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TRANSPORT AND FATE OF NATURAL GAS AND BRINE ESCAPING FROM A
HYDROCARBON RESERVOIR THROUGH A FAILED DEEP-WATER WELL IN THE
OCEANIC SUBSURFACE OF THE GULF OF MEXICO

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26Abstract

27The possibility of broaching, or the release of fluids at the seafloor due to a damaged or faulty
28well, is a hazard that must be assessed in the well-permitting process. This paper describes a
29numerical simulation study of a real-life scenario where a complex, permeable sandy formation,
30connected to the seafloor via known chimneys/seeps, is intersected by a damaged production
31well that drains another deeper, gas-bearing formation. The objective of the study to determine
32the transport and fate of hydrocarbon reservoir fluids (gas and brines) escaping into the sandy
33formation through the casing shoe of the failed well, and to determine the time it takes for these
34contaminants to reach the ocean floor. We conducted a detailed simulation study to represent the
35conditions, properties, and behavior of the system under such failure conditions, and we
36investigated the migration of gas and brine for a range of reservoir and chimney properties. A
37key conclusion is that, for such complex systems, modeling the three-dimensional geometry of
38the system in detail is key to describing transport and assessing the time and magnitude of
39potential releases. For the system studied here, transport times range from under two years
40(highest permeabilities) to many decades, ensuring significant time to respond to potential
41broaching hazards. Under the conditions investigated in this study, we also determine that gas-
42dominated releases associated with low rates of water flow into the sandy formation are likely to
43cause hydrate formation that can reduce permeabilities in the colder, upper regions of the
44chimneys and possibly mitigate releases.

45

46Keywords

47Broaching, Hydrate Formation, Well Failure, Hazard Assessment, Reservoir Simulation

48

491. Introduction

50

51 Oil and gas production in the Gulf of Mexico often requires complex drilling technologies,
52 including the use of directional drilling at great depths. Improperly constructed or damaged well
53 casings could allow the escape of reservoir fluids into formations they intersect, creating the
54 possibility of contaminant releases at the seafloor—a phenomenon referred to as broaching. The
55 Bureau of Ocean Energy Management (BOEM) began performing broaching evaluations after
56 the uncontrolled release of oil and gas from the Macondo reservoir in Mississippi Canyon Block
57 252 (Oldenburg et al., 2012). The sub-seafloor segments of the Macondo well had apparently
58 maintained integrity and the well was eventually capped successfully. The Macondo study
59 addressed the problem of well failure at the ocean floor, and as such could not provide answers
60 to the question of how long it could take for broaching to occur after a well failure deeper in the
61 oceanic subsurface. For scenarios where a well failure creates connectivity between a deep, over-
62 pressured hydrocarbon reservoir and overlying permeable formations, there is no methodology to
63 determine the fate and transport of the released fluids and whether broaching is a possibility.
64 Hazard assessments must be performed to determine whether uncontrolled releases could be
65 contained and stopped in time to prevent escape of hydrocarbons into the environment. For
66 example, how much time would be available to drill a relief well before broaching occurs at the
67 seafloor?

68

69 To answer some of these questions, BOEM completed an Interagency Agreement with the U.S.
70 Department of Energy/Lawrence Berkeley National Laboratory to carry out 3D modeling of
71 possible scenarios of reservoir fluid releases from compromised casing shoes in two real-life

72systems that had been qualitatively judged as either likely or possible to result in broaching. This
73paper describes the numerical simulation work for scenarios associated with the first system: a
74“supra-salt” system where a permeable formation, connected to the seafloor via known
75chimneys/seeps, is intersected by a damaged well producing from a deeper formation.

76

77The system of interest involves a roughly 6 km x 10 km (3.7 x 6.2 mi) region of permeable sand
78(the Purple Sand) intersected by a well producing from a deeper, warmer, higher-pressure
79formation (Figure 1). The Purple Sand is flanked by two chimneys (Figure 2) that are considered
80permeable to fluids escaping from deeper hydrocarbon reservoirs in the Gulf of Mexico -- at
81least one of them has been confirmed to be a prolific hydrocarbon seep (Garcia-Pineda et al.,
822010; Roberts et al., 2010). Portions of the Purple Sand formation climb updip from the potential
83casing-shoe failure point toward the two chimneys, both of which are confirmed active oil and
84gas seep sites. This suggests there are clear pathways for migration to the seafloor of fluids
85released into the Purple Sand. The seeps, seismic wipe out zones, and sea bottom seismic
86amplitude anomalies in association with them provide strong evidence for a leaky system (Figure
872).

88

89The goal of the overall simulation study is to assess the potential transport of reservoir fluids
90through the Purple Sand and the possibility of the emergence of reservoir fluids (gas and brine in
91this case) at the seafloor if a permeable pathway is created between the deeper reservoir and the
92Purple Sand via a failure in the intersecting well. To accomplish this, we conducted a detailed
93simulation study that faithfully represented the conditions, properties, and behavior of the system
94under such a release.

962. Methodology

972.1 Mesh Generation and System Initialization

98The effort to discretize the domain for the simulation effort included two different sets of grids.

99The first set was that of the elevation grids and included elevation data (from the dataset
100represented by Figure 2) to the top of the Purple Sand on a regular pattern (ranging in spacing
101from 150 ft x150 ft to 50 ft x50 ft). The elevation grids served as the basis for the development of
102the simulation grids, i.e., the element system used by the pTOUGH+HYDRATE simulations
103(Moridis et al., 2008; Zhang et al., 2008; Moridis, 2014). Unlike the elevation grids, the
104simulation grids (which were developed by the LBNL team) were irregular in the element size,
105providing higher definition in the vicinity of important features and a coarser discretization at
106locations that did not appear to play key roles in the fluid transport (as determined and confirmed
107by scoping simulations).

108

109BOEM provided mapped data from 3D depth migrated seismic showing the geometry of the
110geosystem over multiple Outer Continental Shelf (OCS) blocks. Mapped data was interpreted on
111a GeoFrame workstation using U.S. feet in Universal Transverse Mercator (UTM) 15 projection.
112Mapped surfaces on the top of salt and on the top of the Purple Sand and 3D seismic arbitrary
113lines provided the system geometry, and GIS shape files of sea floor features and seismic
114amplitude were available for overposting. The mechanical and formation fluid properties for the
115Purple Sand were obtained from well logs by BOEM staff using Log Evaluation System Analysis
116software.

117

118The extreme thinness of the Purple Sand carrier bed compared to the lateral dimensions of this
119formation, the very large height of the chimneys, the undulations of the sand, and the steepness
120of the elevation gradient near the base of the chimneys, coupled with the need to integrate a flow
121pathway to each seep site resulted in frequent vertical discontinuities using conventional
122TOUGH+ mesh generation methods. The MeshMaker 2.0 (MM2) gridding software (Moridis
123and Keen, 2015; Moridis, 2016) was developed specifically to address the challenges of this
124system geometry. MM2 uses the elevations of the top of the formation over its footprint (using
125data provided in tabular form) and powerful 3D interpolation algorithms to create grids that can
126conform to practically any irregular geometry. The new grids created by MM2 eliminated a
127common continuity problem in undulating surfaces by incorporating directional slopes and
128correct volumes and surface areas into elements describing those surfaces. MM2 thus allows
129irregular meshes to represent complex surfaces, while still providing the numerical stability
130afforded by meshes of pure “stacked bricks” without the distortions that such configurations
131impose on fluid flow.

132

133Several realizations of various grid refinement levels were constructed following the process
134described above, including up to 6.7M gridblocks. After comparisons of the 3D images of the
135grids (in an effort to determine which iteration captured key geometric features of the system)
136and of the corresponding execution times, a grid involving 2.65M gridblocks (Figure 3) was
137selected as the most appropriate for the simulations because it reproduced with high fidelity the
138system geometry while requiring reasonable (relatively-speaking) execution times.

139

140The final 3D system, with 2,647,145 gridblocks, was brought to a physically correct initial state
141via numerical simulation, using as fixed reference points the pressures corresponding to the
142seafloor depths at the tops of the two chimneys and temperatures at the seafloor, as well as a
143representative “average” sub-seafloor temperature gradient (see Table 1). A long-term simulation
144then calculated the hydrostatic pressure field and the equilibrium temperature profile for the
145given sediment properties. For systems that lie within the zone of hydrate stability (and thus have
146a potential for hydrate occurrence or hydrate formation if methane becomes available at the
147location under investigation), this is the only physically correct and consistent way to establish
148an initial condition. The results of this simulation provided stable initial conditions for
149subsequent fluid release modeling.

150

151The initial conditions included:

- 152 • A hydrostatic pressure distribution in the entire 3D domain (Purple Sand and chimneys,
- 153 • A specified average temperature gradient from the BOEM dataset: 2.71°C/100m, with an
154 estimated seafloor temperature of 3.9°C at the seafloor (Gallaway et al., 2001; Forrest et
155 al., 2007),
- 156 • An assumption of 1 ppm background methane in the aqueous phase. This was a
157 reasonable assumption (given the lack of information on the subject), and has no impact
158 even if it is incorrect because (a) methane solubility in water is very low and (b)
159 dissolved gas is overwhelmed when gas from the reservoir is released either as free gas.

160

161The physical parameters used were supplied by BOEM, or selected under consultation with
162BOEM, and are given in Table 2 and Table 3. Figures 3 and 4 show the mesh connectivity and
163the initial pressure distribution at the various components of this 2.6M-element system viewed

164from different angles. In Figure 3, an overall view of the system (view at an angle, from above),
165shows the extensive but thin Purple Sand, with the two chimneys, A and B, reaching upward to
166the seafloor. Note that impermeable sediments, represented as heat-transfer boundaries, are not
167shown in this or any other rendering—only permeable formations are visualized. Figure 4 shows
168a rotated view of the system, with a close-up of Chimney B and its mesh connectivity. These
169visualizations of the simulation mesh will be used later in this paper to render the water and gas
170distributions in the failure scenarios. Visualizations were performed using the tough-convert
171postprocessor (Queiruga, 2018).

