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Vegetation fires in the Anthropocene

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Abstract | Vegetation fires are an essential component of the Earth system but can also cause substantial economic losses, severe air pollution, human mortality and environmental damage. Contemporary fire regimes are increasingly impacted by human activities and climate change, but, owing to the complex fire–human–climate interactions and incomplete historical or long-term datasets, it is difficult to detect and project fire-regime trajectories. In this Review, we describe recent global and regional trends in fire activity and examine projections for fire regimes in the near future. Although there are large uncertainties, it is likely that the economic and environmental impacts of vegetation fires will worsen as a result of anthropogenic climate change. These effects will be particularly prominent in flammable forests in populated temperate zones, the sparsely inhabited flammable boreal zone and fire-sensitive tropical rainforests, and will contribute to greenhouse gas emissions. The impacts of increased fire activity can be mitigated through effective stewardship of fire regimes, which should be achieved through evidence-based fire management that incorporates indigenous and local knowledge, combined with planning and design of natural and urban landscapes. Increasing transdisciplinary research is needed to fully understand how Anthropocene fire regimes are changing and how humans must adapt.

Biomass

Non-fossilized organic matter, including living and dead phytomass, organic soils, and animal remains and excrement.

Vegetation fires — also referred to as wildland fires, wildfires, landscape fires, bushfires, biomass burning, forest fires, scrub fires, crop fires and grass fires — are unique Earth-system disturbances that affect the coupled biosphere, hydrosphere, geosphere, cryosphere and atmosphere^{1,2} (FIG. 1). For example, during burning, large quantities of water vapour, CO₂, CH₄, N₂O and aerosols are released, modifying the radiative balance of the Earth³; aerosols reduce transmission of solar energy to the land surface, while greenhouse gases trap solar radiation⁴. Extremely intense fires can also trigger the development of pyrocumulonimbus storms, injecting aerosols into the stratosphere, where they can be transported globally, impacting radiation budgets⁵. Pyrocumulonimbus storms further facilitate extreme fire behaviour by encouraging lightning ground strikes, which ignite new fires⁶, presenting a positive feedback. The high concentration of black carbon⁴ (soot) in smoke similarly acts to influence the Earth system and future fire activity, whereby the prevention of precipitation coalescence inhibits rain cloud formation⁷. The atmospheric transport and subsequent fallout of soot in the cryosphere additionally reduces albedo and increases snow and ice melt^{8,9}.

Fire has been a natural feature of the Earth system for the last 420 million years³. Strong self-reinforcing interactions between climate, vegetation and fire occurrence have led to distinct ‘fire regimes’, defined by the temporal

frequency, spatial extent and pattern, characteristic behaviour and environmental effects of vegetation fires¹⁰. Organisms have evolved specialized strategies to resist, promote or recover from fire disturbance. Some plant species have life histories tied to specific fire regimes, such as fire-stimulated flowering, post-fire seed release from aerial seed-banks and smoke-triggered seed germination¹¹. This specialization means that any departure from the prevailing fire regime, such as fire exclusion or increasing frequency and intensity, can result in population declines or local extinction¹². For example, woody plant species that exclusively regenerate from seed (obligate seeders) experience regeneration failure if fires become too frequent for seedling growth and maturation¹³.

The evolution of fire-wielding hominins around one million years ago introduced much more complexity into the timing, location, extent and behaviour of vegetation fires¹⁴. For example, hunter-gatherers managed natural resources with fire, attracting herbivores to freshly burned areas^{15–17}. Pre-industrial agriculturalists further used fire to clear land and burn crop debris¹⁸. Furthermore, contemporary humans shape fire regimes by suppressing natural ignitions, as well as modifying landscapes and fuel loads through controlled burns (prescribed fires), land clearing, urbanization, cultivation of non-native plants and animal husbandry¹⁹. Humans also accidentally or maliciously set fires that

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Key points

- Vegetation fires are an ancient and essential component of the Earth system, and have shaped the evolution of plants, animals and biogeochemical processes. There are discernible global geographic and temporal patterns of fire activity reflecting the interplay of climate, vegetation and ignitions.
- Anthropogenic influences on fire activity have become more pronounced since the late eighteenth century, reflecting the effects of industrialization and climate change, land clearance, human population growth, replacement of indigenous and traditional fire management, and the subsequent development of large-scale firefighting and fuels management in the twentieth century.
- The human settlement and infrastructure embedded in flammable vegetation contributes to economically disastrous fires.
- Large and frequent fires in boreal and tropical forests have the potential to cause terrestrial carbon stores to become major greenhouse gas sources, amplifying climate change.
- Detecting and predicting changes in fire activity is difficult due to relatively brief fire records, bioregional variability and human involvement. To understand how Anthropocene fire regimes are changing, and how humans must adapt, researchers from biological sciences, physical sciences and humanities and fire-management practitioners must work together.

become uncontrollable (thus, becoming wildfires) and sometimes economically destructive and fatal (fire disasters)²⁰. Globally, very few vegetated environments are unaffected by human fire use²¹.

Past climate change is known to have influenced the extent, frequency and intensity of vegetation fires by affecting vegetation patterns, fuel abundance, and seasonal and interannual drought²². Currently, anthropogenic climate change is altering precipitation patterns and increasing temperatures, resulting in more frequent extreme vegetation-fire events^{23–25}. In recent years, for example, there have been increasing media reports of major fire disasters: the wildfires in Chile in 2017, Portugal in 2018, California in 2018 and Australia in 2019–2020 are prominent examples. Such fires have major economic impacts, affecting life, property and human health^{26–29}, thus, it is important to consider contemporary and future trends in fire activity to inform adaptation and mitigation policies. For instance, the 2019–2020 Australian Black Summer is likely the nation's most costly natural disaster, costing over AUD \$100 billion³⁰, whilst California in the USA saw an estimated USD \$40 billion (NatCatSERVICE) in structure losses alone in a dozen major wildfires in 2017–2018. Importantly, these fire disasters are not simply due to climate factors but also the modification of landscapes during the Anthropocene^{31,32}.

In this Review, we describe how human–environmental interaction shapes fire activity from the viewpoint of pyrogeography³³. The characteristics of global vegetation fires are outlined, followed by a synthesis of current trends and future projections in fire activity. The effects of changing fire regimes and pathways for adaptation are subsequently discussed, ending with the consideration of adaptation measures and future research priorities.

Contemporary fire regimes

Vegetation fires burn an annual global average of 400–500 million hectares (Mha)^{34,35}, with fire patterns driven by the interplay between climate, weather,

vegetation type, ignition and anthropogenic fire management^{36–38} (FIG. 2). Broadly, fire activity has a unimodal relationship with primary productivity^{39–42}. Regions with intermediate levels of primary productivity, such as tropical savannah, burn at a very high frequency (approximately 1–5 years), owing to abundant fuel (vegetation), reliable dry seasons and ample natural and anthropogenic ignitions⁴³. In contrast, in rainforest environments that have high primary productivity, fire is naturally rare (on the timescale of millennia)⁴⁴, limited to climatic periods when fuel is dry enough to burn, although anthropogenic deforestation fires are an annual occurrence⁴⁵ (FIG. 2). In high-biomass boreal and temperate forests, the return time of fires varies between decade and century scales, and is largely controlled by extreme fire weather, ignitions and vegetation cycles⁴⁶. Notably, arid regions are climatically conducive to fires but have a limited capacity to burn due to low primary productivity⁴⁷. However, fire activity in these regions can increase when fuel becomes available, for instance, following interannual wet periods or non-native grass invasions⁴⁸.

Vegetation fires are an important source of greenhouse gases, particularly CO₂ (REF.⁴⁹). Global mean carbon emissions due to fire were approximately 2.2 Pg per year from 1997 to 2016, about 22% of contemporary global annual carbon emissions from fossil fuel combustion. Of fire-related emissions, about 65% are due to savannah and grassland fires, reflecting the high frequency of burning in these environments, and 10% from temperate and boreal forest fires⁵⁰. Approximately 20% are associated with clearing tropical rainforest for pasture, plantations and agriculture, with the remainder associated with agricultural waste burning⁵⁰. Although emissions from flammable vegetation can be balanced by sequestration during post-fire vegetative recovery, emissions from fires associated with permanent deforestation or combustion of organic deposits such as peatlands^{51,52} are net sources of carbon to the atmosphere. Similarly, rapid climate change and increased forest fires could make carbon sinks (forest vegetation and soil) become carbon sources⁵³.

Global trends in fire activity. Despite increased reporting of fire disasters⁵⁴, such as the Australian 2019–2020 bush-fire crisis, it is unclear if these disasters are evidence of increasing global fire activity related to climate change⁵⁵. The uncertainty is partially because historical records of fire activity, even for simple metrics like area burned, are short and available for only a few nations⁵⁵. Prehistorical fire records, which are essential to understanding the fire history of infrequently burned vegetation, rely on palaeoecological proxies such as dendrochronology and analysis of sedimentary charcoal^{19,56–60}. These proxies, unfortunately, do not directly scale to key components of fire regimes, such as frequency or geographic extent of fires, and there are few regions, such as the western USA, with a large number of high-resolution records that extend into the prehistorical period^{60,61}.

Satellite imagery is currently used routinely to monitor fire activity but only became available in the 1970s⁶². This period is much shorter than the natural fire-return interval of many forested biomes and could fail to capture infrequent cyclical events, such as the intense droughts

Pyrocumulonimbus

Intense convective thunderstorms that develop above highly energetic wildfires, which can reach the stratosphere and create localized weather, including rain, hail, lightning and pyro-tornadoes.

Fire regimes

Characteristic syndrome of landscape fire with respect to behaviour, frequency, seasonality, geographic scale and pattern, with predictable biological responses and environmental effects.

Regenerate

The process of plant recovery following fire damage, either from seeds stored in the soil or vegetatively from specialized tissues located in roots, stems and branches.

Ignitions

Sufficient energy to initiate combustion of plant biomass, and can be natural, such as from lightning, or directly set by humans either deliberately, accidentally or indirectly.

Fig. 1 | **Vegetation fire in the Earth system.** Landscape perspective of the multiple factors that influence, interact with and are impacted by vegetation fire. Fires have numerous direct and indirect affects that impact the biosphere (including vegetation cover), geosphere (including soil erosion), hydrosphere (including fluvial sediment and nutrient transport), cryosphere (including soot fallout and changed albedo) and atmosphere (including smoke pollution). PyroCb, pyrocumulonimbus.

Extreme vegetation-fire events

Extreme fire events are characterized by some combination of the following: anomalous fire behaviour, involving extremely high energy releases, very rapid rate of spread, very large flame heights; massive emission of smoke and greenhouse gas pollution; prolonged duration of fires, enormous geographical scale of burned areas, or both; fires causing unusually adverse biological, atmospheric or geomorphological effects.

and floods that control fuel availability⁶³. Moreover, early satellite observations are spatially coarse and have imperfect coverage⁶⁴, with reliable coverage only available since the turn of this century. More uncertainty in fire activity amongst biomes, which demands regional rather than global analyses. As fire disasters are often associated with much smaller burned areas than fires in remote areas, metrics other than burned area need to be incorporated into analyses^{23,65}. Particularly important are estimates of fire intensity, which provides a measure of the energy released from the fires, and fire severity, which is an estimate of the environmental impacts of the fires, such as degree of canopy damage⁶⁶. Despite these observational limitations, there is an emerging picture of changes in global fire activity, which emphasizes the importance of regional-scale variation, climate change and anthropogenic drivers.

