- 1 High-pressure study of dravite tourmaline: Insights into the accommodating nature of the tourmaline structure
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10**Abstract**

The high-pressure behavior of dravite tourmaline [Na(Mg₃)Al₆(Si₆O₁₈)

 $12(BO_3)_3(OH)_3(OH)]$ has been studied using luminescence spectroscopy and synchrotron based 13single-crystal diffraction up to ~65 and ~23.6 GPa, respectively. Two emission bands associated 14with Cr^{3+}/V^{2+} substitution are constant in energy up to ~9.0 GPa, and shift to longer wavelength 15at higher pressures, suggesting that a change in compressional mechanism could occur at this 16pressure. However, single-crystal diffraction data do not show evidence of any structural 17changes, nor any resolvable changes in the deformation mechanism at ~9.0 GPa. Near 15 GPa, a 18splitting of one of the emission bands is observed, suggesting that a phase transition occurs at 19this pressure and that two unique octahedral sites are present in the high-pressure phase. 20Hysteresis is not observed on decompression, which indicates that this is a second order 21transition, and the high-pressure structure appears to be metastable up to ~65 GPa. Single-22crystal diffraction measurements show that a phase transition from rhombohedral *R3m* to 23rhombohedral *R3* occurs at pressures near 15.4 GPa. The high-pressure phase is characterized

24by a distorted Si_6O_{18} ring (e.g. the Si-Si-Si angles deviate from 120°), and the Si, Al, O6, O7, and 25O8 sites of the low-pressure phase split, implying that the high-pressure phase of tourmaline is 26a higher entropy phase. The large X-site exerts the primary control on compressibility, and the 27substitution of larger cations into this site will likely lower the pressure at which this transition 28occurs. Dravite tourmaline shows anisotropic compression with the *c*-axis being more 29compressible thant the *a*-axis. The pressure volume data up to ~15.4 GPa were fit with 2nd and 303rd order Birch-Murnaghan equations of state. We obtain a bulk modulus, K_0 = 109.0(2.9) GPa, 31and a pressure derivative K_0 ' = 4.8(8) GPa, and with the pressure derivative set to 4, a bulk 32modulus of 112.2(9) GPa is derived.

33**Key Words:** dravite; High-pressure; Single-crystal diffraction, luminescence, phase transitions 34**Introduction**

Tourmaline is a supergroup of minerals that is chemically complex and is divided into 36different varieties based on chemical composition (Hawthorne and Henry 1999; Henry et al. 372011). This group of minerals is so complex and chemically variable that an entire subcommittee 38on tourmaline nomenclature exists. These minerals are categorized using the generalized 39tourmaline structural formula $XY_3Z_6(T_6O_{18})(BO_3)_3V_3W$ where X = Na, Ca, K; Y = Fe, Mg, Mn, Al, Li; 40Z = Al, Fe, Mg, Cr; T = Si, Al; B = B; V = (OH), O; and W = (OH), O, F (Hawthorne and Henry 1999; 41Henry et al. 2011). From a compositional perspective, most of the variability occurs in the X, Y, 42and Z sites; more information on the classification of tourmaline supergroup minerals can be 43found in Hawthorne and Henry (1999) and Henry et al. (2011). Tourmaline is also highly sought 44after as a gemstone due to its high hardness and wide range of colors.

- Tourmaline typically crystallizes in the rhombohedral crystal system with R3m symmetry 46(Hawthorne and Henry 1999; Henry et al. 2011), although other crystal systems have been 47reported (Akizuri et al. 2001; Shtukenberg et al. 2007). It is typically considered a ring-silicate 48with six corner-shared TO₄ tetrahedral sites that form a hexagonal ring, a nine-fold X-site, two 49six-fold octahedral Y- and Z-sites, and a trigonal planar B-site (Figure 1). Although the T-sites are 50dominantly occupied by Si, B and Li substitution into them has been reported (Rosenberg and 51Foit 1979; Hughes et al. 2000). Interestingly, the first reported structure of tourmaline 52presented the Si₆O₁₈ ring as ditrigonal in symmetry (Hamburger and Buerger 1948). However, it 53was later shown to be hexagonal in symmetry (Ito and Sadanaga 1951), and all subsequent 54studies agree that the Si₆O₁₈ ring is hexagonal at ambient conditions.
- Tourmaline is the most widespread borosilicate mineral in natural rocks and the 56dominant carrier of B in many rock types (Henry and Dutrow, 1996), and is found in subduction 57zone environments (Nakano and Nakamura 2001; Bebout and Nakamura 2003; Ota et al. 2008). 58B isotope systematics in tourmalines have been used to elucidate crust-mantle recycling 59processes in subduction zones (Nakano and Nakamura, 2001). Petrologic studies have shown 60that dravite tourmaline breaks down at ~7.0 GPa and ~900 °C (Werding and Schreyer 1996) and 61~3-5 GPa and ~1000 °C (Krosse 1995) and that breakdown pressures and temperatures are 62dependent on its precise composition (Ota et al. 2008). The decomposition products associated 63with tourmaline breakdown are complex, but one decomposition product, phengite, is 64important from a deep Earth boron and water cycling perspective; phengite could transport B to 65depths >300 km (Domanik and Holloway 1996; Marschall et al. 2007). Higher pressure studies 66(e.g. > 7.0 GPa) on tourmaline are scarce, and only two studies appear to have been reported: Li

67et al., (2004) and Xu et al., (2016). Li et al. (2004) conducted high-pressure powder diffraction 68experiments on schorl tourmaline [Na(Fe₃)Al₆(Si₆O₁₈)(BO₃)₃(OH)₃(OH)] up to 27.8 GPa, with no 69phase transitions being reported. Xu et al. (2016) reported high-pressure and temperature 70powder diffraction experiments on uvite tourmaline [Ca(Mg₃)MgAl₅(Si₆O₁₈)(BO₃)₃(OH)₃(OH)] up 71to ~18.4 GPa and 723 K and again, no phase transitions were observed. These powder 72diffraction studies provide insights into how the unit cell responds to compression, but they do 73not elucidate the positional changes of the individual atoms in the unit cell (Li et al. 2004; Xu et 74al. 2016).

- This is in contrast to other ring-silicates such as cordierite $[(Mg, Fe)_2Al_3(AlSi_5O_{18})]$ and 76beryl $[Be_3Al_2(Si_6O_{18})]$ which each show extensive high-pressure polymorphism (Prencipe et al. 772011; Miletich et al. 2014; Scheidl et al. 2014; Finkelstein et al. 2015; O'Bannon and Williams 782016b). Both cordierite and beryl undergo high-pressure phase transitions that involve 79distortion of the Si_6O_{18} ring. The main structural difference between tourmaline and these other 80ring silicates is that tourmaline does not have open channels in its structure like cordierite and 81beryl. Thus, it is possible that the denser-packed configuration of the Si_6O_{18} rings in tourmaline 82produces a structure that is more stable under compression than open channel ring silicates like 83cordierite or beryl.
- Tourmaline is also both pyroelectric and piezoelectric, and has been used in industrial 85applications such as pressure gauges and air and water purifiers (Frondel 1948; Lameiras et al. 862010). Growing synthetic tourmalines >1mm in diameter has not yet been achieved, and thus 87industrial applications require natural tourmalines (Shekhar Pandey and Schreuer 2012). Thus, 88despite its widespread use in industrial applications and as a geochemical probe, tourmaline's

89high-pressure stability/polymorphism is poorly understood. Hence, our experiments are 90oriented towards understanding the trace element behavior and the bulk structural response of 91dravite tourmaline to compression by combining the results of luminescence and single-crystal 92diffraction studies to provide an integrated picture of structural changes in dravite tourmaline 93under compression.