172

173**2.2 Reservoir Simulation**

174Following mesh creation and system equilibration, we conducted a set of simulations describing
175several scenarios of casing-shoe failure. We varied the permeability of the sand and chimneys to
176assess the basic sensitivity to system properties. Because of lack of information on its properties
177(and a conservative approach), the fault near the base of the Chimney A was treated as a
178permeable feature, and the elevation offset was described using steeply angled but connected
179gridblocks. In the spirit of the “worst-possible scenario”, this approach allowed the fastest
180possible upward flow toward the chimney.

181

182For each release scenario, constant-pressure and temperature conditions were imposed at the
183casing shoe failure point. These were the conditions in the connecting reservoir, without any
184adjustment for the elevation differences between the reservoir and the casing shoe: $P = 51$ MPa
185and $T = 63.3^\circ\text{C}$. Because the pressure and temperature drops in the failed well could not be
186estimated, this conservative approach was followed. The gas saturation at the failure point was

187initially $S_G = 0.99$, with an alternative case of $S_G = 0.89$ used to study the effects of water release
188in concert with gas release. The point of leakage was also given the properties of an open well, a
189pseudo-medium with $\phi = 1$, a high $k = 5 \times 10^{-9} \text{ m}^2$, a capillary pressure $P_c = 0$, a relative
190permeability that was a linear function of the phase saturations in the wellbore, and a low (but
191nonzero) irreducible gas saturation $S_{irG} = 0.005$. Simulations were run to a maximum duration of
192100 years, or until a logical stopping point was reached, e.g., gas and brine broaching at one or
193both chimneys. During the simulations, we monitored the flow of fluids through the tops of the
194chimneys and monitored T and P conditions at selected points within the chimneys and in the
195Purple Sand.

196

197The parametric variations used for the sensitivity study were:

198Released Gas Saturation, S_G :	0.89, 0.99, with corresponding aqueous saturations
199	$S_A = 0.11$ and 0.01 , respectively
200Chimney Permeability, k_C :	$k_C = 541 \text{ mD}$, 100 mD , 1000 mD for $k_S = 541 \text{ mD}$
201Sand Permeability, k_S :	$k_S = 541 \text{ mD}$, 100 mD , 1000 mD for $k_C = 541 \text{ mD}$

202

203Simulation runs were performed using the pTOUGH+HYDRATE parallel code (Zhang et al.,
2042008), an MPI parallel version of the TOUGH+HYDRATE (T+H) code. T+H can model all the
205known processes involved in the system response of natural CH_4 -hydrates in complex geologic
206media (Moridis et al., 2014; Moridis and Pruess, 2014). T+H is a fully compositional simulator,
207descended directly from the TOUGH family of codes (Pruess, 2003), and it accounts for heat and
208up to four mass components (i.e., H_2O , CH_4 , CH_4 -hydrate, and water-soluble inhibitors such as
209salts) that are partitioned among four possible phases (gas, aqueous liquid, ice, and hydrate). It

210 can describe 15 possible thermodynamic states (phase combinations) of the $\text{CH}_4+\text{H}_2\text{O}$ system
211 and can handle the phase changes, state transitions, strong nonlinearities, and sharp fronts that
212 are typical of hydrate dissociation problems. Runs were performed on NERSC Edison and Cori I,
213 and LBNL's Lawrencium LR3 and LR4 using 384 to 4,096 processors and a total of about 1.1
214 million CPU-hours.

215

216 **3. Results**

217 **3.1. All-Gas Reference Case ($k_c = k_s = 541$ mD)**

218 *Water and Gas Flows.* Figure 5 shows the total upward gas and water flows at the seafloor
219 through the top of Chimney A and Chimney B (flows through A and B were summed in pT+H to
220 give an estimate of the total flow through both chimneys). As gas first approaches the seafloor,
221 upward water flow increases, driven by buoyant gas rise within Chimney B. The gas first reaches
222 the top of Chimney B, where the gas breakthrough at the seafloor begins at $t = 926$ d. Gas
223 escapes continuously from the top of Chimney B until an abrupt cessation of gas escape at $t =$
224 1,287 d. Total water flows decrease during the initial period of gas release, reaching a local
225 minimum just before the cessation of the gas release. Visualizations of the simulation results
226 (shown in the next section) indicate that 100% of the gas release occurs at Chimney B for $t \leq$
227 1,287 d. Water flows increase gradually until $t = 1,959$ d, at which point gas release begins at
228 Chimney A. The simulation was halted after 2,000 d. Water flows decrease once gas appears at
229 the top of Chimney B.

230

231 *Hydrate formation.* The cessation of gas flow and the reduction of water flows are explained by
232 the formation of hydrate near the summit of Chimney B. Figure 6 shows the total system-wide

233rate of hydrate formation vs. time for the entire simulation. Hydrate begins to form at $t = 840$ d,
234just before gas begins to escape into the ocean from the top of Chimney B. As shown by the
235visualization of the simulation results, this hydrate formation occurs entirely in Chimney B when
236the rising released gas arrives at depths with appropriately low temperatures for hydrate
237formation to occur. Hydrate formation continues as gas escapes from Chimney B at the seafloor,
238and proceeds at varying rates (as affected by changing temperature, pressure and salinity at the
239locations where hydrates form). Thus, the rate of hydrate formation Q_H reaches a first peak at
240about $t = 926$ d, and attains a maximum rate at about $t = 1,450$ d. The period from $t = 920$ d to
2411,200 d corresponds to the period of gas release at the seafloor. A column of rising gas forms a
242“tunnel” of solid, low-permeability hydrate with a permeable core, as coupled flow, thermal and
243phase behavior processes temporarily allow an open channel for gas to flow even as hydrate
244accumulates in the surrounding chimney sediments. As the hydrate saturations increase, the
245effective permeability of Chimney B decreases. Eventually, the effective permeability in the
246chimney drops to zero when hydrate saturations exceed 80%. When this occurs at about $t = 1,287$
247d, the top of Chimney B is effectively blocked, and the hydrate formation rate increases as gas
248rising in the chimney is trapped below an impermeable layer of hydrate.

249

250System-wide hydrate formation rates begin to increase again at $t = 1,900$ d, and this increase is
251associated with gas moving upward in Chimney A and encountering suitable pressures and
252temperatures for hydrate formation and stability. The decline in the corresponding water flow
253rates that are evident in Figure 5 are associated with hydrate formation in Chimney A and the
254resulting reduction in permeability. Continuation of gas flow and hydrate formation in Chimney
255A is expected to lead eventually to complete blockage to flow (as is the case in Chimney B).

256

257Figure 7 shows the cumulative hydrate mass formed within the entire simulated domain. We can
258identify the onset of hydrate formation in Chimney B at $t = 840$ d, the formation of hundreds of
259millions of kilograms of hydrate over the course of 1,200 d, and also the upward inflection in the
260cumulative curve at $t = 1,900$ d when hydrate begins forming in Chimney A.

261

262*Visualizations.* New 3-D visualization capabilities, developed in the course of the overall project,
263allow detailed tracking of gas plumes across the simulation domain. For this simulation case, the
264visualizations are renderings of mesh connectivity (each connection between elements as a line
265segment) as previously shown in Figures 3 and 4, with the gas and/or the hydrate saturations
266mapped onto the line segments via color scales. This allows some limited transparency in the
267visualization, helping to better represent the movement of gas within the thin body of the Purple
268Sand and along the sides of the chimneys.

269

270Figure 8 shows the evolution of gas saturation with a view directed toward Chimney B at
271selected simulation times. The intersection of the failed well with the Purple Sand layer is at the
272lower left corner of the domain. At $t = 180$ d and $t = 640$ d, the escaping gas expands outward
273from the point of the casing shoe failure, and is seen to preferentially move toward Chimney B
274(upper left). At $t = 870$ d, the gas has already reached Chimney B, and is seen rising (primarily
275by buoyancy, although there is some pressure differential and upward flow of water) through the
276side of the chimney facing the leak. The panel at $t = 1,000$ d depicts the phase saturation
277distributions just after gas reaches the seafloor (i.e., the boundary at the top of Chimney B). At
278this point, the permeability at the top of the chimney begins to decrease because of hydrate

279formation. The panel at $t = 1,360$ d provides a snapshot of the system just after the cessation of
280the flows of gas and water to the seafloor through the top of Chimney B, with the gas saturation
281at the top of the chimney showing evidence of uniformity—essentially the plugged chimney is
282now filling up with gas rising from below. The last frame, at $t = 2,000$ d, shows a chimney with
283the upper several hundred meters filled with gas saturations $S_G \geq 0.50$.

284

285The advance of gas toward, and release through, Chimney A is illustrated by the four frames in
286Figure 9. At $t = 60$ d and $t = 1,280$ d, the gas is seen moving in a broad front toward the chimney.
287By $t = 1,730$ d, the topographical characteristics of the Purple Sand cause the front to split into
288two plumes moving toward Chimney A, one heading upslope along the upthrown footwall of the
289fault to the chimney on the left, and one heading upslope along the downthrown hanging wall of
290the fault. By $t = 1,910$, the gas has reached Chimney A through both pathways, a lower entry
291point near the downthrown hanging wall to the right and through a more tortuous path over the
292upthrown footwall to the left. Breakthrough, or broaching, at the seafloor occurs soon after, as
293gas rises to the top of the chimney via buoyancy and pressure gradients.

294

295The evolution of hydrate saturation in Chimney B is illustrated by the four frames in Figure 10.
296At $t = 910$ d, hydrate appears along the side of the chimney as the rising gas deposits the first
297layers of solid hydrate. By $t = 1,090$ d, hydrate formation has reached the seafloor, and
298saturations increase rapidly near the top of the chimney, with a corresponding reduction in
299permeability. The frame for $t = 1,280$ d depicts the hydrate saturation distribution just as water
300and gas flows out of the chimney have ceased. At that time and beyond, gas accumulating in the
301chimney leads to hydrate formation at successively lower elevations within the chimney (in

302parallel with the accumulation of gas in those regions, see Figure 8). At $t = 2,000$ d, hydrate
303deposits at saturations of $S_H \geq 0.50$ are extensive throughout Chimney B even at elevations
304hundreds of meters below the seafloor, creating permeability barriers to upward flow.

