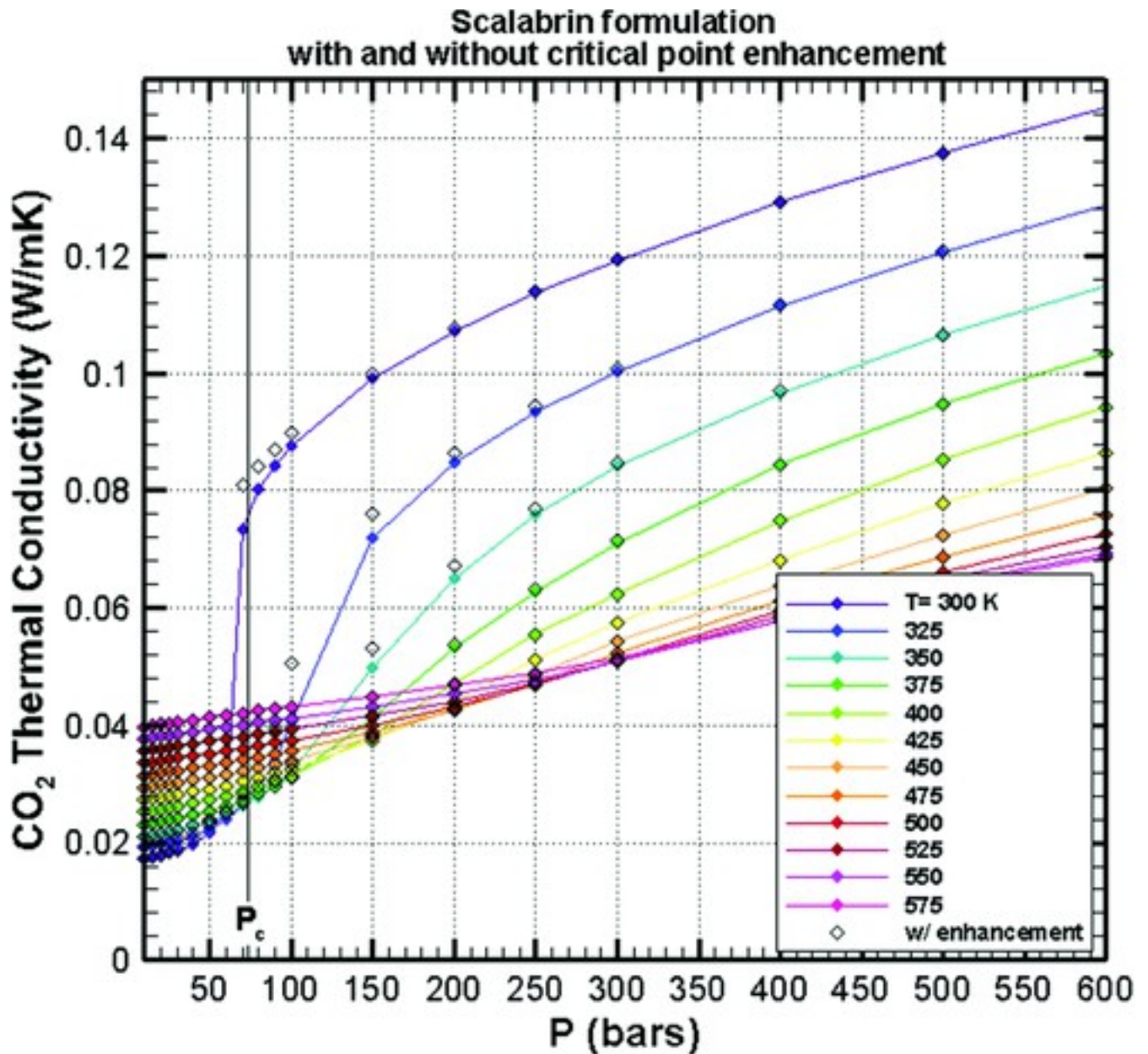


Figure 6

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CO₂ thermal conductivity as a function of temperature and pressure.

[Caption](#)

Example problems

Non-isothermal radial flow from a CO₂ injection well

This is a variation of the test problem #3 in a code inter-comparison project.[23](#), [24](#) The variation is that the flow process is non-isothermal here, i.e., colder CO₂ is injected into a warmer saline aquifer. The problem is formulated as follows. A CO₂ injection well fully penetrates a

homogeneous, isotropic, infinite-acting aquifer of 100 m thickness (Fig. 7), at conditions of 120 bar pressure, 45°C temperature, and a salinity of 15% by weight. Colder CO₂(at 35°C) is injected uniformly at a constant rate of 100 kg/s.