From 1979 to 2013, an average increase of 18.7% in fire-weather-season length has been documented across global burnable lands, with a doubling by long fire-weather seasons across most of the Earth's flammable biomes⁶⁷ (FIG. 3). Extreme fire-weather conditions, alongside drought and fuel dryness, are associated with extreme fire events²³. However, the MODIS burned area record⁶⁸ indicates that, between 1998 and 2015, the area burned by vegetation fires globally declined by around 25%, from over 500 Mha to less than 400 Mha annually⁶⁹. The largest decreases in area burned occurred in African and South American tropical savannahs and Asian semi-arid grasslands (FIG. 3a,b), and were caused by ongoing land-cover conversion, leading to a more fragmented and less flammable landscape⁶⁹. A decline greater than the global average of area burned was also detected in western Australian desert, a change known to be associated with interannual drought cycles and the absence

Fig. 2 | Global patterns of fire and vegetation. **a** | Mean annual burned fraction from 2001 to 2018, based on MODIS burned area³⁴. **b** | 95th percentile of the Fire Weather Index (FWI) from the Canadian Forest Fire Danger Rating System derived from ERA5 (REF.²²⁴), where the 95th percentile is calculated from data pooled over the entire calendar year. Higher values represent an increased potential for fire. **c** | Tree cover from MODIS²²⁵. **d** | Broad spatial pattern of five types of fire regime³⁸: little (or no) fire; surface low frequency; surface medium frequency; surface high frequency; and crown medium frequency. Fire activity and regimes are controlled by the interaction of biomass and climate, as well as human and natural ignitions.

Anthropocene

The geologically novel planetary state resulting from human activities, although the start date of the state is debated.

Pyrogeography

The holistic study of fire on Earth achieved by combining and synthesizing knowledge and methods from the sciences and humanities.

Dendrochronology

Analysis of growth rings in the trunks of suitable tree species can enable reconstruction of past environmental conditions, resolved to annual or seasonal scales.

Fire intensity

The amount of energy released per unit time from a fire front.

Fire severity

A measure of the biological impact of fires, routinely assessed by the degree of canopy or understory defoliation and foliage consumption.

of Aboriginal fire management^{70,71}. These contrasting trends highlight why a sole focus on area burned does not adequately capture trends in the changing risk of highly economically and economically destructive fire events⁶⁵.

Regional trends in fire activity. Regional analyses are essential to reconstruct contemporary fire-activity trends. The most well-studied regions are the western USA (FIG. 3c) and boreal Canada (FIG. 3d), owing to their comparatively high density of researchers, abundant natural archives (lake sediments and fire scars) that record past fires and reliable official fire records since the early twentieth century. Sedimentary charcoal records in the western USA show that, over the last 3,000 years, fire activity was primarily controlled by temperature and drought⁶⁰. During the nineteenth century, however, fire regime became increasingly anthropogenically driven as human population in the region increased and indigenous fire practices waned⁷², and fire activity peaked subsequently mid to late century. Active fire suppression decoupled the fire activity–climate relationships and created a historic landscape ‘fire deficit’ in the twentieth century. Since ~1980, though, warmer, drier summers and earlier spring snowmelt^{73,74} have increased the area burned from vegetation fires, despite sustained investment in industrialized

firefighting⁷⁵ (FIG. 3c). For example, between 1972 and 2018, there has been a 405% increase in total burned area in California⁷⁶, and decadal burned area from 2003 to 2012 increased by 1,200% in forests of the western USA, compared with 1973–1982 (REF.⁷⁷).

In the Canadian boreal forest, fire frequency and extent has generally increased in response to higher regional temperatures^{46,78}, longer fire seasons⁷⁹ and drier fuels⁸⁰ (FIG. 3d). Annual average burned-area in Canada has almost tripled since 1959 from 1 million ha to 2.8 million ha⁸¹. Drier and hotter fire seasons also increase the probability of lightning ignitions, causing a rise in the number of large fires in western Canada^{81,82}. Lightning-caused fires are responsible for 90% of the burned area in Canada, and the number of lightning-caused fires has increased significantly especially over western Canada, resulting in increasing burned area, as burned area by human-caused fires has been decreasing⁸¹. The combination of drought and vegetation fire has also led to a net increase in greenhouse gas emissions, both directly through the combustion of forest biomass and soil carbon stocks (such as in peatlands) mainly in the form of CO₂ (REFS^{51,83}) and indirectly by causing permafrost thaw and changing thermokarst hydrology, where the relative emissions of CH₄ could increase^{84,85}.

Fig. 3 | Trends in vegetation-fire activity. **a** | Linear trends in the number of days per year where the Fire Weather Index (FWI) from ERA5 (REF.²²⁴) exceeded the local 95th percentile over the period 1979–2019. **b** | Annual fire emissions over Africa (1997–2019)⁴². **c** | Burned area in contiguous Western USA, 103 °W to the Pacific (1916–2018) and non-inflation-adjusted US federal fire-suppression costs (blue line). Burned area based on historical reconstruction from 1916 to 1984 (REF.²²⁶), Monitoring Trends in Burn Severity (MTBS) data for 1984–2017 (REF.²²⁷) and MODIS for 2018 (REF.³⁴). Data were bias corrected using periods of overlap to the reference MTBS record (1984–2003 for reconstruction and MTBS, and 2001–2017 for MODIS). **d** | Burned area in Canada (1959–2018)⁸¹. **e** | Burned area in New South Wales, Australia (1979–1980 through 2019–2020 fire seasons) and number of pyrocumulonimbus²⁴ (PyroCb) in Australia (blue line). **f** | Burned area in Tasmania, Australia (1979–1980 through 2018–2019 fire seasons) resulting from all ignitions (unknown and lightning)⁹¹. Blue line represents burned area from fires with lightning ignitions. **g** | Burned area in central Chile (1984–1985 through 2016–2017 fire seasons)⁹⁵. **h** | Burned area in Portugal (1980–2018)²²⁸ (2018 estimated using MODIS data³⁴). **i** | Fire emissions in Indonesia (1960–2019)⁹⁹. **j** | Fire emissions from Brazilian Amazonia (1960–2019)⁹⁹.

In Australia, it is difficult to disaggregate the influences of climate change from the effects of the cessation of >45,000 years of Aboriginal hunter-gatherer fire management following European colonization in the early nineteenth century^{86,87}. However, the effect of climate change is becoming apparent in the increasing numbers of extreme fire events^{24,88,89}. For instance, analysis of historical satellite imagery has demonstrated an increasing trend in pyrocumulonimbus occurrence^{24,90}, with 35 storms in the 2019–2020 fire season, thereby, doubling the known Australian records of these extreme fires (FIG. 3e). In the western Tasmanian wilderness, the number of lightning-ignited fires and the area burned as a result has also sharply increased since 1980–1985, from burning an average of around 100 ha annually⁹¹ to over 200,000 ha in 2019 (FIG. 3f), including rarely burned and fire-sensitive Gondwanan rainforests⁹². Most recently, a globally anomalous 2019–2020 fire season in Australia burned over 5 million ha of *Eucalyptus* forests⁹³ and has been linked with anthropogenic climate change^{93,94}.

Similar to the 2019–2020 fires in Australia, the 2017 Chilean fires were the largest fires on record for that nation, burning over 5,000 km² (FIG. 3g), and were associated with anomalous fire-weather conditions related to drought and high temperatures⁹⁵. However, large expanses of highly flammable, densely stocked monocultures of *Pinus* and *Eucalyptus* plantations contributed to the scale and intensities of the Chilean fires⁹⁵, a situation that also occurred in the 2018 Portugal fires⁹⁶ (FIG. 3h). These cases highlight the importance of factors other than (or in addition to) climate change, such as vegetation changes related to agriculture or invasive species, in fire regime change. For instance, the spread of *Bromus tectorum* (cheatgrass) across the semi-arid, intermountain western USA has facilitated the replacement of a low-frequency, mixed-severity shrubland fire regime with a high-frequency, high-severity grassland fire regime⁹⁷. The introduction of *Andropogon gayanus* (gamba grass) to tropical eucalypt savannah surrounding Darwin in northern Australia has similarly led to the replacement of near-annual high-frequency, low-intensity fires with high-intensity fires, driving the savannah on a trajectory towards a treeless state⁹⁸. Moreover, tropical deforestation fires have remained relatively constant over the last 30 years after increasing during the 1980s in line with increased deforestation^{99,100}, despite climate change. Instead, an important feature of tropical deforestation fires is the strong interannual variation in step with drought cycles, such as those caused by the El Niño–Southern Oscillation^{101,102} (FIG. 3i,j).

In many regions, it is apparent that fire seasons are lengthening and becoming more extreme, and lightning ignitions are increasing, contributing to economically destructive fire events. Yet, simple attribution of these increases to climate change is challenging because there are specific factors, including anthropogenic ones, that shape regional fire activity. These factors can increase the risk of extreme events in some cases, and, in other cases, cause a decline in area burned, highlighting the need to consider and to record more indices than simply area burned.

Future fire regimes

Fire-activity prediction is a rapidly changing field informed by improved resolution of climate-change projections, increased capacity of numerical models and deeper understanding of the climate drivers of fire activity. Over the last 40 years, a variety of modelling approaches have been applied both globally and regionally to discern likely changes in fire activity owing to anthropogenic climate change. TABLE 1 summarizes the rapid-response review on projected changes in fire activity by geographical region based on papers published since the Fifth Assessment Report of the Intergovernmental Panel on Climate Change¹⁰³, augmented with additional key papers discussed here. Overall, there is general agreement amongst the studies that the frequency or severity of fire weather, fire-season length, burned area and fire occurrence are increasing with future climate change. Although some regions are suggested to have no change or decreases in fire activity, particularly in regions with substantial human influence, none of the papers support a widespread decrease in fire risk (TABLE 1). However, it should be noted that there is a limited number of studies for Africa and Asia.

The strong association between anomalous fire weather and extreme fire behaviour suggests that anthropogenic climate change will impact future fire regimes^{23,104}, including in those regions already vulnerable to fire disasters, such as the western USA, Mediterranean and southern Australia²³. A global analysis using 17 global climate models, and highlighted regionally in TABLE 1, suggests there will be a large increase in the occurrence of extreme fire weather for much of the globe, with some of the largest increases the Mediterranean and the Amazon¹⁰⁵. For instance, the number of days exceeding the 95th percentile of the Fire Weather Index is projected to increase almost 130% for the Mediterranean to just under 40% for Alaska when global mean temperatures reach 2 °C above pre-industrial levels relative to contemporary conditions. There is also evidence that increasing atmospheric temperatures will lead to more lightning activity^{82,106}, further adding to fire activity in ignition-limited landscapes. Such changes in environmental factors will be confounded by changes in anthropogenic factors, including human-settlement patterns and land use (such as cropland and pasture), in many areas of the globe that tend to reduce fire activity^{107,108}.

