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1 **The case for combining a large low-band VHF**
2 **transmitter with multiple receiving arrays for**
3 **geospace research: A Geospace Radar**

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Key Points.

1. A radar capable of probing a wide swath of Geospace could be assembled from a HPLA transmitter and a number of radio-array receivers.
2. Applications for such a geospace radar include MST, meteor, ionospheric, plasmaspheric, planetary, and solar research.
3. A geospace radar would promote discovery research and support space-weather applications.

4 We argue that combining a high-power, large-aperture (HPLA) radar trans-
5 mitter with several large-aperture receiving arrays to make a geospace radar -
6 a radar capable of probing near-earth space from the upper troposphere through
7 to the solar corona - would transform geospace research. We review the emer-
8 gence of incoherent scatter radar (ISR) in the 1960s as an agent which unified
9 early, pioneering research in geospace in a common theoretical, experimental,
10 and instrumental framework, and we suggest that a geospace radar would have
11 a similar effect on future developments in space-weather research. **We then dis-**
12 **cuss recent developments in radio-array technology that could be exploited**
13 **in the development of a geospace radar with new or substantially improved**
14 **capabilities compared to the radars in use presently. A number of appli-**
15 **cations for a geospace radar with the new and improved capabilities** are re-
16 viewed including studies of meteor echoes, mesospheric and stratospheric tur-
17 bulence, ionospheric flows, plasmaspheric and ionospheric irregularities, and re-
18 flection from the solar corona and coronal mass ejections (CMEs). We conclude
19 with a summary of technical requirements.

1. Introduction and motivation

20 The object of geospace science is to explore and understand humanity's place in the
21 solar system and the effects of variations in space on natural and technological systems.
22 This encompasses studies of processes in regions ranging from the surface of the Sun
23 through the interplanetary medium to the magnetosphere, radiation belts, plasmasphere,
24 ionosphere, and neutral atmosphere enveloping the Earth. Some of these processes are
25 detrimental to life and society and constitute space weather. The enormous scope of the
26 problem and the breadth of its impacts demand a comprehensive, integrated response from
27 multiple components of the scientific community working with a range of ground- and
28 space-based instruments, the agencies that support them, and the countries in which they
29 operate.

30 Similar comments would have been applicable at the start of the space age before the
31 fundamentals of plasma physics, in space and in the laboratory, were firmly established.
32 Scientists with a range of backgrounds struggled to understand the first observations of the
33 ionosphere from early spacecraft and from radio and radar apparatus designed for other
34 purposes in some instances, using different approaches and formalisms. A landmark event
35 in the history of space physics was the development in the 1960s of a theory to explain iono-
36 spheric scattering of signals transmitted from high-power, large-aperture (HPLA) radars.
37 What emerged, so-called "incoherent scatter theory," remains one of the most successful
38 applications of linear plasma theory to a complicated natural phenomenon and a textbook
39 example of how theory and technology can develop rapidly together. The legacy of the
40 era is a world-wide network of Incoherent Scatter Radars (ISRs) that remains the source

41 of the most incisive and least ambiguous measurements of ionospheric plasmas available
42 through remote sensing. Such radars are the centerpieces of the existing upper-atmospheric
43 facilities and are complemented by radio as well as optical ground-based instruments. The
44 upper-atmospheric facilities are well coordinated in the way they acquire, disseminate, and
45 interpret their observations.

46 While the last half century has seen tremendous advances in radio technology, signal
47 processing, and space-plasma theory, the ISRs function as much as they did in the early
48 days of development. They are configured mainly for monostatic operation, utilizing large
49 antennas or antenna arrays for transmission and reception. Beam steering, mechanical or
50 electronic, remains the paradigm for surveying large volumes. The radars operate at high
51 peak-power levels with relatively small duty cycles. Range resolution is enhanced mainly
52 through pulse coding and pulse compression. Radar frequencies have mainly been chosen
53 in the UHF band **for reasons having to do mainly with sky noise and licensing require-**
54 **ments.** Most of the ISRs concentrate their operating hours into specialized campaigns with
55 durations of a few days or weeks with low-power survey modes filling the gaps. **These**
56 **choices are rooted in history, including funding history and trends,** but are no longer
57 especially well suited for either scientific discovery or space-weather surveillance.

58 In this paper, we explore the outlines of the next generation of HPLA radars for geospace
59 research. The objective is to expand the coverage of the current radars to encompass a
60 larger segment of geospace. The central concept combines a main radio array for transmis-
61 sion with multiple, higher-resolution arrays for reception in the low VHF band, bridging
62 contemporary experimental principles from aeronomy and astronomy. The paper presents

63 a description of the current state of affairs, a discussion of the new capabilities modern ra-
64 dio arrays could provide, and a review of a few possible applications in geospace research.
65 Underlying the discussion is the premise that a geospace radar would serve the same inte-
66 grating role in the international space-physics community as the original ISRs that came
67 before them.

2. Conventional high-power, large-aperture radars

68 Starting with Arecibo in Puerto Rico and Jicamarca in Peru, the upper-atmospheric fa-
69 cilities were established to provide ground-based remote-sensing of near-Earth space to
70 complement observations coming from early spacecraft. The upper-atmospheric facilities
71 were built around pulsed HPLA radars to exploit the principle of incoherent scatter. In-
72 coherent scatter is modified Thompson scatter from fluctuations in free electron density in
73 a plasma. The relationship between the spectra measured by the radar and the state vari-
74 ables in the ionosphere is given by incoherent scatter theory [*Salpeter, 1960; Fejer, 1960;*
75 *Dougherty and Farley, 1960; Farley et al., 1961; Fejer, 1961; Hagfors, 1961; Salpeter,*
76 *1961; Rosenbluth and Rostocker, 1962; Dougherty and Farley, 1963; Perkins et al., 1965;*
77 *Woodman, 1967*]. In principle, incoherent scatter can be used to measure electron number
78 density, electron and ion temperature, ion composition, line-of-sight drifts, and collision
79 frequencies. The energetic electron population can also be inferred from the measure-
80 ments. Incoherent scatter theory becomes complicated in certain limits, notably for inci-
81 dence angles close to perpendicular to the geomagnetic field. Some aspects of the theory,
82 including the effects of Coulomb collisions on the spectra, are still topics of research and

83 further development [*Kudeki and Milla, 2011; Milla and Kudeki, 2011*]. Incoherent scat-
84 ter is nonetheless the most direct and incisive technique we have for ionospheric remote
85 sensing. Radars capable of incoherent scatter observations can presently be found in the
86 United States and Canada, Peru, Scandinavia, Russia, Japan, and China. Other dual-use
87 radars can be used for incoherent scatter such as the ALTAIR radar in the Kwajalein Atoll.

88 The upper-atmospheric facilities measure the most important state parameters in iono-
89 spheric plasmas. Neutral parameters, including temperatures and winds, are not observed
90 directly but can be inferred from the plasma measurements (e.g. *Hysell et al. [2014]*).
91 Higher-level ionospheric processes including chemistry, energetics, and transport can then
92 be inferred from the state parameters through the application of conservation laws rooted
93 in plasma kinetic or fluid theory. Nowadays, the conservation laws are typically encoded
94 in computer models and simulations. Data from the upper-atmospheric facilities are ei-
95 ther used for consistency checks for the models and simulations or incorporated directly
96 through data assimilation. Data assimilation can be applied to very low-level radar data
97 products, for example the spectra or autocorrelation functions of the signals themselves,
98 avoiding some of the distortions inherent in estimating higher-level data products directly.

The sensitivity of an ISR is partly governed by its figure of merit, f , the product of the transmitter average power in Watts and the effective antenna aperture in square meters divided by the product of the sky noise temperature in Kelvin degrees and the receiver bandwidth in Hz:

$$f = \frac{PA}{BT} \quad (1)$$

99 **The signal-to-noise ratio of a given experiment will be proportional to this figure**
100 **[*Beynon and Williams, 1978*] Sensitivity also depends on other experimental charac-**
101 **teristics, including the sampling cadence (which can be enhanced by exploiting fre-**
102 **quency or polarization diversity) and the signal-processing strategy, and so the fig-**
103 **ure of merit is not a comprehensive measure of ISR performance. Another useful**
104 **measure is the speed of an experiment which quantifies the time required to achieve**
105 **measurements with a given statistical accuracy [*Lehtinen, 1986*]. The figure of merit**
106 **nonetheless is a useful tool for comparing and designing HPLA radars.**

107 Arecibo, which has the highest figure of merit of any of the ISRs, operates in the UHF
108 band near a minimum in the sky noise temperature. Jicamarca, meanwhile, operates at 50
109 MHz band where the sky noise temperature is two orders of magnitude greater (although
110 the bandwidth required is also an order of magnitude smaller). The low frequency is disad-
111 vantageous from the point of view of the figure of merit but advantageous in a number of
112 important ways. The figure of merit is, in fact, a limited metric which does not capture im-
113 portant tradeoffs involving the radar frequency, the characteristics of transmit and receive
114 antennas, which could be different, or the characteristics of the emission and the way the
115 signals are processed.

116 An advantage of VHF, for example, is improved immunity from finite Debye-length
117 effects. The bandwidth of the ion line, the dominant incoherent scatter feature (at angles
118 away from the perpendicular to the geomagnetic field direction), is determined by the ion
119 rather than the electron thermal speed, but only if the the radar wavelengths λ is much
120 larger than the plasma Debye length. Measurements in rarefied plasmas necessitate long

121 radar wavelengths. Jicamarca has measured incoherent scatter to altitudes of about 10,000
122 km and is also uniquely sensitive in the mesosphere where the electron number density is
123 also small [*Farley, 1991*].