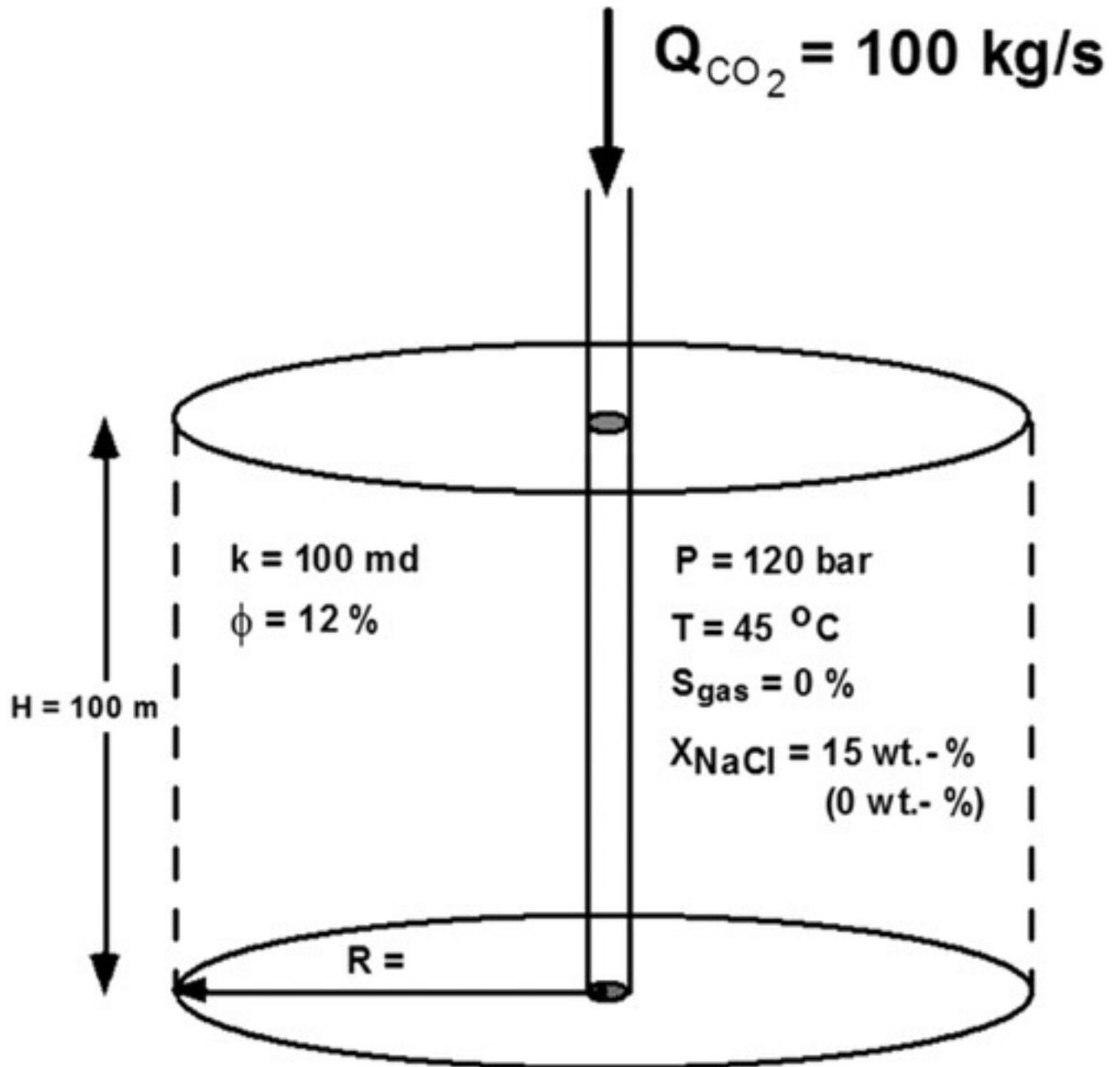


Figure 7

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Schematic of radial flow sample problem.

[Caption](#)

Figure 8 shows the simulated results in terms of the 'similarity variable' R^2/t obtained by O'Sullivan.²⁵ Three distinct regions of gas saturation (Fig. 8(b)), solid salt saturation (Fig. 8(c)),

CO₂ mass fraction in liquid (Fig. 8(d)), and NaCl mass fraction in liquid (Fig. 8(e)), can be found in this non-isothermal CO₂ injection process, namely, dry-out zone, two-phase zone, and single aqueous phase zone. The two-phase zone consists of two sub-regions. In the sub-region near the dry end, the CO₂ mass fraction is higher than near the wet end. The dividing point corresponds to the temperature front formed during injection of colder CO₂ into a warm aquifer (Fig. 9). Behind the front, the temperature is low and more CO₂ can be dissolved in water, while higher temperature and less dissolved CO₂ exist ahead of the front. Interestingly, the temperature in the dry sub-region is slightly lower than the injection temperature, implying that the cooling effect due to water evaporation into the flowing CO₂ is dominating behind the temperature front, whereas the temperature in the wet sub-region is slightly higher than the ambient aquifer temperature, implying that the heating effect due to dissolution of CO₂ into water is dominating ahead of the temperature front (Fig. 8(f)). Note from the temperature profile (Fig. 8(f)) that by accounting for the effects of water in the CO₂-rich phase in the enthalpy calculation, the water-evaporation induced temperature drop predicted by ECO2N V2.0 is smaller than that obtained by the case using pure CO₂ properties for the gas phase (V1.0).

