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Emissions of black carbon, organic, and inorganic aerosols from biomass burning in North America and Asia in 2008

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[1] Reliable assessment of the impact of aerosols emitted from boreal forest fires on the Arctic climate necessitates improved understanding of emissions and the microphysical properties of carbonaceous (black carbon (BC) and organic aerosols (OA)) and inorganic aerosols. The size distributions of BC were measured by an SP2 based on the laser-induced incandescence technique on board the DC-8 aircraft during the NASA ARCTAS campaign. Aircraft sampling was made in fresh plumes strongly impacted by wildfires in North America (Canada and California) in summer 2008 and in those transported from Asia (Siberia in Russia and Kazakhstan) in spring 2008. We extracted biomass burning plumes using particle and tracer (CO, CH3CN, and CH2Cl2) data. OA constituted the dominant fraction of aerosols mass in the submicron range. The large majority of the emitted particles did not contain BC. We related the combustion phase of the fire as represented by the modified combustion efficiency (MCE) to the emission ratios between BC and other species. In particular, we derived the average emission ratios of BC/CO = 2.3 ± 2.2 and 8.5 ± 5.4 ng m⁻³/ppbv for BB in North America and Asia, respectively. The difference in the BC/CO emission ratios is likely due to the difference in MCE. The count median diameters and geometric standard deviations of the lognormal size distribution of BC in the BB plumes were 136–141 nm and 1.32–1.36, respectively, and depended little on MCE. These BC particles were thickly coated, with shell/core ratios of 1.3–1.6. These parameters can be used directly for improving model estimates of the impact of BB in the Arctic.


1. Introduction

[2] Boreal forest fires are one of the most important sources of aerosols transported to the Arctic [e.g., Stohl et al., 2006, 2007; Treffeisen et al., 2007; Eck et al., 2009]. Large amounts of aerosols, predominantly in the form of carbonaceous aerosols, namely black carbon (BC) and organic aerosols (OA), are emitted from biomass burning (BB). These particles strongly absorb and scatter solar visible radiation (downwelling and upwelling), influencing the radiation budget in the Arctic. They can also act as cloud condensation nuclei (CCN), which influence the microphysical properties of clouds [Lubin and Vogelmann, 2006]. Deposition of BC onto snow and ice changes the surface albedo [Warren and Wiscombe, 1980; Clarke and Noone, 1985], contributing to warming of the Arctic [Hansen and Nazarenko, 2004; Flanner et al., 2007, 2009]. Statistically, Russia and Canada have been the largest sources of BC emitted from boreal forest fires [Lavoué et al., 2000; Stocks et al., 2002; Conard et al., 2002; Soja et al., 2004]. Fire activity may increase in the course of significant warming occurring prominently at high latitudes [Stocks et al., 1998].
[3] Reliable estimates of the effects of BC on the climate of the Arctic by global models necessitate quantitative understanding of emissions of BC, aging and removal, and microphysical properties [e.g., Stier et al., 2006; Textor et al., 2006; Koch et al., 2009]. More specifically, calculations of the effects of light absorption by BC require photo-absorption cross sections and size distributions of BC. The cross section depends on diameter, and mixing state, together with the refractive index and shapes of BC [e.g., Moteki et al., 2010; Moteki and Kondo, 2010; Shiraiwa et al., 2010]. The size distributions of BC are controlled by emissions, transport, and deposition of BC. The coating of BC by nonrefractory species also influences the CCN activity of BC [e.g., Riemer et al., 2004; Oshima et al., 2009; Kuwata et al., 2009] and thus the efficiency of wet deposition, partly compensating for the amplification of light absorption [Stier et al., 2006]. The effect of BC on global and regional climate is currently poorly represented by models [e.g., Textor et al., 2006; Koch et al., 2009], owing in large part to the lack of constraints on emissions and the evolution of these parameters.

[4] Previous estimates of emission factors of BC are considered to be highly uncertain, because they are based on measurements with large uncertainties [Reid et al., 2005]. In particular, interference by coemitted organic aerosol (OA) can cause large errors in BC measurements by thermal optical transmittance and filter-based absorption photometers [e.g., Reid et al., 2005; Kondo et al., 2009, 2011]. Reliable measurements of BC in the Arctic in the spring of 2008 using a single particle soot photometer (SP2) were reported by Warneke et al. [2009, 2010]. They observed transport of enhanced BC emitted from BB in Siberia and Kazakhstan over the Alaskan Arctic together with BC-CO correlations.

[5] There is a critical need to characterize the emission ratios and microphysical properties of BC near BB sources before being transported over long distances. We also need to understand the evolution of the microphysical properties of BC after emission for improved representation of these parameters and processes by models. The major objectives of this study are to elucidate these key parameters and processes of BC using a well-calibrated SP2 based on laser-induced incandescence (LII) [e.g., Moteki and Kondo, 2007, 2008; Schwarz et al., 2006, 2008; Kondo et al., 2011].

[6] For this purpose, we made measurements of BC and light-scattering particles (LSP) in the submicron range by the SP2 on board the NASA DC-8 aircraft over Alaska and Canada during the Arctic Research of the Composition of the Troposphere from Aircraft and Satellite (ARCTAS) mission in the spring (ARCTAS-A) and summer (ARCTAS-B) of 2008 [Jacob et al., 2010]. Frequent aircraft samplings were made in plumes from boreal forest fires that occurred in Canada and California in summer and in Asia (Siberia and Kazakhstan) in spring. It is noted that in 2008, the boreal wildfires in southern Siberia and agricultural fires in Kazakhstan both occurred unusually early in spring and were much more intense than usual at that time of year [Warneke et al., 2009].

[7] In Canada, aerosol measurements in large plumes were made in the vicinity of the burning areas 1–2 h after the emissions. We have derived the emission ratios of aerosols from BB based on the measurements with unprecedented accuracy. We sampled emissions from wildfires emitted under varying combustion conditions, namely flaming and smoldering fire phases. We systematically interpret the dependence of the emission ratios on the conditions of combustion.

[8] The air masses impacted by BB in Asia were sampled 2–3 days after the emissions. We identified the locations of fires that influenced concentrations of BC by using kinematic back trajectories. We selected data little influenced by wet deposition during transport to estimate the emission ratios. These observations have enabled detailed comparisons of the aerosol emission ratios and their microphysical properties in the two regions.

[9] The most important parameters and processes presented by this study are the evolution of the size distributions and mixing state (coating thickness) of BC, the emission ratios of BC versus CO and CO₂, and their dependence on the conditions of combustion. We have also elucidated the difference of these parameters between the BB plumes from Asia and North America, which are the two major regions of boreal forest fires. These parameters can be used directly for improving model estimates of the impact of BB on climate, especially in the Arctic.

2. Aircraft Observations and Methodology of Data Analysis

2.1. Instrumentation

[10] The data used for this study are the mass, volume, and number concentrations of BC and LSP, mixing ratios of carbon dioxide (CO₂), carbon monoxide (CO), acetonitrile (CH₂CN), dichloromethane (CH₂Cl₂), sulfur dioxide (SO₂), dimethyl sulfide (DMS), and reactive nitrogen (NOₓ, NOy, PANs, and HNO₃), and submicron aerosol chemical composition. The instruments used for the present study and their measurement accuracies are summarized in Table 1, together with their references.

[11] The number concentrations of BC and LSP and the mixing state of BC (coating thickness of BC by nonrefractory components) were measured by the SP2. The measurement principle and schematic diagram of the SP2 have been described previously [Gao et al., 2007; Moteki and Kondo, 2007], and detailed descriptions of the calibration of the SP2 used for ARCTAS are given elsewhere [Moteki and Kondo, 2010; Kondo et al., 2011]. We summarize the important points below.

[12] Each BC particle sampled by the SP2 was irradiated by a laser beam at a wavelength of 1064 nm and thus heated to incandescence. The SP2 monitors LII signals in two distinct visible bands (λ = 300–550 nm and 580–710 nm). In addition, laser light scattered by individual particles was detected by two avalanche photodiodes to measure the sizes and number concentrations of LSP.

[13] BC particles can be identified by the ratio of LII intensities of the two visible bands, because the ratio is a proxy for the vaporization temperature of the particles. The vaporization temperature of ambient BC was measured to be about 4000 K [Schwarz et al., 2006]. The BC mass for each particle (mSP2) was estimated from the LII peak intensity. We measured the LII peak intensities of monodisperse BC aerosols that were mass selected by an aerosol particle mass analyzer (APM; Model-302, KANOMAX, Inc., Japan) with a 400°C heated inlet [Moteki and Kondo, 2007]. BC in
ambient air was extracted by vaporization of the condensed compounds coating the BC particles by means of the heated inlet. The SP2 calibration using ambient BC was made for several hours on 10 November 2009. The error in extracting refractory BC mass is estimated to be less than about 10% [Kondo et al., 2011]. This extraction procedure means that the definition of the mass of ambient BC is the mass of incandescent aerosols refractory at 400°C (mref*). The uncertainty of mref* is about 10%.

[14] The LII peak intensity increased linearly with mass up to about 20 femtograms (fg). At higher mass, the relationships deviated from linearity. We used the LII peak intensity–mref* relationship for ambient BC to derive the BC size distributions. In this way, mSP2 is fully traceable to the calibration standards (mref*).

[15] The mass equivalent diameters of the BC cores (DpBC) were calculated by assuming a density (ρBC) of 2 g cm−3, which is also close to the value of 1.77 g cm−3 derived by Park et al. [2004]. About 10% uncertainty in ρBC led to errors of a few % in DpBC. The DpBC detection range of the SP2 used for this study was 80–860 nm. The absolute accuracy of the measurements of mass concentration of BC is estimated to be about 10% [Kondo et al., 2011].

[16] The diameters of the coated BC particles (shell diameters, Dp) were derived using the scattering signals for Dp. The coating thickness of BC or shell/core ratio (Dp/DpBC ratio) is a useful conceptual parameter to represent the mixing state of BC. A detailed description of the method to derive Dp is given by Kondo et al. [2011].