94**Methods**

952.1 Sample characterization

96 The natural light green gem quality tourmaline sample used in this study is from the 97Lengenbach Quarry, Switzerland. We characterized the sample with single-crystal X-ray 98diffraction, Raman, and luminescence spectroscopy, which all agree well with previous studies 99(Hawthorne et al. 1993; Gasharova et al. 1997; Gaft et al. 2005). Ambient pressure single-crystal 100X-ray measurements were conducted on Beamline 11.3.1 at the Advanced Light Source (ALS) at 101Lawrence Berkeley National Lab in Berkeley, CA. We modeled our tourmaline as an endmember 102dravite and our ambient structure refines with an R₁ of 1.85% when refined anisotropically and 1032.99% when refined isotropically. Its crystal system, spacegroup and unit cell parameters are: 104rhombohedral R3m, a-axis 15.9370(6) Å, c-axis 7.1749(3) Å, and volume 1578.19(14) Å³, 105respectively. For the ambient structure refinement, we fixed the O1-H1 and O3-H3 bond lengths 106to 0.958(8) and 0.972(2) Å, respectively. These bond lengths were determined from previous 107neutron diffraction results on tourmaline (Gatta et al. 2014). Trace element concentrations of 108dravite tourmaline were measured with a PhotonMachines Analyte 193H, which is a 193-nm 109ArF excimer laser system coupled with a ThermoScientific ElementXR single-collect or magnetic 110sector ICP-MS. The instrument was calibrated with a SRM 610 trace element glass from NIST,

111and Al was used as the internal standard. The following trace elements were measured (in 112ppm): Cr 215 (\pm 19), V 336 (\pm 32), Li 18 (\pm 4.0), Ti 474 (\pm 53), Ca 2183 (\pm 597), and K, Fe, and Mn are 113below detection limits. The Subcommittee on Tourmaline Nomenclature (STN) of the IMA 114Commission on New Minerals, Nomenclature and Classification recommends nomenclature of 115tourmaline supergroup minerals (Novak et al., 2009; Henry et al., 2011; Hawthorne and Dirlam, 1162011). Using our single-crystal X-ray results combined with our trace element results, and 117following STN recommendations, we determined that this tourmaline is of the dravite variety. 118Dravite variety tourmaline has an ideal chemical formula of [Na(Mg₃)Al₆(Si₆O₁₈) 119(BO₃)₃(OH)₃(OH)]. Hence, the structures were refined isotropically and modeled as 120endmember dravite tourmaline: this allowed direct comparisons of the results obtained at each 121pressure.

1222.2 Luminescence spectroscopy

Luminescence spectra were collected from 650-800 nm (15380-12500 cm⁻¹) with a 124Horiba LabRAM HR Evolution Raman spectrometer with a spectrometer focal length of 800 mm. 125Spectra were collected to a pressure of ~65 GPa and on decompression at 300 K using an 126excitation wavelength of 532 nm. A symmetric type DAC with 200 µm type-Ia diamonds was 127used in all high-pressure luminescence experiments, and single crystals of the sample and at 128least two ruby spheres (Chervin et al. 2001) were loaded into the sample compartment. Either 1294:1 methanol:ethanol mixtures or Ne was used as the pressure medium depending on the 130target pressure range of the experiment. An Olympus BXFM-ILHS microscope with a 50x long 131working distance objective was used to focus the laser beam onto the sample. An 1800 132lines/mm grating with a corresponding spectral resolution of ~1 cm⁻¹ (or, equivalently, ~0.05 nm)

133was utilized. Combinations of Gaussian and Lorentzian functions were fit to the luminescence 134spectra with Horiba Labspec6 software.

1352.3 High-pressure single crystal diffraction

High-pressure single-crystal measurements were carried out using a (DESY) BX90 (Kantor 137et al. 2012) type diamond anvil cell (DAC) equipped with type la 500 μm culet Boehler-Almax 138geometry diamond anvils with 85° angular access (Boehler and De Hantsetters 2004). Rhenium 139gaskets with a 300 μm hole were used to contain the samples. Gasket holes were cut using an 1400xford Instruments laser mill at the Advanced Light Source (ALS), Berkeley, CA. Single crystals of 141dravite tourmaline and two ruby spheres were loaded into the sample compartment with neon 142as the pressure medium. The standard ruby fluorescence gauge was used to determine pressure 143(Mao et al. 1986). Neon was loaded using the gas loading apparatus at the ALS. Neon remains 144hydrostatic up to ~15.0 GPa and above that pressure, the pressure gradients remain small (Klotz 145et al. 2009).

High-pressure single-crystal X-ray diffraction measurements were conducted at Beamline 14712.2.2 at the ALS. Measurements were taken at various pressures up to ~23.6 GPa at room 148temperature. The DAC was mounted on a Huber sample stage, and shutterless single-crystal 149diffraction data were collected on a Perkin Elmer amorphous silicon detector using synchrotron 150radiation monochromated by silicon(111) to a wavelength of 0.49594 Å (25 keV). Distance and 151wavelength calibrations were done using a NIST single crystal ruby diffraction standard. Phi 152scans were employed to measure across both diamonds with a 0.25° image width.

1532.4 Data processing

154 Image masks, to avoid integrating signal from detector regions obscured by the DAC,

155were created using the program ECLIPSE (Parsons 2010). The data were integrated using the 156program SAINT v8.34A. A multi-scan correction for absorption was applied using the program 157SADABS-2014/11. Structures were solved by dual space methods (SHELXS-97/ SHELXT) and 158refined by full-matrix least-squares on F² (SHELXL-2014) (Sheldrick 2008) using the graphical 159user interface *ShelXle* (Hübschle et al. 2011). All atoms in the room pressure structure and the 160high-pressure structures were refined isotropically. Because of their small scattering cross-161sections, H atoms were not refined in the high-pressure structures. Complete crystallographic 162information files (CIFs) for each structure can be found in supplementary material, as well as a 163discussion on the selection of our unit cells.

164Results and discussion

165Ambient pressure luminescence spectra

The luminescence spectrum of dravite is characteristic of d^3 ions (e.g. Cr^{3+} , V^{2+}) in an 167intermediate strength crystal field (Tanabe and Sugano 1954). Both sharp and broad emission 168bands are observed at room temperature and 77 K at ambient pressures (Figure 2). Considering 169the similar concentrations of Cr and V in our sample (e.g. Cr: 215 ±19 ppm, V: 336 ±32 ppm), it 170is reasonable to assume that our steady state luminescence spectra contain overlapping bands 171from the emission of both Cr and V. This is commonly observed in emerald, which is the Cr and 172V rich variety of beryl (Ollier et al. 2015; O'Bannon and Williams 2016b). The intense and 173relatively sharp emission bands are commonly referred to as the R-lines and are associated with 174the spin-forbidden 2E - 4A_2 transition, with the 4A_2 state being the ground state and the 2E state 175being split in non-cubic environments (Syassen 2008), and the broad band emission is assigned 176to the spin-allowed 4T_2 - 4A_2 transition (Sugano and Tanabe 1958; Burns 1993; Gaft et al. 2005).