124 Another property of the incoherent scatter spectrum is the drastic narrowing that occurs
125 for incidence nearly perpendicular to the geomagnetic field. In this case, the spectrum
126 becomes much narrower than the ion thermal bandwidth, facilitating very accurate mea-
127 surements of line-of-sight drifts. The phenomenon only occurs when $\lambda > 4\pi\rho_e$, where
128 ρ_e is the electron gyroradius, and can only be observed when the radar beamwidth is very
129 narrow and the backscatter signal is dominated by the part coming from small magnetic
130 aspect angles [*Milla and Kudeki, 2011*].

131 Finally, high-power VHF radars can be used to study backscatter from fluctuations in
132 the refractive index of neutral gases in the middle and lower atmosphere. MST radars are
133 as important for studies of energetics, dynamics, and transport in the neutral atmosphere
134 as incoherent scatter radars are in the ionosphere. With current technology and existing
135 equipment, there need be no observing gap in altitudes between the boundary layer and the
136 plasmopause.

137 The existing ISRs employ vacuum tubes, traveling-wave tubes, klystrons, or transistors
138 for RF power. The power source is lumped in the first three cases and distributed in the
139 last case where transistors are typically situated very close to individual antenna elements.
140 The conventional paradigm involves using pulsed transmitters with high peak-power levels
141 and duty cycles ranging from a few percent to a few tens of percent. The duty cycle can be
142 exploited to the greatest extent by using pulse compression schemes or pulse coding which

143 allows long pulses to behave, in some respects, like shorter but more powerful pulses.
144 Pulse coding can also be used to make the radar instrument function or ambiguity function
145 closer to ideal, facilitating spectral measurements that are at once well resolved in range
146 and frequency. **Both random and deterministic pulse codes are in use at different ISRs**
147 **on the basis, in part, of the availability of signal generation equipment.** The price paid
148 for pulse compression is a reduction in sensitivity due to increased radar clutter.

149 Today's ISRs utilize both phased-array and mechanically-steered dish antennas.
150 **Whereas the former allow for rapid steering, the latter can present attractive cost-**
151 **and performance-tradeoffs with respect to sidelobes and the ability to steer to low el-**
152 **evation angles.** Phased arrays also offer flexibility with regard to the use of spaced-antenna
153 methods, including interferometry and radar imaging. Subdividing the receive array for in-
154 terferometry and imaging comes at the cost of sensitivity and has been used in a limited
155 way with the two EISCAT ESR radars for studying the naturally-enhanced ion acoustic
156 lines (NIELS). Interferometry and imaging are used routinely for studying coherent scatter
157 with the modular phased array at Jicamarca.

158 The only multistatic ISR presently is the EISCAT mainland radar network which is
159 tristatic. Multistatic radars offer the ability to measure vector drifts unambiguously. Since
160 the EISCAT radars are mechanically steered, however, each scattering volume along the
161 transmit radar beam must be interrogated by the receive antennas one altitude at a time.
162 This seriously limits the cadence of experiments and the utility of the multistatic capabil-
163 ity.

3. Overview of radio arrays

164 A new paradigm for radio exploration of geospace comes from the **radio astronomy**
165 community where a new generation of radio arrays is being developed and deployed. The
166 radio arrays exploit spatial diversity to the fullest extent, going beyond traditional beam-
167 forming methods and utilizing all of the spatio-temporal information in radio signals from
168 subjects under study by using large-scale sparse interferometer systems. The geospace
169 radar concept involves combining radio arrays with discrete or distributed transmitters to
170 study volume scatter from the upper atmosphere and beyond. The goal is the production of
171 datasets which reveal the space-time structure of geospace phenomena unambiguously and
172 in a way that can readily be compared with or assimilated into numerical models, including
173 forecast models, as well as machine-learning algorithms.

174 Radio arrays have a long history in the arena of solar research. Reviewing some of the
175 milestones in this area provides context for the geospace-radar concept. The Culgoora
176 array consisted of ninety-six 13m-diameter mechanically steerable dishes, arranged in a
177 circle 3 km in diameter. It operated in dual polarization at 80 and 160 MHz. It was
178 designed for solar work and so did not attempt to suppress grating lobes so long as they
179 were out of the solar field of view. It was responsible for much of our understanding of
180 meter-wavelength radio bursts. The resolution was sufficient to measure scatter broadening
181 of Type III radio bursts by coronal turbulence.

182 In the case of the CLRO, the array was configured as a 3 km “T” of log-spiral antennas,
183 predominantly in a single circular polarization. It was designed for both solar and cosmic
184 work, so grating lobe suppression was important. It was one of the first to use a digital

185 correlation receiver. Its resolution was comparable to Culgoora, but it covered a broader
186 band with the log-spiral elements.

187 In the last couple of decades, astronomical radio arrays have evolved significantly. The
188 resolving capabilities of most modern radio telescopes are realized in two different ways:
189 beamforming and interferometry. Perhaps the simplest way for a telescope to beamform is
190 through the use of a dish (typically parabolic or spherical). Dishes are expensive, both to
191 build and to maintain, and costs increase dramatically with the size of the dish. However,
192 beamforming can also be accomplished using an array of dipoles, where the signal from
193 each element is delayed and the total is coherently summed, effectively synthesizing an
194 aperture. Historically, the delay portion of phased-array beamforming was accomplished
195 with physical delay lines and phase-shifters which are awkward to calibrate and maintain.

196 Unlike beamforming, interferometry by nature requires multiple elements. Voltage sam-
197 ples from each element in an interferometer are cross correlated with the voltages from
198 other elements. The physical distance and orientation of each antenna pair makes them
199 sensitive to different spatial frequencies in different orientations. These spatial-frequency
200 measurements can be Fourier transformed and summed to form an image. This process is
201 called synthesis imaging as the elements in the array are synthesizing points on an aperture.
202 For an interferometer comprised of N elements, there are $N \times (N - 1)/2$ unique cross-
203 correlation products. Although phase shifting can be performed after the cross-correlation,
204 delay compensation is necessary in broadband systems, and this remains awkward and
205 expensive. In the past, the fact that the computational load increased like the number of
206 elements squared prohibited many-element interferometers.

207 Recently, advancements in computational power have opened the door to both digital
208 beamforming and large- N synthesis imaging. In particular, large arrays comprised of hun-
209 dreds of dipoles can digitally steer multiple beams simultaneously across large bandwidths
210 using off-the-shelf CPUs **and, in practice, GPUs. The adoption of GPUs** by the radio
211 astronomy community has allowed for real-time, wide-band synthetic imaging with fields
212 of view limited only by the beam pattern of an individual antenna elements. Delays are still
213 necessary, but they simply imply a digital buffer. Furthermore, digital buffers can be large
214 enough to store the entire RF signal for the duration of an event so that, when an event is
215 detected, the entire buffer can be saved and reprocessed as necessary. Such an event might
216 be a fast radio burst (FRB) in radio astronomy, but many such situations can also be imag-
217 ined in geospace research. These advancements are revolutionizing radio astronomy, but
218 the geospace and space-weather communities for the most part continue to use decades-old
219 facilities that rely on single dish or delay-line technology.

220 **LOFAR:** The Low Frequency ARray (LOFAR), the largest radio telescope in the lower
221 VHF band, was built and is operated by the Netherlands Institute for Radio Astronomy,
222 ASTRON. LOFAR is comprised of individual stations, most of which are located in the
223 Netherlands. Currently, there are 24 core stations located in Exloo, NL, 14 remote stations
224 located across the Netherlands, and 12 international stations located across Europe. Each
225 station includes both low band (10-80 MHz) and high band (120-240 MHz) arrays.

226 The low band antennas consist of two linear, orthogonal dipoles, each made of two cop-
227 per wires. The dipoles are resonant at 60 MHz and drop in sensitivity away from this
228 frequency. The low band arrays of the core and remote station include 48 low band an-

229 tennas, and the international stations include 96. Beamforming with an individual station
230 is done electronically, allowing for fast beam steering with no moving parts. The beam-
231 formed time series are then sent to a correlator for high-resolution imaging [*van Haarlem*
232 *et al.*, 2013].

233 **KAIRA:** The Kilpisjärvi Atmospheric Imaging Receiver Array (KAIRA) is a LOFAR
234 station which was constructed in 2011 in Northern Finland. KAIRA has demonstrated suc-
235 cessfully that modern, low-frequency, digital phased array radio telescopes can be highly
236 versatile instruments for studying various geophysical phenomena in Earth's near space
237 using radio remote sensing [*Vierinen et al.*, 2013; *McKay-Bukowski et al.*, 2015]. The fre-
238 quency range of $\approx 10 - 240$ MHz is ideal for studies of the Earth's near space. At these
239 frequencies, the radar cross-sections of mesospheric echoes, meteor trail, and head echoes
240 are maximized. As stated above, lower frequencies also allow observing lower plasma
241 densities due to Debye-length effects. At lower frequencies, radio propagation effects such
242 as scintillation and absorption are also stronger, facilitating studies of these phenomena.

243 Due to the versatile nature of the underlying digital phased array technology, the same
244 system can be used for a number of different active and passive geophysical radio remote
245 sensing applications including phased array incoherent scatter radar [*Vierinen et al.*, 2013;
246 *Virtanen et al.*, 2014], spectral riometry [*Kero et al.*, 2014], multi-static studies of the meso-
247 sphere [*Chau et al.*, 2018a], and for characterization of wide-band ionospheric scintillation
248 [*Fallows et al.*, 2014].