Approaches have been used to model future changes in global fire activity ranging from the use of process-based coupled Earth System Models to statistical approaches. For example, Kloster and Lasslop¹⁰⁸ examined future changes in global burned area across a set of Earth System Models with integrated fire–vegetation dynamics that were forced with future changes in climate and land use. The models generally showed increases in global burned area, albeit with substantial uncertainty, with some models showing a 58% increase by the end of the twenty-first century under stronger warming scenarios. Statistical-modelling efforts include Moritz et al.¹⁰⁹, who used environmental-niche modelling to project changes in global fire activity. They found heterogeneous changes in fire activity with general

Table 1 | Summary of selected global, biome and regional studies on future changes in fire indicators

Geographic domain	Projected change in Fire Weather Index days >95th percentile (%)	Fire activity indicator	Projected change of fire activity indicator	Ref.
Global studies				
Global	–	Fire weather	Increase	210
Global ^a	–	Fire weather	Increase	105
Global	–	Fire weather	Increase	104
Global	–	Fire weather	No change	211
Global ^a	–	Fire weather	Increase	212
Global ^a	–	Burned area	Mixed	108
Global	–	Burned area	Mixed	107
Global	–	Fire activity	Increase	213
Global ^a	–	Fire activity	Mixed	109
Biome studies				
Mediterranean biome ^a — Europe/S. America/N. America/Australia/Africa	–	Fire activity	Mixed	118
Boreal biome — Central Russia/W. Canada	–	Fire weather	Increase	214
Regional studies				
Europe — Mediterranean	+23.5 (129%)	Burned area	Increase	116
Europe — Mediterranean		Fire activity	Mixed	118
Europe — France		Fire weather	Increase	116
South America — Amazon	+15.5 (84.9%)	Fire activity	Increase	215
South America — Amazon		Burned area	Increase	216
Australia — Southeast Australia	+9 (49.3%)	Fire weather	Increase	217
Australia — Southeast Australia		Fire weather	Increase	90
Australia ^a — Southeast Australia		Fire weather	Increase	24
Australia — Tasmania		Fire weather	Increase	218
North America ^a — Western USA	+10.2 (55.9%)	Burned area	Mixed	115
North America ^a — Western USA		Burned area	Increase	114
North America ^a — Western USA		Fire activity	Increase	112
North America — California		Fire impacts	Increase	219
North America — California		Fire Weather	Increase	220
North America ^a — Central Rocky Mountains		Burned area	Increase	113
North America — Canada/boreal forest	+9.2 (50.4%)	Fire weather	Increase	111
North America — Canada/boreal forest		Fire weather	Increase	221
North America — Canada/boreal forest		Fire weather	Increase	111
North America — Canada/boreal forest		Burned area	Increase	221
North America ^a — Canada/boreal forest		Burned area	Increase	46
North America — Alaska/Western Canada		Burned Area	Increase	78
North America ^a — Alaska	+7.1 (38.9%)	Burned area	Increase	222

Studies include those from Jones et al.²²³ and other studies discussed in the text (marked with^a). Regional studies are summarized by continent and sub-geographic region or country. The primary fire indicator as either fire activity (such as fire occurrence), burned area, fire weather (such as fire danger) or fire impacts, as well as the primary direction of change, is highlighted for each study. For some macroscale regions, also reported are the projected changes in the number of days per year the Fire Weather Index from the Canadian Forest Fire Danger Rating System exceeds the 95th percentile corresponding with global mean temperatures 2 °C above pre-industrial levels versus 1981–2010 (REF.¹⁰⁵). We summarize the median change calculated from 17 climate models, with the percent change from the 1981–2010 baseline in parentheses.

increases in mid-to-high latitudes and declines in the subtropics.

Regional-scale studies of future fire regimes, largely focused on North America and Mediterranean Europe, have examined changes in metrics such as fire-season

length¹¹⁰, extreme fire weather¹¹¹ and burned area. The boreal zone of Canada, for instance, is expected to experience a doubling of burned area by the end of the century, increasing to 5 Mha year⁻¹ (REF.⁴⁶). Similarly, more extreme fire weather is projected to favour more

Fig. 4 | Relationships and feedbacks between climate, fire and vegetation. Linkages and feedbacks between the global climate system, wildfire and vegetation could drive future fire activity in the Anthropocene. The effects of post-colonial anthropogenic changes to global climate, vegetation and wildfire are shown on the red lines. The consequences of these effects on global climate, vegetation and wildfire are shown on the black lines. These interrelationships highlight the complexity of fire in the Earth system and the sensitivity to climate change, underscoring the challenge in predicting future fire regimes. GHG, greenhouse gas; WUI, wildland–urban interface.

very large fires for the western USA¹¹² and, in concert with reduced spring snowpack and drier fuels during the summer, will lead to a substantial increase in burned area for montane forests¹¹³. Process-based land-surface models also show substantial continued increases in burned area over the twenty-first century across forested portions of the western USA¹¹⁴. Projected changes in burned area that incorporate direct flammability controls on fire activity and the indirect effects of fuel productivity, however, show more mixed changes across western USA due to heterogeneity in climate, vegetation and land use¹¹⁵. For example, semi-arid shrub and grasslands that warm further and undergo aridification will see reduced fire due to biomass limitations.

In Mediterranean Europe, projections suggest a 40–100% increase in burned area with a 1.5°C to 3°C warming¹¹⁶, despite recent investment and expansion of fire-suppression efforts that have led to a reduction in fire-activity since the 1980s¹¹⁷. In some Mediterranean ecosystems in Europe, though, burned area could decline if climate-change-associated aridification increases fuel limitation¹¹⁸. However, such changes in fire regimes modelled using space-for-time substitutions are predicated on vegetation equilibrium with climate; trailing-edge disequilibrium of vegetation under

transient climatic changes could delay realization of reduced fire activity in some landscapes.

Projection uncertainty. Although the projections reviewed here support that there will be increased fire activity in the future, these projections carry large uncertainties due to the feedbacks between different aspects of the Earth system and the difficulty in representing human behaviour in models. For example, one approach to predicting future changes in fire regimes is to explore the statistical association between current and projected climate variables and area burned^{40,109}. However, the relationship between vegetation fire and vegetation is likely to change in the future due to dynamic feedbacks, which is not accounted for using statistical associations^{108,119,120}. Another approach, process-based mechanistic modelling, incorporates the complex interplay of climate, CO₂, vegetation and fire, and, sometimes, impacts of land-use change and change in ignitions¹²¹. These factors could amplify, dampen or negate one another⁶³ (FIG. 4). Some global modelling results suggest that the effects of climate change will substantially overwrite human influences on fire regimes¹²², for instance, by limiting the effectiveness of fire suppression^{123,124}. This conclusion, though, hinges on realistic representations of anthropogenic factors, which are particularly difficult to capture when modelling

Aridification

A process where climate change can lead to sustained regional drying with concomitant changes in vegetation, fire regimes and geomorphological processes.

Climate refugia

A landscape setting with an atypical climate where species poorly adapted to ambient environmental conditions are able to persist, for example, a cloudy mountain top.

Fire refugia

A landscape setting that limits the egress of fires, enabling species poorly adapted to ambient fire regimes to persist, such as a deep ravine.

future fire activity, as it relies on plausible scenarios of economic development and numerical relationships between human population density and ignitions, fire suppression and fuel availability^{37,107,108,121}. Thus, regional-scale models are likely to have higher veracity than global models because biophysical and socio-ecological controls of fire activity have distinct geographic patterns^{37,109}. Regardless, comparative analyses of different model outputs and observed patterns in area burned show that one of the greatest barriers to modelling future fire regimes concerns anthropogenic factors¹²⁵.

There is also uncertainty in fire-regime predictions due to the complexity of climate–fire–vegetation interactions and feedbacks¹²⁰. For example, in some regions, climate change could increase the abundance of plants with low flammability, such as through the replacement of coniferous forests by deciduous broadleaf forests in boreal landscapes¹¹⁹. Feedbacks between logging, deforestation, fire, drought and climate change could lead to the replacement of Amazon rainforests with flammable grasslands¹²⁶. Elevated atmospheric CO₂ concentrations increase carbon assimilation and improve plant water-use efficiency¹²⁷, potentially altering the balance of grass and woody plants in tropical savannahs. In this case, shrub and tree invasion reduces the amount of grass that fuels vegetation fires, thus, decreasing flammability^{107,128}. However, climate-change-enhanced drought could negate any effects of carbon fertilization and enhanced water-use efficiency, maintaining or increasing flammability. Most modelling results suggest that negative fire–vegetation feedbacks would limit the extent of projected increases in burned area, based

on fire weather alone^{116,129}, although possible positive fire–vegetation feedbacks, such as invasive annual grass cycles and increased productivity with climate change, could counteract the negative feedbacks¹³⁰.

In theory, interactions between climate change, vegetation and fire could drive a runaway feedback process in which entire landscapes are transformed as fire frequency increases with worsening fire weather, and dry, hot conditions^{13,131} limit vegetation regeneration (FIG. 5). Ecosystems currently located on the fringe of a bioclimatic niche (such as lower tree line forests) and vegetation persisting in climate refugia or fire refugia are particularly vulnerable to transformative fire regimes^{92,131,132}. Landscape-wide conversions from slow-growing, infrequently burned forests to frequently burned grassland and shrublands are possible (FIG. 5), which depletes soil carbon stores, as in regrowth forests in the conifer forests in the western USA¹³³ and in Indonesia¹³⁴. This feedback process is most likely self-correcting because fire regimes will shift from being flammability-limited to fuel-limited with increasing reduced productivity^{40,42,133}. According to this scenario, burned area will eventually decrease and, consequently, more complex vegetation structure will develop, as has been observed in grasslands of the US Great Plains, where the intentional exclusion of fire has facilitated woody tree and shrub expansion¹³⁵.

Another key area of uncertainty in future fire-regime predictions is the range of variability and frequency of extreme fire events, and there is a lack of research projecting these events and linking them to cascading consequences, such as post-fire floods, large-scale erosion and debris flows that alter hydrology and topography, and development of pyrocumulonimbus¹³⁶. The paucity of projections is partially due to the substantial challenges and uncertainties inherent in coupling interdisciplinary models, as well as the increased level of uncertainty and error at the extremes^{90,137,138}. Fire, and its interactions, remains poorly represented in Earth-system models, and their performance struggles with non-analogue, rapidly changing future climates¹³⁹. Thus, predicting ‘black swan’ extreme fire events — which, by definition, have no historical precedent, such as the protracted, enormous and severe Australian 2019–2020 fires — seriously challenge the capacity of Earth-system-model projections. The Australian fires represented a unique constellation of historically anomalous prolonged drought, the unusual conjunction of interannual climate modes (an intense positive-phase Indian Ocean Dipole and negative-phase Southern Annular Mode) that resulted in a prolonged fire weather, localized lightning storms and human ignitions than ignited the fires, combined with intentionally set fires and direct firefighting designed to stop the fires¹⁴⁰. Clearly, capturing such specific combinations of climate events and human agency to predict such an extreme event is beyond the capacities of current predictive models of future fire activity¹³⁹.

Adapting to fire in the Anthropocene

Fire regimes in the Anthropocene require adaptation and effective management to promote environmental sustainability and reduce greenhouse gas emissions^{33,141}.

Fig. 5 | Conversion of infrequently burned forest to a frequently burned, treeless state. Conversion is associated with declines in carbon storage (sequestered C, top panel) and annual global fire-induced emissions (emission rate, middle panel), assuming that both above-ground and below-ground carbon are altered. In most forested systems, particularly in temperate and boreal ecosystems, much of the sequestered carbon is in litter and soil organic layers and the roots, which can become substantial sources of carbon emissions when burned, especially by repeated fires^{188,203,229}.

Fig. 6 | Adaptation to Anthropocene fire regimes on fire-prone landscape. a | Adaptation and management strategies from densely forested areas to the wildland–urban interface, and at the building scale (inset). The intensity of fire management changes proportionally with human population and economic assets, with strategies ranging from allowing fires to burn in wilderness areas to intensive management on the wildland–urban interface. Each type of management solution will be strongly framed by local conditions and social and cultural factors. **b** | Distance of exposure to conductive heat and direct flames, radiant heat, convective heat, embers and brands, and smoke, particulates and aerosols.

Adaptation and management approaches must be tailored to specific environmental and socio-ecological settings (FIG. 6), which we consider in flammable landscapes and the wildland–urban interface (WUI).

Flammable landscapes and the WUI. It is neither possible nor desirable to exclude fire from fire-adapted landscapes. Rather, sustainable fire management demands careful fuel management to conserve biodiversity, sustain ecosystem services (such as the provision of water and clean air) and reduce the risk of wildfire disasters¹⁴².