177 At 300 K, three bands can be fit under the sharp emission bands and two can be fit 178under the broad portion of the spectrum at room temperature. At 77 K, the broad thermally 179populated ⁴T₂-associated bands are not entirely quenched, and four bands can be fit under the 180²E sharp emission bands. A detailed deconvolution of these sharp line emission bands at both 181room temperature and 77 K is given in the supplementary material (Figure S1). Dravite has two 182crystallographically unique octahedral sites, Al and Mg, and emission bands from both sites 183would likely be observed in steady state luminescence spectra. The simplest way for Cr³⁺ and V²⁺ 184to substitute into dravite is through isovalent substitutions (e.g. Cr³⁺ into Al³⁺ site and V²⁺ into 185Mg²⁺). The four deconvolved deconvoluted components at 77 K thus can be attributed to two 186sets of R-lines, one set from Cr³⁺ and one from V²⁺. The Al site in dravite is less distorted than the 187Mg site (Hawthorne et al. 1993, and this study), which would be expected to give rise to a 188narrower R-line splitting. Given the likely relative intensities of R₁ and R₂ and the anticipated 189relative splittings, we propose the following assignments: the Cr³⁺ R₁ band is likely located at 190683.7 nm, with its R_2 band at 681.4 nm, and the V^{2+} R_1 band is at 686.4 nm and its associated R_2 191band at 679.8 nm. This results in an R-line splitting of ~50 cm⁻¹ for Cr³⁺ and ~143 cm⁻¹ for V²⁺. For 192comparison, ruby [Al₂O₃:Cr] has an Al-site that is close to an ideal octahedron with a slight 193trigonal distortion, and has an R-line splitting of ~29 cm⁻¹ (Syassen 2008), while zoisite 194[Ca₂Al₃Si₃O₁₂(OH)] which has a highly distorted Al-site has an R-line splitting of ~340 cm⁻¹ 195(Koziarska et al. 1994). The distortions of the Al and Mg sites in dravite fall intermediate 196between the distortions of the Al sites in ruby and zoisite, so these R-line splittings are plausible. 197To confirm the assignments of these emission bands, luminescence lifetime measurements or 198variably substituted samples would be required.

199High-pressure luminescence spectra

High-pressure luminescence spectra are shown in Figure 3. Under compression the 201intensity of the R-lines decreases: this is likely primarily associated with the transition metal 202absorption bands migrating away from our excitation wavelength of 532 nm. Moreover, the 203disappearance of the broad 4T_2 -associated band shows that a transition from an intermediate to 204strong crystal field occurs in the first few GPa of compression, which is consistent with previous 205observation in other Cr-bearing oxides (Dolan et al. 1986; de Viry et al. 1987; Hommerich and 206Bray 1995; Grinberg and Suchocki 2007; O'Bannon and Williams 2016b). As at 300 K, it is 207difficult to fit four bands under the 2E region under compression, so we fit the spectra with 208three bands: based on our 77 K assignments, these are associated with the R_1 and R_2 peaks of 209Cr ${}^{3+}$ and the R_1 peak of V^{2+} . Near 6.0 GPa, the primary emission feature fit by three bands clearly 210separates into individual peaks (Figure 3).

The pressure shift of two of the original three bands that were fit at ambient pressure 212and temperature shift negligibly--essentially by 0 nm/GPa up to ~9 GPa, while the other band 213shifts by 0.50 ±0.02 nm/GPa (-8.8 (±0.3) cm⁻¹/GPa) up to the same pressure (Figure 4a). The 214observation of effectively stationary emission bands under compression is without precedent 215among Cr³⁺ emission in oxides (e.g., Dolan et al. 1986; Mao et al. 1986; de Viry et al. 1987; 216Hommerich and Bray 1995; Grinberg and Suchocki 2007; O'Bannon and Williams 2016b). Above 217~9.0 GPa, the initially stationary bands begin to shift positively in wavelength, towards lower 218energies (Figure 4b). The two bands that shift by 0 nm/GPa likely are associated with one 219octahedral site while the other band is produced by substitution in the other octahedral site.

220This is in accord with our assignment of the former two bands to emission from Cr³⁺ substituting 221into the Al³⁺ site.

222 Four separate experiments were conducted in the ≤ 10.5 GPa range to assess the 223reproducibility of the unusual negligible pressure shift. It is completely reproducible, and no 224hysteresis is observed on decompression from ~10.5 GPa. The ~0 nm/GPa pressure shift 225 suggests one of three possibilities. First, the crystallographic site that this band is emitting from 226may not be changing in volume under compression. Second, there may be a complex trade-off 227between electron density and volumetric change within the site. Third, an admixture of 228electronic states, rather than a pure ²E -> ⁴A₂ transition, may generate the observed sharp-line 229emission. As described in the single-crystal diffraction section below, there is no structural 230"smoking gun" for this unusual pressure shift. Admixture of states (probably with the ⁴T₂ state) 231is a possibility; however, the troubling aspect of this explanation is that a pressure-induced shift 232in admixture must perfectly balance with the pressure shifts of the two states to produce a net 233zero pressure shift. Stated another way, the separation of the two states, their resulting 234admixture, and their pressure shifts must collectively sum to zero for ~9.0 GPa of compression. 235This scenario is quite unlikely because of the dramatically different pressure-dependences of 236the ⁴T₂ and ²E states (e.g., O'Bannon and Williams 2016a), but could be testable with high-237pressure lifetime measurements. A trade-off between changes in electron density and 238 volumetric compaction is possible, but would require detailed insights into the bonding 239environment/electron density within the octahedral site under compression.

The splitting between these two bands also does not change on compression, suggesting 241that the site distortion is unchanged, as well. The change in slope at 9 GPa is, as described

242below, likely not an indicator of a phase transition, but may represent a change in the
243compression mechanism. Moreover, the other observed emission band shifts positively in
244wavelength, consistent with the volume of this site decreasing with increasing pressure. Since
245we cannot deconvoluteve the expected R₂ component of this band in this pressure range, it is
246difficult to know from the emission spectra if this site becomes more or less distorted. However,
247the peakwidth of the positively shifting band slightly narrows under compression to ~9.0 GPa,
248which is an indicator that its site becomes less distorted under compression.

In the pressure range above 10 GPa, peak splitting is observed near 15 GPa, with new 250emission bands observed; at pressures above 45 GPa, non-linear pressure shifts are observed 251(Figure 3, 4b). No hysteresis is observed in the emission on decompression from 65 GPa, and the 252Raman spectrum of the decompressed sample is identical to the starting material. Peak splitting 253of emission bands under compression has been shown to be an indicator of phase transitions 254(O'Bannon and Williams 2016a,b). This is particularly anticipated when multiple octahedral 255environments are produced by a decrease in symmetry associated with a phase transition. For 256example, in the ring silicate beryl, the Cr³+ R-lines split into two sets of R-lines near 14.7 GPa 257(O'Bannon and Williams, 2016b) which agrees well with a theoretically predicted ~15 GPa soft-258mode transition (Prencipe et al. 2011), as well as high-pressure single crystal diffraction studies 259which show that the Al-site splits into two crystallographically unique sites near 15 GPa (M. 260Merlini, personal communication). The peak splitting observed in dravite tourmaline under 261compression provides clear evidence that one of the octahedral sites splits into two unique sites 262(Figure 5).

263 The additional new bands at longer wavelength that appear above 10 GPa (Figures 3, 4b) 264are likely neighbor lines that are due to paired Cr-Cr emission (e.g., O'Bannon and Williams, 2652017); vibronic lines may also lie at these wavelengths, but these typically do not dramatically 266exceed the zero-phonon lines in amplitude. The appearance of these new bands, and the 267progressive increase in their intensity with pressure may reflect: (1) the intensity of the broad 268⁴T₂ transition has decreased substantially by this pressure and alternate mechanisms for 269emission are enhanced; (2) a change in the nearest neighbor configuration due to the splitting 270of one of the octahedral sites into two unique sites that produces configurations that favor 271energy transfer into and/or the magnetic interactions that generate emission from pairs; and/or 272(3) a pressure-induced shift in absorption that results in preferential excitation of pair-line 273emissions (e.g., O'Bannon and Williams 2017). Within the tourmaline structure, there are a 274wide range of possible prospective pair-wise magnetic interactions that could occur between 275substituents in the octahedral sites through an intervening anion, and it is well-known that the 276strength of such exchanges depends strongly on both the inter-ion distance and on the angle 277(e.g., McCarthy and Gudel 1988). Indeed, such pair-wise interactions between transition metals 278have long been recognized to be important within the optical spectrum of tourmalines (Mattson 279and Rossman 1987). Therefore, while the exact origin of the new bands and the shift in intensity 280above 10 GPa within the emission spectra remain unclear, it is probable that they can be 281attributed to pressure-enhanced pair-line emissions.