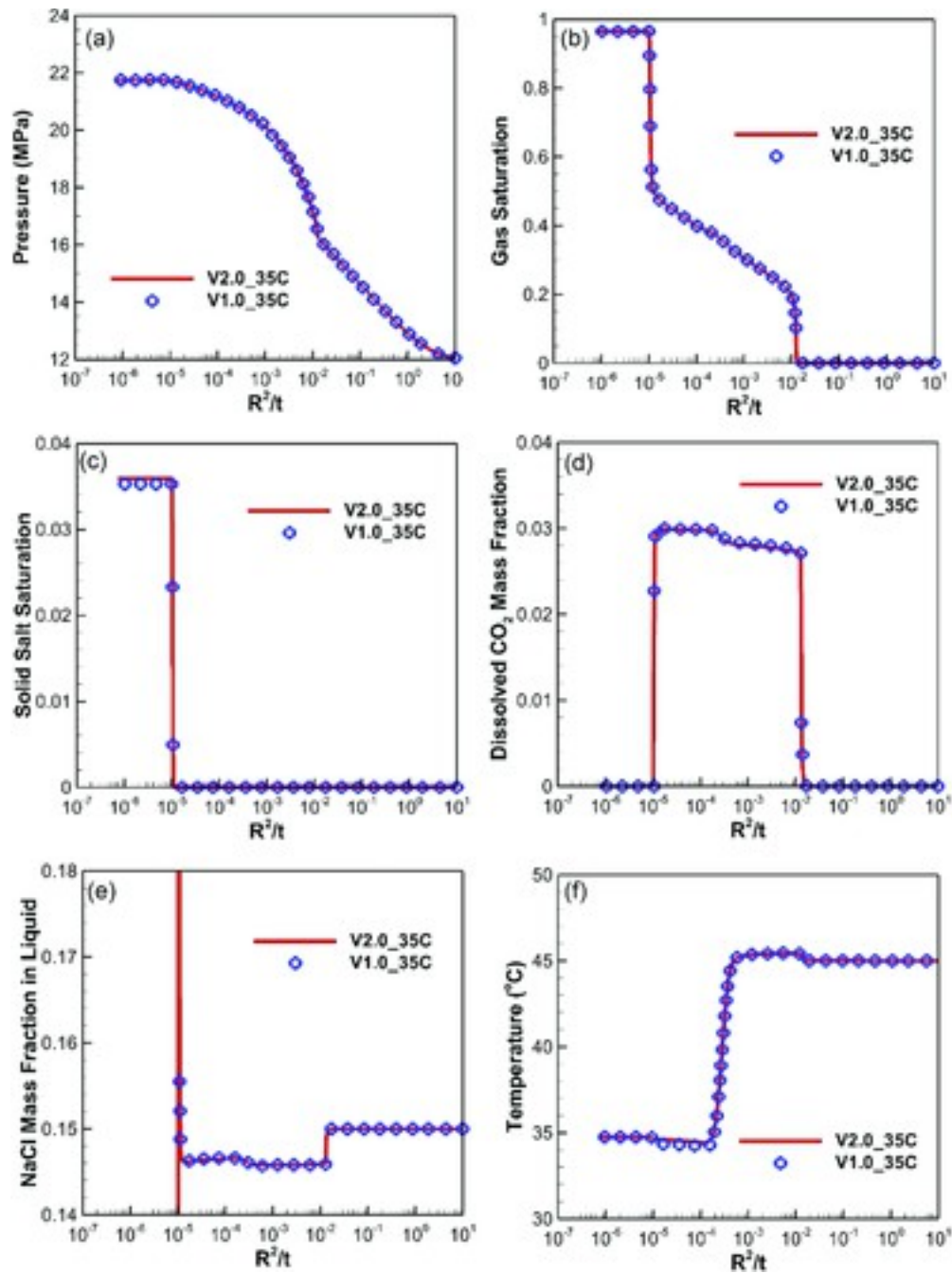


Figure 8

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Simulated pressures (a), gas saturation (b), solid salt saturation (c), dissolved CO_2 mass fraction (d), dissolved NaCl mass fraction (e), and temperature (f) as a function of the similarity variable R^2/t , where R is the radial distance from the well and t is time. The solid red lines represent the results simulated by the new code (ECO2N V2.0), while the blue symbols represent the results simulated by ECO2N V1.0. All results are time series of data for a grid block at a radial distance of $R = 25.25$ m.

[Caption](#)