305

3063.2. Sensitivity to Permeabilities.

307To assess the complex variation in breakthrough times (at the seafloor), multiple full simulations
308of the broached system were conducted using identical initial conditions, but higher or lower
309permeabilities for the Purple Sand and for the chimney media. The range of permeabilities
310included:

311

312Chimney Permeability, k_c : $k_c = 100$ mD, $k_c = 541$ mD, $k_c = 1000$ mD for $k_s = 541$ mD

313Sand Permeability, k_s : $k_s = 100$ mD, $k_s = 541$ mD, $k_s = 1000$ mD for $k_c = 541$ mD

314

315For the sensitivity simulations, simulation length was limited to the point of gas breakthrough at
316the seafloor, as dictated by the need to efficiently utilize limited supercomputer CPU time. For
317these sensitivity simulations, we did not conduct a full analysis of hydrate formation and
318permeability reduction because we expected similar behavior for all cases that exhibit
319breakthroughs. The breakthrough curves for the permeability cases listed above are found in
320Figure 11.

321

322As would be expected, lower-permeability features result in longer times until release at the
323ocean floor. Decreasing the chimney permeability by a factor of five results in a 60% longer
324breakthrough time (625 days) for gas to appear at the seafloor. Doubling the chimney

325 permeability, however, results in only a slightly faster appearance of gas at the seafloor (by 64
326 days only, probably because of the rapid rise of the buoyant gas in the limited vertical dimension
327 of the chimney). Decreasing permeability of the Purple Sand to 100 mD, however, has,
328 unsurprisingly, a large effect—gas takes 3.5X longer (3,858 days) to emerge at the ocean floor.
329 Doubling the sand permeability, conversely, only shifts the gas breakthrough forward by 309
330 days (617 days).

331

332 Plots of the arrival times at the seafloor (solid lines) and of the arrival times at the base of
333 chimney B (dotted line) are presented in Figure 12, which shows the significantly non-linear
334 dependence of these times on the permeabilities of the various components of the system. The
335 red curve gives arrival times at the seafloor for $k_s = 541$ mD and k_c values of 100, 541 (base), and
336 1000 mD. The blue curve shows arrival times at the seafloor for $k_c = 541$ mD, and k_s values of
337 100, 541 (base), and 1000 mD. Figure 12 also shows arrival times at the base of chimney B (first
338 gas exiting the Purple Sand) for k_{sand} values of 100, 541, and 1000 mD with $k_{chimney} = 541$ mD
339 (arrival times at the chimney did not vary significantly with $k_{chimney}$).

340

341 To help with hazard assessment, parametric expressions were created to estimate arrival time for
342 a range of permeabilities. A series of curve fits was used in an attempt to represent the limited
343 data points as a larger response surface. Various functional forms (polynomial, power laws) were
344 tested, and the resulting curves were selected for convenience and ease-of-use. Although the
345 system is very complex (and strongly dependent on the specific geometry of the sand in
346 particular), it is reasonable to expect the arrival times to vary smoothly and gradually with

347respect to permeabilities. Thus, the curves generated through fits to an exponential form gave the
348only smoothly varying representation.

349

350After fitting the data within Igor Pro (our data analysis and plotting environment), we generated
351the following equations:

352

353Equation (1) estimates the arrival time at chimney B for varying Purple Sand permeabilities.

354

355 $t_{chimney} = 418.62 + 4814.5 \exp(-5.2184 \times 10^{-3} k_s)$ (1)

356

357Equation (2) estimates the arrival time at the seafloor for varying chimney permeabilities when
358 $k_s = 541$ mD.

359

360 $t_{seafloor} = 855.51 + 1171.9 \exp(-5.1974 \times 10^{-3} k_c)$ (2)

361

362Equation (3) estimates the arrival time at the seafloor for varying Purple Sand permeabilities
363when $k_c = 541$ mD.

364

365 $t_{seafloor} = 584.77 + 5506.6 \exp(-5.1407 \times 10^{-3} k_s)$ (3)

366

367These single-parameter variations can be expanded to a multiparameter surface. However, one
368additional detail that must be included is that the transit times through the chimneys, while
369largely a function of chimney permeability, will also vary slightly due to Purple Sand

370 permeability. Differences in permeability may cause a different quantity of gas to arrive, or to
 371 arrive at a slightly different location. Additionally, attenuation via hydrate formation may change
 372 transit times. Consequently, a general equation for estimating breakthrough must include time-to-
 373 chimney, $t_{chimney}(k_s)$ and both $t_{chimney-seafloor}(k_c)$, and $t_{chimney-seafloor}(k_s)$. Calculated chimney transit times
 374 and the variation in transit times due to the effect of k_s are shown in Figure 13.

375

376 Now we can construct a parameterization of broaching time as:

377

$$378 \quad t_{seafloor}(k_s, k_c) = t_{chimney}(k_s) + t_{chimney-seafloor}(k_c) + t_{difference}(k_s) \dots\dots\dots(4)$$

379 where

$$380 \quad t_{difference}(k_s) = t_{chimney-seafloor}(k_s) - t_{chimney-seafloor}(k_c = 541 \text{ mD}) \dots\dots\dots(5)$$

381

382 Fitting each of the three terms to exponentials results in the following parametric expression:

383

$$384 \quad t_{seafloor}(k_s, k_c) = 513.95 \dots\dots\dots(6)$$

$$385 \quad + 4814.5 \exp(-5.2184 \times 10^{-3} k_s)$$

$$386 \quad + 1172.9 \exp(-5.1974 \times 10^{-3} k_c)$$

$$387 \quad + 696.78 \exp(-4.6773 \times 10^{-3} k_s)$$

388

389 A visualization of the 3D surface associated with Equation (6) is given in Figure 14.

390

391 Note that these estimates are entirely a function of the specific (complex) geometry of the Purple
 392 Sand (and, to a lesser extent, of the chimneys), as represented in the model. For a system of this

393complexity, the gas arrival times at various locations of interest in the domain are strongly
394dependent on the accuracy of the 3D model. Heuristics that attempt to generalize to other (even
395similar) systems are not likely to be accurate. Note that the potential for hydrate formation and
396permeability reduction must be considered in any analysis of similar systems.

397

3983.3. Water and Gas Release

399*Water and Gas Flows.* Preliminary data provided by BOEM suggested a water saturation of
400approximately 11% within the underlying reservoir. As this is somewhat larger than the
401irreducible saturation of the Purple Sand, we considered an additional gas-release case where the
402released fluids included 11% water (likely an upper limit for water release). All other initial
403conditions were equal to the previous case.

404

405The water-gas case exhibits drastically lower release rates at the shoe failure point, and much
406longer timeframes for gas migration. Figure 15 gives water and gas flow rates and cumulative
407release at the point of the leak vs. time. After an initial surge for $t < 1$ d, the release rates drop
408rapidly and eventually settle at a low rate of $0.03 \text{ m}^3/\text{s}$ of gas and 19 kg/s for water. This is
409because of relative permeability effects, limited pressure gradients and weak buoyancy (as the
410density difference between the released brine and the native water is small).

411

412Monitoring points within the Purple Sand (a point opposite the broach) and points in the center
413of each chimney tracked evolving conditions during the course of the release (Figure 16). We
414observe sharp increases in temperature within the chimneys after a few hundred days of leakage.
415This is caused by the constant flow of the warmer water into the system, and the temperature

416 increases are attributed to both the temperature of the leaking water (63.3 °C) and the fact that
417 warmer water from deeper parts of the Purple Sand is swept upward toward the seafloor. As a
418 result, the simulation was run for an additional 30 yr, to $t = 100$ yr, and yet no hydrate formation
419 occurs anywhere in the system.

420

421 *Visualizations.* Visualizations of the evolving system are shown in Figure 17. At $t = 4$ yr, the
422 expanding plume of gas moves away from the failure point, but much slower than in the all-gas
423 case (Figures 8 and 9). By $t = 25$ yr, the plume has separated into two lobes, each heading toward
424 each of the chimneys via the local high point (saddle) in the Purple Sand that connects the bases
425 of the chimneys. Gas reaches the base of the chimneys at $t = 41$ yr, and reaches the seafloor at
426 both chimneys at $t = 70$ yr. Careful examination of the visualizations shows the appearance of
427 gas throughout and Purple Sand, and in areas away from the main plume as well, after $t = 41$ yr.
428 This is caused by exsolution of dissolved gas from the aqueous phase due to the temperature
429 increases as warmer water moves outward and upward from the release point at the casing shoe.

430

431 To track the movement of the water front, Figure 18 visualizes the salinity of the pore water over
432 the same time intervals as seen in Figure 17. To create a natural tracer, the reservoir water
433 entering at the failed casing shoe was given a concentration of 3.6wt% vs. 3.5wt% for water in
434 the Purple Sand at the initial condition. This allows clear visualization of the water front.
435 Compared to the narrow plumes of gas seen moving along the high points of the Purple Sand
436 topography, the water front expands outward uniformly (because of favorable relative
437 permeability conditions and very small density differences with the native water), initially
438 outpacing the released gas at $t = 4$ yr and 25 yr. The more saline water enters the chimneys

439roughly at the same time as the gas at $t = 41$ yr, and reaches the seafloor at $t = 70$ yr. However, by
440this time the reservoir brine has largely displaced most of the water in the Purple Sand.

441

442The most important conclusion from this alternative scenario is that the release of water with the
443gas shifts the timeframe for potential broaching from a few years to several decades. Release of
444water also adds heat (and, if the reservoir fluids are even more saline, an inhibitor) and prevents
445the formation of hydrate at the tops of the chimneys. Release rates are thus much smaller,
446broaching is delayed significantly, but the mitigating effects of hydrate formation will not be a
447factor.