249 One of the strengths of the all-digital phased array technology is the wide bandwidth.
250 This allows multi-wavelength studies of various plasma physics phenomena in geospace.

251 The refractive index for ionospheric plasma is frequency dependent. The same applies for
252 radar cross-sections of various phenomena such as meteor head echoes and mesospheric
253 echoes. So far, the information from simultaneous, multi-wavelength observations has
254 remained relatively unexplored. Due to the wide band nature, KAIRA can actually be
255 used together with multiple radar transmitters as a phased array radar receiver. So far,
256 it has been used together with the EISCAT VHF transmitter, the MAARSY MST radar,
257 and the Andøya meteor radar. Therefore, another strength of using modern radio arrays is
258 that they can be used together with a range of existing and planned radar transmitters as a
259 multi-static platform.

260 **MWA:** The MWA consists of 2048 dual-polarization dipole antennas optimized for the
261 80-300 MHz frequency range, arranged as 128 “tiles”, each a 4×4 array of dipoles. The
262 array has no moving parts, and all telescope functions including pointing are performed by
263 electronic manipulation of dipole signals, each of which contains information from nearly
264 four steradians of sky centered on the zenith. Each tile performs an analog beamforming
265 operation, narrowing the field of view to a fully steerable 25 degrees at 150 MHz.

266 The majority of the tiles (112) are scattered across a roughly 1.5 km core region, form-
267 ing an array with very high imaging quality, and a field of view of several hundred square
268 degrees at a resolution of several arcminutes. The remaining 16 tiles are placed at loca-
269 tions outside the core, yielding baseline distances of about 3 km to allow higher angular
270 resolution for solar burst measurements.

271 Important ionospheric research has already been performed with the MWA. In particular,
272 *Loi et al.* [2015] recently identified field-aligned ionization ducts between the ionosphere
273 and the plasmasphere in spatially-resolved maps of total electron content (TEC).

274 **LWA:** The Long Wavelength Array (LWA) is a concept for an HF/VHF radio tele-
275 scope comprised of 52 stations spread over the state of New Mexico, providing 0.4 km² of
276 collecting area. With baselines ranging from a few km up to about 400 km, such a radio
277 telescope could achieve down to 5 arc second resolution.

278 A single LWA station is comprised of 256 dual polarization, bow tie antennas pseudo-
279 randomly spread over a 100x110 m ellipse. The antennas are slightly bent to achieve
280 greater sensitivity to a wider range of zenith angles. A single station can act as a 256-
281 element interferometer, capable of imaging nearly the entire visible sky, or as a digital
282 beamformer, capable of producing several beams at once.

283 To date, two stations have been completed in New Mexico. The first station, LWA1,
284 is collocated with the Very Large Array and operates from 10-88 MHz [*Ellingson et al.*,
285 2012]. The second station, LWA-SV, which is located at Sevilleta National Wildlife refuge,
286 is operationally nearly identical to LWA1 except it has modified analog filters allowing
287 observations down to 3 MHz.

288 Furthermore, the digital processor on LWA-SV has been upgraded from LWA1 to make
289 use of graphics processing units (GPUs) through the Python/C++ Framework known as
290 Bifrost [*Cranmer et al.*, 2017]. GPUs are capable of performing large numbers of ba-
291 sic calculations simultaneously, making them ideal for radio arrays comprised of many,
292 non-uniformly distributed elements such as an LWA station. Bifrost is designed to work

293 on streaming data such as radio array voltages in real time and allows for rapid pipeline
294 development with its high-level python interface.

295 LWA at Owens Valley Observatory (LWA-OVRO) is a higher resolution LWA telescope
296 built for monitoring of astrophysical transients. With maximum baselines of 100 m, LWA1
297 and LWA-SV achieve a synthesized beam full width at half the maximum (FWHM) of a
298 few degrees in the lower VHF band. However with maximum baselines of 1.5 km, LWA-
299 OVRO is capable of imaging the entire sky at a resolution of a few 10s of arcmin.

300 **EISCAT 3D:** EISCAT 3D is a next generation all-digital multi-static phased array
301 radar which is currently being constructed in Northern Scandinavia [*McCrea et al.*, 2015].
302 The core transmit-receive site will be located in Skibotn. Two outlier receive-only system
303 are to be located in Sweden and Finland, approximately 150-km distance from Skibotn.
304 Each one of the sites consists of 109 modules, with 91 antennas in each module. The
305 operating frequency is 233 MHz. The transmit power will be 5 MW peak power with a
306 25% duty-cycle.

307 The main novelty of EISCAT 3D is that it will have the capability of observing vector
308 ion velocities simultaneously at all altitudes by using multiple bistatic beams that intersect
309 the transmit beam. Due to the rapid beam switching capability, the radar will be able to
310 perform a volumetric observation quickly. EISCAT 3D will also enable aperture-synthesis
311 radar imaging in order to improve spatial resolution of auroral radar observations, a tech-
312 nique that has been used already for a long time at the Jicamarca Radio Observatory (e.g.
313 *Hysell and Chau* [2006]).

314 Once finished, EISCAT 3D will be the most modern incoherent scatter radar in the world.
315 By virtue of its frequency and its latitude, it will not, however, be able to fulfill all of the
316 objectives of a geospace radar as laid out in this paper. Unlike LWA or LOFAR, the EIS-
317 CAT 3D receiver array is also relatively narrow band (30 MHz), limiting the applicability
318 of the radio array for uses other than 233-MHz radar operations.

4. Capabilities of modern radio arrays paired with HPLA transmitters

319 In this section, we summarize some of the existing and potential capabilities that modern
320 radio arrays would bring to geospace research operating in radar mode, particularly in
321 the low-VHF band. To keep our summary focused, we assume a large radio array for
322 transmission and multiple, spatially separated radio arrays for reception. In this respect,
323 we start with the transmitting capabilities and continue with the receiving capabilities and
324 the motivation for focusing on the low-VHF band.

4.1. Transmitter capabilities

325 Recent developments in technology allow the possibility of large transmitting arrays
326 with solid-state transmitters possibly at each antenna element. Solid-state technology in
327 conjunction with contemporary, compact antenna element designs (e.g., LWA), would al-
328 low operations over a relatively broad band at low VHF frequencies. Moreover, modern
329 designs would allow larger duty cycles and even continuous wave (CW) operation. This
330 implies an increase in average power even for systems that are peak-power limited. The
331 result is the ability to probe a broader range of geospace.

332 Although performance advantages of a distributed transmitter are clear, radar users are
333 well-aware that much of the costs of operating a radar are in operating and maintaining
334 the transmitter. It is this that limits the operating hours of many ISRs. By comparison,
335 a distributed transmitter **might be expected to pose fewer maintenance challenges over**
336 **time** and permit 24/7 operation at little additional cost.

337 The possibility of a transmitter on each antenna, and depending on the RCS of the desired
338 target, allows:

- 339 • Fast beam steering to reduce space-time ambiguity.
- 340 • Multiple, simultaneous beam pointing directions for multi-target, multi-use applica-
341 tions [*Milla et al.*, 2013].
- 342 • Antenna compression to generate beams with different widths [*Woodman and Chau*,
343 2001; *Chau et al.*, 2009].
- 344 • Implementation of multi-frequency approaches to improve range resolution (range
345 imaging).
- 346 • Implementation of coherent MIMO configurations using code diversity [*Urco et al.*,
347 2018a]. Coherent MIMO can be used to take advantage of the spatial information arising
348 from distributed transmitter modules .
- 349 • Implementation of low-power modes, either for long-term operations or for strong
350 targets.

351 In the last 10 years, great progress has been made with **broadband** phased array radio
352 telescopes. There would be great benefit from a matching frequency-agile transmit capabil-
353 ity. This would allow radar studies of space plasma physics phenomena with one additional

354 independent variable: frequency. This information would be useful, e.g., for studies of at-
355 mospheric turbulence spectra and determining the sizes of meteor head echoes. For solar
356 radar, this would allow studies of the solar corona at different depths [*Bastian, 2003, 2004*].

4.2. Receiver capabilities

357 In the case of receivers, based on the recent developments in radio astronomy (e.g.,
358 LOFAR, LWA), monostatic and multistatic configurations would provide significant new
359 capabilities to a geospace radar. Much like a radio telescope, a geospace radar receiver
360 could leverage the power and flexibility of software controlled beamforming and aperture
361 synthesis imaging. Furthermore, a radio astronomy-like digital processor could easily be
362 modified for flexible radar processing and data acquisition. This has already been demon-
363 strated with LOFAR [*Vierinen et al., 2013*] and LWA [*Taylor, 2014*].

364 The new capabilities for reception, related to the transmitter capabilities described above
365 but not limited to them, would include:

- 366 • Many simultaneous independent beam pointing directions available at all times. This
367 allows, e.g., different altitudes to be sampled simultaneously by a multistatic receiver and
368 many other presently impossible operational modes.
- 369 • Multistatic reception capability for resolving three-dimensional flows unambiguously.
- 370 • Antenna compression to generate beams with different widths on reception.
- 371 • Implementation of synthetic aperture imaging by utilizing smaller sections within re-
372 ceiving arrays or combining spatially separated receiving arrays (e.g., LOFAR core).

373 • Implementation of coherent MISO, i.e., single antennas (or group of antennas) could
374 be used for interferometric/imaging purposes if the target is sufficiently strong.

375 • Broadband interference identification and rejection from sources such as lightning
376 and solar radio bursts.