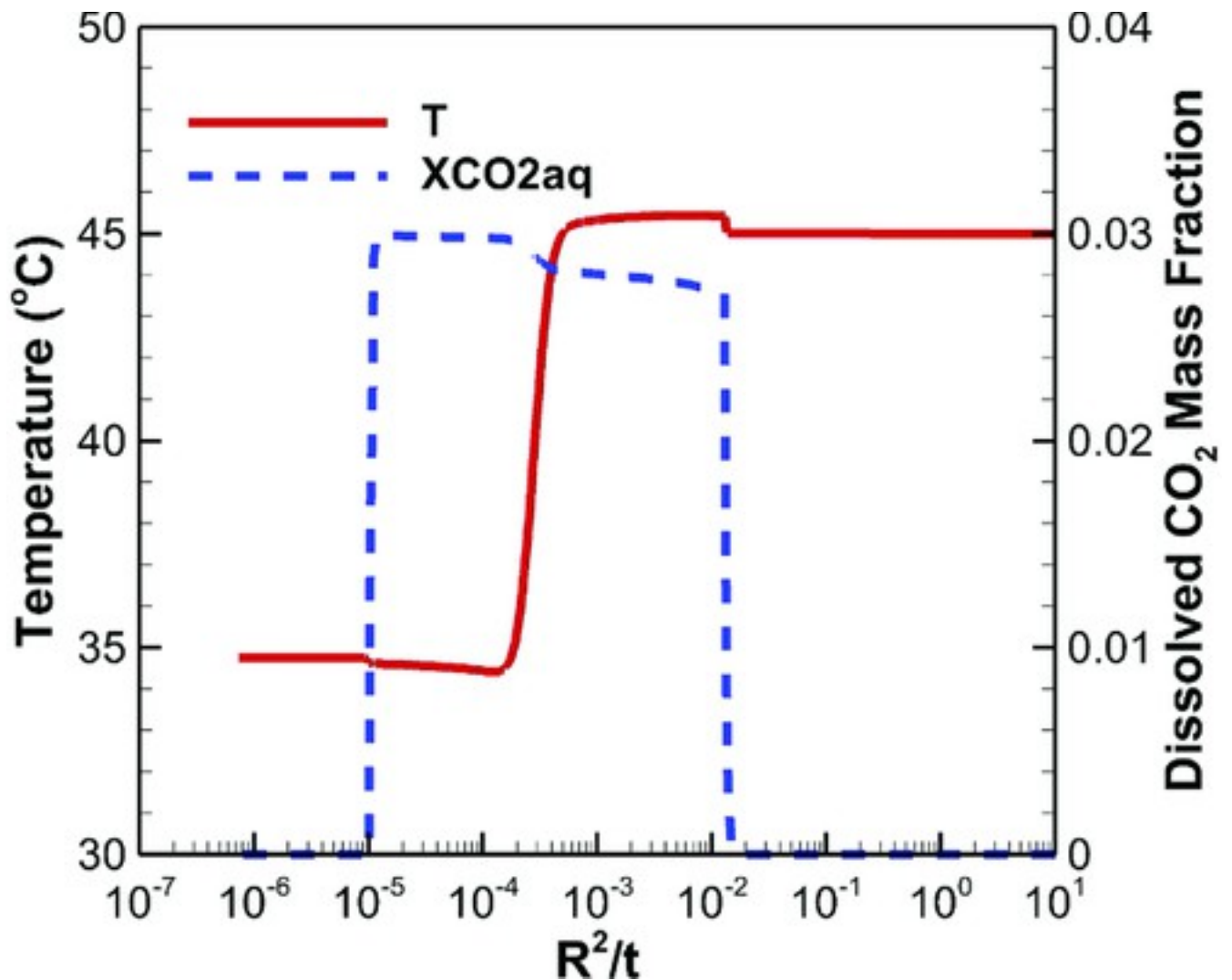


Figure 9

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Simulated temperature and dissolved CO₂ mass fraction as a function of the similarity variable (non-isothermal radial flow). The solid red line represents temperature while the blue dashed line represents the mass fraction of the dissolved CO₂ in aqueous phase. All results are time series of data for a grid block at a radial distance of $R = 25.25$ m.

[Caption](#)

The agreement between ECO2N V2.0 and ECO2N V1.0 is excellent, except for the differences noted, based on more complete physics.

GCS/GHE with a double-porosity reservoir

In this problem, we consider an injection-well/production-well pair representing a 1/8 symmetry element of a five-spot pattern (Fig. 10(a)) that makes up a geothermal heat extraction (GHE) system combined with geological carbon sequestration (GCS). The geothermal reservoir we consider here is an idealized 100 m thick double porosity reservoir,²⁶ whose parameters are

shown in Tables 2 and 3. In the double-porosity model, one continuum represents the higher permeability regions (fractures) through which global flow occurs, while the other represents lower permeability regions (rock matrix), that may exchange fluid and heat with the high-permeability domain locally. The reservoir is assumed to be initially filled with pure CO₂ in the fractures and pure water in the matrix, under the same hydrostatic pressure (29.15 MPa) and temperature (152.2°C). Because the fracture continuum makes up 20% of the reservoir, this initial condition is equivalent to an initial bulk gas saturation of 20%.

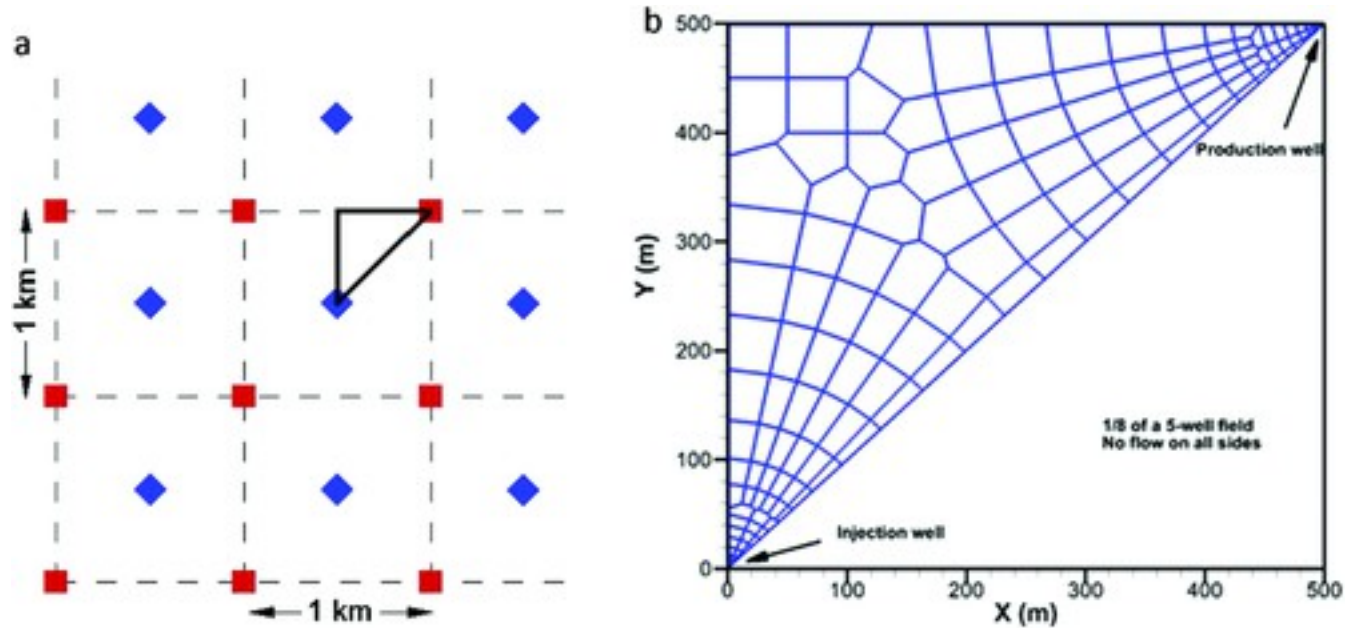


Figure 10

[Open in figure viewerPowerPoint](#)

(a) Diagram of five-spot pattern of geothermal wells (blue-injector; red-producer); (b) map view of the numerical grid used in the simulation. Finer grid resolution is used near the two wells.