Fuel management can involve broad-scale prescribed burning, as it is a cost-effective method to reduce fuel loads and, hence, reduce the risk of uncontrollable fires in frequently burned vegetation, such as savannah and dry *Pinus* and *Eucalyptus* forests¹⁴³. To be effective, however, this approach requires frequent burning because fuels rapidly reaccumulate. Furthermore, large areas must be treated to ensure that enough of the area that is likely to be encountered by a wildfire has been burned¹⁴⁴. Dense, mesic and naturally infrequently burned vegetation are unsuitable for this fuel treatment unless it

Wildland–urban interface
The intermix of urban areas with flammable vegetation that is the locus of the most deadly and economically destructive wildfires.

Biomass smoke

A dynamic mixture of gases and aerosols, made up of organic and inorganic chemical species, emitted during the combustion of biomass.

is combined with mechanical thinning. This latter approach has the potential to increase resilience in US coniferous forests where fire has been excluded for up to a century, with the possible benefit of reduced carbon emissions from subsequent wildfires^{145–147}.

However, prescribed burning has significant constraints, particularly owing to the potential for shrinking safe weather windows due to climate change¹⁴⁸, and the effectiveness of prescribed burning in altering wildfire behaviour sharply declines under extreme fire weather¹⁴⁹. Prescribed burning also causes biomass smoke pollution¹⁵⁰. Smoke from large-scale prescribed burning programmes can cause death and hospitalization, which could potentially exceed the human-health impact of the wildfires that these managed fires are intended to mitigate¹⁵¹. Air-quality regulation by governments has demonstrable public and environmental benefits and has driven technological innovation¹⁵², so, in principle, government regulation of the smoke pollution from both wildfire and prescribed fires could drive innovation in fuels management and achievement of more economical, sustainable and safer strategies¹⁵³.

Fire management should also draw on indigenous and local knowledge about fire regimes^{15,154,155}, because these practices have been refined over millennia and have demonstrably sustained biodiversity^{16,156}. Yet, it must be recognized that wide-scale ‘restoration’ of indigenous fire regimes is no longer possible in many environments, as there have been profound socio-ecological changes to indigenous fire practices and environmental conditions have moved outside of the historical range of variability^{157,158}. Nevertheless, fire-management approaches that combine traditional ecological and local fire knowledge with mainstream fire-management approaches and technologies must be developed to adapt to changing environmental conditions. Such management is particularly important in the flammable WUI, where the most economically destructive vegetation fires on Earth occur²⁷.

The WUI is rapidly expanding as a result of growing urban populations^{23,141,159,160}, particularly in southern Europe, western North America and southern Australia, which have the highest concentration of wildfire fatalities and structural losses^{23,141}. There is rapidly increasing firefighting expenditure to protect the WUI — the US fire season in 2018 exceeded USD \$3 billion in suppression costs, largely owing to the reliance on fleets of specialized aircraft^{159,161,162} (FIG. 3c). Moreover, fires in the wildlands near urban areas can cause extreme air pollution in populated areas. For example, smoke from the 2019–2020 Australian bushfires is thought to have caused 417 deaths (much greater than the 33 recorded direct fatalities) and several thousand hospitalizations for cardiovascular and respiratory problems¹⁶³.

The vulnerability of housing on the WUI reflects a constellation of socio-ecological factors that amplifies fire risk, property damage and loss of life^{141,164–166}. These factors include the large numbers of ignitions from human sources, ineffective fuel treatments, poor planning, design and regulation of the built environment (particularly the careless intermixing of housing and infrastructure in flammable vegetation) and

limited preparation for fire disasters, such as design standards for housing to survive fire, development of evacuation plans and construction of safe places, effective public-education communication, dissemination of advice and emergency warnings. Fundamentally, fire-management challenges in the WUI reflect urban planning and policy failures, which demand reform and the need to enact and enforce national and local laws that control urban development, design and construction^{141,165}.

More effective protection of urban and WUI areas from wildfires should be achieved by managing fuels through prescribed burning, mowing, grazing and browsing ground-cover vegetation, planting strips of fire-resistant trees to create ‘green fire breaks’ (REF.¹⁶⁷) and creating corridors of thinned forest to create ‘shaded fuel breaks’ (REFS^{162,165,168}). China is a world leader in this approach, establishing 364,000 km of green fire breaks to manage landscape fire¹⁶⁹. Communities and individuals must manage fire risk^{165,141,159,162,166} by fire-hardening homes and transforming private property in a partnership model¹⁷⁰, shifting the responsibility from fire-management agencies, which are typically centralized^{171,172}. Dissemination of geographically and temporally accurate information on wildfires and smoke hazards through advanced technology, such as geolocation and smartphone apps, can enable individuals to make informed choices about managing wildfire risks¹⁷³ and their own exposure to high levels of smoke pollution¹⁷⁴, and planning for prospective evacuations in rapidly changing conditions^{175,176}. For instance, the US Environmental Protection Agency has developed a real-time smoke-information smartphone app, Smoke Sense, that combines citizen science and instrumental monitoring of air quality to improve public health¹⁷⁷. Likewise, the AirRater app in Australia was developed by the University of Tasmania, and includes notifications of the health risk posed by biomass smoke to users¹⁷⁴. Additional smartphone apps are being developed by Australian, US, Canadian and South African fire-management agencies to disseminate advice and warning of fire danger¹⁷⁶. A strength of these approaches is that they provide real-time and location-specific information, although they can potentially fail during emergencies if there is a surge of users, communication networks or devices become inoperative and if fires escalate faster than warnings can be updated¹⁷⁸. Warnings of fire danger and smoke hazards need to be combined with a number of other targeted interventions to protect health and wellbeing, such as provision of fireproof and clean air refuges, and community-level support for medically vulnerable individuals¹⁷⁹ (FIG. 6).

Preventing fire-related emissions. Fire regimes are not only impacted by climate change but could lead to net carbon emissions from terrestrial carbon stores. Tropical savannahs are the most fire-prone biomes on Earth and emit the greatest volume of greenhouse gases annually⁵⁰. Emissions can potentially be reduced through concentrating burning in the early dry season, when fire intensity is lower compared with the late

dry season¹⁸⁰. This approach is being attempted by an Australian programme¹⁸¹, with claims it has successfully mitigated the emission of non-CO₂ greenhouse gases, particularly methane (CH₄) and nitrous oxide (N₂O). For instance, application of this approach to 28,000 km² of *Eucalyptus* savannah in Arnhem Land has been estimated to have reduced accountable non-CO₂ emissions by 37.7% relative to emissions in the decade prior to the intervention¹⁸². Careful monitoring and evaluation are required, though, to ensure that managed fire regimes do not have unwanted socio-ecological side effects such as harming biodiversity¹⁸³, promoting a grass-fire cycle¹⁸⁴ with declines in above-ground carbon storage due to increased tree mortality¹⁸⁵ or disempowering indigenous communities¹⁸⁶.

Fires associated with tropical deforestation emit carbon stored in vegetation and soil from areas that otherwise rarely burn⁴⁵. To reduce the impacts of deforestation fires on carbon storage in the tropics, deforestation must be reduced through policy, economic incentives and education. Although some progress has been made, these gains can be quickly lost, as evidenced by the 2019 increase in deforestation fires in the Amazon following the abrupt reversal of the Brazilian government's commitments to control deforestation¹⁸⁷ (FIG. 3j).

The increasing occurrence of fires in huge, circumpolar expanses of coniferous boreal forests also needs careful management, as these fires emit greenhouse gases and create smoke that cross international boundaries¹⁸⁸. These boreal forests and their deep organic soils are one of the largest terrestrial carbon stores, containing somewhere between 367.3 Pg and 1,715.8 Pg of carbon¹⁸⁹. Around 0.2 Pg of carbon is emitted from the boreal zone annually, and emissions are likely to escalate in response to global warming, and past fire and logging disturbances, stresses that have been shown to strongly interact^{51,190}. Reducing boreal forest fire requires a multipronged and international approach, which involves reducing accidental human ignitions through education and enforcement. It will also involve skilful management of natural and anthropogenic fires to create a landscape mosaic of burned and recovering areas, reducing the capacity for extensive fires to develop¹²³.

Large-scale tree planting and natural-forest restoration has recently been suggested as pivotal to drawing down atmospheric CO₂ (REFS^{191–193}), but there are substantial caveats to this idea^{194–198}. Critically, any large-scale reforestation programme to store carbon must have effective and sustainable fire management, as large expanses of young trees are particularly vulnerable to high-severity fires^{199,200}. This vulnerability was exemplified by the widespread destruction of *Pinus* and *Eucalyptus* plantations by fires in Chile, Portugal and Australia during recent fire disasters⁹⁶. Creating fire-resilient timber plantations of flammable but productive timber remains a research challenge, owing to the difficulty and danger of using prescribed burning to reduce fire hazard, and mostly likely will require mechanical removal of fuel⁹⁶. Restoration plantings designed to restore habitat or sequester carbon should focus on either restoring fire-resistant ecosystems¹⁵⁸ or target non-flammable species¹⁶⁷. Regardless,

investment in fire suppression will be mandatory in all cases.

Summary and future perspectives

The flammability of many landscapes is increasing because of the combined effects of changing climate and land-use patterns³. The increased risk of economically and ecologically destructive fires can be reduced using planning and urban-design principles, combined with fuel management and fire management. Development of these fire-management interventions requires transdisciplinary research that combines insights from natural and social sciences, engineering and technology, and humanities^{33,201,202}. Such research is also prerequisite for improving global and regional fire models of future fire activity^{108,121}, which are needed to explore the fire-vegetation-atmosphere feedbacks on terrestrial carbon dynamics (FIG. 4). However, achieving better fire management requires addressing a number of key research challenges.

A basic research need is the description of regional fire regimes that quantify the seasonality, frequency, spatial extent and severity of fires, and then the determination of how these patterns are shaped by climate, terrain and vegetation. Identifying the cause of fires, at the most basic level differentiating natural and anthropogenic ignitions, is also essential. Achieving this basic goal demands nationally and internationally coordinated research that can yield insights through comparative analysis of the development of general ecological principles and evidence-based analysis of the effectiveness of fire-management and adaptation approaches. Within the regional scale, research must address how fire behaviour, particularly uncontrollable extreme fire events, is affected by the interactions between climate, weather, terrain, and the load and type of fuel. Indeed, there are increasing reports from frontline firefighters that the behaviour of many wildfires lies well outside their experience and exhibit unusual characteristics, such as pyro-tornadoes, so it is essential that extreme fire phenomena are systematically documented²⁵.

Better understanding of global and regional trends in fire regimes and vegetation-fire emissions demands improved data acquisition and database assembly. Existing satellite records only allow for coarse-scale mapping, based around a few easily acquired fire-regime metrics, such as seasonality, intensity and area burned³⁶. These data, combined with more accurate vegetation mapping, biomass-structure modelling and emissions factors, are essential for accurate estimates of greenhouse emissions and, hence, global carbon budgets²⁰³. However, the next generation of satellite imagery will enable more complete parameterization of fire regimes²⁰⁴, including the progression and behaviour of individual fires and fine-grain mapping of fire mosaics²⁰⁵. Of prime importance is moving from simplistic counts of fire detection or area-burned mapping to understanding the severity of fires and the associated impacts on vegetation defoliation, plant death, habitat quality and carbon stocks²⁰³. The recently established Global Fire Atlas will provide a central repository on remote sensing of individual fires and their location,

Pyro-tornadoes

Extreme fire behaviours can result in intense localized convection that spawns a violently rotating column of burning gases and debris.

duration, pattern of growth and spatial behaviour²⁰⁴. Furthermore, case studies can be undertaken to understand the effects of land and fuel management and firefighting approaches on wildfire behaviour and document extreme fire behaviour^{25,206}.