[17] LSP scatters laser light but does not incandesce. The peak scattering signal was converted to the scattering cross section of the particles. Calibration was made first by using polystyrene latex (PSL) particles with a refractive index of 1.59 + 0i. The calibration curve of PSL converts to that of (NH4)2SO4 at a refractive index of 1.52 + 0i based on the Mie theory algorithm developed by Bohren and Huffman [1998]. The diameters of LSP (Dp) were estimated from this curve with an uncertainty of about 10% in the size range of about 200–750 nm [Moteki and Kondo, 2008]. We used (NH4)2SO4 in defining Dp because it generally constitutes an important fraction of PM1 and its refractive index is well known. Refractive indices of OA emitted by BB were measured to be in the range of about 1.56–1.64 at visible wavelengths [Levin et al., 2010]. Refractive indices of SOA photochemically generated from α- and β-pinene and toluene were measured to range from 1.4 to 1.5 at λ = 670 nm [Kim et al., 2010].

[18] Figure 1 shows the calculated scattering cross sections of particles with refractive indices of 1.5 ± 0.1 at a wavelength of 1064 nm based on Mie theory. The changes in particle diameters for changes in the refractive index of ±0.1 from 1.5 + 0i for the same scattering cross sections in the diameter range of 200–500 nm is within 10%, as summarized in Table 2. The uncertainty in Dp by possible changes in the refractive indices of LSP is estimated to be about 10%.

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**Table 1. Measurements During ARCTAS**

<table>
<thead>
<tr>
<th>Measurement Technique</th>
<th>Accuracy</th>
<th>References</th>
</tr>
</thead>
<tbody>
<tr>
<td>BC</td>
<td>±10%</td>
<td>Moteki and Kondo [2007, 2010]</td>
</tr>
<tr>
<td>LSP number distribution</td>
<td>±10%</td>
<td>Moteki and Kondo [2007, 2010]</td>
</tr>
<tr>
<td>LSP volume distribution*</td>
<td>±30%</td>
<td>DeCarlo et al. [2006, 2008]</td>
</tr>
<tr>
<td>Aerosol chemistry</td>
<td>±35%</td>
<td>and Dunlea et al. [2009]</td>
</tr>
<tr>
<td>CO</td>
<td>TDLAS</td>
<td>±2%</td>
</tr>
<tr>
<td>CO2</td>
<td>NDIR gas analyzer</td>
<td>±0.25 ppmv</td>
</tr>
<tr>
<td>CH2Cl2</td>
<td>PTR-MS</td>
<td>±10%</td>
</tr>
<tr>
<td>CH3CN</td>
<td>WAS-GC</td>
<td>–</td>
</tr>
<tr>
<td>DMS</td>
<td>WAS-GC</td>
<td>–</td>
</tr>
<tr>
<td>SO2 (CIT)</td>
<td>CIMS</td>
<td>±50%</td>
</tr>
<tr>
<td>SO2 (GIT)</td>
<td>CIMS</td>
<td>–</td>
</tr>
<tr>
<td>NOx and NOy Chemiluminescence</td>
<td>±10% and ±15%</td>
<td>Weinheimer et al. [1994]</td>
</tr>
<tr>
<td>PANS</td>
<td>CIMS</td>
<td>–</td>
</tr>
<tr>
<td>HNO3</td>
<td>CIMS</td>
<td>–</td>
</tr>
<tr>
<td>OH</td>
<td>LIF</td>
<td>±40%</td>
</tr>
</tbody>
</table>

*aEstimated by uncertainty propagation from the number distribution.

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**Figure 1.** Calculated scattering cross sections as a function of particle diameter for different refractive indices at a wavelength of 1064 nm.

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**Table 2. Comparison of Particle Diameter Dp (N) in Units of nm Estimated for Different Refractive Indices (N) and Their Ratios**

<table>
<thead>
<tr>
<th>Dp (1.5)</th>
<th>Dp (1.4)</th>
<th>Dp (1.6)</th>
<th>Dp (1.4)/Dp (1.5)</th>
<th>Dp (1.6)/Dp (1.5)</th>
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<tr>
<td>200</td>
<td>214</td>
<td>189</td>
<td>1.070</td>
<td>0.945</td>
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<tr>
<td>250</td>
<td>269</td>
<td>236</td>
<td>1.076</td>
<td>0.944</td>
</tr>
<tr>
<td>300</td>
<td>324</td>
<td>283</td>
<td>1.080</td>
<td>0.943</td>
</tr>
<tr>
<td>350</td>
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<td>328</td>
<td>1.086</td>
<td>0.937</td>
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<tr>
<td>400</td>
<td>437</td>
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<td>0.935</td>
</tr>
<tr>
<td>450</td>
<td>494</td>
<td>418</td>
<td>1.098</td>
<td>0.929</td>
</tr>
<tr>
<td>500</td>
<td>551</td>
<td>463</td>
<td>1.102</td>
<td>0.926</td>
</tr>
</tbody>
</table>

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considering the reported refractive indices of inorganic and organic aerosols. The total mass concentrations of BC ($M_{BC}$) and volume concentrations of LSP ($V_{SC}$) were derived by integrating the size distributions with respect to diameter. The propagated uncertainty of $V_{SC}$ from the uncertainty of $D_p$ is estimated to be about 30%. We did not convert $V_{SC}$ to the mass concentration of LSP because the densities of LSP are not provided by the SP2 measurements. The values of $M_{BC}$ and $V_{SC}$ are given in the units of mass and volume per unit volume of air at standard temperature and pressure (STP; 273.15 K and 1013 hPa), respectively.

Sample air was provided to the instruments using a forward-facing, window-mounted inlet probe. Real-time measurements of air speed, static pressure, and temperature were used to maintain isokinetic flow through the inlet tip to minimize inertial enhancement or depletion of coarse particles. Detailed measurements made during the DC-8 Inlet Characterization Experiment (DICE) in the spring of 2003 showed that the inlet efficiently transmits both dust and sea salt particles smaller than 4 μm in dry diameter [McNaughton et al., 2007].

2.2. Air Sampling of Fire Plumes

Figure 2 shows the flight tracks of the NASA DC-8 aircraft, flown from Cold Lake (54°N, 110°W), Alberta, Canada, between 29 June and 10 July 2008, and the transit flight to Palmdale, CA on 13 July. On 29 June, 102 fires were burning and more than half of those ignitions resulted from lightning strikes that occurred within the previous 24 h (A. J. Soja et al., ARCTAS: The perfect smoke, in The Canadian Smoke Newsletter, pp. 2–7, Fall 2008). Most of the plumes were encountered over Saskatchewan Province, while some of them were measured over the Yukon Territory and Manitoba. A substantial portion of the flight time was devoted to sampling forest fire plumes within a few hours after emission mostly in the boundary layer (BL) and the lower free troposphere (FT). Figure 2 shows the locations of the sampling of the BB plumes used for the present study that includes data obtained in forest fire plumes in California during the transit flight.

Figure 2. Flight track of the DC-8 over Canada during ARCTAS-B. Phases of BB of the air masses sampled are shown with circles.

[21] Large-scale BB occurred in Siberia (boreal forest fires) in Russia and in Kazakhstan (agricultural fires) in March and April 2008 [Warneke et al., 2009, 2010]. The air masses impacted by these fires were transported from these regions and were sampled by the local flights from Fairbanks, Alaska, in April 2008 during ARCTAS-A. The flight tracks of the DC-8 during ARCTAS-A are shown by Jacob et al. [2010]. These data are used for comparison with the ARCTAS-B data.

[22] Satellite-derived column amounts of CO from the Measurements of Pollution in the Troposphere (MOPITT) instrument [Drummond and Mand, 1996; Bowman, 2006] and aerosol optical depth (AOD) from the Moderate resolution Imaging Spectroradiometer (MODIS) [Remer et al., 2005] during ARCTAS-A were similar to monthly averages for individual months over the Alaskan Arctic (the region including both measurements and transport pathways of measured air parcels) [Matsui et al., 2011]. This suggests that ARCTAS-A data are representative of the North American Arctic in April 2008.

[23] Air masses impacted by forest fires that occurred in California were also sampled by the DC-8 ARCTAS instrument payload on 18–24 June 2008 for the California Air Resources Board (CARB) [Singh et al., 2010]. We used relevant data obtained during ARCTAS-CARB for comparison with the ARCTAS-B data.

2.3. Backward Trajectories

[24] We used 10 day backward trajectories calculated every minute along the flight tracks using a kinematic model [Fueldberg et al., 2010, and references therein]. The meteorological data were from runs of the Weather Research and Forecasting (WRF) model with a horizontal resolution of 45 km and 50 vertical layers [Skamarock et al., 2008]. Source regions were identified as the regions (projected onto the surface) where the backward trajectories encountered and remained in altitudes below 700 hPa (about 3 km above sea level), which is the approximate altitude of the upper boundary of the BL over the land during daytime.

[25] Air masses with origins outside Canada were excluded for the analysis of the ARCTAS-B data, which focuses on near-field emissions. Some of the air masses sampled during ARCTAS-A were estimated to have originated from two different biomass burning regions in Asia from 10 day back trajectories.

[26] We calculated the accumulated precipitation along each trajectory (APT) using Global Precipitation Climatology Project (GPCP) global precipitation data [Huffman et al., 2001; Adler et al., 2003]. We used daily data (version 1.1) with a horizontal resolution of 1 x 1 degree. The GPCP data were derived solely from observations based on gauge measurements and estimates of rainfall by geostationary and polar orbiting satellites. Precipitation at the surface (i.e., vertically integrated) was summed up along the trajectories from the sampling location to the identified source regions. We used the APT data as a measure of wet removal of water-soluble species during transport. The uncertainties in the estimates of source regions and APT were calculated by using two meteorological data sets for trajectory calculations and three precipitation data sets, including the GPCP data. The comparison has shown that the general features of the statistics and conclusions obtained in this study do
not change with the choice of trajectory model and meteorological and precipitation data sets. More detailed descriptions of the uncertainties in the estimation of the source regions and APT are given elsewhere [Matsui et al., 2011].

2.4. Extraction of Fire Plumes

[27] We first determined species background concentrations to quantify the effect of BB on the concentrations of the species measured on board the aircraft. They were determined as the 5th percentiles of the values in every 10 K range of potential temperature, and they were linearly interpolated vertically. The background values of most of the species were close to zero, except for CO2 and CO. The differences between the measured and background concentrations (Δ values) were used for the analysis. The 1 min averaged data with ΔCO > 20 parts per billion by volume (ppbv) were used for the statistical analysis to obtain emission ratios from individual data points with sufficient precision maintaining a sufficient number of the data points for statistical analysis.