377 The theme of these capabilities is the exploitation of spatial diversity to resolve features
378 and flows in three dimensions, expanding the domain of geospace radar observations be-
379 yond simple profiling and conventional monostatic beam forming. The applications iden-
380 tified below exemplify the need for the capabilities specifically.

4.3. Why low-band VHF?

381 **There are clear disadvantages associated with operating in the VHF band. Lower**
382 **frequencies imply larger array sizes, greater material costs, and greater demands**
383 **on real estate. Licensing in the VHF band may also be difficult in some countries. Most**
384 **importantly, increased galactic noise at VHF compared to UHF implies challenges for**
385 **system sensitivity and a potential requirement for greater transmitter power. The sen-**
386 **sitivity problem is mitigated partly by reduced atmospheric absorption and cabling**
387 **losses.**

388 Despite these problems, most of the geospace-radar applications above would benefit
389 from using low-band VHF. In terms of feasibility, the technology needed for a low-band
390 VHF geospace radar is already proven; the required antennas, digital receivers, beam for-
391 mers (based on FPGA), and solid-state transmitters are available commercially now. A
392 number of other factors make low-band VHF advantageous. For example, radio frequency
393 interference (RFI) in the low band VHF can be a limiting factor for receiving arrays. How-

394 ever, as shown by both LOFAR [*van Haarlem et al.*, 2013] and LWA1 [*Ellingson et al.*,
395 2012], the spectrum between 30 and 80 MHz is relatively clean, even in densely populated
396 countries like the Netherlands. Other noteworthy factors related to low-band VHF radar
397 are summarized in the following list:

398 • Radio propagation effects of interest such as Faraday rotation and scintillation are
399 stronger at low VHF frequencies.

400 • Array element patterns are broader in the low VHF band. This allows any fully-digital
401 low-band receiving array to have an unlimited number of beams covering the entire visible
402 hemisphere at no penalty in signal-to-noise ratio.

403 • The radar cross-sections of many of the phenomena of interest are also larger.

404 • Finite Debye-length effects for ISR impose less severe limitations on the minimum
405 detectable electron number density.

406 • Planetary (Lunar) studies would benefit from deeper subsurface penetration at VHF.

407 • The power spectral density of galactic synchrotron emission as a function of fre-
408 quency is proportional to $f^{-2.5}$ [*Rogers and Bowman*, 2008]. However, the bandwidth
409 of the incoherent scatter ion line, and therefore the minimum receiver noise bandwidth,
410 is proportional to f . This means that the signal to noise ratio for the ion-line is not as
411 strongly dependent on frequency as one would expect by just looking at the sky noise:

412 $\text{SNR}(f) \propto f^{1.5}$.

5. Geospace radar applications

413 In this section, we describe some of the possible applications of a new geospace radar,
414 starting from the lower atmosphere and ending at the Sun. The applications are meant to be
415 illustrative rather than exhaustive. New applications will emerge as novel remote sensing
416 approaches designed for the current applications mature.

5.1. MST applications

417 The neutral atmosphere can be also studied with VHF radars in so-called MST mode.
418 The first application of such radars for neutral atmospheric dynamics was conducted at
419 Jicamarca [*Woodman and Guillén, 1974*]. Since then, the MST technique proliferated
420 world-wide, with small systems covering the altitudes between a few hundred meters to
421 lower stratospheric altitudes (i.e., less than 20 km). The altitudinal coverage of these sys-
422 tems depends on the frequency of operation, size of the antennas, and transmitting power.
423 The corresponding systems are called boundary layer radars, wind profilers, and ST radars
424 [*Hocking, 2011*]. In the case of the mesosphere, there is a handful of HPLA radars work-
425 ing between 45 and 55 MHz, able to receive echoes from irregularities embedded in neu-
426 tral turbulent layers (e.g., Jicamarca, Gadanki, MU, MAARSY). At high latitudes, due to
427 the presence of charged ice particles during the summer, smaller systems are being used
428 to study the polar summer mesosphere from so-called polar mesospheric summer echoes
429 (PMSE). In all these cases, low VHF frequencies have been used to study the neutral at-
430 mosphere from a few hundred meters to the lower stratosphere and in some parts of the
431 mesosphere depending on the latitude, time of the day, and season.

432 The region between the lower stratosphere and the lower mesosphere (i.e., between 20
433 and 55 km) has been known as the “radar gap region” since this region typically cannot
434 be observed with existing radars. Using long integration times and dual polarizations at
435 Jicamarca, *Maekawa et al.* [1993] reported on atmospheric echoes from this region for the
436 first time.

437 A modern and powerful geospace radar should be able to sample the whole MST re-
438 gion by combining low-power modes for the lower atmosphere with high-power modes
439 for the stratosphere and mesosphere. The stratosphere and lower mesosphere is the region
440 where primary gravity waves are expected to break and dissipate, modify the background
441 conditions, and generate so-called secondary gravity waves.

442 The daytime exploration of the mesosphere could be also complemented with ISR obser-
443 vations of the D region. Although radar scattering is from electrons, the collision frequency
444 at these altitudes is so high that the dynamics of the measured spectra are dominated by the
445 neutrals. Using the Poker Flat Incoherent Scatter Radar (PFISR) during a strong auroral
446 precipitation event, *Nicolls et al.* [2010] were able to study inertial gravity waves that had
447 originated a few days before at the surface a few thousand kilometers away. In the case of
448 a powerful geospace radar, such atmospheric events could be studied during the daytime
449 routinely, i.e., without the need of strong electron density enhancements.

450 Besides exploring a new altitudinal region, a modern geospace radar could be used to
451 explore the MST region with larger horizontal coverage and greater altitudinal, horizon-
452 tal, and temporal resolution than is presently possible. Such improvements would allow
453 the combination of radar imaging, range imaging, multi-static observations, and MIMO

454 implementations. Recently, *Urco et al.* [2018b] observed PMSE with unprecedented hor-
455 izontal resolution using a combination of Maximum Entropy and MIMO techniques. The
456 resolution achieved with the MAARSY 90-m diameter antenna was equivalent as to hav-
457 ing a 450-m diameter array. The resulting images show that the summer polar mesosphere
458 presents rich spatial and temporal structures associated to a combination of vertically prop-
459 agating waves from below, horizontally drifting waves, and instabilities and waves gener-
460 ated in situ. Figure 1 shows an example of such 3D images obtained with 40 second inte-
461 grations: (a) a horizontal cut at 85.80 km, (b) an altitude vs zonal direction cut, and (c) an
462 altitude vs meridional cut, color-coded with Doppler velocity, spectral width and intensity
463 information. A powerful geospace radar would allow similar mesospheric observations at
464 low and middle latitudes where mesospheric echoes are mainly dependent on atmospheric
465 turbulence and therefore have an RCS a couple of orders of magnitude smaller than PMSE.

466 The horizontal coverage can be extended, in the case of lower atmospheric altitudes
467 where the echoes are stronger, by using smaller sections of the transmitter array, different
468 pointing directions, and multi-static configurations.

5.2. Meteor science and applications

469 In the mesosphere and lower thermosphere (MLT), i.e., between 70 and 120 km, echoes
470 from meteors can be also observed with low VHF-band radars. There are three main types
471 of meteor echoes: (a) specular trail echoes due to Fresnel scattering when the radar points
472 perpendicular to the meteor trajectory, (b) Head echoes coming from plasma as it forms
473 in front of the meteoroid, and (c) non-specular trail echoes, also known as range-spread

474 trail echoes resulting from the combination of field-aligned irregularities and charged dust
475 particles.

476 Meteor echoes have been used to study meteoroid composition, origin, and masses, and
477 to explore the MLT region. Traditional specular meteor echoes have been used since the
478 1950s to explore the MLT dynamics with small radars systems called specular meteor
479 radars (SMRs). They are typically composed of one wide-beam transmitting antenna and
480 a receiver interferometer operating between 30 and 50 MHz. Recently, such systems have
481 been improved using multi-static geometries, coded CW transmissions, and MIMO con-
482 figurations [Stober and Chau, 2015; Vierinen et al., 2016; Chau et al., 2018b]. The im-
483 provements allow such systems to measure the MLT wind fields with much larger horizon-
484 tal coverage and resolution as well as improved vertical and temporal resolutions [Stober
485 et al., 2018]. Current efforts are being devoted to determining the second order statistics
486 of the wind field as function of wavenumber (horizontal and vertical) as well as frequency.

487 A powerful geospace radar would allow the detection of much weaker signals due to
488 meteoroids with less mass or smaller entry speeds. One of the outstanding questions in
489 aeronomy is how much meteor mass is deposited in the atmosphere? So far, different
490 methods provide estimates with a couple of orders of magnitude spread. A geospace radar
491 could greatly improve the measurements of the lower mass/speed end of the spectrum,
492 from both head as well as specular meteor echoes.

493 In the case of head echoes, only HPLA radars with interferometry are able to determine
494 precisely the meteoroid trajectory and where in the beam such echoes occur, something
495 crucial for the scattering meteor mass determination. So far, such measurements have been

496 conducted only at Jicamarca, MU, and MAARSY using narrow beam configurations [*Chau*
497 *et al.*, 2007; *Kero et al.*, 2011; *Schult et al.*, 2017]. The use of a wider beam, longer duty
498 cycle sequences, and MIMO configurations would improve significantly the detection and
499 characterization of such echoes.