Improved understanding and modelling of interactions between fire and the hydrosphere, geosphere, cryosphere and atmosphere is necessary. Fire and fuel cannot be managed without impacts to these other Earth systems, and integrating dynamic feedback processes across systems will be critical to predicting both consequences and beneficial outcomes of future anthropogenic fire regimes. For example, municipal and agricultural water supplies are often dependent on the delivery of water of a consistent quality and quantity from fire-impacted vegetated watersheds²⁰⁷, so fire management of that vegetation must also consider potential outcomes on projected water storage and delivery²⁰⁸.

As anthropogenic fire regimes are influenced by human activity, there is a considerable need for data collection to improve quantification and modelling of fire activity and human populations, including their socio-economic status and historical and cultural-political legacies. Anthropological research into 'fire cultures' can improve our understanding of human relationships with fire use and illuminate why some societies have effectively coexisted with flammable landscapes for millennia¹⁵. Social science is critical in both understanding how contemporary societies relate to fire, smoke and fire-management strategies and identifying policy barriers, in order to establish pathways for transforming landscapes and achieving management objectives. Risk and disaster sciences are similarly important, particularly in exploring linkages between fire and post-fire geophysical disasters, as well as modifying existing established risk frameworks to include fire. Economic models are essential to informing investment in sustainable fire management and understanding the full costs of fire disasters, including human health impacts. A global database (similar in form to the Global Fire Atlas) on the economic costs of wildfire management and economics losses from fire disasters is required⁵⁵, as, presently, these data are only available for a few regions (such as the USA

and Canada)^{54,209}, and methods, terminology and data sources are inconsistently applied²⁶. Such economic modelling and policy development are also required to design fire-management programmes to reduce, and, ideally, reverse, greenhouse gas emissions from wildfires and tropical deforestation.

Transdisciplinary pyrogeographic research remains difficult to implement because of numerous intellectual and institutional barriers, such as the limited interaction between the humanities and physical sciences²⁰². A critical challenge for wildfire adaptation is the capacity to undertake transdisciplinary, applied research that is on the landscape scale and co-designed and funded by affected communities, multiple government and non-government agencies and landowners. Studies of the impacts and effectiveness of firefighting and emergency management responses to major unprecedented fire events, such as the recent California fires and the 2019–2020 Australian bushfire season, are, by definition, difficult to conduct because they cannot be prospectively planned, funded nor approved by regulators. Furthermore, the capacity of the global fire-science community is being increasingly stretched by the surge of anomalous fire events. Ongoing local, regional, national and international investment in training, institutional capacity building and fostering diversity amongst researchers and practitioners and their approaches is urgently required. This needs to be strategically combined with research and development and evaluation of wildfire mitigation and adaptation strategies. Without such investment, wildfire management and adaptation cannot be evidence-based or cost-effective.

Achieving sustainable stewardship of fire regimes requires acknowledging that fire is an inherent feature of the Earth system that has long been modified by humans. Climate change and numerous other anthropogenic effects on fire regimes present a new, urgent challenge for effective human adaptation to vegetation fire. Coordinated transdisciplinary research can lead to the development of fire-management strategies and avoid adverse economic and environmental impacts of fire disasters.

1. Scott, A. C., Bowman, D. M., Bond, W. J., Pyne, S. J. & Alexander, M. E. *Fire on Earth: An Introduction* (Wiley, 2013).
2. Ward, D. et al. The changing radiative forcing of fires: global model estimates for past, present and future. *Atmos. Chem. Phys.* **12**, 10857–10886 (2012).
3. Bowman, D. M. et al. Fire in the Earth system. *Science* **324**, 481–484 (2009).
4. Carslaw, K. et al. A review of natural aerosol interactions and feedbacks within the Earth system. *Atmos. Chem. Phys.* **10**, 1701–1737 (2010).
5. Peterson, D. A. et al. Wildfire-driven thunderstorms cause a volcano-like stratospheric injection of smoke. *NPJ Clim. Atmos. Sci.* **1**, 30 (2018).
6. McRae, R. H., Sharples, J. J. & Fromm, M. Linking local wildfire dynamics to pyroCb development. *Nat. Hazards Earth Syst. Sci.* **15**, 417–428 (2015).
7. Rosenfeld, D. TRMM observed first direct evidence of smoke from forest fires inhibiting rainfall. *Geophys. Res. Lett.* **26**, 3105–3108 (1999).
8. Thomas, J. L. et al. Quantifying black carbon deposition over the Greenland ice sheet from forest fires in Canada. *Geophys. Res. Lett.* **44**, 7965–7974 (2017).
9. Skiles, S. M., Flanner, M., Cook, J. M., Dumont, M. & Painter, T. H. Radiative forcing by light-absorbing particles in snow. *Nat. Clim. Change* **8**, 964–971 (2018).
10. Krebs, P., Pezzatti, G. B., Mazzoleni, S., Talbot, L. M. & Conedera, M. Fire regime: history and definition of a key concept in disturbance ecology. *Theory Biosci.* **129**, 53–69 (2010).
11. Keeley, J. E. & Fotheringham, C. Role of fire in regeneration from seed. *Seeds* **2**, 311–330 (2000).
12. Noble, I. R. & Slatyer, R. O. The use of vital attributes to predict successional changes in plant communities subject to recurrent disturbances. *Vegetatio* **43**, 5–21 (1980).
13. Enright, N. J., Fontaine, J. B., Bowman, D. M., Bradstock, R. A. & Williams, R. J. Interval squeeze: altered fire regimes and demographic responses interact to threaten woody species persistence as climate changes. *Front. Ecol. Environ.* **13**, 265–272 (2015).
14. Glikson, A. Fire and human evolution: the deep-time blueprints of the Anthropocene. *Anthropocene* **3**, 89–92 (2013).
15. Huffman, M. R. The many elements of traditional fire knowledge: synthesis, classification, and aids to cross-cultural problem solving in fire-dependent systems around the world. *Ecol. Soc.* **18**, 3 (2013).
16. Trauernicht, C., Brook, B. W., Murphy, B. P., Williamson, G. J. & Bowman, D. M. Local and global pyrogeographic evidence that indigenous fire management creates biodiversity. *Ecol. Evol.* **5**, 1908–1918 (2015).
17. Scherjon, F. et al. Burning the land: an ethnographic study of off-site fire use by current and historically documented foragers and implications for the interpretation of past fire practices in the landscape. *Curr. Anthropol.* **56**, 314–315 (2015).
18. Mertz, O. et al. Swidden change in Southeast Asia: understanding causes and consequences. *Hum. Ecol.* **37**, 259–264 (2009).
19. Bowman, D. M. et al. The human dimension of fire regimes on Earth. *J. Biogeogr.* **38**, 2223–2236 (2011).
20. Calkin, D. E., Stonesifer, C. S., Thompson, M. P. & McHugh, C. W. Large airtanker use and outcomes in suppressing wildland fires in the United States. *Int. J. Wildland Fire* **23**, 259–271 (2014).