Thus, the major effects of compression on the luminescence spectrum are the change in 283pressure shift near 9.0 GPa and the peak splitting and increased intensities of new peaks 284observed at ~15 GPa. The lack of hysteresis on decompression from ~65 GPa indicates that any

285phase transition associated with these changes is second order. We do not believe that the 286change observed at 9.0 GPa is produced by a pressure-induced structural phase transition, but 287 is likely generated by either a change in compression mechanism or a progressive change in the 288electron density associated with the site. The changes observed at ~15 GPa are, however, 289consistent with a phase transition that involves a subtle distortion of the overall structure of 290dravite tourmaline that splits the Al-site into two unique sites. By analogy to high-pressure 291phase transitions in cordierite and beryl (Prencipe et al. 2011; Miletich et al. 2014; Finkelstein et 292al. 2015; O'Bannon and Williams 2016b), the structural change likely involves a modest 293distortion of the Si₆O₁₈ ring away from hexagonal symmetry. Importantly, the overall topology of 294the structure appears to be preserved, and a lowering of symmetry from rhombohedral is 295unlikely. The presence of such a modest structural change implies that dravite tourmaline's 296metastability under room temperature compression to ~65 GPa is generated by the topology of 297the structure that stabilizes the Si₆O₁₈ ring at high compressions. Moreover, a structural 298transition that involves a modest distortion of the Si₆O₁₈ ring with no change in crystal system 299would be difficult to constrain with powder diffraction techniques (e.g., Li et al. 2004; Xu et al. 3002016).

301High-pressure single-crystal diffraction

To characterize the changes observed using luminescence spectroscopy, single crystal X-303ray diffraction measurements were done at various pressures up to ~23.6 GPa. Dravite 304tourmaline indexes to a rhombohedral unit cell across the pressure range of the X-ray 305measurements. Structures solved as rhombohedral R3m up to ~15.4 GPa, and above this 306pressure the space group changed to R3. In the following sections, we discuss the equation of

307state (EoS) for the low pressure phase of dravite tourmaline, and the high-pressure crystal 308structure of this material.

309Dravite tourmaline EoS

310 Dravite tourmaline shows highly anisotropic compression (Figure 6), with the c-axis 311being substantially more compressible than the *a*-axis, which is in good agreement with the 312previous study of uvite tourmaline reported by Xu et al. (2016). Interestingly, Liu et al. (2004) 313report that the c-axis of schorl tourmaline is less compressible than the a-axis, which is opposite 314of what is observed in dravite and uvite tourmaline. But, the lattice parameters and unit cell 315 volumes reported by Liu et al. (2004) are quite scattered. For clarity, we do not include these 316 results in Figure 6, but they are included in supplementary Figure S2. The reason for the scatter 317in their data, and the difference in compressibility of the a- and c-axes is not clear, but could be 318a result of non-hydrostatic conditions of methanol:ethanol mixtures above ~10.5 GPa. We fit 319our pressure volume data with both 2nd and 3rd order Birch-Murnaghan equations of state (Table 3201) using the EOSFit7GUI (Angel et al. 2014), and our results are shown in Table 2. An F-f plot 321based on the Birch-Murnaghan EoS fit of the volume data is shown in Figure S3. Our results 322indicate that dravite tourmaline is more compressible than uvite tourmaline. For comparison, 323the uvite sample measured by Xu et al. (2016) is not an endmember, but has ~10 % Na in the X-324site, and pure uvite is likely less compressible than their measurements. The relatively large 325 value for K_0 reported for schorl tourmaline by Li et al. (2004) does not appear compatible with 326our results, nor with those of Xu et al. (2016).

327High-pressure crystal structures

Structures were solved at various high pressures, and results of the structure

329refinements are shown in Table 3. The site volumes as a function of pressure show that the Na

330site is the most compressible (Figure 7). The Si site is the least compressible, while the two

331octahedral sites behave essentially identically to one another, and are intermediate in their

332compressibility between the Si and Na sites. Clearly, neither octahedral site volume is constant

333under compression. So, the unusual observation of a negligible pressure shift of two of the

334luminescence bands cannot be explained via a site volume argument alone. A more nuanced

335explanation for the shift of the energy levels, involving a shift in covalency/electron density of

336the site that counterbalances the volumetric compaction, appears necessary. The anisotropic

337compression of dravite tourmaline can be straightforwardly explained by the location of the Na

338site in the dravite structure and its large volume change under compression (Figure 7).

339Speculatively, the large volume change in the Na site could produce a pressure-induced

340reduction in the electron density around the Al-sites, such that their electron density remains

We calculate the distortion parameters for the octahedral and tetrahedral sites as a 343function of pressure (Supplementary Figure S4). Interestingly, the quadratic elongation (QE) and 344angle variance (AV) of the Mg site both decrease under compression, indicating that this site 345becomes less distorted under compression: this is in accord with our inference based on the 346width of the emission peak that we have assigned to this site. The Al site distortion changes by 347very little, and the Si site becomes slightly more distorted under compression. The Al-site 348distortion change is negligible, which also provides confirmation for our assignment of the two

31 16

349emission bands with invariant positions that show no change in their splitting up to ~9.0 GPa to 350this site.

351 The X-site in tourmaline (the Na site in dravite) is typically considered a 9-fold 352coordinated site. Among the bond lengths associated with this site as a function of pressure 353(Figure 8), there are two unique bonds between the Na and the Si₆O₁₈ ring (O5 and O4), and 354there is one unique bond between the Na and the Mg octahedra (O2). There is also a potential 35510th coordinating oxygen atom (O1), and the distance to this oxygen atom is also shown in 356Figure 8. The O1 atom is, however, one of the two proton sites in tourmaline, although the H1 357site has a very low reported occupancy (Gatta et al. 2014). F substitution also occurs only at the 358O1 site, and this site is influenced by cation occupancy of the X and Y sites (Henry and Dutrow 3592011). Hence, O1 is not simply a simple oxygen site that appears to become bonded to the Na 360atom at high-pressure: rather, it is a site that is partially occupied by O, OH, and F. Interestingly, 361the two unique Na-O bonds to the Si₆O₁₈ ring show different behaviors. Under compression to 362~15.4 GPa, the Na-O4 bonds change by ~0.05 Å while the Na-O5 bonds change by ~0.25 Å. Over 363the same pressure range, the Na-O2 bonds change by ~0.11 Å. In this sense, the Na site appears 364to undergo the bulk of its compaction through six neighbors: three bonds into the Si₆O₁₈ ring 365and three bonds into the Mg octahedra. This pseudo-six-fold behavior, with compacting 366connections to both the Mg polyhedra and the Si-ring, also sets up a natural way to distort the 367Si₆O₁₈ ring away from hexagonal symmetry.

The B-O distances are less well-resolved, and change very little under compression but 369show a trend towards being more equidistant up to ~15.4 GPa (Supplementary Figure S5). In the 370high-pressure phase, B is surrounded by three crystallographically unique oxygen atoms (O8)

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371splits into O8a and O8b). There is no indication that the BO $_3$ groups deviate from a planar 372configuration up to at least 23.6 GPa, which is expected given the spacegroups of the low- and 373high-pressure phases.