[28] We selected data strongly influenced by BB using the concentrations of trace hydrocarbon species, CH3CN and CH2Cl2, as described in detail by Matsui et al. [2011]. CH3CN is known to be emitted from forest fires [de Gouw et al., 2003]. CH2Cl2 is emitted from industrial solvents, but not emitted from combustion of fossil fuels [Chen et al., 2007; Choi et al., 2003]. However, it can be used as a tracer of anthropogenic emissions of aerosols because the emissions of solvent (CH2Cl2) generally coexist with emissions of species by fossil fuel combustion. The BB-impacted air masses transported from Asia were defined as those with ΔCH3CN > 50 parts per trillion by volume (pptv) and ΔCH2Cl2 < 5–10 pptv. The threshold of ΔCH2Cl2 increased from 5 pptv at ΔCH3CN = 50 pptv to 10 pptv at ΔCH3CN > 100 pptv, as depicted in Figure 3 of Matsui et al. [2011]. For the Canadian BB, the data selection was made using the thresholds of ΔCH3CN > 100 pptv and ΔCH2Cl2 < 10 pptv.

2.5. Modified Combustion Efficiency

[29] The phase of BB has been shown to be represented by the combustion efficiency (CE) or modified combustion efficiency (MCE) [e.g., Delmas et al., 1995; Yokelson et al., 1999, 2008]. The calculation of CE requires measurement of incompletely oxidized compounds (CO) and reduced compounds such as methane, nonmethane hydrocarbons, and ammonia, among others [Ward et al., 1991], while MCE is estimated only from the ΔCO/ΔCO2 ratio.

\[
\text{MCE} = \frac{\Delta \text{CO}}{\Delta \text{CO2} + \Delta \text{CO}} = 1/(1 + \Delta \text{CO}/\Delta \text{CO2}).
\] (1)

MCE is useful in interpreting observed emission ratios as a function of MCE, irrespective of the relative amount of data in different combustion phases [Yokelson et al., 1999; Goode et al., 2000]. A MCE of 0.9 is sometimes treated as the value classifying air masses predominantly influenced by combustion in the flaming phase (MCE > 0.9) whereas, MCE < 0.9 represents the smoldering phase [Reid et al., 2005]. The 1 s time resolution of CO and CO2 data led to the correspondingly high time resolution of the MCE values. This parameter was used for the interpretation of the data impacted by BB.

3. Aerosols in BB Plumes in Canada and California During ARCTAS-B

[30] In Figure 3 (top), we show the time periods of sampling BB plumes identified by the methods discussed in sections 2.3 and 2.4. Figures 3 (top) and 3 (middle) show the time series plots of CO, MBC, VSC, MCE, and NOx/NOy ratio measured during Flight 18 (F#18) on 1 July. The APT and altitude of the DC-8 are shown in Figure 3 (bottom). The MBC and VSC values were greatly enhanced during encounters of wildfire plumes. In Figures 3 (top) and 3 (bottom), we show the time periods of sampling BB plumes identified by the methods discussed in sections 2.3 and 2.4. Generally, the MCE was lower than 0.95 in the relatively fresh plumes (NOx/NOy > 0.4) sampled during this flight.

[31] Figure 4 shows the ΔMBC/ΔCO correlation versus the estimated APT that the air masses experienced within 24 h prior to sampling. The integration time in calculating the APT was fixed at 24 h, independent of the age of the BB plumes. Most of the Canadian and Californian air masses experienced little precipitation (APT < 3 mm). For the analysis of the ARCTAS-B and CARB data, we selected data with APT < 1 mm to minimize the effect of wet deposition on the analysis of aerosols.

[32] The first panel in Figure 5 (left) shows the ΔCO-ΔCO2 correlations used to derive the MCE for individual plumes. Depending on the ΔCO/ΔCO2 ratios, the phases of combustion of these plumes were classified as smoldering, flaming, and others. It should be noted that the Δ values were used for the statistical analysis throughout this study, and the linear regressions are used only for supplementary analysis. The locations of the measurements of the plumes with the different combustion phases are shown in Figure 2. Smoldering phase plumes were measured mainly on 1 July during F#18 (Figure 3), and flaming phase plumes were measured on 4 July (F#19) and 13 July (F#24 in California). The flaming phase data obtained in California are shown with crosses. No smoldering phase data were obtained in California during ARCTAS-B. The correlations of the flaming phase data measured in Canada and California showed no statistical differences, as discussed below. Considering this, the data obtained in these regions were grouped as North American for statistical analysis. Figure 6 shows the profiles of MBC and VSC measured over North America during ARCTAS-B. The data points are color coded by the phase of BB. The color-coded data are also with APT < 1 mm.

[33] Figure 7 shows the average number and mass (or volume) size distributions of BC and LSP observed in layers impacted by wildfires over North America. A small discontinuity in the BC mass size distributions near 200 nm is a numerical artifact in constructing the size distributions from the data in each size bin. The count median diameters (CMD) of the lognormal fitted BC size distributions were unvarying at 136 ± 6 nm, with a geometric standard deviation (σg) of 1.32, although the CMD slightly increased with the increase in MCE, as discussed in section 8. The size distributions of the LSP were also unvarying, and their
average CMD and $\sigma_{bc}$ are summarized in Table 3. The MMD of BC and volume median diameters (VMD) of LSP were 187 ± 10 nm and 294 ± 42 nm, respectively. Table 3 summarizes these values, together with the geometric standard deviations ($\sigma_{gm}$ or $\sigma_{gv}$) of these size distributions. The average fractions of the unmeasured $M_{BC}$ and $V_{SC}$ values ($D_{BC} < 80$ nm and $D_{scat} < 185$ nm) were estimated to be 4% and 18%, respectively.

[34] We also derived the median shell/core ratios ($D_p/D_{BC}$) of BC for $D_{BC} > 200$ nm every 1 min. The median value of the $D_p/D_{BC}$ was 1.48 (1.42 for smoldering phase and 1.62 for flaming phase (Table 3)). Possible causes of the dependence of $D_p/D_{BC}$ on MCE are discussed in section 8. The estimated value using the above median of $V_p/V_{BC}$ ($=D_p/D_{BC}$) is 3.24 (2.86 for smoldering phase and 4.25 for flaming phase). Thus BC represents about 0.31 (=1/3.24) of the volume (ranging from 0.23 to 0.35 for flaming and smoldering, respectively) of the particles that contain BC. We made more detailed analysis of the dependence of the microphysical parameters of BC on MCE in section 8.

![Figure 3](image1.png)

**Figure 3.** (top) Time series plots of CO, BC mass concentration ($M_{BC}$), volume concentration of LSP ($V_{SC}$), and identification of BB plumes. (middle) MCE and NO$_x$/NO$_y$ ratio. (bottom) Precipitation (APT) and altitude of the DC-8 on 1–2 July.

![Figure 4](image2.png)

**Figure 4.** $M_{BC}/\Delta CO$ ratios plotted versus APT during ARCTAS-B. The APT was calculated for 24 h prior to aircraft sampling.
Figure 5. Correlations between $M_{BC}$, $V_{SC}$, $\Delta CO_2$, $\Delta CO$, and CH$_3$CN in wildfire plumes over Canada (circles) and California (crosses) during ARCTAS-B with APT $< 1$ mm. The data points are color coded for smoldering phase ($MCE < 0.9$; blue), flaming phase ($MCE > 0.95$; red), and others ($0.90 < MCE < 0.95$; green).

[35] Figure 8 and Tables 4 and 5 show the average mass concentrations and mass fractions of submicron non-refractory chemical components and BC in the extracted BB plumes as measured by the High-Resolution Time-of-Flight Aerosol Mass Spectrometer (HR-ToF-AMS) [DeCarlo et al., 2006; Dunlea et al., 2009] and SP2 on board the DC-8. The mass concentrations are in the units of $\mu g$ STP m$^{-3}$. The average $V_{SC}/M_{BC}$ and $V_{SC}/V_{BC}$ ratios in the North American
plumes are summarized in Table 6. The uncertainties are given as the errors in the determinations of the slopes of the correlations shown in Figures 5 and 11. The prescribed measurement errors are not included in this uncertainty estimate.

[36] On average, OA constituted about 83% of the total mass concentration of BC and nonrefractory aerosol with a \( D_{\text{va}} \) (vacuum aerodynamic diameter [see DeCarlo et al., 2004]) smaller than about 1 \( \mu \text{m} \) (PM1). We therefore infer that \( V_{\text{SC}} \) in the plumes approximately represents the volume concentration of OA. \( M_{\text{BC}} \) was converted to BC volume concentration (\( V_{\text{BC}} \)) assuming \( r_{\text{BC}} = 2 \text{g cm}^{-3} \). The average \( V_{\text{SC}}/M_{\text{BC}} \) and \( V_{\text{SC}}/V_{\text{BC}} \) ratios in the PM1 range are summarized in Table 6. The average volume fraction of BC was about 0.02. Together with the result above that the volume fraction of BC was 0.31 for particles containing BC, this indicates that the large majority of the scattering particles measured by SP2 in the size range of 200 nm < \( D_p < 750 \text{ nm} \) did not contain BC with \( D_{\text{BC}} > 80 \text{ nm} \), which is consistent with the particle number observations with the SP2.

Figure 6. Vertical profiles of (left) \( M_{\text{BC}} \) and (right) \( V_{\text{SC}} \) for all data points measured over Canada during the ARCTAS-B campaign. The mass concentration of aerosols with density of 1 g cm\(^{-3}\) is 1 \( \mu \text{g m}^{-3} \) for a volume concentration of 1 \( \mu \text{m}^3 \text{ cm}^{-3} \).