21. Le Page, Y., Oom, D., Silva, J. M., Jönsson, P. & Pereira, J. M. Seasonality of vegetation fires as modified by human action: observing the deviation from eco-climatic fire regimes. *Glob. Ecol. Biogeogr.* **19**, 575–588 (2010).
22. Abatzoglou, J. T., Williams, A. P., Boschetti, L., Zubkova, M. & Kolden, C. A. Global patterns of interannual climate–fire relationships. *Glob. Change Biol.* **24**, 5164–5175 (2018).
23. Bowman, D. M. et al. Human exposure and sensitivity to globally extreme wildfire events. *Nat. Ecol. Evol.* **1**, 0058 (2017).
24. Sharples, J. J. et al. Natural hazards in Australia: extreme bushfire. *Clim. Change* **139**, 85–99 (2016).
25. Tedim, F. et al. Defining extreme wildfire events: difficulties, challenges, and impacts. *Fire* **1**, 9 (2018).
26. Ladds, M., Keating, A., Handmer, J. & Magee, L. How much do disasters cost? A comparison of disaster cost estimates in Australia. *Int. J. Disaster Risk Reduct.* **21**, 419–429 (2017).
27. Kramer, H. A., Mockrin, M. H., Alexandre, P. M. & Radeloff, V. C. High wildfire damage in interface communities in California. *Int. J. Wildland Fire* **28**, 641–650 (2019).
28. Thomas, D., Butry, D., Gilbert, S., Webb, D. & Fung, J. *The Costs and Losses of Wildfires. NIST Special Publication 1215* (NIST, 2017).
29. Fann, N. et al. The health impacts and economic value of wildland fire episodes in the US: 2008–2012. *Sci. Total Environ.* **610**, 802–809 (2018).
30. Read, P. & Denniss, R. With costs approaching \$100 billion, the fires are Australia's costliest natural disaster. *The Conversation* <https://theconversation.com/with-costs-approaching-100-billion-the-fires-are-australias-costliest-natural-disaster-129433> (2020).
31. Steffen, W. et al. Trajectories of the Earth System in the Anthropocene. *Proc. Natl Acad. Sci. USA* **115**, 8252–8259 (2018).
32. Steffen, W. et al. The emergence and evolution of Earth system science. *Nat. Rev. Earth Environ.* **1**, 54–63 (2020).
33. Bowman, D. M., O'Brien, J. A. & Goldammer, J. G. Pyrogeography and the global quest for sustainable fire management. *Annu. Rev. Environ. Resour.* **38**, 57–80 (2013).
34. Giglio, L., Boschetti, L., Roy, D. P., Humber, M. L. & Justice, C. O. The Collection 6 MODIS burned area mapping algorithm and product. *Remote Sens. Environ.* **217**, 72–85 (2018).
35. Randerson, J. T., Chen, Y., Van Der Werf, G. R., Rogers, B. M. & Morton, D. C. Global burned area and biomass burning emissions from small fires. *J. Geophys. Res.* **117**, G04012 (2012).
36. Archibald, S., Lehmann, C. E., Gómez-Dans, J. L. & Bradstock, R. A. Defining pyromes and global syndromes of fire regimes. *Proc. Natl Acad. Sci. USA* **110**, 6442–6447 (2013).
37. Kelley, D. I. et al. How contemporary bioclimatic and human controls change global fire regimes. *Nat. Clim. Change* **9**, 690–696 (2019).
38. Lavorel, S., Flannigan, M. D., Lambin, E. F. & Scholes, M. C. Vulnerability of land systems to fire: Interactions among humans, climate, the atmosphere, and ecosystems. *Mitig. Adapt. Strateg. Global Change* **12**, 33–53 (2007).
39. Van Der Werf, G. R., Randerson, J. T., Giglio, L., Gobron, N. & Dolman, A. Climate controls on the variability of fires in the tropics and subtropics. *Global Biogeochem. Cycles* **22** (2008).
40. Krawchuk, M. A., Moritz, M. A., Parisien, M.-A., Van Dorn, J. & Hayhoe, K. Global pyrogeography: the current and future distribution of wildfire. *PLoS ONE* **4**, e5102 (2009).
41. Pausas, J. G. & Ribeiro, E. The global fire–productivity relationship. *Global Ecol. Biogeogr.* **22**, 728–736 (2013).
42. McKenzie, D. & Littell, J. S. Climate change and the eco-hydrology of fire: Will area burned increase in a warming western USA? *Ecol. Appl.* **27**, 26–36 (2017).
43. Bowman, D. M., Murphy, B. P., Williamson, G. J. & Cochrane, M. A. Pyrogeographic modes, feedbacks and the future of global fire regimes. *Global Ecol. Biogeogr.* **23**, 821–824 (2014).
44. Haberle, S. G., Hope, G. S. & van der Kaars, S. Biomass burning in Indonesia and Papua New Guinea: natural and human induced fire events in the fossil record. *Palaeogeogr. Palaeoclimatol. Palaeoecol.* **171**, 259–268 (2001).
45. Cochrane, M. A. Fire science for rainforests. *Nature* **421**, 913–919 (2003).
46. Flannigan, M. D., Logan, K. A., Amiro, B. D., Skinner, W. R. & Stocks, B. Future area burned in Canada. *Clim. Change* **72**, 1–16 (2005).
47. Abatzoglou, J. T. & Kolden, C. A. Climate change in western US deserts: potential for increased wildfire and invasive annual grasses. *Rangel. Ecol. Manag.* **64**, 471–478 (2011).
48. Balch, J. K., Bradley, B. A., D'Antonio, C. M. & Gómez-Dans, J. Introduced annual grass increases regional fire activity across the arid western USA (1980–2009). *Glob. Change Biol.* **19**, 173–183 (2013).
49. Andreae, M. O. & Merlet, P. Emission of trace gases and aerosols from biomass burning. *Global Biogeochem. Cycles* **15**, 955–966 (2001).
50. Van Der Werf, G. R. et al. Global fire emissions estimates during 1997–2016. *Earth Syst. Sci. Data* **9**, 697–720 (2017).
51. Walker, X. J. et al. Increasing wildfires threaten historic carbon sink of boreal forest soils. *Nature* **572**, 520–523 (2019).
52. Cramer, W. et al. Tropical forests and the global carbon cycle: impacts of atmospheric carbon dioxide, climate change and rate of deforestation. *Philos. Trans. R. Soc. B Biol. Sci.* **359**, 331–343 (2004).
53. Kurz, W. et al. Quantifying the impacts of human activities on reported greenhouse gas emissions and removals in Canada's managed forest: conceptual framework and implementation. *Can. J. For. Res.* **48**, 1227–1240 (2018).
54. Doerr, S. H. & Santin, C. Global trends in wildfire and its impacts: perceptions versus realities in a changing world. *Philos. Trans. R. Soc. B Biol. Sci.* **371**, 20150345 (2016).
55. Bowman, D. Wildfire science is at a loss for comprehensive data. *Nature* **560**, 7–8 (2018).
56. Foreman, P. W. A framework for testing the influence of Aboriginal burning on grassy ecosystems in lowland, mesic south-eastern Australia. *Australian J. Botany* **64**, 626–642 (2016).
57. Van Wagner, C. Age-class distribution and the forest fire cycle. *Can. J. For. Res.* **8**, 220–227 (1978).
58. Larsen, C. P. S. Fire and climate dynamics in the boreal forest of northern Alberta, Canada, from AD 1850 to 1989. *Holocene* **6**, 449–456 (1996).
59. Bergeron, Y., Flannigan, M., Gauthier, S., Leduc, A. & Lefort, P. Past, current and future fire frequency in the Canadian boreal forest: implications for sustainable forest management. *AMBIO* **33**, 356–360 (2004).
60. Marlon, J. R. et al. Long-term perspective on wildfires in the western USA. *Proc. Natl Acad. Sci. USA* **109**, E535–E543 (2012).
61. Marlon, J. R. et al. Climate and human influences on global biomass burning over the past two millennia. *Nat. Geosci.* **1**, 697–702 (2008).
62. Chuvpilo, E. et al. Historical background and current developments for mapping burned area from satellite Earth observation. *Remote Sens. Environ.* **225**, 45–64 (2019).
63. Forkel, M. et al. Recent global and regional trends in burned area and their compensating environmental controls. *Environ. Res. Commun.* **1**, 051005 (2019).
64. Schultz, M. G. et al. Global wildland fire emissions from 1960 to 2000. *Global Biogeochem. Cycles* **22**, GB2002 (2008).
65. Clode, D. & Elgar, M. A. Fighting fire with fire: does a policy of broad-scale prescribed burning improve community safety? *Soc. Nat. Resour.* **27**, 1192–1199 (2014).
66. Keeley, J. E. Fire intensity, fire severity and burn severity: a brief review and suggested usage. *Int. J. Wildland Fire* **18**, 116–126 (2009).
67. Jolly, W. M. et al. Climate-induced variations in global wildfire danger from 1979 to 2013. *Nat. Commun.* **6**, 7537 (2015).
68. Justice, C. et al. The MODIS fire products. *Remote Sens. Environ.* **83**, 244–262 (2002).
69. Andela, N. et al. A human-driven decline in global burned area. *Science* **356**, 1356–1362 (2017).
70. Burrows, N., Ward, B. & Robinson, A. Fuel dynamics and fire spread in spinifex grasslands of the Western Desert. *Proc. R. Soc. Queensland* **115**, 69–76 (2009).
71. Bird, R. B., Coddling, B. F., Kauhane, P. G. & Bird, D. W. Aboriginal hunting buffers climate-driven fire-size variability in Australia's spinifex grasslands. *Proc. Natl Acad. Sci. USA* **109**, 10287–10292 (2012).
72. Taylor, A. H., Trouet, V., Skinner, C. N. & Stephens, S. Socioecological transitions trigger fire regime shifts and modulate fire–climate interactions in the Sierra Nevada, USA, 1600–2015 CE. *Proc. Natl Acad. Sci. USA* **113**, 13684–13689 (2016).
73. Westerling, A. L., Hidalgo, H. G., Cayan, D. R. & Swetnam, T. W. Warming and earlier spring increase western US forest wildfire activity. *Science* **313**, 940–943 (2006).
74. Abatzoglou, J. T. & Williams, A. P. Impact of anthropogenic climate change on wildfire across western US forests. *Proc. Natl Acad. Sci. USA* **113**, 11770–11775 (2016).
75. Balch, J. K. et al. Switching on the Big Burn of 2017. *Fire* **1**, 17 (2018).
76. Williams, A. P. et al. Observed impacts of anthropogenic climate change on wildfire in California. *Earths Future* **7**, 892–910 (2019).
77. Westerling, A. L. Increasing western US forest wildfire activity: sensitivity to changes in the timing of spring. *Philos. Trans. R. Soc. B Biol. Sci.* **371**, 20150178 (2016).
78. Balshi, M. S. et al. Assessing the response of area burned to changing climate in western boreal North America using a Multivariate Adaptive Regression Splines (MARS) approach. *Glob. Change Biol.* **15**, 578–600 (2009).
79. Jain, P., Wang, X. & Flannigan, M. D. Trend analysis of fire season length and extreme fire weather in North America between 1979 and 2015. *Int. J. Wildland Fire* **26**, 1009–1020 (2018).
80. Flannigan, M. et al. Fuel moisture sensitivity to temperature and precipitation: climate change implications. *Clim. Change* **134**, 59–71 (2016).
81. Hanes, C. C. et al. Fire-regime changes in Canada over the last half century. *Can. J. For. Res.* **49**, 256–269 (2018).
82. Veraverbeke, S. et al. Lightning as a major driver of recent large fire years in North American boreal forests. *Nat. Clim. Change* **7**, 529–534 (2017).
83. Turetsky, M. R. et al. Global vulnerability of peatlands to fire and carbon loss. *Nat. Geosci.* **8**, 11–14 (2015).
84. O'Connor, F. M. et al. Possible role of wetlands, permafrost, and methane hydrates in the methane cycle under future climate change: A review. *Rev. Geophys.* **48**, RG4005 (2010).
85. Gibson, C. M. et al. Wildfire as a major driver of recent permafrost thaw in boreal peatlands. *Nat. Commun.* **9**, 3041 (2018).
86. Bowman, D. M., Walsh, A. & Prior, L. D. Landscape analysis of Aboriginal fire management in Central Arnhem Land, north Australia. *J. Biogeogr.* **31**, 207–223 (2004).
87. Bird, R. B., Bird, D. W., Coddling, B. F., Parker, C. H. & Jones, J. H. The “fire stick farming” hypothesis: Australian Aboriginal foraging strategies, biodiversity, and anthropogenic fire mosaics. *Proc. Natl Acad. Sci. USA* **105**, 14796–14801 (2008).
88. Cruz, M. et al. Anatomy of a catastrophic wildfire: the Black Saturday Kilmore East fire in Victoria, Australia. *For. Ecol. Manag.* **284**, 269–285 (2012).
89. Ndalila, M. N., Williamson, G. J. & Bowman, D. M. Geographic patterns of fire severity following an extreme *Eucalyptus* forest fire in southern Australia: 2013 Forcett-Dunalley fire. *Fire* **1**, 40 (2018).
90. Di Virgilio, G. et al. Climate change increases the potential for extreme wildfires. *Geophys. Res. Lett.* **46**, 8517–8526 (2019).
91. Styger, J., Marsden-Smedley, J. & Kirkpatrick, J. Changes in lightning fire incidence in the Tasmanian Wilderness World Heritage Area, 1980–2016. *Fire* **1**, 38 (2018).
92. Bowman, D. M., Bliss, A., Bowman, C. J. & Prior, L. D. Fire caused demographic attrition of the Tasmanian palaeoendemic conifer *Athrotaxis cupressoides*. *Austral Ecol.* **44**, 1322–1339 (2019).
93. Boer, M. M., de Dios, V. R. & Bradstock, R. A. Unprecedented burn area of Australian mega forest fires. *Nat. Clim. Change* **10**, 171–172 (2020).
94. van Oldenborgh, G. J. et al. Attribution of the Australian bushfire risk to anthropogenic climate change. *Nat. Hazards Earth Syst. Sci. Discuss.* **2020**, 1–46 (2020).
95. Bowman, D. M. et al. Human–environmental drivers and impacts of the globally extreme 2017 Chilean fires. *AMBIO* **48**, 350–362 (2019).
96. Gómez-González, S., Ojeda, F. & Fernandes, P. M. Portugal and Chile: Longing for sustainable forestry while rising from the ashes. *Environ. Sci. Policy* **81**, 104–107 (2018).
97. Fusco, E. J., Finn, J. T., Balch, J. K., Nagy, R. C. & Bradley, B. A. Invasive grasses increase fire occurrence and frequency across US ecoregions. *Proc. Natl Acad. Sci. USA* **116**, 23594–23599 (2019).
98. Setterfield, S. A., Rossiter-Rachor, N. A., Hutley, L. B., Douglas, M. M. & Williams, R. J. Biodiversity research: turning up the heat: the impacts of *Andropogon*

