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RADIANCE DISTRIBUTION AS A FUNCTION OF DEPTH IN THE

SUBMARINE ENVIRONMENT

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INTRODUCTION

During the decade from 1935 to 1945 there seems to have been considerable 1/2/4/5/10/11 interest in the angular distribution of natural light underwater. Ingenious although somewhat crude instruments were devised including the "shadowingscreen" photometer (described by Pettersson and also by Johnson ans Liljequist), and the Gershun-tube photometer which directly limited the solid angle of $h\ell$ published by Poole-/ \mathcal{P} observations were made on compared variables were made on compared variables variables were made on compared variables were made on compared variables were made on compared variables were mad plane in the sun's direction and in the vertical plane at right and in the vertical plane at right angles to the vertical plane at \sim -and Whitney \rightarrow expansion \sim the light field with azimuth sweeps taken at various zenith angles.

As a result of this research it was observed that the angular distribution pattern of the natural light field changed in shape with depth, and it was surmised that at some unknown depth an equilibrium shape might be reached. It was also noted that the direction of the "bright" spot in the submarine light. field tended to approach the zenith as depth was increased.

Although the data obtained during this period was adequate for the purpose for which it was intended it does not now lend itself to detailed investigations \clubsuit of the submarine light field or to the computation of water constants. The resolving power of the instruments used was never better than 15° ($\rlap{/}$ /nis l . refers to the apex angle of the circular cone of collection), Large angles of $\mathcal O$ collection such as this average too much information, especially in the direction $\mathcal S$ of the sun and tend to broaden the shape of the distribution diagram in that $\mathcal A$ direction and at the same time to lower the apparent value of the peak radiance.

$$
\frac{1}{2} \sum_{i=1}^n \frac{1}{2} \
$$

The data from the various publications cannot be combined because of differences in geographical location and instrumentation details (for example_{Λ} $\mathcal A$ in the selection of filters and photo cells). The papers individually have insufficient information on the homogeneity of the water sample which was being meausred, and for computational purposes have insufficient data points covering $\mathcal{L}^{\mathcal{M}}$ A the radiance of the submarine light field.

The importance of the radiance distribution of the natural submarine light field as a primary means for documenting the optical properties of large bodies of water was first recognized by Dr. S. Q. Duntley in 1949 while he was conducting research on visibility problems in scattering-absorbing media. -At-that -time Dr. Duntley initiated a program of instrument development and study which $\overline{}$ has been continued to the present time. t *then*

Some of the theoretical work accomplished by this program has been published in the Journal of the Optical Society of America, and in reports. $\frac{6}{7}$ from the Visibility Laboratory, and some is being brought to the attention of the Optical \mathcal{A} Society by Dr. R. W. Preisendorfer. $\frac{12/13}{12}$

The development of instrumentation for the measurement of radiance distribution was first described to the Optical Society by Dr. Duntley in a paper presented at the October 1955 meeting in Pittsburgh and the program of work in progress at that time was discussed by Mr. Tyler. Following the summer of 1956, major design changes were made in the instrumentation, and in the spring of \mathcal{F} 1957 field operations were conducted in deep water at the United States Navy Electronics Laboratory Calibration Station on lake Pend Oreille in Northern Idaho.

The objective of the 1957 field work was to obtain detailed data on radiance distribution as a function of depth in homogeneous water under "clear sunny sky"

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and "overcast sky" lighting conditions. Since the data was to provide new input to the theory of radiative transfer through an hydrosol, measurements were confined within a narrow wavelength band.

Pertinent Description of Lake Pend Oreille

Lake Pend Oreille is in many ways ideally situated for submarine optical investigations. The field station is a barge moored at the south end of the lake about two miles from the village of Bayview and about one-half mile from the nearest shoreline. Cape Horn Peak to the north, rises 16.5° above the horizon, \mathcal{L} and peaks of the western edge of the Bitterroot mountains block out, at most, and peaks of the western edge of the Bitterroot mountains block out, a t most, 12.5⁰ of the sky to the south. To the east and west the sky line is much lower 12.5° of the sky to the sky to the sky line is much lower than α line is much lower low than this. Although the effect of the sky line on the submarine light field
is not yet known, it is presumed to be small due to Fresnel reflection at these owned is not yet known, i t is presumed to be small determined to $\mathcal A$ **A** angular subtense of the sky line so that the 16.5° between sky line and horizontal becomes only 2.7[°] in the underwater scene.

 $\sqrt{\int$ *The major inlet to Lake Pend Oreille is Clark Fork* some twenty miles to the ∂ north of the station and practically in line with the Pend Oreille River, which is the major outlet. Thus the largest source of particulate matter has little effect on the water near the station. The eleven streams and creeks along the south and east shore line that were examined at the time of the experiment were all carrying clear water from melting snow over streambeds of clean boulders.

The lack of currents in the lake, the somewhat stagnant location of the station, and the absence of silt-laden drainage into the southern end of the lake all helped to minimize the possibility of stratification or inhomogenvity in $\mathcal{L}_{\mathcal{A}}$ the water,

a.