Figure 7. Mean distributions for (top) BC in number and mass space and (bottom) LSP in number and volume space, averaged for eight major forest fire plumes encountered in Canada during ARCTAS-B and those in BB fire plumes transported from Asia. Solid fitted lines are lognormal distribution functions.
Table 3. Size Distributions of BC and LSP

<table>
<thead>
<tr>
<th></th>
<th>Asia</th>
<th>Canada</th>
</tr>
</thead>
<tbody>
<tr>
<td>CMD</td>
<td>141 ± 8 nm</td>
<td>136 ± 6 nm</td>
</tr>
<tr>
<td>σ_ge</td>
<td>1.37 ± 0.04</td>
<td>1.32 ± 0.01</td>
</tr>
<tr>
<td>MMD</td>
<td>207 ± 31 nm</td>
<td>187 ± 10 nm</td>
</tr>
<tr>
<td>σ_gm</td>
<td>1.54 ± 0.1</td>
<td>1.48 ± 0.04</td>
</tr>
<tr>
<td>D_p/D_Bc</td>
<td>1.52</td>
<td>1.48</td>
</tr>
</tbody>
</table>


[37] Figure 5 shows the correlations between ΔCO2, ΔCO, ΔCH3CN, ΔM_BC, and ΔV_{SC}. The correlations (ΔCO-ΔCO2, ΔCH3CN-ΔCO, ΔM_BC-ΔCO2, ΔV_{SC}-ΔCO2, ΔM_BC-ΔCO, ΔV_{SC}-ΔCO, and ΔV_{SC}-ΔM_BC) are color coded by the phase of BB. The least squares fits were made for each phase and were constrained through the origin. The flaming phase data obtained in Canada (red circles) and California (red crosses) showed no significant differences. The variability in the slopes was smallest for the ΔCH3CN-ΔCO correlation. The correlation depended little on the phase of BB.

4. Aerosols in Plumes Transported From Asia (Siberia and Kazakhstan)

[38] Figure 9 shows time series plots of CO, M_BC, V_{SC}, and altitude of the DC-8, measured on 12–13 April (F#8). The M_BC and V_{SC} values were enhanced during the encounters with wildfire plumes. The time periods of sampling plumes of different origins in Asia are also shown in Figure 9. The plumes of the different origins appear as a number of interleaved spikes.

[39] We extracted data from air masses strongly influenced by biomass burning in Asia (Siberia and Kazakhstan). Most of the air masses enhanced in ΔM_BC and ΔV_{SC} were transported from Asia in 4.5 ± 2.1 days according to the calculated 10 day back trajectories.

[40] Figure 10 shows the ΔM_BC/ΔCO correlation versus the estimated APT that the air masses experienced. Generally BB plumes transported from Asia in spring experienced small amounts of APT. The observed ΔM_BC/ΔCO correlation tends to decrease with APT, especially for APT > 20 mm. We used the data with APT < 10 mm for the analysis of the emission ratios to minimize the possible effect of wet removal during transport. The slopes changed little by selecting data with APT < 5 mm, although the number of data points was greatly reduced.

[41] Figure 8 and Tables 4 and 5 show the average mass concentrations and mass fractions of PM1 nonrefractory components of aerosol and BC in the Asian BB plumes measured by HR-ToF-AMS and SP2. OA constituted the dominant fraction (about 60%) of the nonrefractory components, similar to the Canadian BB plumes. However, the sulfate mass fraction was as large as about 30%, which is much larger than that for the North American plumes (3%–8%). It is very likely that SO2 emitted from biomass burning was oxidized to form sulfate during transport. Formation of sulfate is discussed in more detail in section 7.

[42] Table 6 shows the average V_{SC}/M_BC and V_{SC}/V_{BC} ratios in the Asian plumes, restricted to low APT. These ratios were about 2 times lower than those in the Canadian plumes. The conclusion that most of the light-scattering particles did not contain BC is still valid. The difference in the BC fractions is discussed in terms of the differences in the conditions of combustion in section 6.

[43] The total mass concentrations of nonrefractory components measured by the AMS (M_{AMS}) were highly correlated with V_{SC}, with r^2 = 0.89 and 0.84 for ARCTAS-A and B, respectively. The average densities of nonrefractory components derived from the least squares fits of the M_{AMS}-V_{SC} were 1.33 and 1.54 g cm^{-3} for ARCTAS-A and B, respectively.

[44] Aerosols strongly influenced by BB have been shown to act as cloud condensation nuclei (CCN) [e.g., Hallett et al., 1989; Vestin et al., 2007; Fors et al., 2010] due to the water solubility of some organic compounds. Therefore, the BC particles (thickly coated) in the Canadian and Asian plumes will act as CCN. It is likely that BC particles from Asia were also coated by sulfate, leading to higher CCN activity because of the high hygroscopicity of sulfate [e.g., Petters and Kreidenweis, 2007]. However, actual BC removal by precipitation depends on the timing of precipitation with respect to the formation of sulfate during transport. It is therefore difficult to compare the efficiency of wet removal BC particles in Canadian and Asian plumes.

[45] Figure 11 shows the correlations between ΔCO2, ΔCO, CH3CN, ΔM_BC, and ΔV_{SC} of the plumes (with

Figure 8. Fractions of major chemical components of aerosols measured by AMS in major forest fire plumes in (left) Canada and (right) Asia.
APT < 10 mm) transported from Siberia and Kazakhstan during the ARCTAS-A period. We defined mixed air masses as those in which the trajectories passed over both regions. The plumes from the different source regions in Asia were classified and are color coded in Figure 11. The correlations of air masses transported from Kazakhstan and Siberia did not show significant differences. The results were the same even if we used the data set without discrimination by APT. Because of the small differences between theirs and the present results are not understood, although it may be partly due to the difference in identifying the source regions by using different methods or to variability within each source region coupled with the sampling of different plumes by the two aircraft.

[47] It should be noted that the $\Delta CO/\Delta CO_2$ ratios can change during the course of dilution of the BB plumes through mixing with ambient air [e.g., Takegawa et al., 2003, 2004]. The effect of dilution can be larger for longer-lived species with higher background values, such as CO$_2$. Initial changes in the $\Delta CO/\Delta CO_2$ ratios of the BB plumes occur when they are injected into the FT. However, generally $\Delta CO_2$ and CO in the BB plumes are much larger than the difference in the background CO$_2$ and CO mixing.

Table 4. Concentrations of Species for BB Plumes During ARCTAS and CARB

<table>
<thead>
<tr>
<th>Species Units</th>
<th>NA BB, ARCTAS-B</th>
<th>Canada BB, ARCTAS-B</th>
<th>CA BB, CARB</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>MCE &lt; 0.90, Smoldering (1 July)</td>
<td>MCE &gt; 0.95, Flaming (4 and 13 July)</td>
<td>MCE &lt; 0.90, Smoldering (4 July)</td>
</tr>
<tr>
<td>Number</td>
<td>69</td>
<td>353</td>
<td>125</td>
</tr>
<tr>
<td>$\Delta CO$ ppbv</td>
<td>40.8</td>
<td>124</td>
<td>328</td>
</tr>
<tr>
<td>$\Delta CO_2$ ppmv</td>
<td>2.79</td>
<td>2.17</td>
<td>1.33</td>
</tr>
<tr>
<td>$\Delta CH_3CN$ pptv</td>
<td>85.3</td>
<td>194</td>
<td>536</td>
</tr>
<tr>
<td>BC ng m$^{-3}$</td>
<td>273</td>
<td>331</td>
<td>517</td>
</tr>
<tr>
<td>LSP $\mu m^3$ cm$^{-3}$</td>
<td>3.51</td>
<td>8.78</td>
<td>13.6</td>
</tr>
<tr>
<td>Shell/core ratio</td>
<td>1.52</td>
<td>1.48</td>
<td>1.42</td>
</tr>
<tr>
<td>Cl$^-$</td>
<td>0.0192</td>
<td>0.0253</td>
<td>0.0653</td>
</tr>
<tr>
<td>NH$_4^+$ $\mu g$ m$^{-3}$</td>
<td>0.398</td>
<td>0.674</td>
<td>0.812</td>
</tr>
<tr>
<td>NO$_3^-$ $\mu g$ m$^{-3}$</td>
<td>0.138</td>
<td>0.259</td>
<td>0.847</td>
</tr>
<tr>
<td>SO$_4^{2-}$ $\mu g$ m$^{-3}$</td>
<td>1.67</td>
<td>1.79</td>
<td>1.19</td>
</tr>
<tr>
<td>SO$_2$ pptv</td>
<td>2.02</td>
<td>1.78</td>
<td>1.75</td>
</tr>
<tr>
<td>$\Delta CO_2$ pptv</td>
<td>3.69</td>
<td>12.2</td>
<td>40.5</td>
</tr>
<tr>
<td>PM$_1$ mg m$^{-3}$</td>
<td>6.19</td>
<td>14.69</td>
<td>43.93</td>
</tr>
<tr>
<td>PM$_1$ + Av SO$_2$</td>
<td>6.54</td>
<td>15.27</td>
<td>44.49</td>
</tr>
<tr>
<td>NO ppmv</td>
<td>15.4</td>
<td>51.3</td>
<td>46.7</td>
</tr>
<tr>
<td>NO$_2$ pptv</td>
<td>16.5</td>
<td>153</td>
<td>171</td>
</tr>
<tr>
<td>NO$_y$ pptv</td>
<td>881</td>
<td>1158</td>
<td>1598</td>
</tr>
<tr>
<td>NO$_2$ pptv</td>
<td>15.6</td>
<td>110</td>
<td>43.6</td>
</tr>
<tr>
<td>PAN pptv</td>
<td>47.4</td>
<td>457</td>
<td>1387</td>
</tr>
<tr>
<td>NO$_y$/NO$_2$</td>
<td>0.0362</td>
<td>0.176</td>
<td>0.136</td>
</tr>
<tr>
<td>SO$_2$(CT) pptv</td>
<td>91.9</td>
<td>144</td>
<td>131</td>
</tr>
<tr>
<td>SO$_2$ pptv</td>
<td>73.6</td>
<td>126</td>
<td>129</td>
</tr>
<tr>
<td>Average SO$_2$ pptv</td>
<td>83</td>
<td>135</td>
<td>130</td>
</tr>
<tr>
<td>SO$_2$-CO$_2$ pptv/ppbv</td>
<td>2.0</td>
<td>1.1</td>
<td>0.39</td>
</tr>
<tr>
<td>SO$_2$-CO$_2$ pptv/ppbv</td>
<td>0.35</td>
<td>0.58</td>
<td>0.56</td>
</tr>
<tr>
<td>DMS pptv</td>
<td>NA</td>
<td>6</td>
<td>8</td>
</tr>
</tbody>
</table>

Table 5. Averages and Standard Deviations of the Ratios of Mass Concentrations of BC, Inorganic, and Organic Aerosols to PM1 in wt % for the Smoldering (MCE < 0.90) and Flaming Phases (MCE > 0.95) During ARCTAS and CARB

<table>
<thead>
<tr>
<th>Species</th>
<th>ARCTAS-B</th>
<th>CARB</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>MCE &lt; 0.90 (1 July)</td>
<td>MCE &gt; 0.95 (4 and 13 July)</td>
</tr>
<tr>
<td>BC</td>
<td>4.9 ± 1.1</td>
<td>2.7 ± 2.9</td>
</tr>
<tr>
<td>Cl$^-$</td>
<td>0.3 ± 0.2</td>
<td>0.2 ± 0.2</td>
</tr>
<tr>
<td>NH$_4^+$</td>
<td>6.1 ± 1.7</td>
<td>3.6 ± 2.6</td>
</tr>
<tr>
<td>NO$_3^-$</td>
<td>2.6 ± 1.6</td>
<td>1.8 ± 1.7</td>
</tr>
<tr>
<td>SO$_4^{2-}$</td>
<td>29.6 ± 9.1</td>
<td>7.7 ± 7.1</td>
</tr>
<tr>
<td>SO$_2$ + SO$_2$</td>
<td>31.1 ± 9.6</td>
<td>11.9 ± 1.3</td>
</tr>
<tr>
<td>Org</td>
<td>56.3 ± 8.6</td>
<td>84.2 ± 11.4</td>
</tr>
</tbody>
</table>
ratios between the BL and FT. Therefore, the changes in the slopes of the CO–CO2 correlations by this process should be small.