374Compression mechanism

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The compression mechanism of tourmaline involves a complex interplay of bond length 377shortening and oxygen atom displacement/rotation relative to its neighbors. Under 378compression the Y-site (MgO₆) volume decreases through Mg-O bond shortening and the 379distortion parameters of this site also decrease. As the volume of the Y-site decreases, the Si₆O₁₈ 380ring becomes more puckered due to the displacement of the O6 atom towards the threefold 381rotation axis (Figure 9a). This increase in ring puckering has been observed in tourmaline when 382smaller cations are substituted into the Y-site. Bosi and Lucchesi (2007) report that the 383configuration of the six-membered ring is strongly affected by the average Y-O distance <Y-O> 384and, as <Y-O> decreases due to the substitution of smaller cations, the tetrahedra rotate around 385the O4-O5 edge and the O6 atom is displaced towards the threefold axis which produces 386puckering of the tetrahedral ring. Thus, chemical substitution of smaller cations into the Y-site 387and pressure appear to have the same effect on ring puckering.

As ring puckering increases under compression the O7 atom rotates towards the Z-site $389(AlO_6)$, which shortens the Al-O7 bond, and compresses the Z-site. This displacement of the O7 390atom contributes to the increase in observed ring crimping under compression (Figure 9b), and 391the increase in ring crimping changes the O7-O8 edge distance. Crimping is defined as $\Delta Z = [_2O5-392(_zO4+_zO7)/2]*c$, where z is the coordinate of the oxygen atoms and c is the length of the c-axis 393(Gorskaya et al. 1982). From a chemical substitution perspective, it has been shown that

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394substitution into the X-site primarily affects the crimping and ditrigonality of the ring and 395distortion of the tetrahedra, while substitution into the Y-site primarily affects ring puckering 396(Foit, 1989). Ditrigonality increases across the pressure range of this single-crystal study (Figure 3979c). It is defined as $\delta = (r_1 - r_s)/r_s$, where r_1 and r_s are the distances from O4 and O5, respectively, 398to the threefold axis (Barton 1969). Under compression, the Na site moves towards the ring, 399which also contributes to the increase in ring crimping and ditrigonality. Moreover, the overall 400distortion of the Z-site is primarily controlled by the size of the Y-site cation (Foit 1989). 401Therefore, it is not surprising that the distortion parameters of the Z-site show very little change 402under compression. Above ~15.4 GPa, tetrahedral rotation and ring puckering continue to 403increase while ring crimping decreases above the transition and then remains almost constant 404to the highest pressure of this study.

405High-pressure phase transition

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At pressures above 15.4 GPa, the structure of dravite tourmaline is solved as 408rhombohedral R3. The only difference in symmetry associated with this transition is the loss of a 409mirror plane in the high-pressure structure. The net result is that the Si, Al, O6, O7, and O8 sites 410all split into unique sites (Figure 10), which implies that the high-pressure phase of tourmaline is 411likely a higher entropy phase. Since the 3-fold rotation axis is preserved across this phase 412transition there is no change in the Mg, Na, or B sites. Importantly, this subtle crystallographic 413change is completely consistent with our luminescence observations of a second order phase 414transition occurring near \sim 15.4 GPa. Additionally, our assignment of the emission bands with an 415initial \sim 0 nm/GPa pressure shift to the Al site is a robust assignment. Interestingly, the slight 416distortion of the Si₆O₁₈ ring preserves the overall topology of the tourmaline structure, while

417providing a new deformation mechanism for the high-pressure phase. A plot of the Si-Si-Si
418angles shows that this angle deviates away from 120° above ~15.4 GPa and increases with
419increasing pressure up to ~23.6 GPa (Figure 11a). There is also a discontinuous shift in
420tetrahedral rotation, as well as a change in the pressure dependence of ring puckering and
421crimping above ~15.4 GPa (Figures 9b, 11). The deformation mechanism of the high-pressure
422phase primarily involves tetrahedral rotation and changes in the Si-Si-Si angles while ring
423crimping and puckering play less of a role. These subtle changes in ring distortion demonstrate
424the subtlety of this phase transition, and why it would be very difficult to resolve this type of
425phase transition with high-pressure powder diffraction techniques. Our luminescence data
426show that this high-pressure phase remains metastable under room temperature compression
427up to at least 65 GPa.

Speculatively, the high-pressure transition appears to be driven by steric/geometric 429effects caused by the decrease in volume of the Y-site and displacement of the X-site towards 430the Si_6O_{18} ring. As the Y-site volume decreases, the O6 atom is displaced towards the threefold 431axis causing ring puckering, which displaces the O7 atom towards the Z-site (AI), causing the AI-432O7 distance to shorten which increases ring crimping. This causes a change in the tehtrahedral 433rotation of the SiO_4 tetrahedra, which breaks the mirror plane symmetry for O6 and O7. This 434change in rotation induces an additional twist on the MgO_6 and AIO_6 octahedra, which in turn 435tilts the BO_3 plane out of the plane normal to the rotation axis, and breaks the mirror symmetry 436of the O8 atom.

437Na site coordination change?

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The possible coordination change of the Na site merits discussion, as the topology of the 439tourmaline structure shows that the incipient Na1-O1 bond (Figure 8) provides the only 440straightforward means to further polymerize this structure. The distance to the potential 10th 441oxygen decreases by ~0.43 Å to 15.4 GPa, and the pressure dependence of its distance is 442constant above this pressure. Multiple phenomena could be at work here in changing the 443behavior of the Na-O bonds above the phase transition: (1) Na1 and O1 behave like they are 444bonded near the phase transition pressure (e.g. a coordination change of Na takes place); (2) 445repulsion increases between the O1 and O4, and O5 atoms; (3) an increase in repulsion 446between H1 and Na1 occurs near 15.4 GPa, which results in a shortening of the O1-H1 bond, 447and/or a change in H-bonding configuration; and/or (4) the compression mechanism of the 448high-pressure phase is different than the low pressure phase.

Considering the mixed occupancy of the O1 site (e.g. Gatta et al., 2014; Henry and 450Dutrow, 2011; Bosi, 2018), it is unlikely a bond is forming between Na and O1. A plot of O-O 451distances reveals that there are no systematic trends that would suggest that the change in 452pressure dependence of the Na-O1 distance is due to an increase in O-O repulsion 453(supplementary Figure S6). It has been reported that an increase in size of the *X*-site cation 454results in an expansion of the X-site coordination polyhedron, which impinges on the hydroxyl at 455the O1 site, leading to a shortening of the O1-H1 bond (Berryman et al., 2016). It is possible that 456above ~15.4 GPa, an increase in Na-H repulsion occurs that results in a shortening of the O-H 457bond which changes the pressure dependency of the Na1-O1 distance. However, due to the low 458reported occupancy of H1 this seems unlikely. As discussed above, ring puckering and crimping 459also behave differently in the high-pressure phase, indicating that compression mechanisms

460shift in the high-pressure phase. Thus, a change in compression mechanism is the most 461plausible explanation for the change in Na1-O1 pressure dependence that is observed above 462~15.4 GPa.

463Bond Valence Analysis

We quantitatively analyze the stability of the dravite tourmaline structure using the 464 465bond valence approach (Brown, 2002). The bond valence approach has been extensively 466deployed to examine the structural changes in tourmalines due to chemical variability at 467ambient conditions (Fortier and Donnay, 1975; Hawthorne et al., 1993; Hawthorne, 1996; 468Hawthorne, 2002; Bosi and Lucchesi, 2007). We first use the relative deviation of the bond 469 valence sum from the ideal value $(\Sigma(S_i) - v_i)/v_i)$ for all cations and anions to quantify the steric 470stress of the structure at ambient condition. Here, S_{ij} is the bond valence between atoms i and j, 471and v_i is the absolute value of atom i's valence. The bond valence sums are shown in Figure 12 472on the ordinate (i.e. the zero-pressure values). It is striking that even at ambient conditions, the 473Mg cation is overbonded by almost 20 %. This is due to the MgO₆ octahedron sharing two edges 474in cis-conformation with the AlO₆ octahedron. The relaxed cation-oxygen distances for Al³⁺ and 475Mg²⁺ differ by about 10% (1.9 vs 2.1 Å). Thus, each MgO₆ shares two edges with an AlO₆ 476octahedron, whereas each AlO₆ octahedron shares one edge with a MgO₆ octahedra. As a 477consequence, Mg shows an almost 20% overbonding while AlO₆ is underbonded, albeit to a 478smaller extent (~ 3%).