DESCRIPTION OF EXPERIMENTAL SET-UP

The Pend Oreille Calibration Station is a two-story, 40 by 40 foot barge floated on T6 pontoons and held in place by mooring cables which are designed

Radiance Distribution $-4 -$

to minimize yaw. The siding, at the time of the experiment, was painted light *Q* green and as a result, the barge, at noon, exhibited high positive contrast as seen from the instrument. The new underwater photometer was suspended from a single supporting cable at the end of a 30 foot boom on the south side of the building. The inboard end of the boom was fastened to the barge near the water line and the outboard end was about six to eight feet above the water. The instrument was thus 30 feet or more from the barge at all times. The maximum horizontal angles subtended by the barge at the point of immersion were 45.5° to the east, 8.5° to the west and about 33.7° in the vertical direction¹. The image of the barge and the submarine shadow created by the barge have a noticeable effect on the natural light field near the surface and a detectable effect even at the deeper stations. The depth of the water at the station is 750 feet which \mathcal{A} assured the absence of bottom reflection effects in the measurements. a_{2} •aaoupoeL the absence of bottom reflection effects in the measurements. *.£**yeS* $\overline{\text{A}}$

All of the radiance distribution data presented herein were taken with the underwater photometer shown in Fig. 1 . This instrument was suspended on a single cable and powered by a 31-conductor electric cable looped into the underside at the vertical axis of rotation. The instrument was thus free to rotate around this vertical axis. Rotation about this axis was controlled by means of a gyrosyn compass assembly in the instrument with control transformer and indicator ${\mathscr{G}}$ on the main control panel. The error signal resulting from an azimuth mismatch between the control transformer and the gyro heading was used to drive a propeller which rotated so as to minimize the error signal. With this mechanism it was possible to maintain an azimuth setting to plus or minus 1° or better.

* From a point below the surface the barge subtended a vertical angle of about $10^{\rm o}$ at the instrument, and appears between tilt angles of 38 $^{\rm o}$ and 48 $^{\rm o}$.

 $**$ This instrument will be described in detail in a later publication.

Figure 1

The Visibility Laboratory's Photometer mounted and ready for lowering into Lake Pend Oreille, Idaho. The measuring head with its Gershun tubes is on the right and the propeller and damping fin on the left.

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For all the data presented here the azimuth error is less than plus or minus 1° .

The optical system of the underwater photometer measuring head, shown in Fig. 2, is a dual detecting system, each channel consisting of a 931A multiplier phototube wrapped with black tape except for a window covered with a Wratten No. 45 blue-green gelatine filter.

The measured transmittance characteristics of this filter are shown in Table I, Since the readings obtained are proportional to

 $EST_FT_W d\lambda$ **A,**

 E = Energy distribution of the light just below the surface

 $S =$ Spectral sensitivity of phototube T_{p} = Transmittance of No. 45 filter T_{bl} = Transmittance of water path

and since the combination $\text{ST}_{1,j}$ limits the value of λ , to 700 millimicrons or less, it can be seen that this data is tagged with an energy distribution having half band width of 64 millimicrons and extreme limits between 430 millimicrons and 546 millimicrons.

In order to extend the range of measurement, each tube is surrounded by a cylinder containing openings covered by filters having neutral densities of about 0, 1, 2, 3, 4, and infinity. The window in the pressure housing is frosted on the inside to assure a constant coupling factor between the collection system

^{*} Deep water measurgments in this region of the spectrum would not necessarily require an auxiliary filter since the combination of water and phototube sensitivity would provide all the "filtering" action required. In this work the Wratten No. 45 filter was used simply to make the shallow water measurement consistent in bandwidth with the deep water measurements.

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MEASURED TRANSMITTANCE OF WRATTEN NO.45 GELATINE FILTER

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203 10 20 40 50 60 70 80	O 0 $\mathsf O$ \circ \circ $\mathsf O$ $\mathsf O$ \circ	490 500 10 20 30 536 40 546 50 60	16.9 14.4 10.0 5.3 2.0 1.0 0.32 0.10 0 $\mathsf O$	770 80 90 800 10 20 30 40	84.0 86.0 87.0 87.5 87.8 88.5 88.8 89.0	1500 1600 1700 1800 1900 200C	56.5 44.5 35.0 27.0 22.0 20.5
90	\circ	70	Ω	50	89.5		
300 10 20 30 40 50 60 70 80 90	$\mathsf O$ O \circ $\mathsf O$ $\mathsf O$ $\mathsf O$ \circ O O \overline{O}	80 90 600 10 20 30 40 50 60 70	\hbox{O} 0 \hbox{O} $\mathsf O$ O $\mathsf O$ \hbox{O} 0 $\mathsf O$ \circ	60 70 80 90 900 10 ∞ 30 40 50	89.6 89.8 89.5 89,8 89.8 89.8 89.8 89.8 89.8 89.8		
400 10 20 30	$\mathsf O$ $\mathsf O$ $\mathsf O$. 0.1 1.0	80 90 700 10	$\overline{0}$ 0 0.25 3.3	60 70 80 90	89.5 89.8 89.8 90.0		
433 40 50 60 70 80	3.4 10.0 15.0 17.0 17.6	20 30 40 50 60	19.0 42.5 62.0 74.5 80.8	1000 1100 1200 1300 11,00	90.2 90.8 90.5 86.5 76.5		