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4504. Conclusions

451The key conclusion from this investigation is that, for a system as complex as the Purple
452Sand/Chimney system, the three-dimensional geometry of the system is key to assessing the time
453and magnitude of releases. Any release pathway will have a length determined by a potentially
454tortuous gas or water pathway dependent on the topography of the sand body. The volume of this
455pathway will be dependent on the thickness of the sand. Changes in the geometry or topography,
456changes in connectivity between features, and any known heterogeneity must be included in a
457detailed and realistic model. Multiple release pathways between the shoe failure point and the
458chimneys may form depending on the contours of the permeability zone or zones. The
459multiphase flow regime (proportion of gas/water released) will also affect release rate, release
460travel time, and also alter the shape and length of the release pathway. All of these factors must
461be considered in hazard assessments. For this particular scenario, we expect timeframes from

462two years to many decades between casing shoe failure and potential broaching, depending on
463system permeabilities and the composition of the release. It is thus likely that there would be
464sufficient time to drill a relief well and control fluid releases before broaching of hydrocarbons at
465the sea floor could occur.

466

467Once the geometries are established, parametric variations within those geometries result in
468fairly straightforward changes in expected travel times, with the unsurprising conclusion that the
469permeability of the pathway will greatly control the release timeline. Relative permeability
470effects related to the composition of the release would have even greater consequences, reducing
471leakage rates at the well and delaying releases at the seafloor. Alterations of the estimated
472thickness of the Purple Sand would likely have similarly large effects.

473

474For systems with gas release and low rates of water injection into the system (water flow limited
475to displacement by the gas phase) it is likely that hydrate formation will reduce permeabilities in
476the colder, upper regions of the system and possibly mitigate release under the conditions seen in
477the simulations. One question that we cannot address at this point is how a system with all outlets
478blocked by hydrate would eventually behave if releases continued for longer periods of time.

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4805. References

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 529 Columbia, Canada, July 6-10, 2008

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Table 1: Reservoir Properties Summary

Properties of Sand Layer and Chimneys			
Location	Parameters	Value	Unit
General	Geothermal gradient	0.0271	°C/m
	Temperature at ocean floor	3.9(*)	°C
	Porosity:		
	Average	0.25	fraction
	Low	0.1	fraction
	High	0.33	fraction
Purple Sand	Permeability:		
	Average	541/5.3E-13	mD/m ²
	Low	0.7/7.0E-16	mD/m ²
	High	1416/1.41E-12	mD/m ²
	Thickness	10/3	Ft/m
	Temperature (3,020 m)	127/52.8	°F/°C
	Well (x,y,z) at Purple Sand		
	x	9659.42	m
y	4050.77	m	
z	-3136.39	m	
Chimneys	Porosity	Same as Purple Sand	
	Permeability	Same as Purple Sand	
	Depth at ocean floor:		
	Chimney A	-1,135.38	m
Chimney B	-1,181.10	m	
Casing shoe	(x,y,z) of casing shoe		
	x	9799.3	m
	y	3810.0	m
	z	-2915.1	m
Escaping Fluids	P	7385	Psi
	T	146/63.3	°F/°C
	X _{CH4}	100	%

534(*) Gallaway et al., 2001; Forrest et al., 2007

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Table 2. Reservoir parameters

Region	k (mD)	ϕ	S_H	S_w
0. Boundary Layers	0	0.0	0.0	1.0
1. Chimney A	541	0.25	0.0	1.0
2. Chimney B	541	0.25	0.0	1.0
3. Purple Sand	541	0.25	0.0	0.0
4. Failed Well Casing	500-5000	1.0	0.0	varies

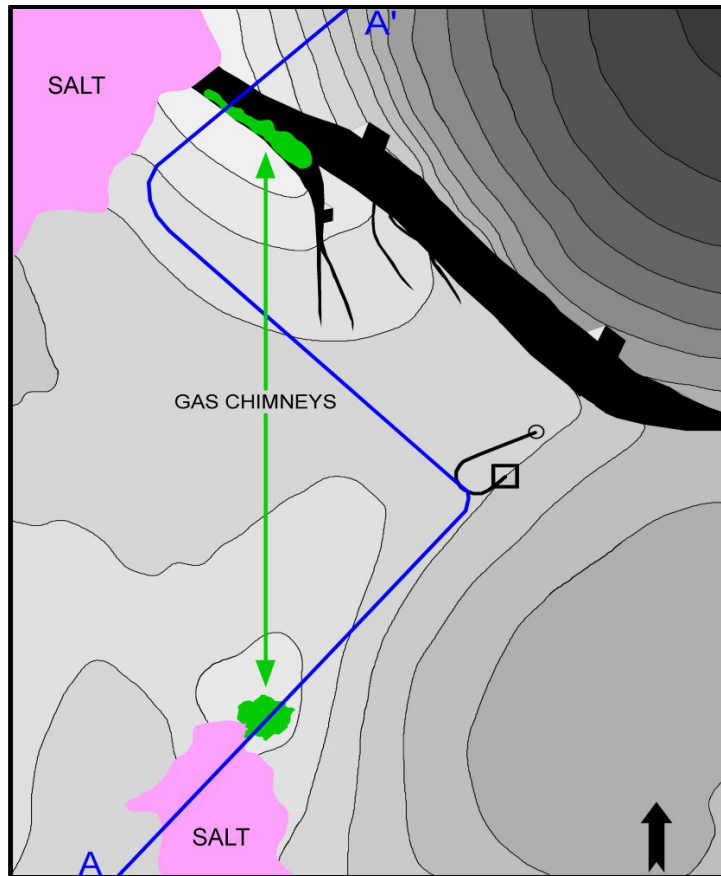
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Table 3. Other simulation parameters.

Water salinity (mass fraction)	0.035
Grain density ρ_R (all formations)	2600 kg/m ³
Dry thermal conductivity k_{QRD} (all formations)	0.5 W/m/K*
Wet thermal conductivity k_{QRW} (all formations)	3.1 W/m/K*
Composite thermal conductivity model (Moridis et al., 2008)	$k_C = k_D + (S_A^{1/2} + S_H^{1/2})(k_W - k_D) + \phi S_I k_I$
Capillary pressure model (van-Genuchten, 1980; Moridis et al., 2008)	$P_{cap} = -P_0 \left[(S^*)^{-1/\lambda} - 1 \right]^{1-\lambda} S^* = \frac{(S_A - S_{irA})}{(S_{mA} - S_{irA})}$
S_{irA}	0.24
$1/P_0$	0.6
3-Phase (Aqueous-Gas-Hydrate) Relative Permeability Model (Moridis et al., 2008)	$5.3 \times 10^{-5} \text{ Pa}^{-1}$ $k_{rA} = (S_A^*)^n$ $k_{rG} = (S_G^*)^n$ $S_A^* = (S_A - S_{irA}) / (1 - S_{irA})$ $S_G^* = (S_G - S_{irG}) / (1 - S_{irA})$ OPM model (Moridis et al., 2008)
n	3.0
S_{irG}	0.02
S_{irA}	0.25

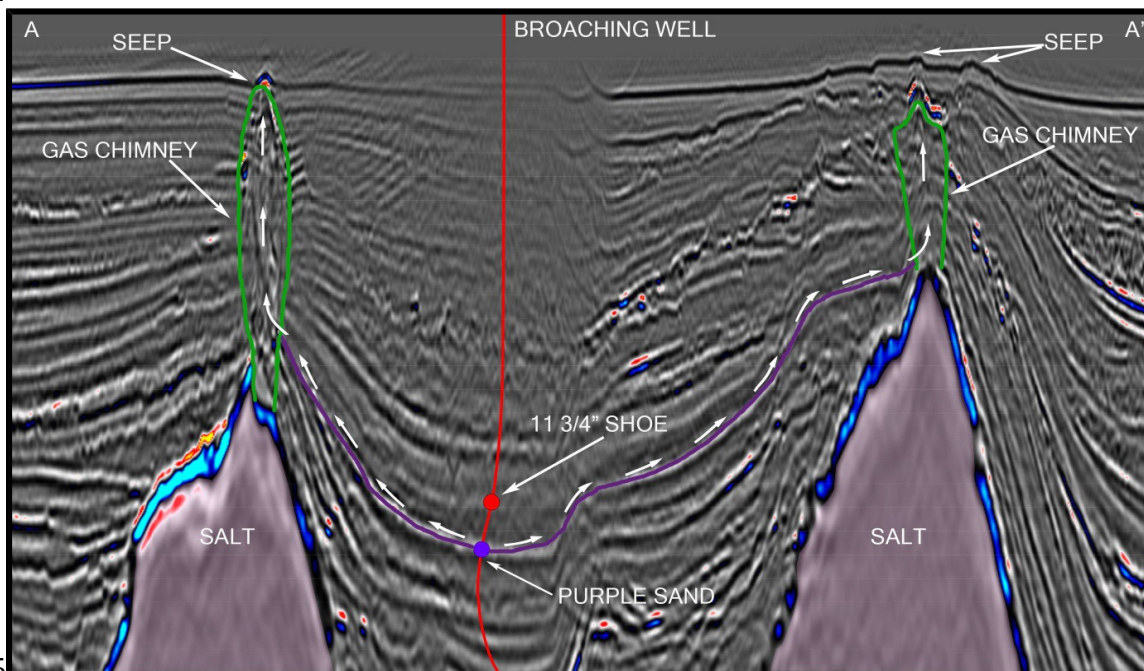
546(*) Moridis and Reagan, 2007.
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 550 Figure 1. Structure map on the purple sand showing relative depth subsurface) in grey-shade,
 551 gas chimney locations, and seismic traverse A-A'. The well surface location is designated by a
 552 square and the bottom-hole location is designated by a circle (proprietary data courtesy of
 553 Western Geco/Schlumberger).