In terms of the bond valence sums as a function of pressure (Figure 12), two features 480stand out: Na (which at ambient pressure is slightly underbonded) increases its overbonding by 481a remarkable 70%. Mg, which starts out at almost 20% overbonded increases to about 40% by

48215 GPa, at which pressure it then levels out. Notably, Na and Mg are the two cations with the 483weakest individual bonds. All other cations show a shallow increase in overbonding, which 484increase moderately to a maximum of 20% at the highest pressures. The same general increase 485holds for the oxygen atoms, which are all tightly bunched together at values below 20 % (we 486disregard O1 and O3 here because the poorly constrained hydrogen positions make their bond 487valence sums ill defined).

The weakly bonded Na and Mg cations seem to be dramatically overbonded at high 489 pressures, without affecting significantly effect on the stability of the tourmaline structure 490 (beyond the modest descent of symmetry near 15 GPa). This stability of the structure may be 491 rooted in none of the oxygen atoms being overbonded by more than 20 % up to the highest 492 pressure of \sim 24 GPa. Consequently, we hypothesize that this high-pressure stability of the 493 tourmaline structure is closely tied to its extraordinary structural tolerance with respect to 494 chemical substitutions on the X, Y, and Z sites. The bond network of tourmaline seems to have 495 the ability to balance the modest overbonding of the oxygen atoms which hold the framework 496 together, in spite of the significant overbonding on the most weakly bonded metal ions.

Indeed, the subtle mirror plane breaking phase transition observed above ~15.4 GPa is
498not reflected in the bond valence analysis. Instead, this phase transition is likely due to a
499steric/geometric effect caused by a complex interplay of bond length shortening and oxygen
500atom rotation/displacement which alter ring puckering, ring crimping, tetrahedral rotation, and
501ditrigonality. As the ring distortion parameters change under compression, mirror plane
502symmetry is broken for O6, O7, and O8. Hence, there does not appear to be an obvious

503villain(s) that drives the transition from *R*3*m* to *R*3; it is perhaps best viewed as a cooperative 504instability of the (initially) hexagonal ring configuration, via its interactions with the Y (Na) 505cation.

506 Implications

507 Other compositions of tourmaline that are rhombohedral R3m at room temperature and 508pressure will likely behave similarly, and undergo a similar transition, under compression. The 509main difference is likely to be the pressure at which the R3m to R3 transition occurs. Tourmaline 510is a complex mineral group, and the high-pressure behavior of tourmalines with extensive Li, Cr, 511or V substitution into the Y and Z sites, Pb substitution into the X-site, or Al, Li substitution into 512the T-sites could be quite different than that of tourmalines with predominantly Al and Mg in 513their Y and Z sites. Nevertheless, the main chemical control over the pressure at which this 514transition occurs is likely the identity of the cation in the X-site. Here, simple systematics would 515 indicate that a larger cation would generate this distortion at lower pressure conditions. Thus, 516the pressures of this transition in uvite tourmaline [Ca(Mg₃)MgAl₅(Si₆O₁₈)(BO₃)₃(OH)₃(OH)] and 517maruyamaite tourmaline [K(MgAl $_2$)(Al $_5$ Mg)(BO $_3$) $_3$ (Si $_6$ O $_1$ 8)(OH) $_3$ O] are expected to be higher 518and lower than in dravite tourmaline, respectively. Moreover, our high-pressure results show 519that massive overbonding of the X and Y sites can be accommodated by the tourmaline 520structure. This unexpected result may explain the extraordinary structural tolerance with 521respect to chemical substitution on the X, Y, and Z sites. Hence, any over/underbonding due to 522chemical substitutions is balanced by the topology of the tourmaline structure, and perhaps in 523particular the oxygen atoms.

524 The crystallization conditions of metamorphic tourmaline are complex, but one means of 525crystallization involves consuming B that is released during clay and mica mineral dehydration 526reactions (Ota et al. 2007). Hence, tourmaline is likely present in subduction zones where 527sedimentary rocks are being subducted. Petrologic studies have shown that dravite tourmaline 528breaks down at ~7.0 GPa and ~900 °C (Werding and Schreyer 1996) and ~3-5 GPa and ~1000 °C 529(Krosse 1995). However, tourmaline stability is highly dependent on compositional variations 530(e.g. its Mg/[Fe+Mg] ratio, Na, Al and B contents), coexisting fluids and the host rock (Henry and 531Dutrow, 1996; Dutrow et al., 1999; Schreyer, 2000; Ota et al. 2008). The higher pressure/lower 532temperature stability of dravite tourmaline is not well constrained (see figure 2b in Ota et al. 5332007). Moreover, tourmaline has a greatly diminished stability range in SiO₂ saturated systems 534with common pelitic compositions, and even in relatively cold subduction zones, tourmaline 535breaks down by ~150 km depth (Ota et al. 2008). Hence, the commonly invoked "cold 536subduction zone" argument for the stability of tourmaline appears to be invalid. A complete 537understanding of the host rock chemistry, redox conditions, and pressure and temperature 538conditions are required to truly constrain the stability field of tourmaline supergroup minerals in 539subduction zones. The stability fields of dravite tourmaline and other varieties of tourmaline 540are not well constrained, and neither the onset pressure of the symmetry-lowering transition in 541other varieties of tourmaline, nor its temperature-dependence is known. Thus, it is unclear 542whether the high-pressure phase of dravite or other compositions of tourmaline would be 543 found at depth within subduction zones.

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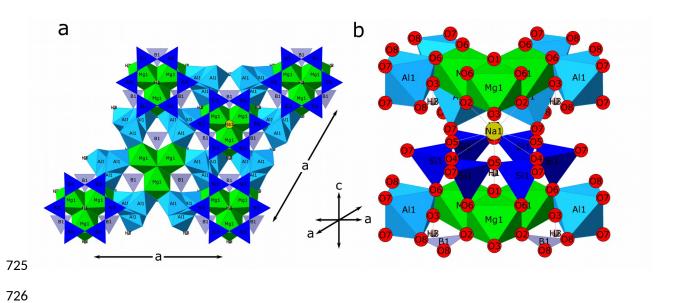
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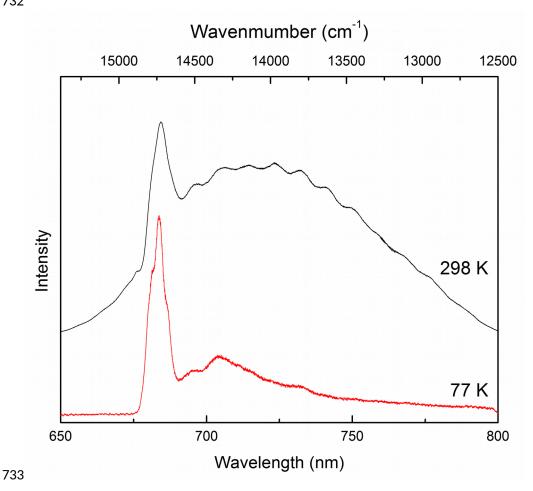
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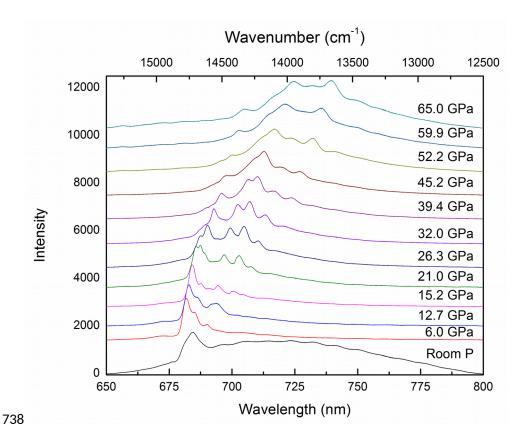
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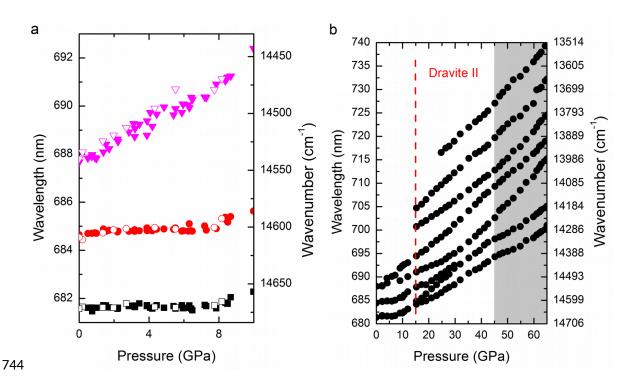
727Figure 1. Our room pressure crystal structure diagram of dravite tourmaline. (a) looking down 728the c-axis with oxygens removed for clarity. (b) oblique view showing the configuration of the X-729site within the Si $_6$ O $_{18}$ ring. Modified after Hawthorne and Dirlam (2011). Diagrams were 730generated in CrystalMaker $^{\circ}$ v8.7.6.