Larger changes in the DCO/DCO2 ratios are considered to occur by mixing of the diluted BB plumes with ambient air during the long-range transport in the FT. The average CO2 mixing ratios at the surface and in the FT were observed to increase by about 1 ppmv between 50°N and 70°N [e.g., Komhyr et al., 1985; Tans et al., 1989; Anderson et al., 1996]. In addition, the mean CO2 mixing ratios of the 1 km averaged values between 2 and 7 km were 391.7 ± 0.7 ppmv during ARCTAS-A. If 50% of the BB plumes are mixed with ambient air during transport from 50°N to 70°N, the CO2 mixing ratio of the mixed air is anticipated to increase by 0.5 ppmv on average. This leads to an uncertainty of the ΔCO/ΔCO2 ratios of about 15%–20% for the average ΔCO2 of 2.8 ppmv (Table 4), giving a measure of the uncertainties in the derived ΔCO/ΔCO2 for the Asian BB plumes.

The size distributions of BC and LSP in the plumes from Asia (APT < 10 mm) are shown in Figure 7. The CMD and MMD (VMD) of the lognormal fitted BC and LSP size distributions together with the values of the standard deviations are summarized in Table 3. The variability of these parameters was small. The median Dp/DBC ratios for the Asian plumes was 1.52 (corresponding

Table 6. Emission Ratios of BC, CO, CO2, and CH3CN for Biomass Burninga

<table>
<thead>
<tr>
<th>Species</th>
<th>Asia BB, ARCTAS-A</th>
<th>NA BB, ARCTAS-B</th>
<th>CA BB, CARB, ARCPAC</th>
<th>Kazakhstan,b</th>
<th>Lake Baikal,b</th>
<th>Andreae and Merlet</th>
<th>Tokyo FF,c</th>
</tr>
</thead>
<tbody>
<tr>
<td>CO/CO2 (ppbv/ppmv)</td>
<td>15 ± 5</td>
<td>80 ± 110</td>
<td>222 ± 118</td>
<td>26 ± 10</td>
<td>114 ± 23</td>
<td>50 ± 25</td>
<td>42 ± 19</td>
</tr>
<tr>
<td>MBC/CO (ng m⁻³/ppbv)</td>
<td>8.5 ± 5.4</td>
<td>2.3 ± 2.2</td>
<td>1.7 ± 0.8</td>
<td>3.4 ± 1.6</td>
<td>3.4 ± 1.4</td>
<td>10 ± 5</td>
<td>7 ± 4</td>
</tr>
<tr>
<td>MBC/CO2 (ng m⁻³/ppmv)</td>
<td>129 ± 67</td>
<td>180 ± 269</td>
<td>369 ± 257</td>
<td>88 ± 80</td>
<td>236 ± 103</td>
<td>NA</td>
<td>NA</td>
</tr>
<tr>
<td>MBC/CH3CN (ng m⁻³/ppmv)</td>
<td>3.9 ± 2.1</td>
<td>1.2 ± 1.4</td>
<td>0.88 ± 0.6</td>
<td>2.2 ± 1.3</td>
<td>1.7 ± 0.6</td>
<td>NA</td>
<td>NA</td>
</tr>
<tr>
<td>CH3CN/CO (pptv/ppmv)</td>
<td>1.6 ± 1.6</td>
<td>1.8 ± 0.4</td>
<td>1.9 ± 0.4</td>
<td>1.7 ± 0.4</td>
<td>1.7 ± 0.4</td>
<td>NA</td>
<td>NA</td>
</tr>
<tr>
<td>VSC/MBC (cm³ mg⁻¹ g⁻¹)</td>
<td>12 ± 3</td>
<td>24 ± 6</td>
<td>26 ± 7</td>
<td>24 ± 6</td>
<td>23 ± 6</td>
<td>NA</td>
<td>NA</td>
</tr>
<tr>
<td>VSC/VBC (cm³/cm³)</td>
<td>24 ± 6</td>
<td>48 ± 12</td>
<td>52 ± 14</td>
<td>48 ± 12</td>
<td>46 ± 12</td>
<td>NA</td>
<td>NA</td>
</tr>
<tr>
<td>VSC/CO (μm² cm⁻³/ppbv)</td>
<td>10 ± 2.3</td>
<td>5.4 ± 1.6</td>
<td>4.4 ± 2.3</td>
<td>8.3 ± 4.3</td>
<td>–</td>
<td>NA</td>
<td>NA</td>
</tr>
<tr>
<td>VSC/CO2 (μm² cm⁻³/ppmv)</td>
<td>1.6 ± 0.7</td>
<td>4.4 ± 4.0</td>
<td>9.6 ± 8.0</td>
<td>2.1 ± 0.9</td>
<td>–</td>
<td>NA</td>
<td>NA</td>
</tr>
</tbody>
</table>

aThe mass concentrations are in units of μg STP m⁻³.
bWarneke et al. [2009].
cFossil fuel combustion in California.
dFossil fuel combustion in Tokyo [Kondo et al., 2010].

Figure 9. Time series plots of CO, BC mass concentration (MBC), and volume concentration of LSP (VSC) measured over Alaska on 12–13 April. Bars indicate the periods when the plumes from the respective region were sampled.
to a BC volume fraction of 0.28 for BC–containing particles, indicating that the BC particles were thickly coated. The size distributions of BC and LSP and the $D_p/D_{bc}$ ratios of the Asian fire plumes were very similar to those of the North American fire plumes (Figure 7), despite the significant differences in the emission ratios, including the $\Delta CO/\Delta CO_2$ ratios.

5. Average Emission Ratios

[50] We derived the emission ratios between $M_{bc}$, $V_{sc}$, CH$_3$CN, CO, and CO$_2$ averaged over the measurement period during ARCTAS-B. The averaging method described by Andreae and Merlet [2001] was used for this purpose. Namely, the net increases of species $X$ and $Y$ above their background values in the $j$th plume were first calculated ($\Delta X_j$ and $\Delta Y_j$). Then the sums for all the plumes were calculated, i.e., $\Delta X = \sum_j \Delta X_j$ and $\Delta Y = \sum_j \Delta Y_j$. The average emission ratio of $X$ versus $Y$ was given as $ER_{av} = \Delta Y/\Delta X$. We note that for $V_{sc}$ the value derived is not strictly an “emission” ratio, as secondary inorganic (e.g., sulfate, as discussed above) and organic aerosol species [e.g., DeCarlo et al., 2010] are formed during transport and solar exposure.

[51] We derived the emission ratios of CO/CO$_2$, $M_{bc}/CO$, $M_{bc}/CO_2$, $M_{bc}/CH_3CN$, CH$_3$CN/CO, $V_{sc}/CO_2$, $V_{sc}/CO$, and $V_{sc}/M_{bc}$, and they are summarized in Table 6. The large variability in $M_{bc}/CO$, $M_{bc}/CO_2$, and $M_{bc}/CH_3CN$ are readily seen. The lower variabilities in the CH$_3$CN/CO emission ratio is also seen.

[52] The emission ratios for BB in Asia, discussed in section 4, were derived in the same way and are given in Table 6. In Table 6, we added the reported values for biomass burning in Asia for April 2008 during the ARCPAC Mission [Warneke et al., 2009]. In addition, we also included the data for forest fires in California during ARCTAS-CARB in Table 6 for comparison.

[53] The average CO/CO$_2$ (or MCE) and $M_{bc}/CO$ emission ratios were similar for wildfires during ARCTAS-B (Canada and California) with $\Delta CO/\Delta CO_2 = 80 \pm 100$ ppbv/ppmv and ARCTAS-CARB (California) with $\Delta CO/\Delta CO_2 = 114 \pm 23$ ppbv/ppmv. In the Asian plumes, the average CO/CO$_2$ emission ratios were much smaller with $\Delta CO/\Delta CO_2 = 15 \pm 5$ ppbv/ppmv, and the $M_{bc}/CO_2$ emission ratios were a few times larger, although the $M_{bc}/CO_2$ emission ratios were similar.

[54] Estimates of the emission ratios for biomass burning in previous studies were compiled by Andreae and Merlet [2001]. The compiled values for temperate forests are also listed in Table 6. The present $M_{bc}/CO$ and $M_{bc}/CO_2$ emission ratios for wild fires in North America are lower than the compiled value by a factor of 2–4, although the CO/CO$_2$ emission ratio was close to their estimates. The derived $M_{bc}/CO$ emission ratio for BB plumes from Asia was much closer to the compiled value, although the observed $M_{bc}/CO_2$ emission ratio was about 5 times lower.

[55] Table 6 also shows the $V_{sc}/M_{bc}$ ratios observed in plumes strongly influenced by emissions by combustion of fossil fuel in California (ARCTAS-CARB) and in Tokyo in August–September 2009 [Kondo et al., 2011]. These data were taken from the same SP2 calibrated in the same way, enabling direct comparison. The comparison shows that much larger volumes of LSP relative to those of BC were emitted with BC from the wildfires than those from fossil fuel combustion. Generally, the emission ratios of primary organic aerosols to BC have been reported to be of the order of 1 when scaled with the mass concentrations [Takegawa et al., 2006] for fossil fuel combustion, consistent with the $V_{sc}/V_{bc}$ ratio for Tokyo in Table 6. The variability in the $V_{sc}/M_{bc}$ ratios in Tokyo was caused mainly by the formation of secondary aerosols (organic and inorganic) [e.g., Kondo et al., 2011], consistent with results at many other locations [de Gouw and Jimenez, 2009].