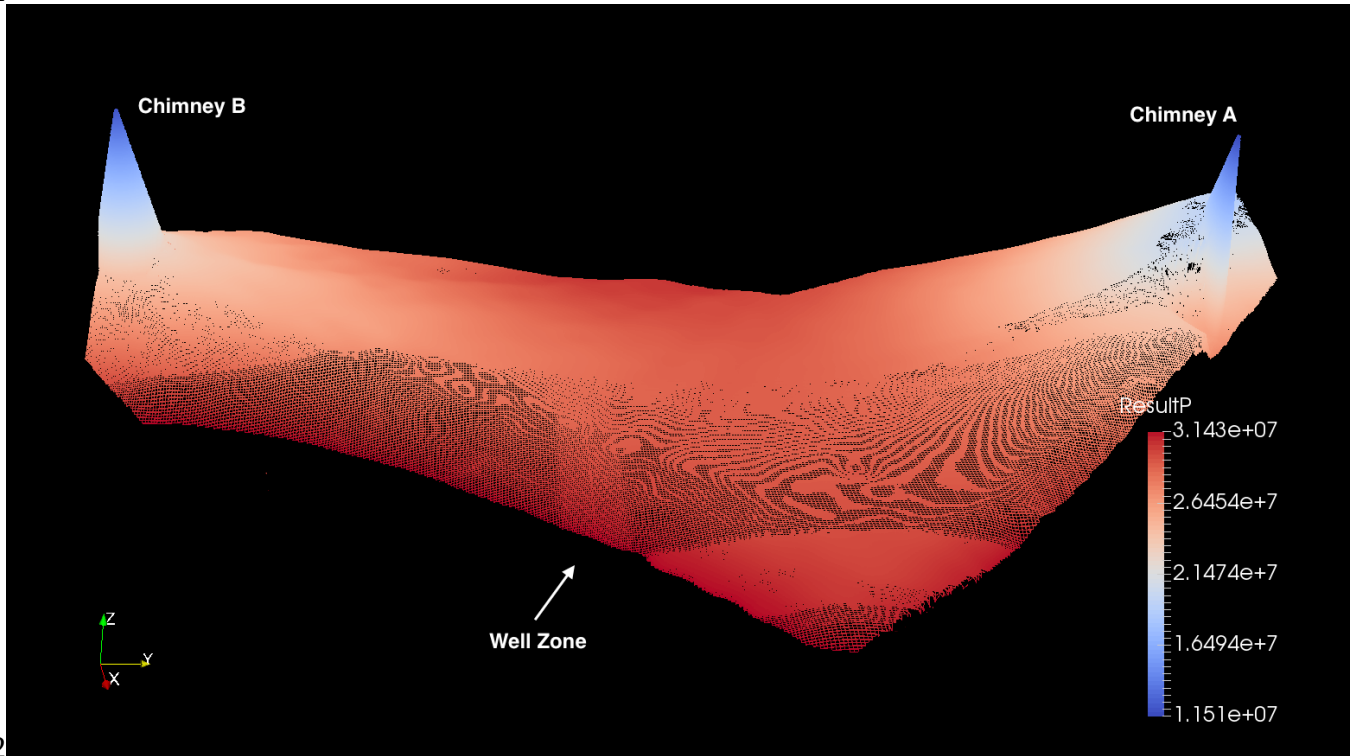
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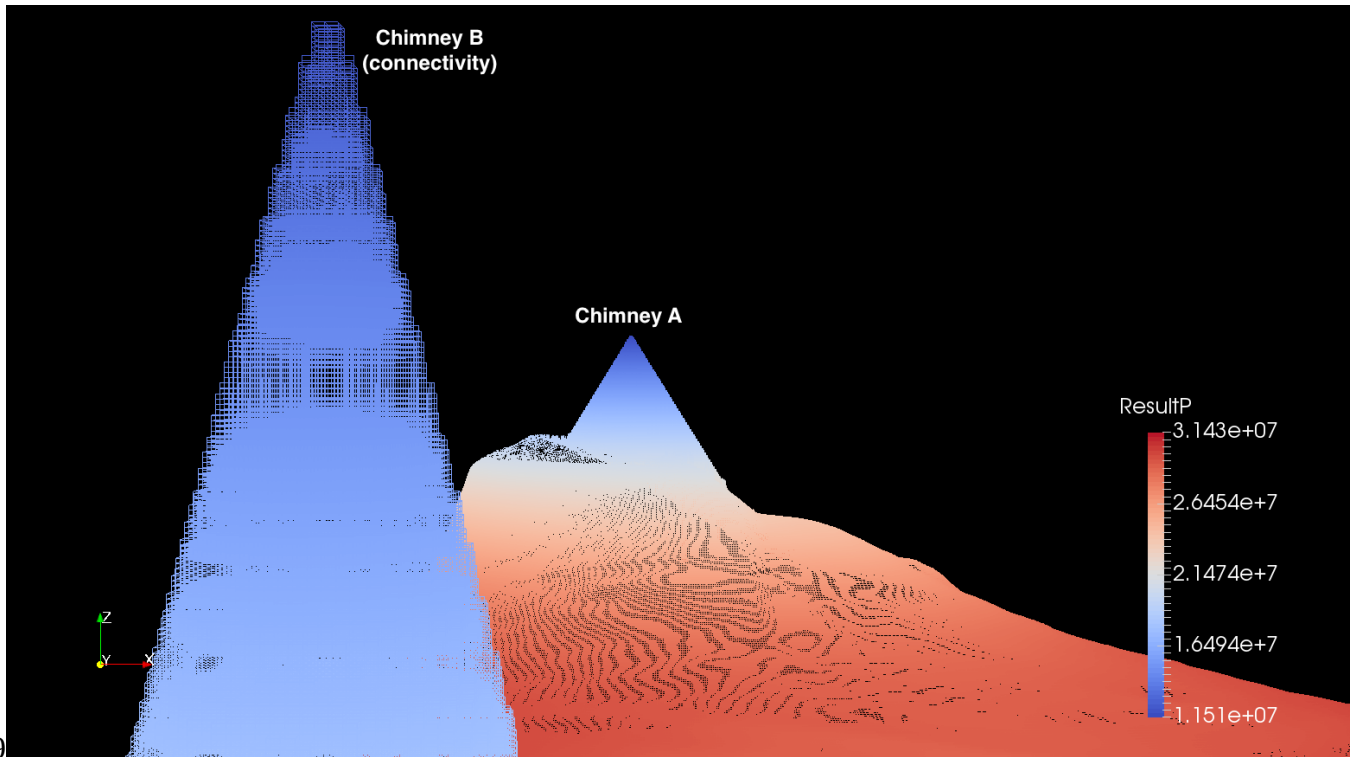
556 Figure 2. Purple sand seismic traverse A-A' that shows well path, 11.75-in casing shoe, purple
557 sand, broaching pathways, adjacent salt structures, natural seeps, and gas chimneys
558 (proprietary seismic courtesy of WesternGeco, LLC).
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Figure 3. Grid connectivity and initial pressure distribution in the entire domain. "Well zone" points toward the region of the sand intersected by the well, and thus the point at which fluids from a failed casing shoe would emerge into the formation. Impermeable sediments overlying the sands are not visualized.



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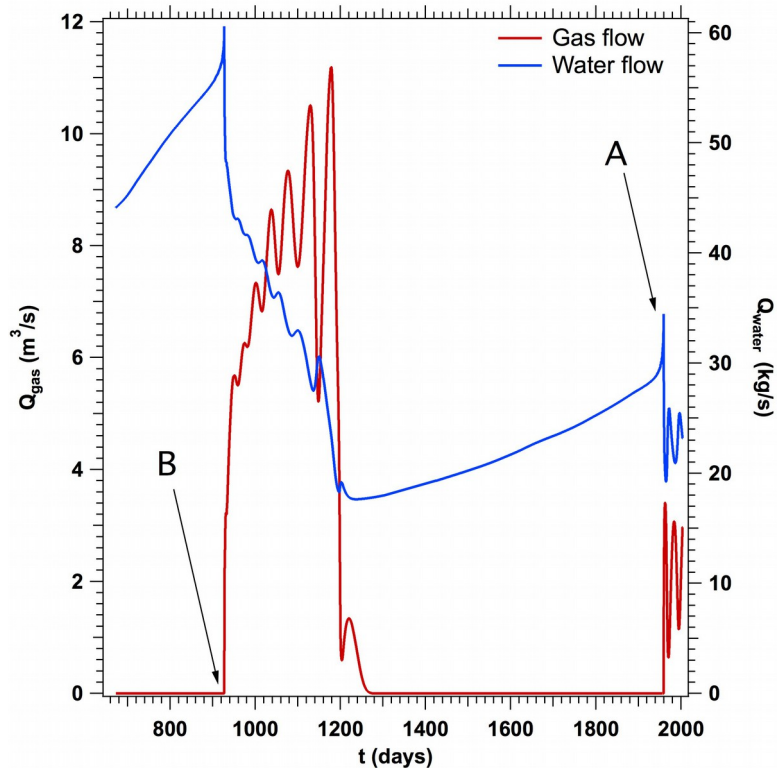
571 Figure 4. Connectivity and initial pressure distribution in the vicinity of Chimney B. Impermeable
 572 sediments overlying the sands are not visualized.