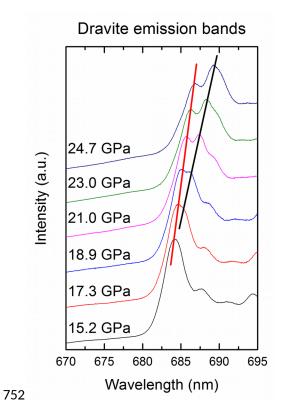
734Figure 2. Steady state luminescence spectra of dravite tourmaline at 300 and 77 K using 532nm 735excitation. Note that the thermally populated 4T_2 bands are not entirely quenched at 77K. The 736modulation in the broad band emission band at 300 K is an artifact.



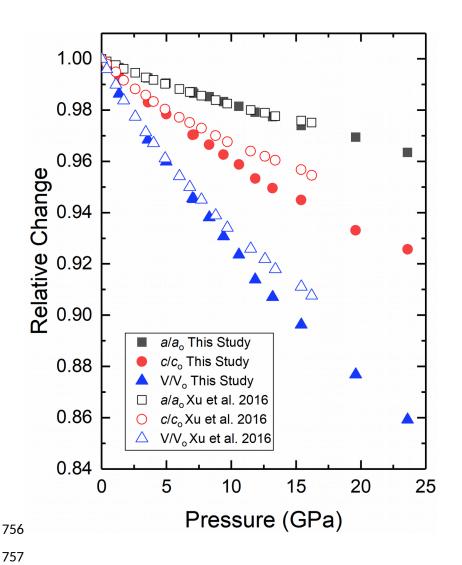
739Figure 3. High-pressure luminescence spectra of dravite tourmaline up to 65 GPa in a Ne 740pressure medium. The transition from intermediate to strong crystal field occurs between room 741pressure and \sim 6.0 GPa, as manifested by the decline in intensity of the broad-band 4T_2 emission 742centered near 725 nm.



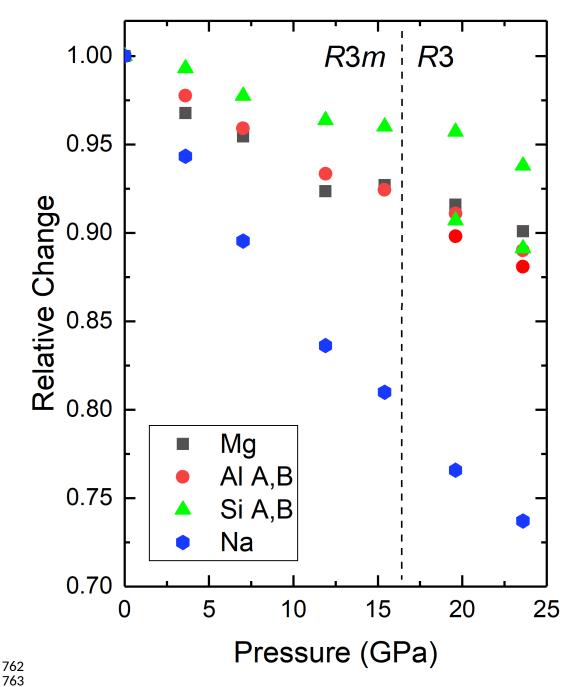
745Figure 4. Pressure shift of the observed luminescence bands in dravite (a) low pressure region, 746showing the essentially 0 nm/GPa pressure shift of two of the emission bands; and (b) up to 65 747GPa. The grey region indicates a possible change in compression mechanism. Closed symbols 748are data collected on compression, and open symbols are on decompression (for clarity, these 749are not included in b). Error bars are smaller than the symbols.



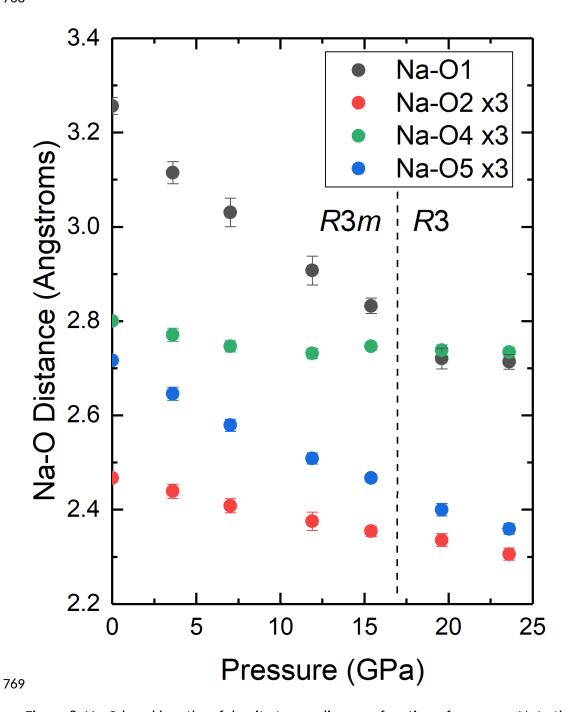
753Figure 5. Detail of dravite R-related emission band illustrating the splitting observed under 754compression.



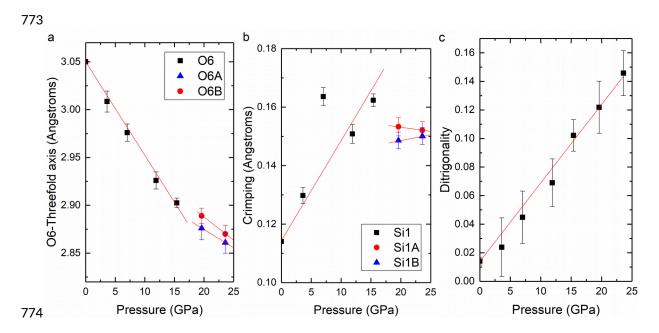
758Figure 6. Relative change of lattice parameters and unit cell volume for uvite (Xu et al. 2016), 759and dravite (this study). The results of Li et al. (2004) are not plotted for clarity (see 760Supplementary Information). Error bars for our study are smaller than the symbols. 761