500 Non-specular meteor echoes are still enigmatic since they appear to result from a com-
501 bination of field-aligned irregularities and charged dust particles. A clear indication of the
502 former comes from observations at low and mid latitudes where these relatively long-lived
503 range-spread echoes come primarily when the beam points perpendicular to *B*. The latter
504 come from high latitude observations where perpendicular-to-*B* observations are not possi-
505 ble. Besides the interest in understanding the physics behind these echoes, they can be used
506 to get precise altitudinal profiles of MLT winds with high temporal and spatial resolution
507 [*Oppenheim et al.*, 2009]. The proposed geospace radar, if located at low or mid latitudes,
508 would be able to provide such measurements on routine basis, contributing further to the
509 complicated atmospheric dynamics of the MLT region. Besides routine observations of
510 these profiles, the horizontal coverage would be improved by using multi-static and multi-
511 beam approaches.

5.3. Ionospheric research

512 Since its inception, the incoherent scatter technique has provided the most incisive mea-
513 surements of ionospheric state parameters available through ground-based remote sensing.
514 These include plasma number densities, temperatures, composition, and line-of-sight drifts
515 as well as information about the electron energy distribution and inferences about the state
516 of the neutral atmosphere. Most of the information has come in the form of profiles ob-

517 tained along the beam pointing direction of the radar. Profiles are useful in view of the fact
518 that the ionosphere is vertically stratified. Volumetric information came mainly from me-
519 chanical beam steering before the emergence of the Advanced Modular Incoherent Scatter
520 Radars (AMISRs) currently deployed in Alaska and Canada which employ electrically-
521 steered phased arrays.

522 The importance of obtaining volumetric information is highlighted by Figs. 2 and 3
523 which represent radar observations of plasma density irregularities associated with plasma
524 convective instability at low magnetic latitudes and so-called “equatorial spread F .” Fig. 2
525 was obtained by the ALTAIR radar in the Marshall Islands and shows an east-west scan
526 acquired at pointing angles a few degrees away from perpendicular to the geomagnetic
527 field. The figure exhibits large-scale undulations in the F -region bottomside along with
528 depletion plumes penetrating through the F peak and into the topside. The irregularities
529 evolve on timescales of a few minutes or tens of minutes which is comparable to the time
530 required to complete a scan. Although informative, the figure is not a true image and is
531 highly distorted.

532 Fig 3 shows data acquired a few minutes earlier along antenna pointing positions pre-
533 cisely perpendicular to the geomagnetic field. Patches of intense backscatter concentrated
534 in the depleted regions in Fig. 2 represent coherent scatter or Bragg scatter from small-
535 scale field-aligned plasma density irregularities. Like the large-scale irregularities evident
536 in Fig. 2, the small-scale irregularities are a consequence of plasma instability and signify
537 the presence of free energy in the ionosphere. They are a vivid manifestation of space

538 weather and a hazard for radio communication, navigation, and imaging systems operating
539 at low geomagnetic latitudes.

540 The coherent scatter conveys information about the the underlying plasma instabilities
541 that is highly complementary with the incoherent scatter which conveys information about
542 plasma state variables. Ideally, we would like to acquire signals from the entire ionospheric
543 volume within the figure simultaneously, including from both large and small magnetic
544 aspect angles. This would be true imaging and would permit direct, unambiguous compar-
545 ison with direct numerical simulations.

546 One of the major shortcomings of the ALTAIR data in Figs. 2 and 3 is the absence
547 of reliable drifts measurements. Plasma drifts at the geomagnetic equator are relatively
548 small, and the incoherent integration time necessary to acquire and formulate accurate
549 drifts measurements in this case is longer than the timescale along which the large-scale
550 plasma irregularities evolve. The measurement is not stationary at UHF frequencies. At
551 low VHF frequencies, and for beam pointing angles very close to perpendicular to the
552 magnetic field, accuracy of ISR drifts measurements improves dramatically, and the mea-
553 surements become stationary. This is a strong argument for favoring low VHF frequencies
554 for a geospace radar.

555 The high quality of ISR drifts obtained with beams pointing perpendicular to the mag-
556 netic field is due to the narrowness of the ISR spectrum at low VHF frequencies allowing
557 the application of Doppler-shift detection to estimate the drifts. The narrowness of the
558 spectrum is caused by the relatively long temporal correlation of the field-aligned electron
559 density fluctuations that diffuse slowly across the magnetic field lines. Such diffusion is

560 not dominated by Landau damping but by Coulomb collisions. Incoherent scatter theory
561 has been reformulated recently to incorporate the effects of Coulomb collisions on particle
562 dynamics by describing the statistics of electron and ion trajectories using a Fokker-Planck
563 kinetic equation with speed-dependent friction and diffusion coefficients. Measurements
564 with the Jicamarca radar at 50 MHz have been shown to agree with the proposed theory.
565 However, the validation is limited since only a single radar frequency has been used in a
566 particular configuration. A modern geospace radar would let us to conduct similar tests to
567 study the effects of Coulomb collisions on ISR signals at multiple radar frequencies, mag-
568 netic aspect angles, and antenna beam widths for different ionospheric altitudes including
569 topside regions where plasma composition would impose another level of complexity. This
570 study is of general relevance as a test for theoretical plasma physics.

5.4. Plasmaspheric research

571 A problem was discovered recently with the standard high-altitude ISR mode at Jica-
572 marca. Samples acquired from very high altitudes were being used as the basis for noise
573 estimates. From time to time, however, the noise estimates became contaminated. Upon
574 closer examination, the contamination was found to be due to coherent scatter from high
575 altitudes just above 2000 km. This is the altitude where the on-axis antenna pointing
576 position at Jicamarca was perpendicular to B at the time. High-altitude field-aligned irreg-
577 ularities were the cause. The irregularities occurred mainly in the pre-dawn sector. The
578 immediate problem was solved by using samples taken during transmitter-off intervals for
579 noise estimates. The problem of the origin of the high-altitude field-aligned irregularities
580 remains.

581 An example of the phenomenon in question is shown in Fig. 4. The figure was generated
582 by computing the modulus of the longest lag in the measured ISR lag-product profiles, 1.5
583 ms, which should be essentially zero for incoherent scatter. The experiment in question is
584 very sensitive compared to more conventional coherent scatter experiments, and the echoes
585 were not very strong compared to what is received from the electrojet or equatorial spread
586 *F*. The coherent scatter occurred in layers that migrated in altitude slowly in time. Since
587 the experimental geometry favored field-aligned echoes between about 2000-2500 km in
588 this case, we do not know the true altitude span of the irregularities.

589 Subsequent discussions revealed that this phenomenon had been known since the earliest
590 days of Jicamarca (personal communication, D. T. Farley, J. P. McClure). However, it was
591 not deemed to be as interesting as the other ionospheric phenomena being discovered in
592 the 1960s and was never pursued. Contemporary ISR experiments designed to remove
593 outliers arising from different kinds of radar clutter masked the phenomenon which was
594 only rediscovered by accident.

595 The high-altitude coherent scatter is significant since it signifies the presence of an un-
596 known source of free energy and an unknown plasma instability. The echoes are not merely
597 high-altitude examples of FAIs caused by equatorial spread *F* which can occur above 2000
598 km altitude during periods of high solar flux under geomagnetically-disturbed conditions.
599 The echoes in Fig. 4 occurred during a period of low solar flux and quiet conditions.
600 They were not preceded by spread *F* and occurred during June solstice when spread *F* is
601 relatively uncommon. That they occur mainly in the pre-dawn sector suggests that beam

602 instabilities driven by photoelectrons may be playing a role. To test the hypothesis, dedi-
603 cated rather than serendipitous observations are required.

5.5. Planetary research

604 Low frequencies are especially useful for probing the subsurface of dielectric bodies.
605 The penetration depth of an electromagnetic field is inversely proportional to frequency.
606 Lower frequencies are used in Earth remote sensing for applications such as hydrology
607 and foliage-penetrating radar. In the case of planetary exploration, low frequency observa-
608 tions are used to probe the subsurface of planetary bodies, e.g., in search of subsurface ice
609 deposits on Mars or ice on permanently shadowed craters on the Moon. In general, radar
610 observations across a number of different frequencies can yield a great deal of information
611 on the geological surface and subsurface composition and structure of planetary bodies.

612 With a geospace radar, the only planetary body that can be mapped with good resolution
613 is the Moon. The Moon is close enough that high signal-to-noise ratios can be achieved. It
614 is also large enough that the target size in range and Doppler shift permits a large number
615 of measurement points. Contemporary 6-meter wavelength radar mapping efforts achieved
616 approximately a 15x15 km resolution which is limited by ionospheric scintillations and the
617 limited tracking time possible. With a more flexible low frequency radar, it is plausible to
618 obtain range-Doppler resolution of approximately **1.5 km** × **1.5 km** per pixel with good
619 ionospheric conditions from a non equatorial location.

620 A low frequency polarimetric inverse synthetic aperture radar image of the Moon ob-
621 tained at Jicamarca [Vierinen *et al.*, 2017] is shown in figure 5. This 49.92 MHz radar
622 image is the lowest frequency polarimetric synthetic aperture map of the Moon that has

623 been made, allowing us to peer deeper beneath the surface of the Moon than before. This
624 frequency is close to the lowest possible frequency that can be used for a ground based ob-
625 servation due to deleterious effects of ionospheric scintillation and other radio propagation
626 effects.

627 The 6-meter wavelength radar map provides a unique view of the Moon. Resonant
628 scattering from Breccia in the vicinity of newer impact craters allows us to infer the amount
629 of large blocky material. These regions appear radar bright. Examples of such regions
630 include the area in and around the Copernicus and Tycho craters. The bulk radar brightness
631 on the other hand allows us to infer the dielectric properties of the subsurface. Lossy FeO
632 and TiO₂ rich mare regions of the Moon have a larger loss tangent and appear radar dark
633 while low loss terrae regions appear radar bright.