[56] OA/BC ratios were observed for aerosol particles emitted from crown fires in Canadian wildfires in a previous study [Conny and Slater, 2002]. These authors report that BC/TC (OC + BC) ratios were 0.085 and 0.0087 for the flaming and smoldering stages, respectively. The corresponding OC/BC ratios were about 10 and 100 for these stages, consistent with the present results. OA/BC and OC/BC mass ratios of about 6 and 4 were observed in agricultural fire plumes in the Yucatan [Yokelson et al., 2009]. These ratios are much lower than the present results and those reported by Conny and Slater [2002].

6. Dependence of the Emission Ratios on MCE

6.1. Injection Height

[57] The injection heights of the plumes formed in different phases of burning can vary because of the difference in the thermal energy release as a function of the burning phase. This point was suggested by Andreae and Merlet [2001] but has not been fully demonstrated by observed atmospheric species. To investigate this possible effect, the sampling altitudes of the plumes are plotted versus the $\Delta CO/\Delta CO_2$ ratios and MCE in Figure 12. It is seen that the $\Delta CO/\Delta CO_2$ ratios decreased and MCE increased with altitude, indicating a strong dependence of these parameters on the phase of burning.

[58] The thermal energy received by the plumes during burning increases their potential temperature (PT) from an average surface value of $\theta_0$ to $\theta_{air}$. The plumes with $\theta_{air} (>\theta_0)$
Figure 11. Correlations between $M_{BC}$, $V_{SC}$, $\Delta CO_2$, $\Delta CO$, and CH$_3$CN observed in major wildfire plumes that experienced little precipitation during transport from Asia to Alaska (ARCTAS-A). The least squares linear regression fits are shown by solid lines.
gain buoyancy and ascend to the altitude where the $PT$ of the ambient air ($\theta_a$) is equal to $\theta_{BC}$. Cloud formation will not substantially alter $\theta_{BC}$, if clouds evaporate without precipitation. Therefore $\Delta \theta = \theta_a - \theta_0$ is considered to be proportional to the thermal energy gained per unit mass of the plume near the surface. We estimated $\Delta \theta$ by assuming that $\theta_a$ is the $PT$ of the sampled plume and $\theta_0$ is the average $PT$ near the surface during each flight, for plumes sampled near the point of emission. Active mixing in the BL during daytime by solar heating of the surface will smooth out the vertical structure of $\Delta \theta$ produced by the thermal energy of combustion. Therefore the derived $\Delta \theta$ should be interpreted as a representation of the injection height that accounts for the vertical thermodynamic structure near the BB sources in North America.

Figure 12 (bottom) shows the relationships of the $\Delta CO/\Delta CO_2$ ratio and MCE with $\Delta \theta$ for the North American plumes. The sharp increase in $\Delta \theta$ at MCE > 0.98 is clearly seen. Namely, a substantial increase in $\Delta \theta$ was detected only at the highest combustion efficiencies.

For comparison, the data for the Asian plumes are also plotted in Figure 12. We estimated $\theta_0$ for these plumes by combining the calculated trajectories and the surface $PT$ of the burning areas in Asia estimated using NCEP data. The uncertainty of $\theta_0$ was estimated from the variability in the temperature of the source region. Because we cannot narrow down the source region from the trajectory analysis, the uncertainty in $\theta_0$ was as large as about 10 K. However, because the sampled plumes were injected directly into the FT, the uncertainty in estimating $\Delta \theta$ caused by mixing in the BL should be smaller than for the North American plumes measured in the BL. Despite this uncertainty, the increase in $\Delta \theta$ corresponded with an MCE as high as 0.98, similarly to the North American plumes.

It is likely that plumes formed during the flaming phase were selectively injected to higher altitudes than those during the smoldering phase over Asia. The higher-altitude plumes were transported in shorter time to the sampling regions near Alaska, associated with higher wind speeds [Fuelberg et al., 2010]. In other words, the $\Delta CO/\Delta CO_2$ ratios of the Asian plumes measured during ARCTAS-A may not represent the vertically averaged emission values in Asia. The estimation of relative contributions of the flaming and smoldering phases in producing aerosols and trace gases including CO will require data in the BL over Asia.

### 6.2. Emission Ratios

We have attempted to interpret the large variability of the slopes of the $\Delta M_{BC}/\Delta CO$, $\Delta M_{BC}/\Delta CO_2$, $\Delta M_{BC}/\Delta CH_3CN$, $\Delta V_{SC}/\Delta CO_2$, $\Delta V_{SC}/\Delta CO$, and $\Delta V_{SC}/\Delta M_{BC}$ correlations in terms of MCE. These ratios are plotted versus MCE in Figure 13. Standard deviations (1σ) of the 1 min data used to derive individual data points were calculated first. Table 7 summarizes the averages of the 1σ values and the average ratios for BB plumes in Asia, Canada, and California shown in Figure 13. The least squares fits were derived using the data sets with APT < 1 mm for Canada and California and APT < 5 mm for Asia. It should be noted that some of the emission ratios ($\Delta V_{SC}/\Delta CO_2$, $\Delta V_{SC}/\Delta CO$, $\Delta V_{SC}/\Delta M_{BC}$) are not independent of MCE, because the estimate of MCE contains $\Delta CO_2$ and $\Delta CO$.

The $\Delta M_{BC}/\Delta CO$ ratio increased with the increase in MCE. This relation is important because the ratio is often used in estimating BC emissions and BC removal during
transport using CO as a tracer. Because of the large variability in the $\Delta M_{BC}/\Delta CO$ ratios, the estimates of BC removal based on downwind measurements of $\Delta M_{BC}/\Delta CO$ will have large uncertainty unless $\Delta M_{BC}/\Delta CO$ was measured at the source. In contrast, the $\Delta M_{BC}/\Delta CO_2$ and the $\Delta V_{SC}/\Delta CO_2$ ratios decreased with the increase in MCE. The $\Delta V_{SC}/\Delta M_{BC}$ ratios also decreased with the increase in MCE.

For flaming fires, the lower $\Delta V_{SC}/\Delta M_{BC}$ ratios tend to lower the single scattering albedo (SSA) of the ensemble of aerosols. The thicker coating of BC for the flaming phase leads to amplified absorption as compared with that for externally mixed BC. The SSA of individual BC particles increases with the increase in the shell/core ratios because of the increase in the scattering cross section [Oshima et al., 2009]. However, the amplified BC absorption leads to a decrease in the SSA of the ensemble of aerosols [e.g., Shiraiwa et al., 2008; Seinfeld and Pandis, 2006, Figure 21.11]. For this reason, we expect a lower (higher) SSA in plumes of the flaming (smoldering) phase. In fact, the SSA at a wavelength of 538 nm was observed to decrease substantially with the increase in MCE in BB plumes in Brazil [Reid and Hobbs, 1998].

The $\Delta M_{BC}/V_{SC}$ ($\Delta V_{SC}/\Delta M_{BC}$) ratio showed an almost exponential increase at MCE > 0.95. The rapid increase in BC/OC at MCE > 0.95 was observed by combustion facility experiments [Christian et al., 2003], consistent with the present results, considering that OA is the dominant component of $\Delta V_{SC}$.

The $\Delta V_{SC}/\Delta CO$ ratios increased with the increase in MCE, although the correlation was not tight. The interpre-

Table 7. Averages of the 1σ Values ($\pm$) and the Average Ratios, MCE, and Shell/Core Ratios for BB Plumes in Asia, Canada, and California

<table>
<thead>
<tr>
<th>Ratios</th>
<th>Asia</th>
<th>Canada</th>
<th>California</th>
</tr>
</thead>
<tbody>
<tr>
<td>$\Delta V_{SC}/\Delta CO_2$</td>
<td>1.88 ± 0.40</td>
<td>7.97 ± 2.30</td>
<td>2.50 ± 1.16</td>
</tr>
<tr>
<td>$\Delta V_{SC}/\Delta M_{BC}$</td>
<td>0.0116 ± 0.0011</td>
<td>0.0261 ± 0.0053</td>
<td>0.0246 ± 0.0019</td>
</tr>
<tr>
<td>$\Delta V_{SC}/\Delta CO$</td>
<td>11.8 ± 4.5</td>
<td>3.25 ± 0.67</td>
<td>2.86 ± 0.35</td>
</tr>
<tr>
<td>$\Delta M_{BC}/\Delta CO_2$</td>
<td>0.1033 ± 0.051</td>
<td>0.0749 ± 0.0116</td>
<td>0.0699 ± 0.0053</td>
</tr>
<tr>
<td>$\Delta M_{BC}/\Delta M_{BC}$</td>
<td>167 ± 37</td>
<td>357 ± 117</td>
<td>100 ± 48</td>
</tr>
<tr>
<td>$\Delta M_{BC}/\Delta CH_3CN$</td>
<td>5.16 ± 1.44</td>
<td>2.10 ± 0.56</td>
<td>2.12 ± 0.34</td>
</tr>
<tr>
<td>MCE</td>
<td>0.985 ± 0.0023</td>
<td>0.846 ± 0.060</td>
<td>0.961 ± 0.021</td>
</tr>
<tr>
<td>Shell/core ratio</td>
<td>1.54 ± 0.033</td>
<td>1.45 ± 0.026</td>
<td>1.51 ± 0.033</td>
</tr>
</tbody>
</table>
Table 8. Comparison of SO2 and Sulfate Concentrations in Background Air and BB Plumes From Asia During ARCTAS-A With and Without Selection by APT

<table>
<thead>
<tr>
<th>Instrument</th>
<th>All</th>
<th>APT &lt; 10 mm</th>
</tr>
</thead>
<tbody>
<tr>
<td>SO2 (ppbv)</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Background CIT</td>
<td>98</td>
<td>116</td>
</tr>
<tr>
<td>GIT</td>
<td>78</td>
<td>101</td>
</tr>
<tr>
<td>Asian BB CIT</td>
<td>108</td>
<td>140</td>
</tr>
<tr>
<td>GIT</td>
<td>81</td>
<td>105</td>
</tr>
<tr>
<td>Sulfate (µg m⁻³)</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Background AMS</td>
<td>0.13</td>
<td>0.14</td>
</tr>
<tr>
<td>Asian BB AMS</td>
<td>0.46</td>
<td>0.49</td>
</tr>
</tbody>
</table>

A number of studies have shown the particle emission factor (g/kg) to decrease with the increase in MCE by burning tropical forest fuels in a combustion facility [Ferek et al., 1998; Yokelson et al., 2007, 2008]. The particle mass and fuel mass are approximately proportional to ΔFSC and ΔCO/ΔCO₂ ratios. Their results are qualitatively consistent with the present observations, although the types of fuels were different, namely tropical versus boreal forests. The agreement in the trend between our study and those of Yokelson et al. [2007, 2008] provides additional support for the methodology in using MCE in the interpretation of field data. The present analysis demonstrates the necessity of taking into account the dependence of the emission ratios of aerosols on the burning phase (e.g., MCE) in improving the uncertainty in estimating aerosol emissions from BB.