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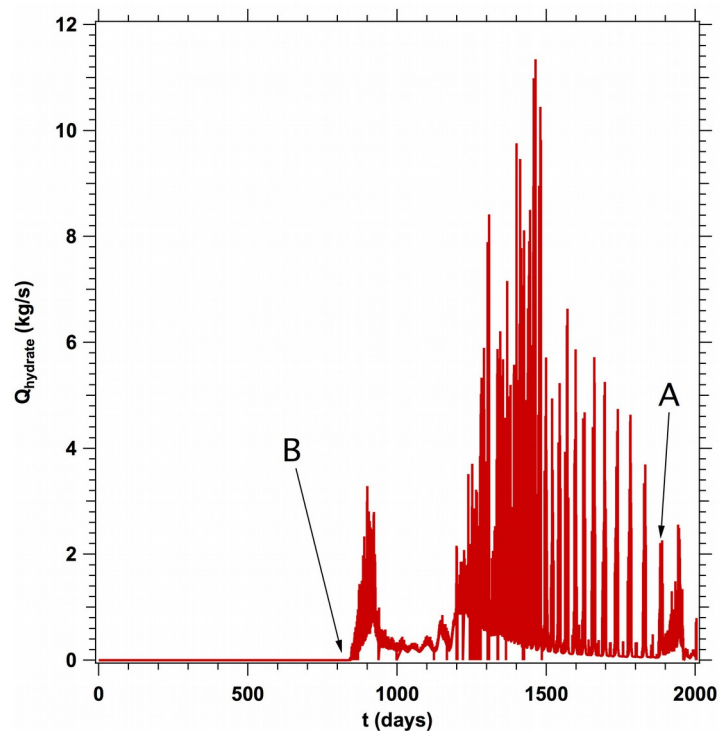
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579 Figure 5: Gas and water flow rates at the seafloor (Chimney A + Chimney B) for the all-gas
580 release case. Arrows indicate gas arrival times for Chimneys A and B.

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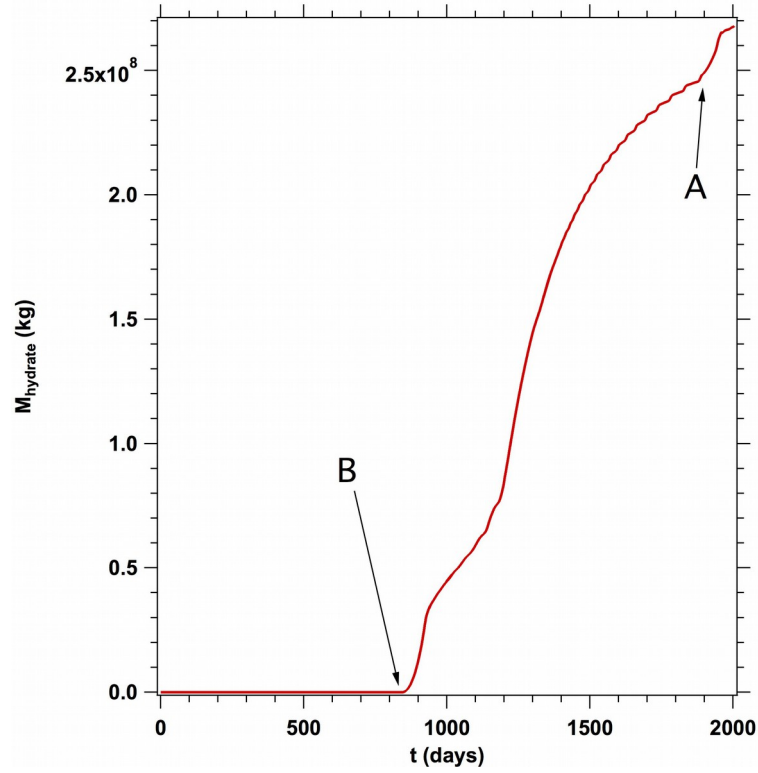


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585 Figure 6: Total hydrate formation rate, Q_H , for the overall system for the all-gas release case.

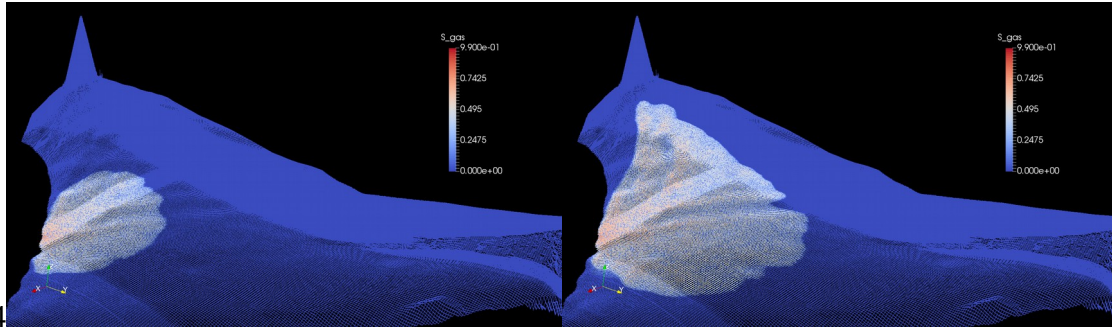
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589 Figure 7: Cumulative hydrate formation, M_H , for the overall system in the all-gas release case.

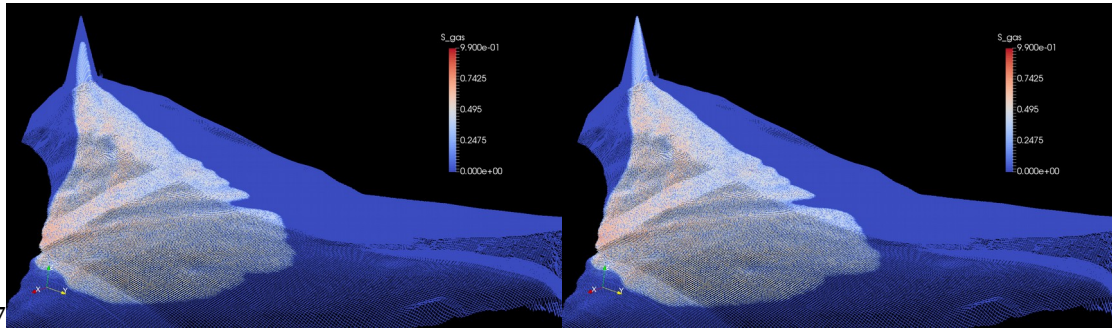
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t = 180 d

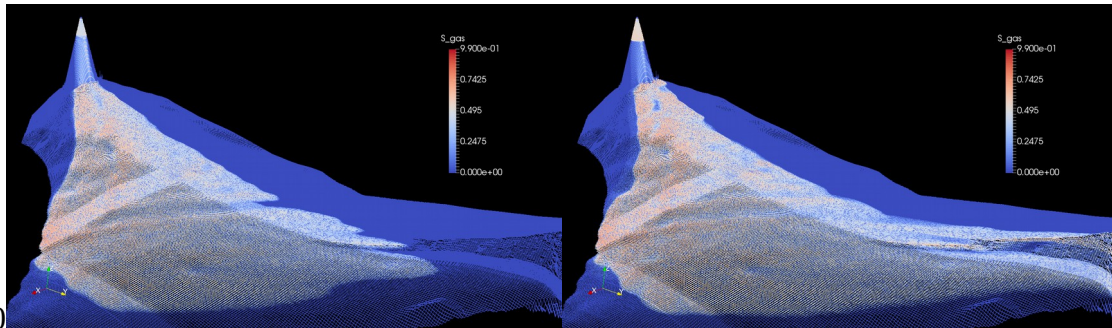
t = 640 d



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t = 870 d

t = 1,000 d



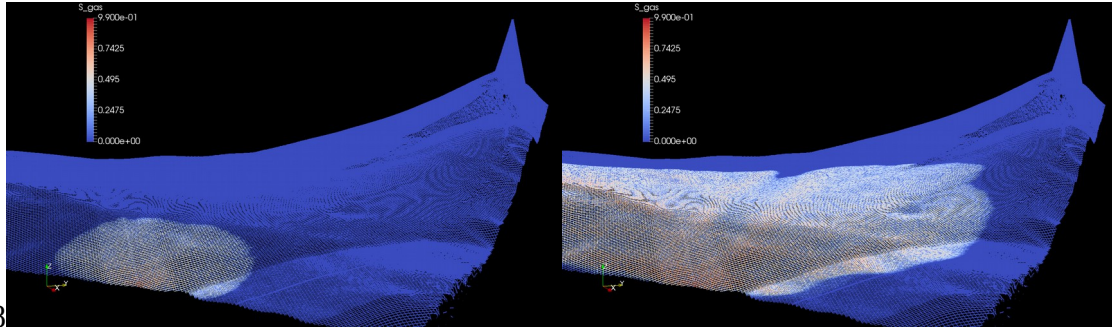
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t = 1,370 d

t = 2,000 d

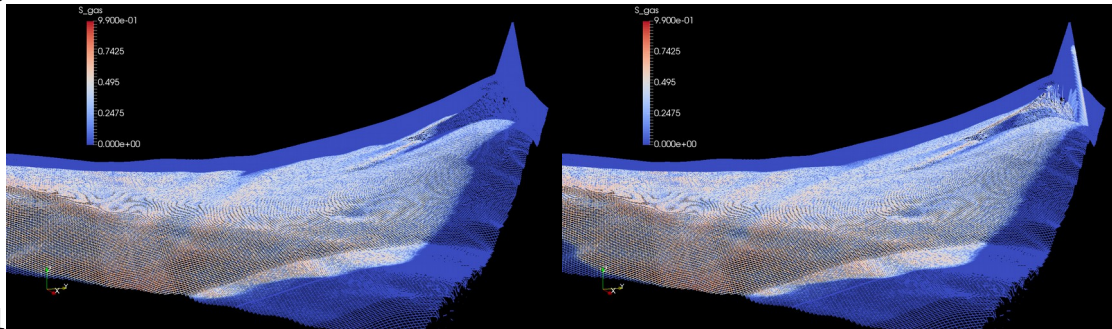
603 Figure 8: Visualizations of S_G , showing gas approach to, and release through, Chimney B.