634 One of the key findings of the 6-meter wavelength study [*Vierinen et al.*, 2017] was the
635 discovery of a large radar dark region which joins Mare Frigoris and Mare Imbrium. This
636 finding supports the hypothesis that Mare Imbrium and Mare Frigoris are part of the same
637 impact basin – one of the largest impact basins on the Moon. The other key finding was
638 the radar dark Schiller-Zucchius impact basin on the southern hemisphere of the Moon.
639 This can be seen on the top-right panel of Figure 5. The radar observation supports the
640 hypothesis that this old impact basin is associated with basaltic flows which are now mostly
641 covered by optically bright terrae material produced by newer impact ejecta. These recent
642 findings show that low frequency planetary radar, which peers beneath the visible surface
643 of the Moon, can provide important new information about the geological formation of the
644 Moon which is still poorly understood.

5.6. Solar research

645 The cause celebre for geospace radar could well be solar research. The ability to receive
646 soundings from the sun with ground-based radar would represent a new avenue for study-
647 ing solar physics. It could expedite the creation of an operational space-weather forecast
648 strategy involving tracking along with prediction of CME trajectories. Finally, it would
649 unify the solar, magnetospheric, and ionospheric research communities under a common
650 observational paradigm. The concept is not new and is not simple but appears to be plau-
651 sible using available dual-use technology.

652 5.6.1. Historical background

653 Since the start of the planetary-radar era, multiple attempts have been made to detect
654 radar reflections (soundings) from the solar corona. The first, in 1959, was by a Stanford
655 University group working at 25.6 MHz [*Eshleman et al.*, 1960]. Eshleman reported detect-
656 ing echoes at 1.7 solar radii after 36 min. of integration using a monostatic radar system
657 with 25 dB of antenna gain and 40 kW average power. The results were deemed to be
658 consistent with expectations for an ideal, conducting sphere at the time but could not be
659 repeated.

660 Subsequently, a major, multi-year effort was undertaken by a group from MIT using a
661 dedicated 38.25 MHz radar built in El Campo, Texas [*James*, 1964, 1970] and operated
662 continuously from 1960 to 1969. The average power and gain of the El Campo radar
663 were 500 kW and 33–36 dB, respectively. These are very impressive specifications for
664 what would be a temporary, short-lived radar facility. Highly variable echoes (in terms
665 of power, Doppler shift, and bandwidth) were reported. Several analysts attributed the

666 variable echoes not to critical-frequency reflection but to coherent scatter from one or more
667 electrostatic plasma waves (e.g. [*Khotyaintsev*, 2005]). Despite the unexpected variability
668 of the signals, at the time and for decades thereafter, the El Campo results were considered
669 to be reliable and definitive if not particularly well documented proof of the solar-radar
670 concept. **The results have not been replicated, in part due to the absence of a suitable**
671 **facility for performing the experiments.**

672 Another attempt was made at Arecibo in 1966 and 1967, using a 40-MHz, 50/100 kW
673 average power transmitter [*Parrish*, 1968]. The gain of the Arecibo system at 40 MHz was
674 37 dB. The investigators duplicated the experimental mode utilized earlier at Stanford.
675 While preliminary work suggested that solar echoes were detectable, and while positive
676 results were again obtained when the experiments were repeated, the results were never
677 published. The perceived lack of novelty rather than any shortcomings in the research kept
678 the results from the literature (D. Campbell, 2016, personal communication.) The 40-MHz
679 transmitter was decommissioned, and no further attempts have been made at Arecibo.

680 In 1996, solar radar experiments were attempted in the FSU using the SURA heater in
681 Russia as a transmitter and the UTR-2 radio telescope in Ukraine as a receiver [*Rodrigues*,
682 2013]. The study was not comprehensive, and the findings were neither conclusive nor
683 well documented.

684 Most recently, solar radar experiments were attempted repeatedly at Jicamarca during a
685 multi-year study [*Coles et al.*, 2006]. An unsuccessful attempt to observe solar echoes at
686 Jicamarca was actually made back in 1964, but it was not documented (personal communi-
687 cation, K. Bowles, B. Balsley, and D. Farley). In the recent experiments, the average power

688 and antenna gain were about 112 kW and 41.6 dB, respectively. After a thorough statistical
 689 analysis of the data, no clear, unambiguous echoes were found, and the upper bound on the
 690 solar cross section was reduced to a figure below that implied El Campo observations. The
 691 implication of the study was that the echoes reported from El Campo might have been spu-
 692 rious and associated with solar radio bursts rather than radar reflections. This conclusion
 693 is controversial and does not explain the Arecibo results.

694 **5.6.2. Experimental demands**

695 As mentioned above, for solar, magnetospheric, and plasmaspheric studies, the radar
 696 wavelength must be longer than the plasma Debye length. This places a premium on low
 697 radar frequencies which overrides the penalty of increased sky noise. However, the radar
 698 frequency should not fall below the maximum usable frequency (MUF) since that would
 699 invite radar clutter from sky waves. The ideal frequency is therefore between 40–50 MHz.
 700 Some additional frequency considerations are discussed below.

The simple figure of merit described below is not appropriate for optimizing solar radar
 experiments in which the transmit and receive antenna specifications need to be be consid-
 ered separately. Optimum array sizes are determined by the power budget for solar echoes.
 The radar equation can be used to estimate the received signal flux:

$$S_r = P_t G_t L_p \sigma_r / (4\pi r^2)^2 \quad (2)$$

where P_t is the transmitted power, G_t is the gain of the transmitting antenna, L_p is the
 two-way loss factor for absorption in the lower corona, σ_r is the solar radar cross section,

and r is the solar distance. The noise flux is similarly given by

$$S_s = KT_s\Omega_s B/\lambda^2 \quad (3)$$

where K is Boltzmann's constant, T_s is the solar noise temperature, B is the bandwidth, and λ is the wavelength. The symbol Ω_s is the solid angle subtended by the Sun. It is assumed here that the Sun nearly (but not entirely) fills the field of view of the radar and that all of the noise sources outside the Sun may be neglected. Taking the ratio of Eqs. 2 and 3 gives the anticipated signal-to-noise ratio

$$\text{SNR} = P_t A_t L_p \epsilon_s / (4\pi r^2 K T_s B) \quad (4)$$

701 in which A_t is the transmitting antenna effective area and ϵ_s is the scattering efficiency of
 702 the Sun (**the ratio of the scattering cross section to the physical cross section**), which is
 703 assumed to be fully illuminated.

704 **Estimating the factors in Eq. 4 is challenging since solar echo detection remains**
 705 **to be demonstrated and since a mode capable of detecting them remains to be de-**
 706 **finied. The effective bandwidth of the experiments at Jicamarca following coherent**
 707 **processing is 1 kHz. James [1970] estimated the loss factor L_p to be about 3 dB, al-**
 708 **though that figure is probably an underestimate (see Coles [2004]; Coles et al. [2006]**
 709 **and references therein). The noise temperature of the quiet sun at 50 MHz is of the**
 710 **order of 10^6 K, although the actual system noise is dominated by solar radio bursts**
 711 **as discussed below.**

712 The crucial performance metric is the transmitter power-aperture product which sets
 713 the flux that can be delivered to the Sun. In order to optimize this flux, the antenna for

714 transmission should be a steerable aperture or filled array at least comparable in size to
715 Jicamarca's. Steerability is necessary to keep the radar beam trained on the Sun, facilitating
716 long incoherent integration times.

717 Remarkably, the receive-array size does not enter into the calculation. In fact, the receive
718 array must be large enough that most of the noise it receives comes from the solar disk itself
719 and not from the galactic background. This assumption, which is not difficult to satisfy in
720 practice, is built into eq. 4. Moreover, we have to consider that the main source of noise in
721 solar radar experiments will be solar radio bursts.

722 Observations of the Sun at meter wavelengths with Culgoora and CLRO have shown
723 that the various radio burst (Types I, II, III, IV, etc) have at least 20 dB and often 30 dB
724 higher brightness than the quiet Sun. The most common burst, type III, comes from a small
725 region and is broadened by scattering in the corona to angular size of order 3 arc min at
726 80 MHz. When the radar receiver beam covers the whole Sun, as it did in the early solar
727 radar experiments, these bursts dominate the system noise temperature by a large factor.
728 Accordingly, a solar radar receiver must be able to resolve type III bursts so that the receive
729 beam can be directed to the quiet Sun between the bursts. In fact it would be preferable
730 to design the receive beam to have nulls at the position of the type III bursts, implying
731 the need for a maximum likelihood beam-former. However the antenna must be capable
732 of resolving the burst in order to put a null on it and to perform imaging with very high
733 dynamics range [*Oberoi et al.*, 2018; *Mondal et al.*, 2018; *Sharma et al.*, 2018].

734 These considerations suggest that a suitable receiving antenna will be a phased array
735 with an aperture of order 10 km in diameter and a beam-former that can form multiple

736 simultaneous beams with very low sidelobes. However, special efforts to suppress grating
737 lobes will not be necessary since the Sun is by far the most powerful noise source in the
738 Sky, and very low-noise receivers will not be required. Instead, the receivers must be
739 optimized for high dynamic range and good stability. Good calibration and good stability
740 will be required to minimize sidelobes.