[Caption](#)

Table 2. Parameters of the double porosity model.

Parameter	Value
Percentage of fracture pore space (%)	20
Permeability of the fracture continuum (m ²)	2E-14
Permeability of the matrix (m ²)	2E-17

Parameter	Value
Percentage of matrix pore space (%)	80
Fracture/matrix interface area per unit volume (m ² /m ³)	0.2
Characteristic fracture/matrix distance (m)	5.0

Table 3. Other properties of the reservoir (both continua).

Parameter	Value	Note
Porosity	0.254	Uniform
Thermal conductivity	2.51 W m ⁻¹ K ⁻¹	
Pore compressibility	10 ⁻¹⁰ Pa ⁻¹	
Parameters for relative permeability:		
Residual gas saturation	0.01	Liquid relative permeability using van Genuchten-Mualem ²⁷ model and gas relative permeability using Corey ²⁸ model

Parameter	Value	Note
m_{vc}	0.65	
Residual liquid saturation	0.05	
Saturated liquid saturation	1.0	
Parameters for capillary pressure:		
Residual liquid saturation	0.03	Capillary pressure using van Genuchten ²⁷ model
m_{vc}	0.4118	
$1/P_0$	6.08E-5 Pa ⁻¹ (fracture continuum)	
	1.216E-6 Pa ⁻¹ (matrix continuum)	
Maximum capillary pressure	6.4×10 ⁷ Pa	
Saturated liquid saturation	1.0	

A 2D, irregular, dual-continuum grid was created to represent the reservoir, in which each continuum is represented by a 2D mesh having the same geometry (Fig. 10(b)), except that the

matrix continuum mesh does not have global flow connections. The two overlapping meshes are connected locally. In other words, fluid can flow from the injection well to the production well through the fracture continuum only, whereas the matrix continuum plays a passive role through mass and heat exchange with the fractures. Grid resolution varies from 0.1 m near the wells to 50 m at far field to capture the important details of the flow field. Both the injection and the production wells are fully perforated in the reservoir (connected to the fracture continuum only). The parameters for the double porosity model used in this study are shown in Table [2](#).

With the exception of capillary strength, the parameters for relative permeability and capillary functions are the same for both continua, as shown in Table [3](#).

No-flow boundaries are assigned on all sides except for conductive heat flow through the reservoir/basement rock interface, which is calculated using a semi-analytical solution implemented in TOUGH2.[29](#) Injection of CO₂ is simulated as a source term at the injection well cell with a rate of 6.25 kg/s (1/8 of 50 kg/s for the full well) at a constant temperature of 75°C. The same flow rate is assigned for the mass produced at the production well.

Figure [11](#) shows six snapshots of pressure change (from the initial pressure) in the fracture continuum during the production. The reservoir pressure drops quickly at early time and then slowly recovers to some degree. As a result, the pressure drop after one year is the biggest among the six snapshots. This implies that the reservoir pressure loss is mainly caused by the volume imbalance due to production of hot CO₂ and injection of cold CO₂. Such volume loss is gradually compensated by the expansion of the injected cold CO₂ with time. Figure [12](#) shows the temperature distribution in both continua at various times. The cold front advances with time from the injection well to the production well. There is a time-delay in the matrix continuum in such propagation, especially at early time.