7. Formation of Sulfate

A number of studies have shown the particle emission factor (g/kg) to decrease with the increase in MCE by burning tropical forest fuels in a combustion facility [Ferek et al., 1998; Yokelson et al., 2007, 2008]. The particle mass and fuel mass are approximately proportional to ΔFSC and ΔCO/ΔCO₂ ratios. Their results are qualitatively consistent with the present observations, although the types of fuels were different, namely tropical versus boreal forests. The agreement in the trend between our study and those of Yokelson et al. [2007, 2008] provides additional support for the methodology in using MCE in the interpretation of field data. The present analysis demonstrates the necessity of taking into account the dependence of the emission ratios of aerosols on the burning phase (e.g., MCE) in improving the uncertainty in estimating aerosol emissions from BB.

8. Microphysical Properties of BC and LSP

8.1. Size Distribution

The number and mass (volume) size distributions of aerosols in the BB plumes from North America and Asia were fitted to a good approximation with lognormal size distribution functions, with typical examples shown in Figure 7. To a first approximation, these parameters were very similar in all the BB plumes. For more detailed investigations,
Figure 14 compares the average CMD and standard deviation ($\sigma_{gc}$) of BC and LSP and the median $D_p/D_{BC}$ ratios in the North American and Asian plumes as a function of MCE. Table 7 summarizes the averages of the 1 $\sigma$ values of the $D_p/D_{BC}$ ratios together with the average $D_p/D_{BC}$ ratios. The least squares fits were derived using the data sets with APT $<$ 1 mm for Canada and California and APT $<$ 5 mm for Asia. Some of the plumes in Canada were measured after 1–2 h after emissions, suggested from the NOx/NOy ratios (Table 4).

[75] As a whole, the CMD and $\sigma_{gc}$ of BC did not change substantially without depending on the location of the fires, MCE, and age, although the CMD increased and $\sigma_{gc}$ decreased slightly with the increase in MCE. The rate of growth BC by coagulation depends strongly on the initial BC concentration [e.g., Hinds, 1999; Seinfeld and Pandis, 2006]. After this initial growth, BC particles thickly coated by mainly organic compounds probably grew much more slowly by coagulation under largely diluted conditions. Coagulation generally increases the CMD and decreases $\sigma_{gc}$ [Kajino and Kondo, 2011]. Figure 14 shows that the CMD and $\sigma_{gc}$ of BC slightly increased and decreased, respectively, with the increase in MCE. The observed anticorrelation of CMD and $\sigma_{gc}$ of BC is qualitatively consistent with this effect.

[76] The small variability of the size distributions of LSP measured by SP2 is also noteworthy, considering the differences in the vegetation and therefore different chemical compounds constituting the OA particles. The CMD and $\sigma_{gc}$ increased slightly with the increase in MCE. These are the first extensive measurements of the size distributions of BC and LSP (mainly OA) emitted from boreal forest fires.

[77] The CMD and MMD of BC impacted by BB near the Gulf of Mexico were about 140 nm and 200 nm [Schwarz et al., 2008], confirming the small variability of the size distributions of BC from BB. The ages of the BB plumes were estimated to be about 0.5–1.5 h. These observations simplify modeling efforts in expressing the size distributions of primary particles emitted from boreal forest fires. Namely, the BC and LSP number size distributions with ages older than a few hours are represented to a good approximation by lognormal distribution functions, with an average CMD = 138 nm and $\sigma_{g} = 1.35$ and an average CMD = 230 nm and $\sigma_{g} = 1.32$, respectively.

Figure 14. CMD and $\sigma_{g}$ of (top) BC and (middle) LSP and (bottom) median shell/core ratios plotted versus MCE for the Asian and North American BB plumes. Linear regression fits using all the data points are also shown.
8.2. Evolution of the Mixing State of BC

8.2.1. Observed Features

We investigated the evolution of the mixing state of BC for different phases of combustion mainly using the Canadian data, which contain data with largely different air mass ages. We used NOx/NOy and NOy/DCO ratios as a measure of aging with different time scales. Figure 15 shows the median Dp/DbC (SC) ratio plotted versus NOx/NOy and NOy/DCO ratios in BB plumes during ARCTAS-B. The data points are color coded for the smoldering phase (MCE < 0.9; blue), flaming phase (MCE > 0.95; red), and others (0.90 < MCE < 0.95; green).

For smoldering phase emissions, the Dp/DbC ratio increased to 1.45 at NOx/NOy ratios of 0.05–0.1. The ages of plumes are roughly estimated to be about 12–24 h. The increase in the Dp/DbC ratio from 1.35 to 1.45 corresponds to a 40 ± 30% increase of the volume of the coating materials. It should be noted that the estimate in the increase depends sensitively on the choice of the Dp/DbC ratio. For the plumes with MCE between 0.9 and 0.95 (green points in Figure 15), the increase in volume can be as high as 80%, although the scatter is substantial.

For the flaming phase plumes, we used the NOy/DCO ratios as indicators of photochemical age, because these plumes were more aged than the smoldering plumes. The Dp/DbC ratios clearly increased with the decrease in the NOy/DCO ratios shown in Figure 15 (bottom). The increase in the Dp/DbC ratio from 1.4 at NOy/DCO = 30 to 1.66 at NOy/DCO = 10 corresponds to a 205 ± 40% increase of the volume of the coating materials.

In very fresh BB plumes in Australia, the NOy/DCO ratios were observed to increase with the increase with MCE [Takegawa, 2001]. They were about 40 pptv/ppbv and 20 pptv/ppbv at MCE = 0.96 and 0.92, respectively, similar to the present observations. NOx emitted is converted to PAN and HNO3, followed by conversion of PAN to HNO3 and dry deposition of HNO3 [e.g., Mauzerall et al., 1998; Takegawa et al., 2003; Alvarado et al., 2010]. About 60% of NOx was observed to decrease within 2–3 days [Takegawa et al., 2003]. Most of the data shown in Figures 15 and 17 were obtained in the BL. The main conclusions discussed in this section are unchanged even if the BL data were selectively used.

The changes in the NOx/NOy, PAN/NOy, and HNO3/NOy ratios in the plumes in the flaming phase are shown in Figure 16, in order to understand the processes of the changes in the NOy. It can be seen that about 70% of NOx was lost by conversion to PAN and HNO3 already at an NOy/DCO ratio of 30 pptv/ppbv. The decrease in the NOy/DCO ratios occurred mainly through the deposition of HNO3, produced from NOx. It is estimated that the increase in the Dp/DbC ratios continued longer than about 2–3 days, considering the time scale of the NOy loss, through HNO3 deposition, as discussed above.
Figure 17 shows the $\Delta V_{SC}/\Delta M_{BC}$ and OA/$\Delta M_{BC}$ ratios versus the NO$_y$/$\Delta$CO ratio. Figure 17 indicates no substantial increase in $V_{SC}$ and OA relative to $M_{BC}$ with age. Therefore it is unlikely that condensation of secondary aerosol species formed from precursor gases emitted directly from BB caused the increase in the $\Delta p/\Delta BC$ ratios. Since the fraction of $V_{SC}$ accounted for by the BC-containing particles is very small, a potential explanation for this apparent discrepancy is that OA was evaporating from particles that contained OA but not BC and then recondensing (possibly after gas phase oxidation) on the BC-containing particles in the SP2 size range [Marcolli et al., 2004; Robinson et al., 2007; Grieshop et al., 2009].

8.2.2. Effect of Coagulation

Brownian coagulation may also explain the observed evolution of the BC mixing state. Figure 18 shows the total number concentration of LSP particles ($N_{sc}$) versus the NO$_y$/$\Delta$CO ratio. Generally $N_{sc}$ was less than 3000 cm$^{-3}$ and 1000 cm$^{-3}$ in the BB plumes in smoldering and flaming phases, respectively. We use these data in calculating the effect of coagulation, as detailed below.

A simple coagulation model was used to assess the effect of coagulation between LSP and BC on the evolution of the BC mixing state. The model solves for the intramodal coagulation and intermodal coagulation between the two modes, namely LSP (mode $i$) and BC (mode $j$). The evolution of number size distribution functions of LSP ($n_i(v_p)$) and BC ($n_j(v_p)$) due to intramodal coagulation can be expressed as [Whitby and McMurry, 1997],

$$\frac{d}{dt} n_i(v_p) = \frac{1}{2} \int_0^{v_p} \beta(v_p, v_p - \tilde{v}_p) n_i(v_p) n_j(v_p - \tilde{v}_p) d\tilde{v}_p - \int_0^{\infty} \beta(v_p, \tilde{v}_p) n_i(\tilde{v}_p) n_j(v_p - \tilde{v}_p) d\tilde{v}_p$$

(2)

where $v_p$ and $n_i(v_p)$ are the volume of a particle and number size distribution of mode $i$, respectively, and $\beta$ is the Brownian coagulation kernel [Binkowski and Shankar, 1995]. Here $v_p = \pi D_p^3 / 6$, assuming that all particles are spherical. The evolutions of $n_i(v_p)$ and $n_j(v_p)$ due to intermodal coagulation are expressed as follows.

$$\frac{d}{dt} n_i(v_p) = -\int_0^{v_p} \beta(v_p, v_j) n_i(v_p) n_j(v_j) dv_j$$

$$\frac{d}{dt} n_j(v_p) = \int_0^{v_p} \beta(v_p, v_p - v_j) n_i(v_p - v_j) n_j(v_j) dv_j$$

$$- \int_0^{\infty} \beta(v_p, \tilde{v}_p) n_i(\tilde{v}_p) n_j(v_p - \tilde{v}_p) d\tilde{v}_p$$

(3)

Figure 17. $\Delta V_{SC}/\Delta M_{BC}$ and OA/$\Delta M_{BC}$ ratios plotted versus the NO$_y$/$\Delta$CO ratio in the flaming phase plumes shown in Figure 15.