- gayanus* (gamba grass) invasion on fire behaviour in northern Australian savannas. *Divers. Distrib.* **16**, 854–861 (2010).
99. Van Marle, M. J. et al. Historic global biomass burning emissions based on merging satellite observations with proxies and fire models (1750–2015). *Geosci. Model Dev.* **10**, 3329–3357 (2017).
100. Le Quéré, C. et al. Global carbon budget 2018. *Earth Syst. Sci. Data* **10**, 2141–2194 (2018).
101. Wooster, M. J., Perry, G. L. W. & Zoumas, A. Fire, drought and El Niño relationships on Borneo (Southeast Asia) in the pre-MODIS era (1980–2000). *Biogeosciences* **9**, 317–340 (2012).
102. Chen, Y. et al. A pan-tropical cascade of fire driven by El Niño/Southern Oscillation. *Nat. Clim. Change* **7**, 906–911 (2017).
103. Stocker, T. F. et al. (eds) *Climate Change 2013: the Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change* 1535 pp (Cambridge Univ. Press, 2013).
104. Flannigan, M. et al. Global wildland fire season severity in the 21st century. *For. Ecol. Manag.* **294**, 54–61 (2013).
105. Abatzoglou, J. T., Williams, A. P. & Barbero, R. Global emergence of anthropogenic climate change in fire weather indices. *Geophys. Res. Lett.* **46**, 326–336 (2019).
106. Romps, D. M., Seeley, J. T., Vollaro, D. & Molinari, J. Projected increase in lightning strikes in the United States due to global warming. *Science* **346**, 851–854 (2014).
107. Knorr, W., Arneith, A. & Jiang, L. Demographic controls of future global fire risk. *Nat. Clim. Change* **6**, 781–785 (2016).
108. Kloster, S. & Lasslop, G. Historical and future fire occurrence (1850 to 2100) simulated in CMIP5 Earth System Models. *Glob. Planet. Change* **150**, 58–69 (2017).
109. Moritz, M. A. et al. Climate change and disruptions to global fire activity. *Ecosphere* **3**, 1–22 (2012).
110. Wotton, B. M. & Flannigan, M. D. Length of the fire season in a changing climate. *Forestry Chron.* **69**, 187–192 (1993).
111. Wotton, B., Flannigan, M. & Marshall, G. Potential climate change impacts on fire intensity and key wildfire suppression thresholds in Canada. *Environ. Res. Lett.* **12**, 095003 (2017).
112. Barbero, R., Abatzoglou, J. T., Larkin, N. K., Kolden, C. A. & Stocks, B. Climate change presents increased potential for very large fires in the contiguous United States. *Int. J. Wildland Fire* **24**, 892–899 (2015).
113. Westerling, A. L., Turner, M. G., Smithwick, E. A., Romme, W. H. & Ryan, M. G. Continued warming could transform Greater Yellowstone fire regimes by mid-21st century. *Proc. Natl Acad. Sci. USA* **108**, 13165–13170 (2011).
114. Buotte, P. C. et al. Near-future forest vulnerability to drought and fire varies across the western United States. *Glob. Change Biol.* **25**, 290–303 (2019).
115. Kitzberger, T., Falk, D. A., Westerling, A. L. & Swetnam, T. W. Direct and indirect climate controls predict heterogeneous early-mid 21st century wildfire burned area across western and boreal North America. *PLoS ONE* **12**, e0188486 (2017).
116. Turco, M. et al. Exacerbated fires in Mediterranean Europe due to anthropogenic warming projected with non-stationary climate-fire models. *Nat. Commun.* **9**, 3821 (2018).
117. Turco, M. et al. Decreasing fires in Mediterranean Europe. *PLoS ONE* **11**, e0150663 (2016).
118. Battlori, E., Parisien, M. A., Krawchuk, M. A. & Moritz, M. A. Climate change-induced shifts in fire for Mediterranean ecosystems. *Global Ecol. Biogeogr.* **22**, 1118–1129 (2013).
119. Mekonnen, Z. A., Riley, W. J., Randerson, J. T., Grant, R. F. & Rogers, B. M. Expansion of high-latitude deciduous forests driven by interactions between climate warming and fire. *Nat. Plants* **5**, 952–958 (2019).
120. Harris, R. M., Remenyi, T. A., Williamson, G. J., Bindoff, N. L. & Bowman, D. M. Climate–vegetation–fire interactions and feedbacks: trivial detail or major barrier to projecting the future of the Earth system? *Wiley Interdiscip. Rev. Clim. Change* **7**, 910–931 (2016).
121. Hantson, S. et al. The status and challenge of global fire modelling. *Biogeosciences* **13**, 3359–3375 (2016).
122. Pechony, O. & Shindell, D. T. Driving forces of global wildfires over the past millennium and the forthcoming century. *Proc. Natl Acad. Sci. USA* **107**, 19167–19170 (2010).
123. Flannigan, M., Stocks, B., Turetsky, M. & Wotton, M. Impacts of climate change on fire activity and fire management in the circumboreal forest. *Glob. Change Biol.* **15**, 549–560 (2009).
124. Podur, J. & Wotton, M. Will climate change overwhelm fire management capacity? *Ecol. Model.* **221**, 1301–1309 (2010).
125. Teckentrup, L. et al. Response of simulated burned area to historical changes in environmental and anthropogenic factors: a comparison of seven fire models. *Biogeosciences* **16**, 3883–3910 (2019).
126. Nepstad, D. C., Stickler, C. M., Filho, B. S. & Merry, F. Interactions among Amazon land use, forests and climate: prospects for a near-term forest tipping point. *Philos. Trans. R. Soc. B Biol. Sci.* **363**, 1737–1746 (2008).
127. Swann, A. L., Hoffman, F. M., Koven, C. D. & Randerson, J. T. Plant responses to increasing CO₂ reduce estimates of climate impacts on drought severity. *Proc. Natl Acad. Sci. USA* **113**, 10019–10024 (2016).
128. Bond, W. J. & Midgley, G. F. Carbon dioxide and the uneasy interactions of trees and savannah grasses. *Philos. Trans. R. Soc. B Biol. Sci.* **367**, 601–612 (2012).
129. Hurteau, M. D., Liang, S., Westerling, A. L. & Wiedinmyer, C. Vegetation–fire feedback reduces projected area burned under climate change. *Sci. Rep.* **9**, 2838 (2019).
130. Liu, Z. & Wimberly, M. C. Direct and indirect effects of climate change on projected future fire regimes in the western United States. *Sci. Total Environ.* **542**, 65–75 (2016).
131. Stevens-Rumann, C. S. et al. Evidence for declining forest resilience to wildfires under climate change. *Ecol. Lett.* **21**, 243–252 (2018).
132. Wilkin, K., Ackerly, D. & Stephens, S. Climate change refugia, fire ecology and management. *Forests* **7**, 77 (2016).
133. Kashian, D. M., Romme, W. H., Tinker, D. B., Turner, M. G. & Ryan, M. G. Carbon storage on landscapes with stand-replacing fires. *Bioscience* **56**, 598–606 (2006).
134. Wiggins, E. B. et al. Smoke radiocarbon measurements from Indonesian fires provide evidence for burning of millennia-aged peat. *Proc. Natl Acad. Sci. USA* **115**, 12419–12424 (2018).
135. Donovan, V. M., Wonkka, C. L. & Twidwell, D. Surging wildfire activity in a grassland biome. *Geophys. Res. Lett.* **44**, 5986–5993 (2017).
136. Bladon, K. D. Rethinking wildfires and forest watersheds. *Science* **359**, 1001–1002 (2018).
137. Cannon, S. H. & DeGraff, J. in *Landslides–Disaster Risk Reduction* (eds Sassa, K. & Canuti, P.) 177–190 (Springer, 2009).
138. Garfin, G. et al. *Managing for Future Risks of Fire, Extreme Precipitation, and Post-fire Flooding. Report to the U.S. Bureau of Reclamation, from the Project Enhancing Water Supply Reliability* (Institute of the Environment, 2016).
139. Sanderson, B. M. & Fisher, R. A. A fiery wake-up call for climate science. *Nat. Clim. Change* **515**, 175–177 (2020).
140. King, A. D., Pitman, A. J., Henley, B. J., Ukkola, A. M. & Brown, J. R. The role of climate variability in Australian drought. *Nat. Clim. Change* **10**, 177–179 (2020).
141. Moritz, M. A. et al. Learning to coexist with wildfire. *Nature* **515**, 58–66 (2014).
142. Kolden, C. A. We're not doing enough prescribed fire in the Western United States to mitigate wildfire risk. *Fire* **2**, 30 (2019).
143. Fernandes, P. M. & Botelho, H. S. A review of prescribed burning effectiveness in fire hazard reduction. *Int. J. Wildland Fire* **12**, 117–128 (2003).
144. Price, O. F., Penman, T. D., Bradstock, R. A., Boer, M. M. & Clarke, H. Biogeographical variation in the potential effectiveness of prescribed fire in south-eastern Australia. *J. Biogeogr.* **42**, 2234–2245 (2015).
145. Hurteau, M. D., Koch, G. W. & Hungate, B. A. Carbon protection and fire risk reduction: toward a full accounting of forest carbon offsets. *Front. Ecol. Environ.* **6**, 493–498 (2008).
146. Stephens, S. L. Evaluation of the effects of silvicultural and fuels treatments on potential fire behaviour in Sierra Nevada mixed-conifer forests. *For. Ecol. Manag.* **105**, 21–35 (1998).
147. Campbell, J. L., Harmon, M. E. & Mitchell, S. R. Can fuel-reduction treatments really increase forest carbon storage in the western US by reducing future fire emissions? *Front. Ecol. Environ.* **10**, 83–90 (2012).
148. Clarke, H. & Evans, J. P. Exploring the future change space for fire weather in southeast Australia. *Theor. Appl. Climatol.* **136**, 513–527 (2019).
149. Price, O. F. & Bradstock, R. A. The efficacy of fuel treatment in mitigating property loss during wildfires: Insights from analysis of the severity of the catastrophic fires in 2009 in Victoria, Australia. *J. Environ. Manag.* **113**, 146–157 (2012).
150. Williamson, G., Bowman, D. M. S., Price, O. F., Henderson, S. & Johnston, F. A transdisciplinary approach to understanding the health effects of wildfire and prescribed fire smoke regimes. *Environ. Res. Lett.* **11**, 125009 (2016).
151. Broome, R. A., Johnston, F. H., Horsley, J. & Morgan, G. G. A rapid assessment of the impact of hazard reduction burning around Sydney, May 2016. *Med. J. Aust.* **205**, 407–408 (2016).
152. U.S. Environmental Protection Agency, Office of Air and Radiation. *The Benefits and Costs of the Clean Air Act from 1990 to 2020: Final Report — Rev. A* [U.S. Environmental Protection Agency, Office of Air and Radiation, 2011].
153. Bowman, D. et al. Can air quality management drive sustainable fuels management at the temperate wildland–urban interface? *Fire* **1**, 27 (2018).
154. Mistry, J. & Berardi, A. Bridging indigenous and scientific knowledge. *Science* **352**, 1274–1275 (2016).
155. Reyes-García, V. & Benyei, P. Indigenous knowledge for conservation. *Nat. Sustain.* **2**, 657–658 (2019).
156. Bird, R. B., Taylor, N., Coddling, B. F. & Bird, D. W. Niche construction and Dreaming logic: aboriginal patch mosaic burning and varanid lizards (*Varanus gouldii*) in Australia. *Proc. R. Soc. B Biol. Sci.* **280**, 20132297 (2013).
157. Jackson, S. T. & Hobbs, R. J. Ecological restoration in the light of ecological history. *Science* **325**, 567–569 (2009).
158. Bowman, D. M. & Legge, S. Pyrodiversity — why managing fire in food webs is relevant to restoration ecology. *Restor. Ecol.* **24**, 848–853 (2016).
159. Schoennagel, T. et al. Adapt to more wildfire in western North American forests as climate changes. *Proc. Natl Acad. Sci. USA* **114**, 4582–4590 (2017).
160. Strader, S. M. Spatiotemporal changes in conterminous US wildfire exposure from 1940 to 2010. *Nat. Hazards* **92**, 543–565 (2018).
161. Balch, J. K. et al. Human-started wildfires expand the fire niche across the United States. *Proc. Natl Acad. Sci. USA* **114**, 2946–2951 (2017).
162. Fischer, A. P. et al. Wildfire risk as a socioecological pathology. *Front. Ecol. Environ.* **14**, 276–284 (2016).
163. Borchers Arriagada, N. et al. Unprecedented smoke-related health burden associated with the 2019–20 bushfires in eastern Australia. *Med. J. Aust.* <https://doi.org/10.5694/mja2.50545> (2020).
164. Fischer, A. P. et al. Wildfire risk as a socioecological pathology. *Front. Ecol. Environ.* **14**, 276–284 (2016).
165. Smith, A. M. et al. The science of fireescapes: achieving fire-resilient communities. *Bioscience* **66**, 130–146 (2016).
166. McWethy, D. B. et al. Rethinking resilience to wildfire. *Nat. Sustain.* **2**, 797–804 (2019).
167. Curran, T., Perry, G., Wyse, S. & Alam, M. Managing fire and biodiversity in the wildland-urban interface: A role for green firebreaks. *Fire* **1**, 3 (2018).
168. Bowman, D. M. J. S. & Stoof, C. Diversity helps fight wildfires. *Nature* **571**, 478 (2019).
169. Cui, X. et al. Green firebreaks as a management tool for wildfires: Lessons from China. *J. Environ. Manag.* **233**, 329–336 (2019).
170. Kolden, C. A. & Henson, C. A socio-ecological approach to mitigating wildfire vulnerability in the wildland urban interface: a case study from the 2017 Thomas fire. *Fire* **2**, 9 (2019).
171. Eriksen, C. *Gender and Wildfire: Landscapes of Uncertainty* (Routledge, 2013).
172. Huffman, M. R. Making a world of difference in fire and climate change. *Fire Ecol.* **10**, 90–101 (2014).
173. Pratt, M. et al. The implications of megatrends in information and communication technology and transportation for changes in global physical activity. *Lancet* **380**, 282–293 (2012).
174. Johnston, F. et al. Using smartphone technology to reduce health impacts from atmospheric environmental hazards. *Environ. Res. Lett.* **13**, 044019 (2018).
175. Lovreglio, R., Kuligowski, E., Gwynne, S. & Strahan, K. A modelling framework for household decision-making