741 **5.6.3. Solar-radar opportunity**

742 The possibility of observing not only the sun but also solar arcs and coronal mass ejec-
743 tions (CMEs) with ground-based radar remains an attractive scientific objective and an
744 important impetus for this project. A solar-radar capability could possibly be used to es-
745 timate the range, bearing, and speed of a CME. Such information could form the basis
746 of practical space weather event forecasts. It could also supply critical information to
747 more conventional model-based forecasts, including information about the coronal mag-
748 netic field inferred from Faraday rotation.

6. Facility capabilities and requirements

749 Table 1 summarizes the capabilities of a geospace radar which would be able to pursue
750 the research objectives described above. The categories in the table are broad, and the
751 delineations are somewhat artificial, as most of the radar's capabilities would be exploited
752 in every avenue of research over time as the research matures. The table does show how
753 progress across a broad span of geospace science would follow from the pursuit of the
754 geospace-radar concept.

755 Most every scientific application requires high peak power of the order of one to several
756 MW. Additional sensitivity would come from high average power or even CW operations
757 involving the transmission of long pulse codes. Sensitivity will be critical for applications
758 where either transmission or reception occurs on subarrays of the main antenna array, for
759 example in MIMO and imaging applications. High average power is important especially
760 for estimating meteoric mass flux and for probing distant targets, the solar corona most
761 acutely. High duty cycles are a cost-effective strategy for achieving high average power
762 and good sensitivity in view of the fact that costs are driven largely by peak power.

763 Multistatic observations are the only way to measure vector drifts unambiguously and
764 are important for every avenue of geospace research except perhaps planetary radar. The
765 contemporary practice of inferring three-dimensional flow fields from monostatic line-of-
766 sight drift measurements is impeding our understanding of the complex dynamics found
767 throughout the upper atmosphere. The ability to track CMEs with multistatic solar radar
768 would provide an incisive tool for space-weather forecasting.

769 Distributed, modular receive arrays supported by multichannel receivers are central to
770 the radio-array concept and the foundation of true radio and radar imaging. More than any
771 other capability, the modern geospace radar would rely upon distributed spatial sensing
772 to unravel space-time ambiguity and reveal the three-dimensional structure of targets and
773 features in the upper atmosphere, the plasmasphere, and the solar corona. While planetary
774 radars derive imaging information from the Doppler shift, interferometry is essential for
775 disambiguating echoes from the northern and southern hemispheres.

776 Both multi-static and distributed modular arrays could be implemented with current tech-
777 nology used in modern radio arrays, e.g., LOFAR, LWA, MWA. To use such systems as
778 receivers of the proposed main transmitter array, the capability to record signals centered
779 around the frequencies of interest is needed. Such capability has been proven already at
780 KAIRA, LOFAR International, and LWA-SV, receiving signals from transmitters operat-
781 ing from a few MHz to 54 MHz. In the case of solar-radar applications, an array of arrays
782 with spacings up to 10 km or so, like LOFAR-core (e.g., *Kontar et al.* [2017]), would be
783 needed. To be able to get both high resolution images of solar radio burst and solar radar
784 echoes, the pre-combined complex signals of each array need to be recorded so that spatial
785 autocorrelation functions are available and, from them, the brightness of solar burst as well
786 as of solar echoes. Complex voltages would be needed to accommodate the transmitter
787 modulation.

788 Full polarization diversity is absent in all of the current ISRs with the exception of Jica-
789 marca. Polarimetry is used there now to measure Faraday rotation which serves as means of
790 calibrating ISR power measurements absolutely. Echoes from meteors and from the plan-
791 ets can also be polarized due to incidences of multiple scatter. The polarization of meteor
792 echoes in particular is an important source of information about the scattering mechanism
793 that has been largely overlooked. Looking forward, polarimetry could be a rich new source
794 of information about geospace and could potentially reveal the magnetic structure in the
795 corona and the solar wind in the case of solar radar. Work along these lines is already
796 underway *Salah et al.* [2005]; *Bisi et al.* [2017]; *Lenc et al.* [2017].

797 Wide-band and multi-frequency capabilities represent another new frontier for discovery
798 science. Solar-radar modes are inherently wide band because the modulation schemes
799 involved use frequency-shift keying (FSK).

800 In cases where the target RCS is large enough, e.g., MST, Meteor, FAIs, and perhaps
801 the Moon, coherent and non-coherent MIMO implementations will be useful. The avail-
802 able high power could be diversified to study such targets with unprecedented horizon-
803 tal and altitude resolution and horizontal coverage. For example, FAIs could be studied
804 with simultaneous multi-static links and aperture synthesis radar imaging by employing
805 interferometric configurations on transmission with diversity (e.g., code) and receiving on
806 multiple receiving sites consisting of single antennas (MISO configurations) or more an-
807 tennas (MIMO). This would facilitate studies of FAIs at different scattering points along
808 the Earth's magnetic field in addition to traditional imaging in just the transverse direction.

7. Concluding Remarks

809 A geospace radar as described in this paper would lead not only to a better understanding
810 and practical utilization of the natural echoes studied with existing radars but also to an
811 appreciation of more challenging targets and regions that are waiting to be explored. Here,
812 we have considered just a few research opportunities starting from the lower atmosphere
813 and arriving at the Sun.

814 Although special emphasis has been devoted to applications related to space-weather
815 research and operations like the early detection of the sources of the most extreme events
816 (i.e., CMEs) and efforts to forecast the day-to-day variability of ionospheric irregularities

817 (e.g., ESF), a geospace radar would contribute to other basic and applied research areas.
818 To name a few: plasma instabilities in the plasmasphere as well as the valley region, turbu-
819 lence and layering processes in the MST region, meteor composition, meteor mass, solar
820 plasma processes, etc. Furthermore, the receive array portion of the radar could be used
821 passively for radio astronomy or be paired with HF sounders for multistatic ionospheric
822 sounding experiments. **The solar-radar application remains the most compelling sin-
823 gle science objective, however, being a compelling new avenue of remote sensing in its
824 own right and holding the promise for end-to-end space-weather forecasting.**

825 As with any large endeavor, the implementation and subsequent operation of a geospace
826 radar would require a multidisciplinary approach. The effort would necessarily involve,
827 for example, signal processing, statistical inverse theory, software radio, big data han-
828 dling/processing, machine learning, energy efficient practices, etc. That is, it would require
829 the participation and interaction of a workforce encompassing key disciplines of society
830 drawn from science, technology, engineering, and mathematics (STEM) working toward
831 societal security and prosperity.

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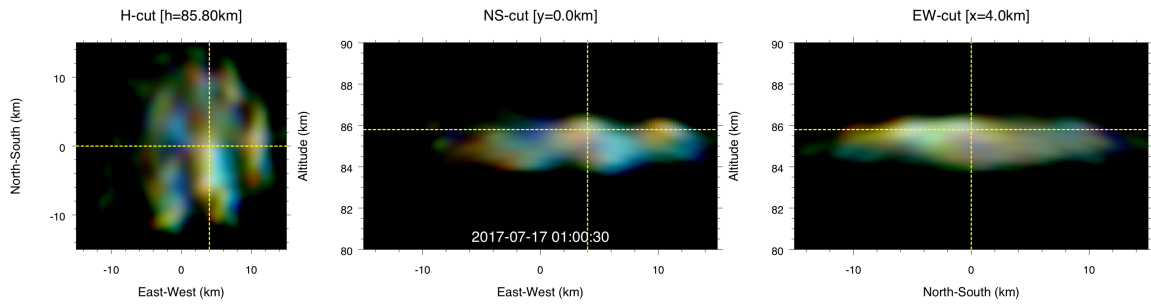


Figure 1. Example of 3D PMSE images obtained with a combination of MaxENT radar imaging and time-diversity MIMO techniques using the MAARSY radar: (a) horizontal cut at 85.80 km, (b) altitude versus zonal cut, and (c) altitude versus meridional cut. Doppler velocity, spectral width, and intensity are color-coded in the images. Adapted from *Urco et al.* [2018b]

Mon Sep 4 09:12:44 2017

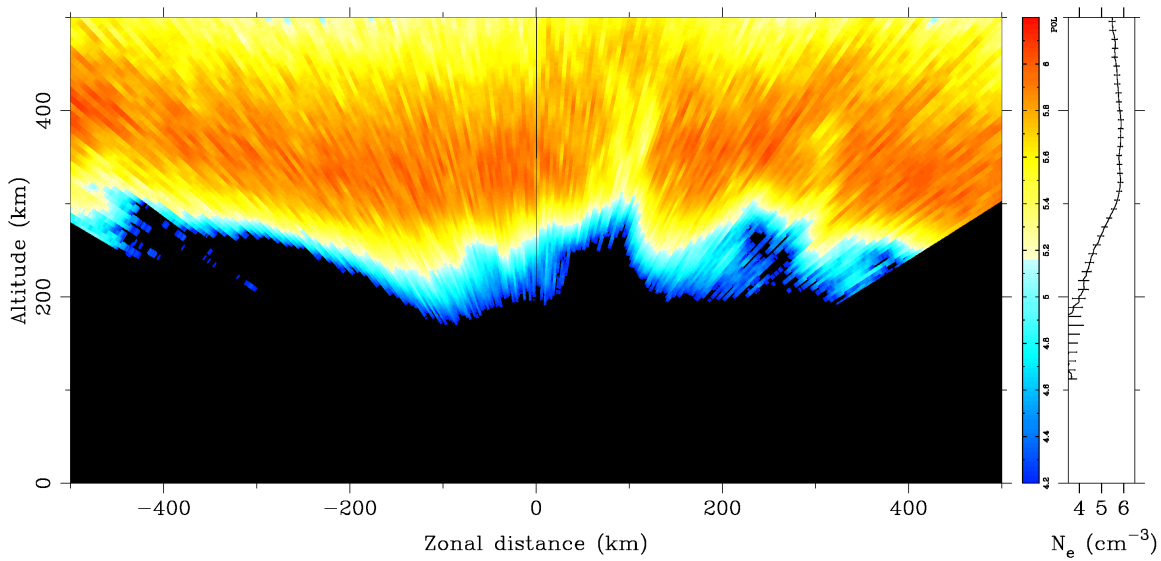


Figure 2. Incoherent scatter observations of large-scale plasma density irregularities and plume-like depletions associated with equatorial spread F made with the ALTAIR radar. The data were acquired with antenna scans directed a few degrees away from perpendicular to the geomagnetic field.