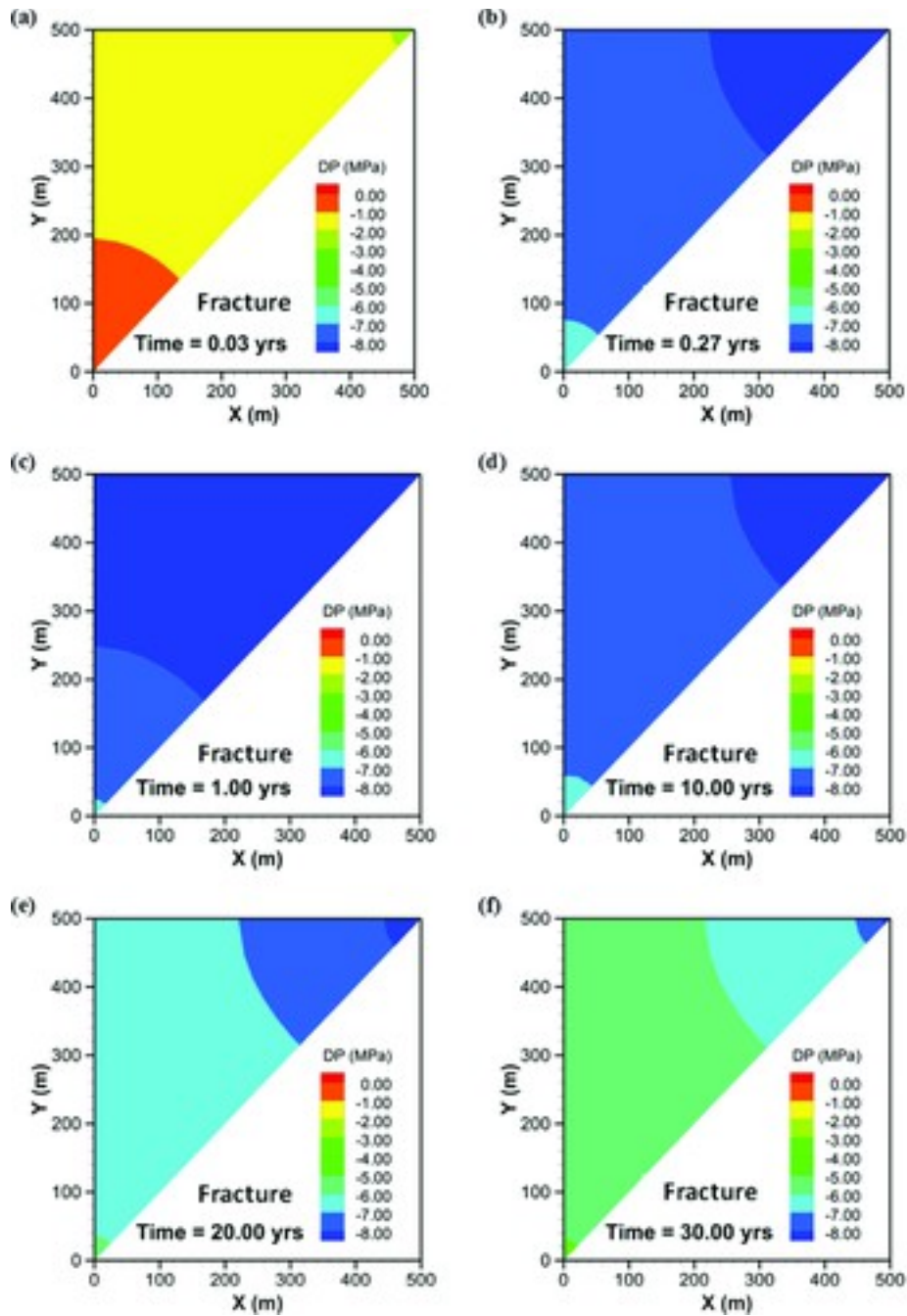


Figure 11

[Open in figure viewerPowerPoint](#)

Simulated pressure drop (from the initial reservoir pressure) in the reservoir (a) 10 days, (b) 100 days, (c) 1 yrs, (d) 10 yrs, (e) 20 yrs, and (f) 30yrs.

[Caption](#)