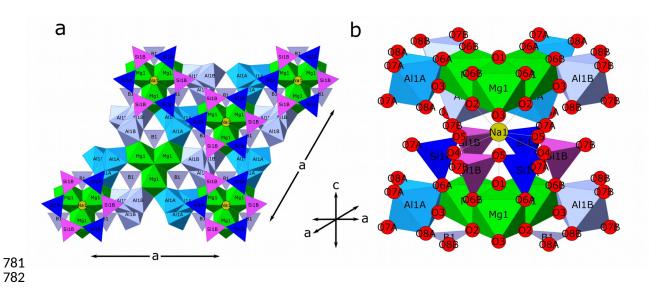
764Figure 7. Site volumes in dravite tourmaline as a function of pressure. There is a large difference 765in compressibility between the Na site and the Si, Al, and Mg sites. Note that the Na site is 766considered 9-fold across the pressure range of these measurements. 767



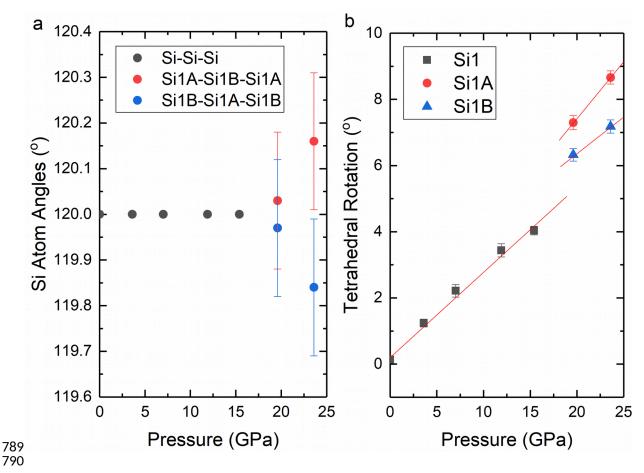
770Figure 8. Na-O bond lengths of dravite tourmaline as a function of pressure. Note the different 771behavior of the two sets of three Na-O-Si bonds, and the change in slope of the Na-O1 distance 772above ~15.4 GPa.



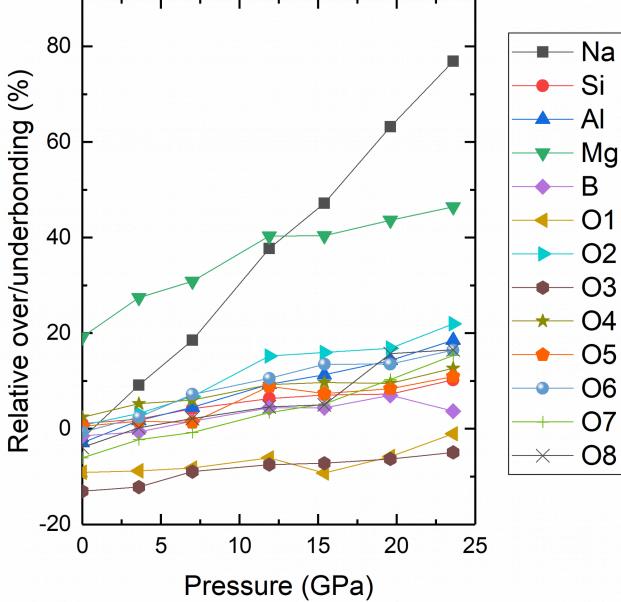
775Figure 9. Ring distortion parameters as a function of pressure: (a) O6-threefold axis distance as a 776function of pressure (e.g. ring puckering). This distance decreases as pressure increases, 777indicating that ring puckering increases as pressure increases. (b) Ring crimping (see text for 778definition), showing the change in ring crimping in the high-pressure phase. (c) Ditrigonality (see 779text). The change in puckering and crimping in the high-pressure phase indicate that the 780compression mechanism changes in the high-pressure phase.



783Figure 10. Our 23.6 GPa R3 structure of dravite tourmaline. (a) looking down the c-axis with 784oxygens removed for clarity. (b) oblique view showing the configuration of the X-site within the 785Si₆O₁₈ ring. Note that Al, Si, O6, O7, and O8 all split into unique sites in the high-pressure phase. 786Diagrams were generated in CrystalMaker $^{\circ}$ v8.7.6.



791Figure 11. (a) Si-Si-Si atom angles and (b) tetrahedral rotation as a function of pressure. Note 792that the errors on bond angles are typically quite large; however, by ~23.6 GPa the deviation 793away from 120° is larger than the errors. The discontinuous shift in tetrahedral rotation near 79415.4 GPa shows that the phase transition occurs near this pressure, but the pressure 795dependence of rotation is only slight altered in the high-pressure phase.



800Figure 12. Bond valence calculations for different ions within the dravite tourmaline structure as 801a function of pressure.

Pressure (GPa)	Pressure (GPa) a-axis (Å)		Volume (ų)
*0.0001	15.9370(6)	7.1749(4)	1578.19(14)
1.4(1)	15.8820(8)	7.1258(4)	1556.60(22)
*3.6(2)	15.8180(7)	7.0534(3)	1528.38(22)
5.0(1)	15.7860(7)	7.0198(3)	1514.95(20)
*7.0(1)	15.7323(7)	6.9617(4)	1492.20(20)
7.1(1)	15.7219(8)	6.9628(4)	1491.83(22)
8.3(2)	15.7015(8)	6.9344(4)	1480.55(23)
9.4(2)	15.6711(7)	6.9066(4)	1468.91(21)
10.6(2)	15.6418(10)	6.8794(5)	1457.66(29)
*11.9(2)	15.6040(4)	6.8401(17)	1442.30(70)
13.2(3)	15.5765(14)	6.8127(7)	1431.53(40)
*15.4(4)	15.5214(21)	6.7796(11)	1414.48(58)
*19.6(4)	15.4491(13)	6.6948(6)	1383.80(32)
*23.6(5)	15.3537(15)	6.6419(6)	1355.97(37)
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804

805Table 1. Unit cell parameters of dravite tourmaline at various pressures. *Crystal structures 806solved at these pressures 807

Sample	v ₀ (Å ³)	K _O (GPa)	K' _O (GPa)	Method	Reference
Schorl	1595.52(1.98)	183.5(4.2)	4(fixed)	Powder XRD	Li et al. (2004)
Uvite	1537.1(11)	96.6(9)	12.5(4)	Powder XRD	Xu et al. (2016)
		120(2)	4(fixed)	Powder XRD	Xu et al. (2016)
Dravite	1578.19(14)	109.0(2.9)	4.8(8)	Single crystal	This study
		112.2(9)	4(fixed)	Single crystal	This study

809Table 2. Pressure-volume data fit up to ~15.4 GPa using both 2nd and 3rd order Birch-810Murnaghan equations of state.

811

Pressure (GPa)	0.0001	3.6(2)	7.0(1)	11.9(2)	15.4(4)
Temperature (K)	298(2)	298(2)	298(2)	298(2)	298(2)
Crystal System,	rhombohedral	rhombohedral	rhombohedral	rhombohedral	rhombohedral
space group	R3m	R3m	R3m	R3m	R3m
a (Å)	15.9370(6)	15.8180(7)	15.7323(7)	15.6040(4)	15.5214(21)
c (Å)	7.1749(3)	7.0534(3)	6.9617(4)	6.8401(17)	6.7796(11)
Volume (ų)	1578.19(14)	1528.38(22)	1492.20(20)	1442.30(70)	1414.48(58)
R ₁ (%)	2.99	6.57	4.89	4.07	3.07
$WR_2(\%)$	7.35	16.23	11.86	9.83	8.33
Goodness-of-fit	1.079	1.158	1.238	1.267	1.193
No. refined parameters	44	44	44	44	44

23.6(5)		
298(2)		
rhombohedral		
R3		
15.3537(15)		
6.6419(6)		
1355.97(37)		
7.35		
18.50		
1.056		
70		

814Table 3. Crystal structure refinement details for dravite tourmaline at various pressures. 815Complete CIFs can be found in the supplementary material.