Mon Sep 4 09:03:02 2017

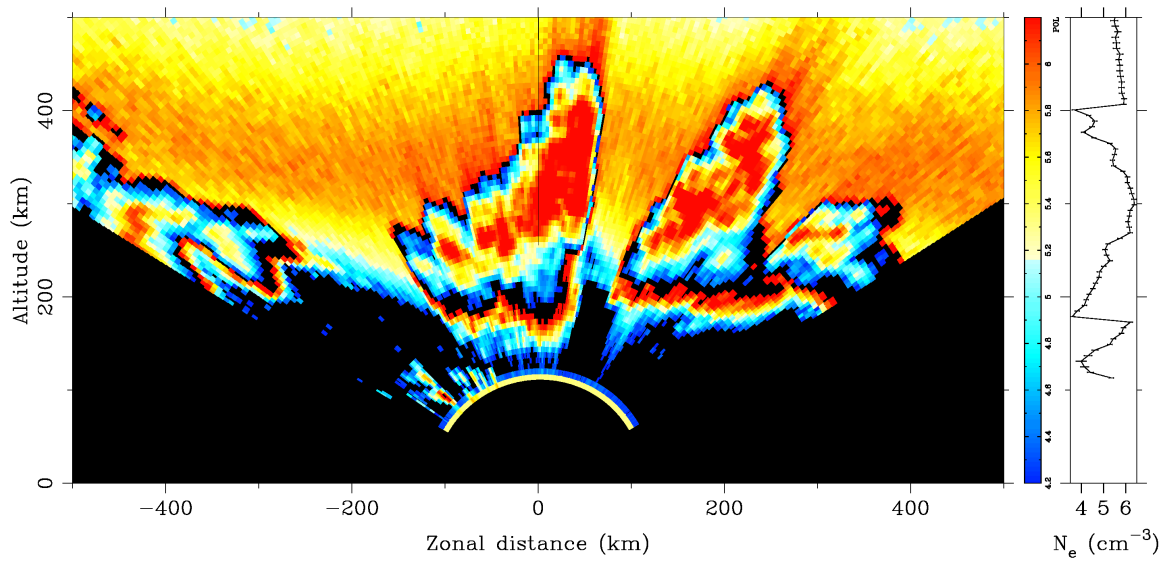


Figure 3. Radar observations of irregularities associated with equatorial spread F made with ALTAIR antenna scans directed perpendicular to the geomagnetic field. Coherent scatter can be observed in the most deeply depleted parts of the F region. In this figure, the coherent scatter is automatically attenuated by 20 dB for plotting.

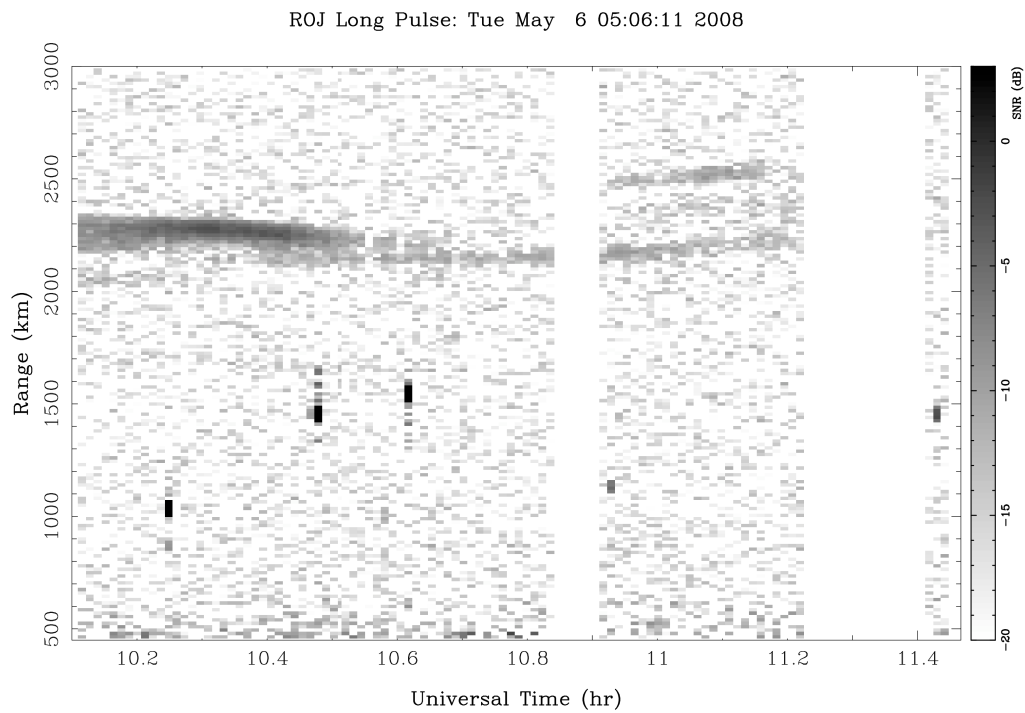


Figure 4. Coherent scatter signal-to-noise ratio derived from topside observations at Jicamarca in the pre-dawn sector. For this figure, radar clutter due to satellites and debris most evident between 1000-1500 km has not been removed.

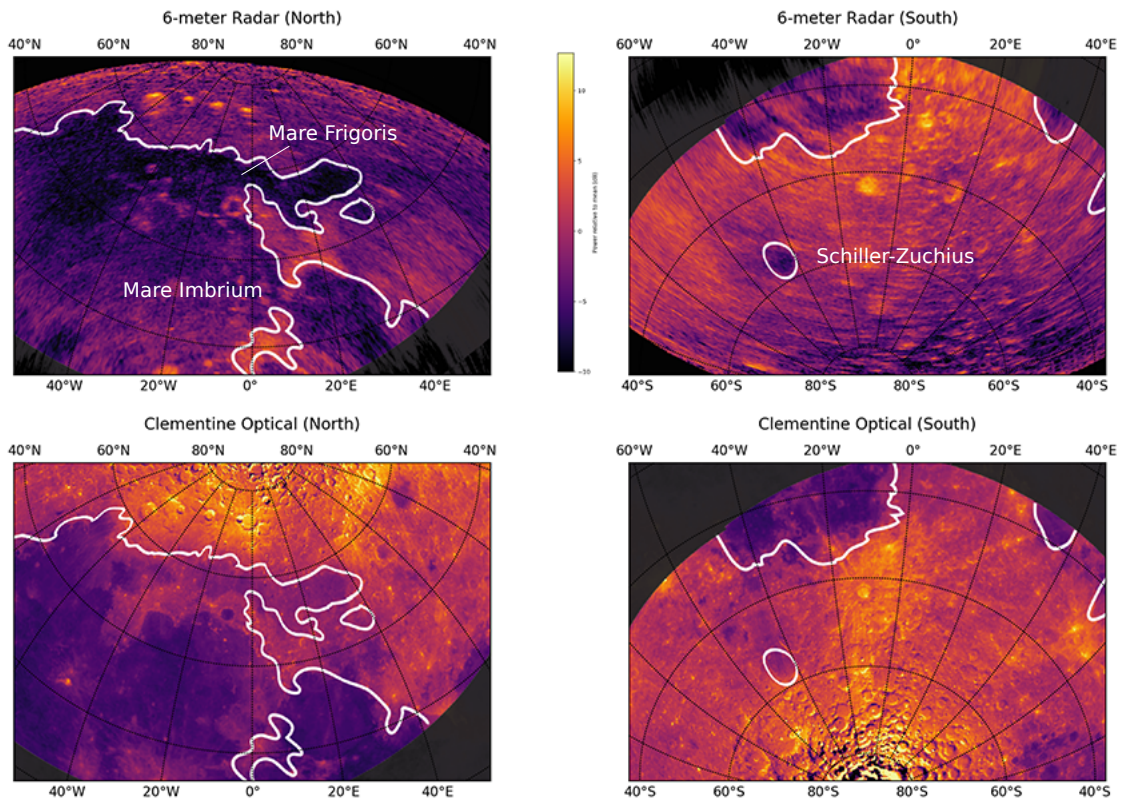


Figure 5. Top: 6-meter wavelength radar map of the Moon derived from Jicamarca data with polarization orthogonal to the specular polarization. Bottom: Clementine optical image of the Moon.

Capability	MST	Meteor	ISR	FAIs	Plasmasphere	Planetary	Solar
High peak power	✓	✓	✓		✓	✓	✓
High duty cycle	✓	✓	✓	✓	✓	✓	✓
Multistatic	✓	✓	✓	✓	✓		✓
Distributed	✓	✓	✓	✓	✓	✓	✓
Polarimetric		✓	✓			✓	✓
Wide band	✓	✓					✓
MIMO	✓	✓		✓			

Table 1. Capabilities versus potential applications.