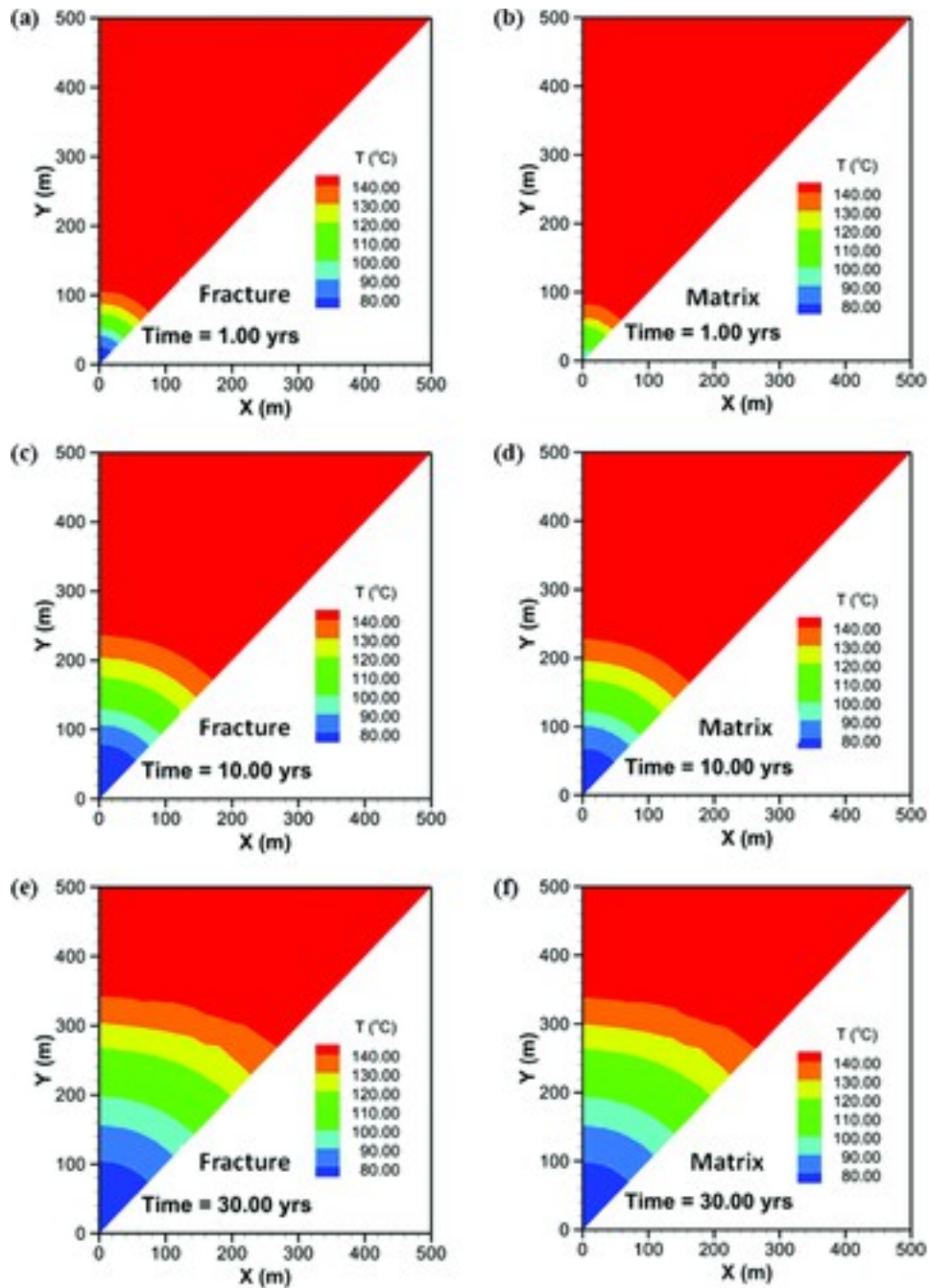


Figure 12

[Open in figure viewerPowerPoint](#)

Simulated temperature in the reservoir after 1 year ((a) and (b)), 10 years ((c) and (d)), and 30 years ((e) and (f)) in the two continua.

[Caption](#)

Figure 13 shows three snapshots of gas saturation in each continuum during production. The gas saturation in the matrix continuum slowly increases with time as CO_2 enters from the fracture continuum. The gas saturation in the fracture continuum first drops over the entire domain and then increases near the injection well as injection continues, forming a significant gradient from

the injection well to the production well. Water accumulates in the region close to the production well (Fig. 13(e)). However, the liquid phase production rate is small for most times (Fig. 14(a)), and the CO₂ component in the total production is larger than 97% (Fig. 14(b)).

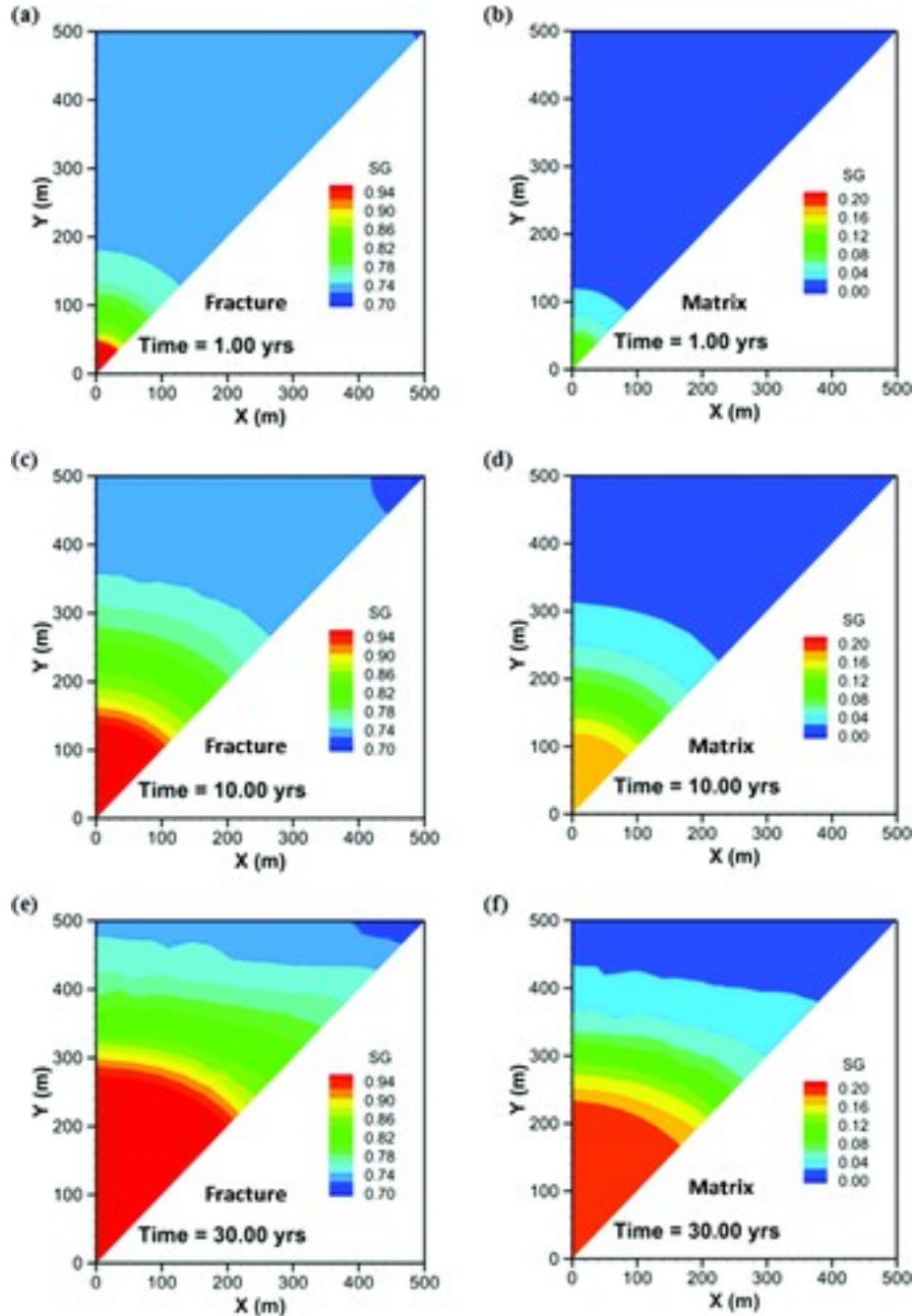


Figure 13

[Open in figure viewer](#) [PowerPoint](#)

Simulated gas saturation in the reservoir after 1 year ((a) and (b)), 10 years ((c) and (d)), and 30 years ((e) and (f)) in the two continua. Different color scales are used for each continuum.