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t = 60 d

t = 1,280 d



t = 1,730 d

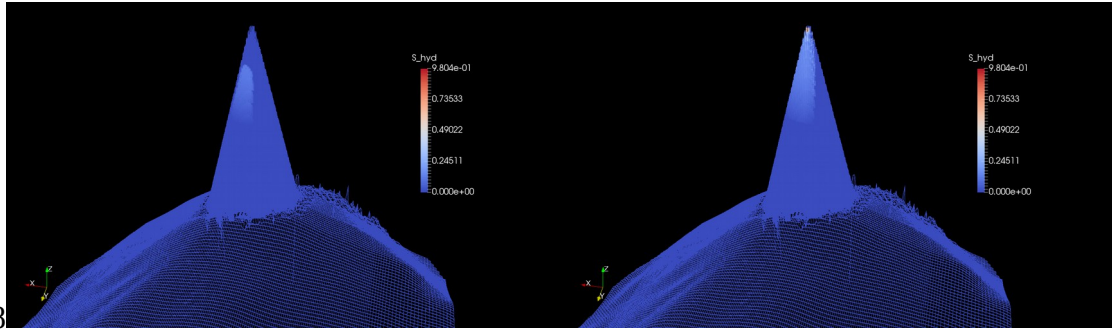
t = 1,910 d

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Figure 9: Visualizations of S_G , showing gas approach and release through Chimney A.

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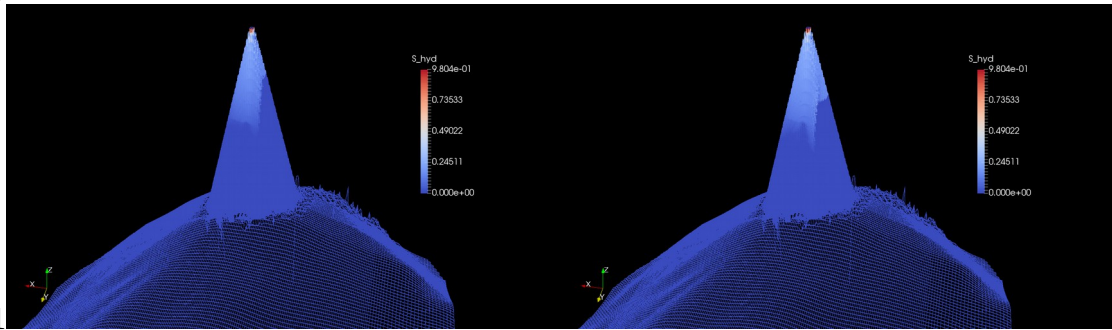


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t = 910 d

t = 1,090 d

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t = 1,280 d

t = 2,000 d

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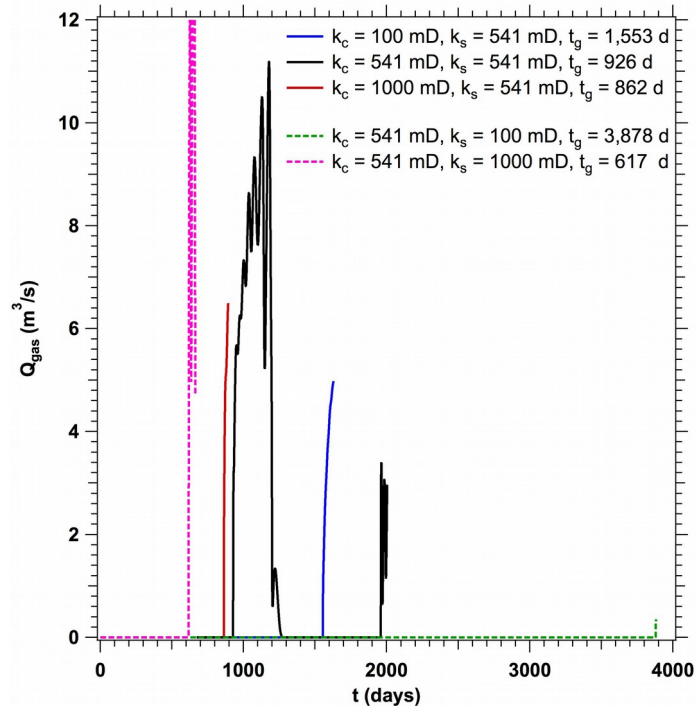
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624 Figure 10: Visualizations of S_H , showing hydrate formation in Chimney B.

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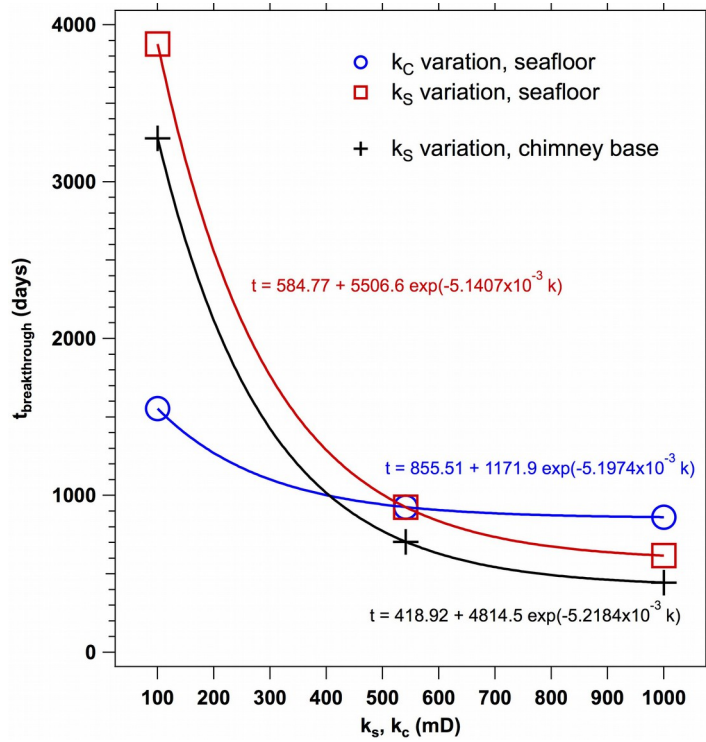
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630 Figure 11: Variations in breakthrough curves, Q_G vs. t , as a function of permeabilities.

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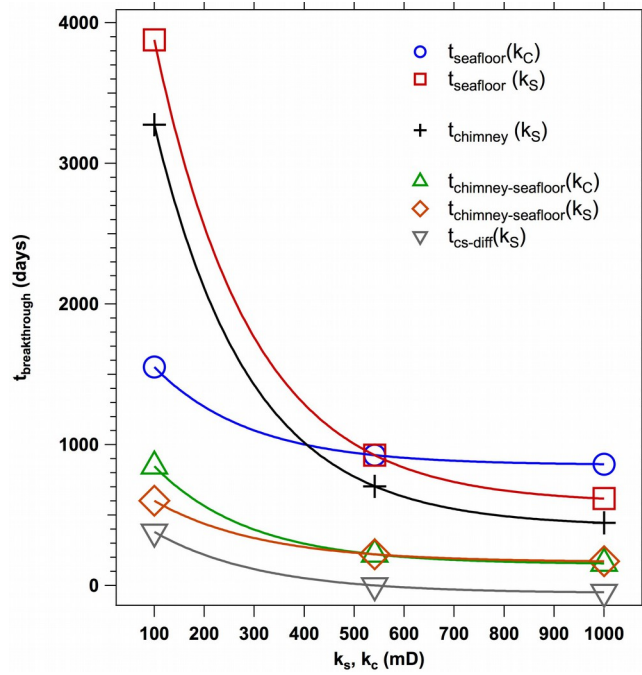
635 Figure 12: Gas arrival times at Chimney B and the seafloor as a function of permeabilities.

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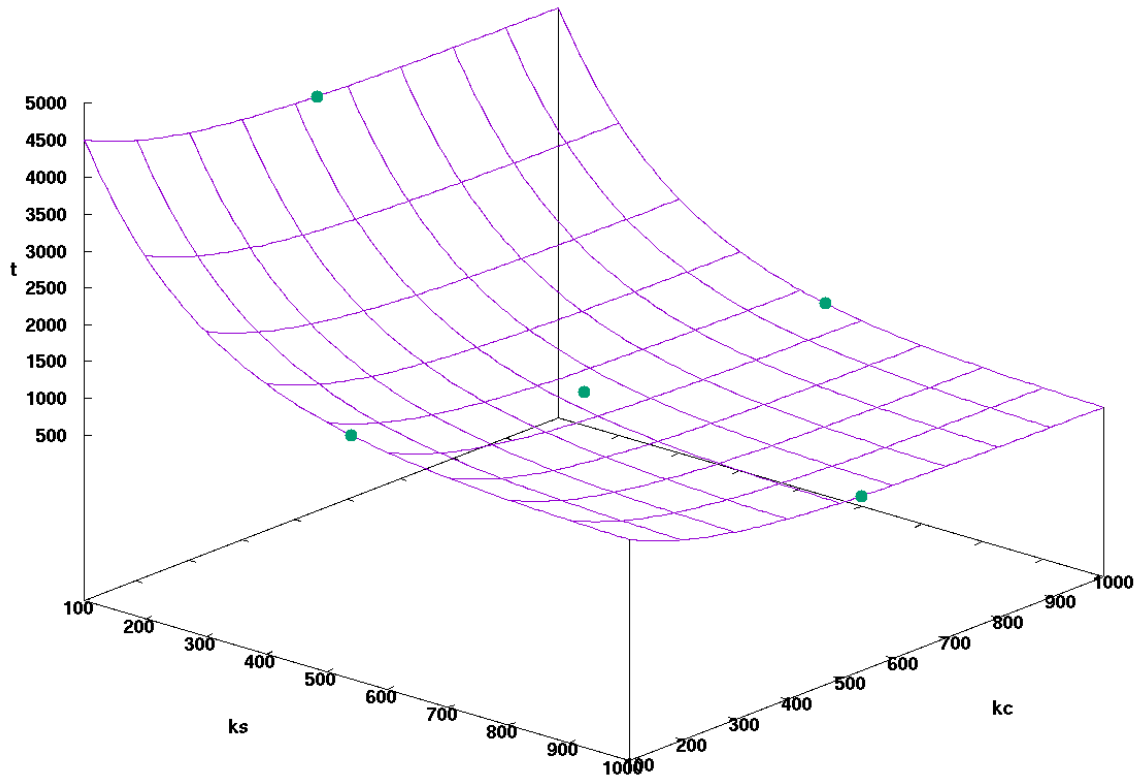
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641 Figure 13: Gas arrival times at Chimney B and the seafloor, transit times through the chimney
642 and k_s/k_c adjustment as a function of permeabilities.