Figure 18. Total number concentrations of LSP particles plotted versus the NO$_y$/$\Delta$CO ratios during ARCTAS-B for the smoldering phase (MCE < 0.9; blue), flaming phase (MCE > 0.95; red), and others (0.90 < MCE < 0.95; green).
We assumed that the collision of a LSP particle (diameter of $D_{p}$) with a BC particle (shell diameter of $D_{p}$) forms a spherical BC particle with a shell diameter of $(D_{p}^{1/3} + D_{p}^{1/3})^{1/3}$, considering the conservation of volume concentration. Initially, $n_{p}$ was assumed to have a lognormal size distribution (LNSD) function. The coagulation processes represented by equations (2) and (3) were calculated using the difference (bin) method, allowing deviations from the LNSD function.

As initial conditions, the LNSD parameters were given covering the range of values shown in Figures 14 and 18. Namely, $N_{SC} = 500, 1500, and 2500\, \text{cm}^{-3}$; $N_{BC} = 0.1 \times N_{LSP}$; CMD$_{LSP} = 200, 230, and 250\, \text{nm}$; CMD$_{BC} (\text{shell}) = 120, 160, and 200\, \text{nm}$; $\sigma_{LSP} = 1.2$ and 1.4; and $\sigma_{BC} = 1.2$ and 1.3. The detailed model settings are the same as BIN100 of Kajino [2011], in which the size distribution was divided into 100 bins between 1 nm and 1 $\mu$m, and 100 bins between 1 $\mu$m and 1 mm. The high accuracy of BIN100 is shown in this study.

As discussed above, the rates of coagulation are insensitive to the parameters that define LNSD. Therefore, it is the difference in the size of LSP that contributes primarily to the difference in the rate of increase in the SC ratio or $V(t)$. Generally, larger $\sigma_{SC}$ and larger differences between $\sigma_{SC}$ and $\sigma_{BC}$ result in faster increases in the SC ratio because of this mechanism. In fact, the rate of increase in SC is largest for $\sigma_{SC} = 1.3$ shown in Figure 20 are the lowest values.

The number of LSP colliding with BC particles in 1–2 days is nearly the same as the number of BC particles per unit volume of air. Namely, one BC particle collides with one LSP particle once in 1–2 days, on average. The collision of one BC and one LSP particle with similar sizes results in a 1.26 times ($\sim 2^{1/3}$) increase in the SC ratio ($D_{p}/D_{BC}$). The SC ratio is predicted to increase from 1.2 to 1.5 by coagulation in 1–2 days. This result generally points out the importance of coagulation with LSP in increasing SC ratios with aging, in addition to condensation.

Now we translate the results of Figure 19 into the changes of the volumes of shells coating BC cores. Figure 20 shows time series of the ratio of the total shell volume of BC at time $t (V(t) = \int \nu_{p} n(\nu_{p}, t) d\nu_{p})$ to the initial $V(t = 0)$ for some selected cases. The initial LNSD parameters are shown in Figure 20. Figure 20a is the result of calculations for conditions close to those for the observations. CMD$_{BC} (\text{shell}) = 160\, \text{nm}$ corresponds to a SC ratio of about 1.33 for CMD$_{BC} (\text{core}) = 120\, \text{nm}$. The lines with triangles, squares, and circles are the results for initial CMD$_{SC} = 250, 230, and 200\, \text{nm}$, respectively. Figure 20b is the same as Figure 20a but for CMD$_{BC} (\text{shell}) = 120\, \text{nm}$ (i.e., SC ratio = 1 for CMD$_{BC} (\text{core}) = 120\, \text{nm}$). Figures 20c and 20d are the same as Figure 20a, but for an initial $N_{SC}$ of 1500 and 2500 $\text{cm}^{-3}$. The increase in $V(t)$ is faster for larger CMD$_{SC}$ in Figures 20a–20d.
This can largely explain the observed increase in $D_p/D_{BC}$ ratio by a factor of about 1.2 in the flaming phase plumes. BC was thickly coated 1–2 h after emission, as discussed in section 8.2.1 (Figure 15). Near BB sources, $N_{SC}$ can be substantially higher than $2500 \text{ cm}^{-3}$. In these areas, the coating of BC may occur in 1–2 h, considering the results shown in Figure 20d. Coagulation can play an important role in coating BC from early stages after BC emission from BB.

Previous studies considered only condensation process to explain the evolution of the BC mixing state in plumes strongly impacted by anthropogenic emission [e.g., Moteki et al., 2007; Shiraiwa et al., 2007, 2008]. For the first time, we suggest that coagulation of LSP with BC can also effectively increase $D_p/D_{BC}$ ratios based on the detailed comparison of observations and model calculations.

9. Conclusions

We have presented multiple analyses of the ARCTAS chemical data combined with meteorological data to quantify the emissions (including secondary aerosol formation) of different types of aerosols (black carbon (BC) and organic and inorganic species) from biomass burning (BB) that occurred in North America (Canada and California) and Asia, especially in terms of modified combustion efficiency (MCE). The microphysical properties of aerosols (size distribution and mixing state) emitted from BB were also quantified. Data strongly impacted by BB were unambiguously extracted using tracers, namely CH$_3$CN and CH$_2$Cl$_2$. We also selected data with low accumulated precipitation along trajectory (APT) values to minimize the effect of wet removal of aerosol on the statistical analysis of the aerosol emissions from BB. Below, we summarize the most important parameters and processes presented above, which can be used directly for improving model estimates of the impact of BB in the Arctic.

In North American plumes, organic aerosol (OA) constituted about 83% of the total PM$_1$, and the sulfate fraction was only 8% on average. By contrast, in the Asian plumes of the flaming phase, the sulfate fraction was as high as about 27%, with an OA fraction of about 60%, suggesting large SO$_2$ emissions from BB in Asia followed by oxidation during transport. In support of this interpretation, the sulfate + potential sulfate mass fractions were as high as 15%–20% in the BB plumes of the flaming phase in North America.

In all types of BB plumes, the BC mass (volume) fractions of the total PM$_1$(LSP) were 1%–5% (2%–4%), while the BC volume fractions were about 30% for the particles containing BC, consistent with the fact that most particles

Figure 20. Time series of the ratios of the total shell volume of BC at time $t$ ($V(t) = \int \pi [D_p (t)]^2 dD_p / 6$) to the initial $V(t=0)$ shown together with the initial LNSD parameters. (a) Calculations for the conditions of $CMD_{BC}$ (shell) = 160 nm. The lines with triangles, squares, and circles are the results for initial $CMD_{LSP} = 250, 230, \text{ and } 200 \text{ nm}$, respectively. (b) Same as Figure 20a but for $CMD_{BC}$ (shell) = 120 nm. (c and d) Same as Figure 20a but for initial $N_{LSP}$ of 1500 and 2500 cm$^{-3}$, respectively.
did not contain BC. The $\Delta M_{BC}/\Delta V_{SC}$ ratios were much smaller than those for BC particles emitted from fossil fuel combustion.

[97] We derived the average emission ratios of $M_{BC}/CO = 2.3 \pm 2.2$ ng m$^{-3}$ppbv and $M_{BC}/CO_2 = 180 \pm 269$ ng m$^{-3}$/ppmv for BB in North America. For BB plumes from Asia, the average ratios were $M_{BC}/CO = 8.5 \pm 5.4$ ng m$^{-3}$/ppbv and $M_{BC}/CO_2 = 129 \pm 67$ ng m$^{-3}$/ppmv. The much higher $M_{BC}/CO$ ratios for the Asian plumes were likely due to higher MCE. The $M_{BC}/CO$ ratios in North America were lower than the values compiled by Andreae and Merlet [2001] by a factor of 2–4, although the ratios for the Asian plumes were similar. The $M_{BC}/CO_2$ emission ratios for North America and Asia were 4–5 times lower than their estimates.

[98] We related the MCE to the emission ratios between $\Delta M_{BC}$, $\Delta V_{SC}$, $\Delta CO_2$, and $\Delta CO$ for BB in North America and Asia. The $\Delta M_{BC}/\Delta CO_2$ ratios decreased with the increase in MCE, whereas the $\Delta M_{BC}/\Delta CO$ ratios increased with the increase in MCE. The dependence of the ratios on MCE is useful in improving emissions estimate of aerosols from BB and estimating the transport efficiency of aerosols by using tracers, especially CO.

[99] The shapes of the size distributions of BC were similar for plumes originating from BB in North America and Asia. The distribution functions were represented to a good approximation by lognormal functions. The count median diameters (CMDs) of BC were 136–141 nm (±6–8 nm), with standard deviations of 1.32–1.36 (±0.01–0.04) and little dependence on the phase of combustion.

[100] These BC particles were thickly coated by organics, with $D_p/D_{BC}$ ratios of about 1.4 (0.6 < MCE < 0.8) 1–2 h after emission. In the smoldering phase, the $D_p/D_{BC}$ ratios increased from 1.35 to 1.45 within about 12 h, with a corresponding increase in the volume of nonrefractory species on BC by 15%–40%. In aged flaming phase plumes, the increase in $D_p/D_{BC}$ ratios continued for a few days, leading to an increase of the volume of coating materials by a factor of 2. A potential explanation is that OA was being redistributed via the gas phase [e.g., Marcolli et al., 2004] by evaporating from particles that did not contain BC and recondensing on BC-containing particles.

[101] We assessed the effect of the Brownian coagulation of LSP with BC on the evolution of the BC mixing state. For the observed conditions, coagulation can increase $D_p/D_{BC}$ ratios by factors of 1.2–1.5 in 1–2 days, depending on various parameters, influencing the rate of coagulation. Previous studies considered only condensation process to explain the evolution of the BC mixing state. For the first time, we suggest that coagulation of LSP with BC can effectively increase $D_p/D_{BC}$ ratios, as well as condensation.

References


Ferek, R. J., et al. (2010), Hyperscopic properties of Amazonian biomass burning and European background HULIS and investigation of their effects on surface tension with two models linking H-TDMA to CCNC data, *Atmos. Chem. Phys.*, 10, 10187–10197, doi:10.5194/acp-10-10187-2010.


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