[Caption](#)

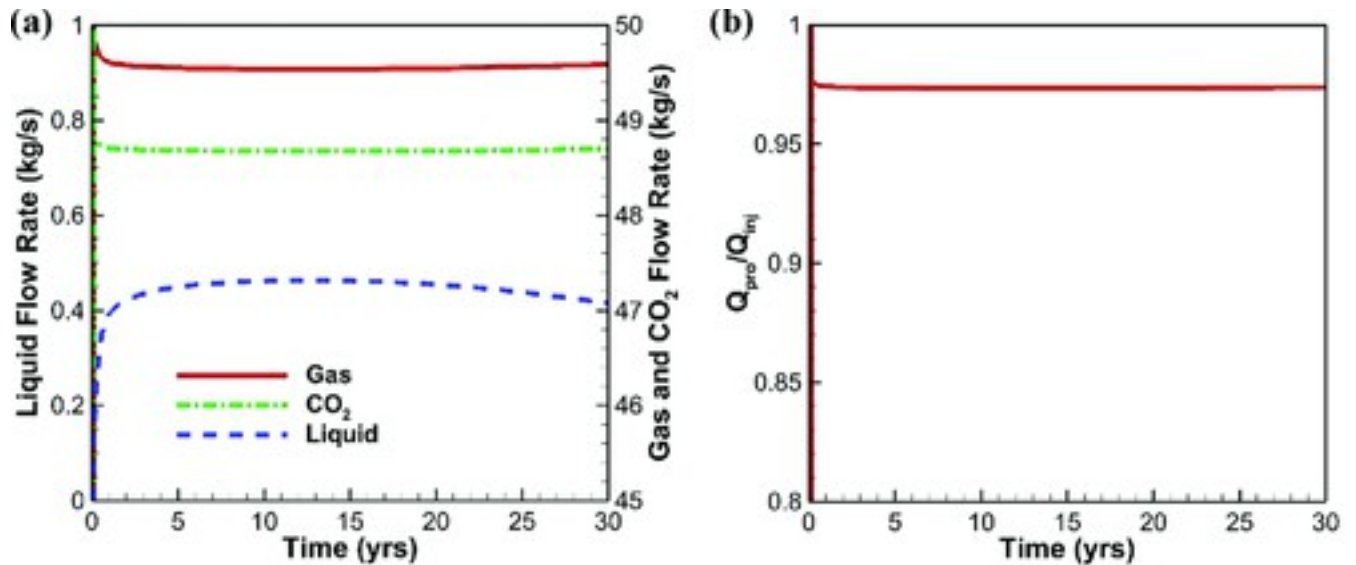


Figure 14

[Open in figure viewer](#) [PowerPoint](#)

(a) Simulated gas and liquid phase flow rates as well as CO₂ component flow rate and (b) Ratio of CO₂ injection rate and production rate. Because the total injection (pure CO₂) rate and the total production (mixture) rate are equal, this ratio is also a measure of how much CO₂ enters the production stream.

[Caption](#)

Conclusions

ECO2N V2.0 is an enhanced version of ECO2N V1.0, a fluid property module for the multiphase, multicomponent simulator TOUGH2, Version 2.1. The new code provides capabilities for modeling advective and diffusive flow and transport in multidimensional heterogeneous systems containing H₂O-NaCl-CO₂ mixtures with temperatures from ambient up to 300°C. Process capabilities include coupling between fluid and heat flow, partitioning of H₂O and CO₂ among different phases, and precipitation/dissolution of solid salt. The code represents thermophysical properties of brine-CO₂ mixtures generally within experimental accuracy for the range of conditions of interest in geologic disposal of CO₂ and CO₂ enhanced geothermal reservoirs. Supercritical as well as subcritical conditions may be modeled, but the code currently has no provisions to treat separate liquid and gas CO₂ phases, or transitions between them.

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Biographies



- **Lehua Pan** Dr Lehua Pan has been working at Lawrence Berkeley National Laboratory since 1997 and is an expert in computer modeling of Earth systems and processes. He received his PhD in Soil Physics/Hydrology from University of Arizona in 1995, his MSc in Soil Physics from Zhejiang Agricultural University in 1986, and his BSc in Geology from Zhejiang University in 1982.



- **Nicolas Spycher** Nicolas Spycher has over twenty years' research experience in aqueous geochemistry and water/rock/gas interactions, including development of multicomponent geochemical and reactive transport models. His current research focuses on developing biogeochemical conceptual and numerical models to understand water/gas/rock/sediment interactions in the subsurface.



- **Christine Doughty** Christine Doughty is a staff scientist at Lawrence Berkeley National Laboratory. Her research interests include mathematical modeling of multi-component, multi-

phase fluid flow and heat transport in heterogeneous geologic settings, including fractured rock. She obtained a PhD in Hydrogeology from the University of California, Berkeley in 1995.



- **Karsten Pruess** Karsten Pruess is a senior scientist emeritus of the Lawrence Berkeley National Laboratory, which he joined in 1977. He has conducted research in multiphase, non-isothermal, and chemically reactive flows in porous media, including mathematical modeling, analysis of field data, and laboratory experiments. He received a PhD in Physics from the University of Frankfurt, Germany in 1972.