- for wildfire emergencies. *Int. J. Disaster Risk Reduct.* **41**, 101274 (2019).
176. Kulemeka, O. A review of wildland fire smartphone applications: a classification study from Australia, USA, Canada and South Africa. *Int. J. Emerg. Serv.* **4**, 258–270 (2015).
177. Rappold, A. et al. Smoke Sense initiative leverages citizen science to address the growing wildfire-related public health problem. *GeoHealth* **3**, 443–457 (2019).
178. Maryam, H., Shah, M. A., Javaid, Q. & Kamran, M. A survey on smartphones systems for emergency management (SPSEM). *Int. J. Adv. Comput. Sci. Appl.* **7**, 301–311 (2016).
179. Vardoulakis, S., Jalaludin, B. B., Morgan, G. G., Hanigan, I. C. & Johnston, F. H. Bushfire smoke: urgent need for a national health protection strategy. *Med. J. Aust.* **212**, 349–353.e1 (2020).
180. Lipsett-Moore, G. J., Wolff, N. H. & Game, E. T. Emissions mitigation opportunities for savanna countries from early dry season fire management. *Nat. Commun.* **9**, 2247 (2018).
181. Russell-Smith, J. et al. Deriving multiple benefits from carbon market-based savanna fire management: An Australian example. *PLoS ONE* **10**, e0143426 (2015).
182. Russell-Smith, J. et al. Managing fire regimes in north Australian savannas: applying Aboriginal approaches to contemporary global problems. *Front. Ecol. Environ.* **11**, e55–e63 (2013).
183. Andersen, A. N., Woinarski, J. C. Z. & Parr, C. L. Savanna burning for biodiversity: Fire management for faunal conservation in Australian tropical savannas. *Austral Ecol.* **37**, 658–667 (2012).
184. Bowman, D. M., MacDermott, H. J., Nichols, S. C. & Murphy, B. P. A grass–fire cycle eliminates an obligate-seeding tree in a tropical savanna. *Ecol. Evol.* **4**, 4185–4194 (2014).
185. Murphy, B. P., Russell-Smith, J. & Prior, L. D. Frequent fires reduce tree growth in northern Australian savannas: implications for tree demography and carbon sequestration. *Glob. Change Biol.* **16**, 331–343 (2010).
186. Petty, A. M., deKoninck, V. & Orlove, B. Cleaning, protecting, or abating? Making indigenous fire management “work” in northern Australia. *J. Ethnobiol.* **35**, 140–163 (2015).
187. de Oliveira Andrade, R. Alarming surge in Amazon fires prompts global outcry. *Nature* <https://doi.org/10.1038/d41586-019-02537-0> (23 Aug 2019).
188. Kasischke, E. S., Christensen, N. Jr & Stocks, B. J. Fire, global warming, and the carbon balance of boreal forests. *Ecol. Appl.* **5**, 437–451 (1995).
189. Bradshaw, C. J. & Warkentin, I. G. Global estimates of boreal forest carbon stocks and flux. *Glob. Planet. Change* **128**, 24–30 (2015).
190. Dieleman, C. M. et al. Wildfire combustion and carbon stocks in the southern Canadian boreal forest: Implications for a warming world. *Glob. Change Biol.* <https://doi.org/10.1111/gcb.15158> (2020).
191. Bastin, J.-F. et al. The global tree restoration potential. *Science* **365**, 76–79 (2019).
192. Bastin, J.-F. et al. Response to comments on “The global tree restoration potential”. *Science* **366**, eaay8108 (2019).
193. Lewis, S. L., Wheeler, C. E., Mitchard, E. T. A. & Koch, A. Restoring natural forests is the best way to remove atmospheric carbon. *Nature* **568**, 25–28 (2019).
194. Veldman, J. W. et al. Comment on “The global tree restoration potential”. *Science* **366**, eaay7976 (2019).
195. Friedlingstein, P., Allen, M., Canadell, J. G., Peters, G. P. & Seneviratne, S. I. Comment on “The global tree restoration potential”. *Science* **366**, eaay8060 (2019).
196. Lewis, S. L., Mitchard, E. T., Prentice, C., Maslin, M. & Poulter, B. Comment on “The global tree restoration potential”. *Science* **366**, eaaz0388 (2019).
197. Grainger, A., Iverson, L. R., Marland, G. H. & Prasad, A. Comment on “The global tree restoration potential”. *Science* **366**, eaay8334 (2019).
198. Luedeling, E. et al. Forest restoration: Overlooked constraints. *Science* **366**, 315 (2019).
199. Lindenmayer, D. B., Hobbs, R. J., Likens, G. E., Krebs, C. J. & Banks, S. C. Newly discovered landscape traps produce regime shifts in wet forests. *Proc. Natl Acad. Sci. USA* **108**, 15887–15891 (2011).
200. Anderegg, W. R. et al. Climate-driven risks to the climate mitigation potential of forests. *Science* **368**, eaaz7005 (2020).
201. Nerini, F. F. et al. Connecting climate action with other sustainable development goals. *Nat. Sustain.* **2**, 674–680 (2019).
202. Castree, N. Speaking for the ‘people disciplines’: Global change science and its human dimensions. *Anthropocene Rev.* **4**, 160–182 (2017).
203. Stenzel, J. E. et al. Fixing a snag in carbon emissions estimates from wildfires. *Glob. Change Biol.* **25**, 3985–3994 (2019).
204. Andela, N. et al. The global fire atlas of individual fire size, duration, speed and direction. *Earth Syst. Sci. Data* **11**, 529–552 (2019).
205. Meng, R. et al. Using high spatial resolution satellite imagery to map forest burn severity across spatial scales in a Pine Barrens ecosystem. *Remote Sens. Environ.* **191**, 95–109 (2017).
206. Filkov, A., Duff, T. & Penman, T. Improving fire behaviour data obtained from wildfires. *Forests* **9**, 81 (2018).
207. White, I. et al. The vulnerability of water supply catchments to bushfires: impacts of the January 2003 wildfires on the Australian capital territory. *Australas. J. Water Resour.* **10**, 179–194 (2006).
208. Kliskey, A. et al. Planning for Idaho’s waterscapes: A review of historical drivers and outlook for the next 50 years. *Environ. Sci. Policy* **94**, 191–201 (2019).
209. Stocks, B. & Martell, D. L. Forest fire management expenditures in Canada: 1970–2013. *Forestry Chron.* **92**, 298–306 (2016).
210. Burton, C., Betts, R., Jones, C. & Williams, K. Will fire danger be reduced by using solar radiation management to limit global warming to 1.5 C compared to 2.0 C? *Geophys. Res. Lett.* **45**, 3644–3652 (2018).
211. Bedia, J. et al. Global patterns in the sensitivity of burned area to fire-weather: Implications for climate change. *Agric. For. Meteorol.* **214**, 369–379 (2015).
212. Liu, Y., Stanturf, J. & Goodrick, S. Trends in global wildfire potential in a changing climate. *For. Ecol. Manag.* **259**, 685–697 (2010).
213. Huang, Y., Wu, S. & Kaplan, J. O. Sensitivity of global wildfire occurrences to various factors in the context of global change. *Atmos. Environ.* **121**, 86–92 (2015).
214. de Groot, W. J., Flannigan, M. D. & Cantin, A. S. Climate change impacts on future boreal fire regimes. *For. Ecol. Manag.* **294**, 35–44 (2013).
215. Fonseca, M. G. et al. Effects of climate and land-use change scenarios on fire probability during the 21st century in the Brazilian Amazon. *Glob. Change Biol.* **25**, 2931–2946 (2019).
216. Le Page, Y. et al. Synergy between land use and climate change increases future fire risk in Amazon forests. *Earth Syst. Dynam.* **8**, 1237–1246 (2017).
217. Dowdy, A. J. et al. Future changes in extreme weather and pyroconvection risk factors for Australian wildfires. *Sci. Rep.* **9**, 10073 (2019).
218. Fox-Hughes, P., Harris, R., Lee, G., Grose, M. & Bindoff, N. Future fire danger climatology for Tasmania, Australia, using a dynamically downscaled regional climate model. *Int. J. Wildland Fire* **23**, 309–321 (2014).
219. Syphard, A. D. et al. The relative influence of climate and housing development on current and projected future fire patterns and structure loss across three California landscapes. *Global Environ. Change* **56**, 41–55 (2019).
220. Yoon, J.-H. et al. Extreme fire season in California: a glimpse into the future? *Bull. Am. Meteorol. Soc.* **96**, S5–S9 (2015).
221. Wang, X. et al. Projected changes in daily fire spread across Canada over the next century. *Environ. Res. Lett.* **12**, 025005 (2017).
222. Young, A. M., Higuera, P. E., Duffy, P. A. & Hu, F. S. Climatic thresholds shape northern high-latitude fire regimes and imply vulnerability to future climate change. *Ecography* **40**, 606–617 (2017).
223. Jones, M. W. et al. Climate change increases the risk of wildfires. *ScienceBrief* <https://sciencebrief.org/briefs/wildfires> (2020).
224. Vitolo, C., Di Giuseppe, F., Krzeminski, B. & San-Miguel-Ayanz, J. A 1980–2018 global fire danger re-analysis dataset for the Canadian Fire Weather Indices. *Sci. Data* **6**, 190032 (2019).
225. DiMiceli, C. et al. MOD44B v006. MODIS/Terra Vegetation Continuous Fields Yearly L3 Global 250 m SIN Grid (NASA EOSDIS Land Processes DAAC, 2015).
226. Littell, J. S., McKenzie, D., Peterson, D. L. & Westerling, A. L. Climate and wildfire area burned in western US ecoprovinces, 1916–2003. *Ecol. Appl.* **19**, 1003–1021 (2009).
227. Eidsenink, J. et al. A project for monitoring trends in burn severity. *Fire Ecol.* **3**, 3–21 (2007).
228. Turco, M. et al. Climate drivers of the 2017 devastating fires in Portugal. *Sci. Rep.* **9**, 13886 (2019).
229. Bowman, D. M., Murphy, B. P., Neyland, D. L., Williamson, G. J. & Prior, L. D. Abrupt fire regime change may cause landscape-wide loss of mature obligate seeder forests. *Glob. Change Biol.* **20**, 1008–1015 (2014).

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Author contributions

D.M.J.S.B. coordinated the project, led the writing and contributed to the design of the graphics. C.A.K. led the design of the graphics and contributed to the writing. J.T.A. led the climate analyses and contributed to the graphic design and writing. F.H.J. contributed to the project development, graphic design and writing. G.R.v.d.W. led the analysis of carbon emissions and contributed to the graphic design and writing. M.F. contributed to the project development, graphic design and writing.

Competing interest

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