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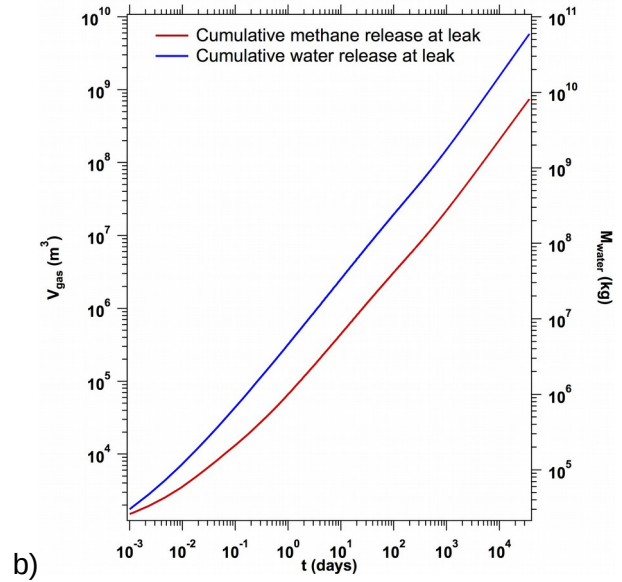
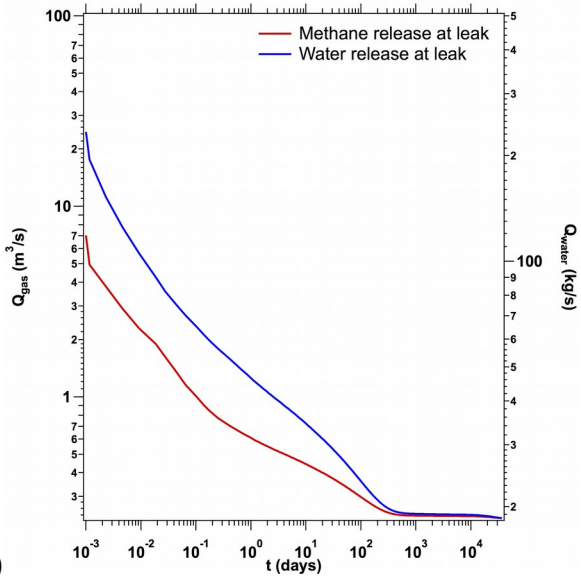
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648 Figure 14: Gas arrival time surface as a function of k_s and k_c .

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652 a)

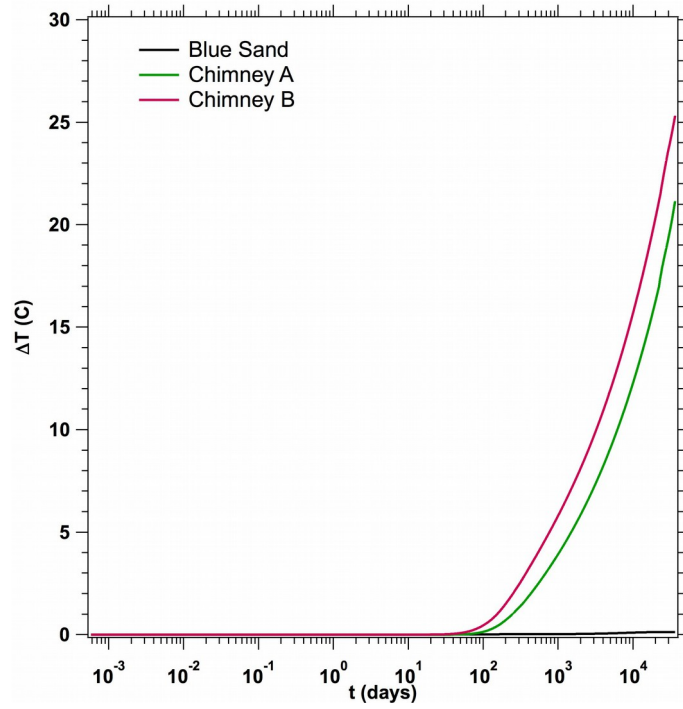
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654 Figure 15: Rates of water and gas released at the leak (a) and cumulative volume/mass (b).

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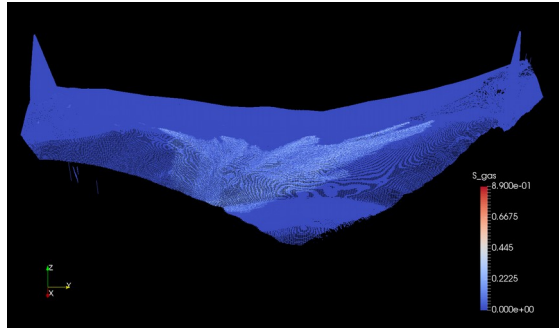
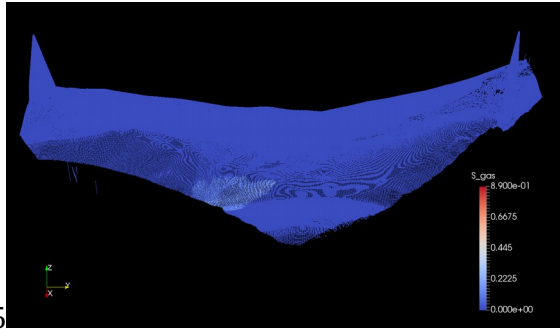
660 Figure 16: Change in temperature at monitoring points within the chimneys and in the Purple

661 Sand

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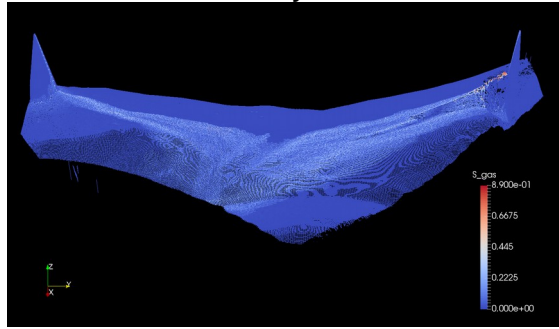
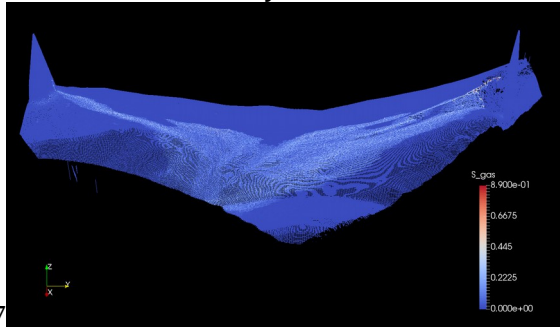


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t = 4 yr

t = 25 yr



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t = 41 yr

t = 70 yr

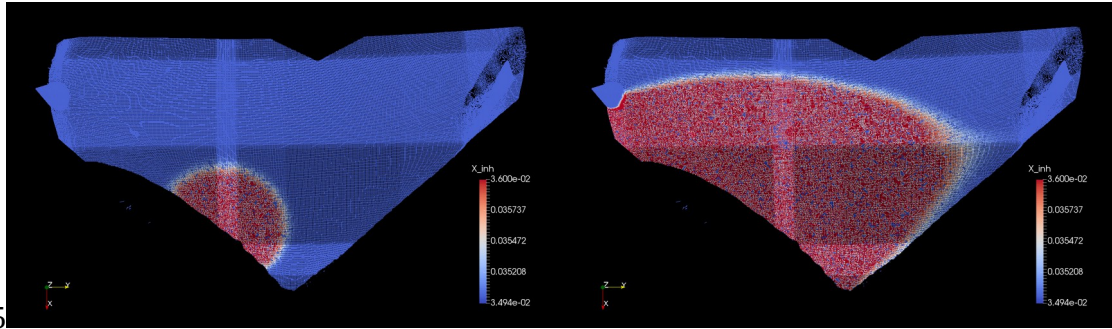
669 Figure 17: Visualizations of S_G , showing gas transport and escape through the chimneys for a
670 mixture of water and gas.

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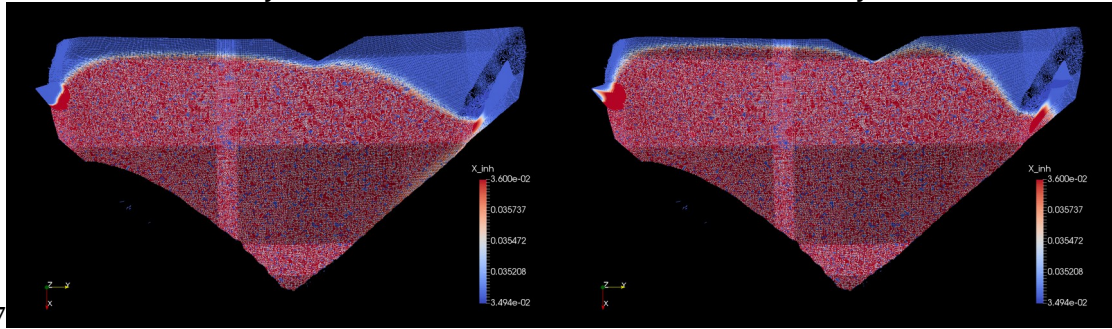


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t = 4 yr

t = 25 yr



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678

t = 41 yr

t = 70 yr

679 Figure 18: Visualizations of X_{NaCl} , showing brine transport and escape through the chimneys for
680 a mixture of water and gas.

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