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Figure 2

Exploded view of the measuring head of the Visibility Laboratory's Underwater Photometer. The two multiplier phototubes with the cylinders containing neutral density steps can be seen in the central assembly. The Gershun tubes (not shown) are mounted over the windows, one of which is shown in the housing on the right.

and the multiplier phototube. The internally baffled radiance tubes limit the angle of acceptance to 6.6° . The whole head can be tilted through a range of somewhat more than 180[°] by means of a synchronous motor. Event marks are automatically recorded at a specific angle at the beginning of each tilt sweep.

HOMOGENEITY OF THE WATER

Considering the intended application of this data, to examine the ability of current theory to describe and predict the passage of natural light through a homogeneous hydrosol, it was of the utmost importance for the water sample to be homogeneous. The favorable location of the barge and the absence of conditions likely to produce inhomogeneity in the wator have already been mentioned. The water itself was also examined for evidence of inhomogeneity. This examination consisted of:

A. Bathythermograph measurements. It is well known that certain lakes, including lake Pend Oreille, undergo seasonal changes in their temperature profile, that warm weather will develop a layer of warm water at the surface and that various forms of microorganisms will "bloom" in or near this warm layer. Stratification of this type means discontinuous changes in the structure of the light field as a function of depth, a condition which would not yield suitable data. The absence of a thermocline implas that the bloom has not occurred and that this ic A source of inhomogeneity is absent.

Bathythermograph measurements were made at frequent intervals during the \mathscr{A} entire operation and consistently showed no thermocline. The record for March 16, for example, shows a surface temperature of 2.35° C rising at constant rate \sim to 3.20°C at 122 M depth. The record for 29 April shows a constant temperature of 3.6°C from the surface to 137 M. More critical examination (on 29 April 1957) of the top two meters with a bucket thermometer showed 4.5° C at the surface, $\frac{1}{2}$ $\frac{4}{9}$

4.0°C at 1.25n_A and 4.0°C below 1.25 n_A indicating a calibration error in the $\frac{1}{3}$ ³
bathy_cthermograph that is not significant to these light measurements, and possibly a slight surface heating due to full sunshine on the 28th of April. All other bathythermograph records were, happily, monotonously the same.

B. Beam Transmittance. Beam transmittance and its variability with depth were measured with a beam transmissometer. Total transmittance for a collimated beam of light was 67.3% per meter at the surface and increased at a *d* steady rate to about 68.6\$ per meter at 6lm.</ (Measurements of 29 April 1957) *e* Other determinations taken during the period of field operations showed no significant variation from this condition.

•a- $C.$ Particulate Matter. One and two liter samples were taken with a Nansen bottle at descrete depths from $6n$ to $60n$ and filtered through Millipore H A filters having a pore size of 0.45 microns plus or minus 0.02 microns. Results are shown in Table II.

Although each of the above tests seems to exhibit a trend with increasing depth, the trends are conflicting and the variation of the parameter in each $\mathscr P$ case is close to the expected error. Definitive evidence of stratification or inhomogeneity was considered to be absent during the period of the experiment.

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Further evidence of the homogeneity of the lake water is provided by the results of calculations from radiance distribution measurements.

Constancy of the Surface Lighting

The natural lighting on the surface of the water can undergo changes in ambient level and also in structure, that is, in the position of bright spots, such as the sun or single white clouds, or in the relative proportions of specular to diffuse light. The effect of very slow changes in ambient level can be successfully corrected but no method is known at present which will properly \mathcal{A} correct for the effect of rapid changes in ambient level or for the effect of changes in the structure of the light field. One must wait for desirable lighting conditions and even then may be forced to reject some data because of these uncontrollable variables!

For several years $\overline{\text{#}}$ it has been the practice of this laboratory to monitor the natural lighting at the surface with an instrument sensitive to both the light level and to its structure and on the basis of this record to sort and reject data. For the data taken on 28 April no rejection was indicated.

PLAN OF OPERATION

The plan of operation was dictated to a large extent by azimuth and elevation changes in the sun's position, by the proposed computations from the data, and by the operating characteristics of the instrument. For computational purposes,

* A description of this instrument will be the subject of a separate paper.

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data at equally spaced azimuth intervals was needed. A 20[°] interval would give **A** 18 continuous tilt sweeps with each measuring channel, or two complete radiance α for fixed depth station in \mathcal{L} , in the problem in \mathcal{L} , the problem in \mathcal{L} run a sweep by one radiance tube would be repeated 12 (12 minutes later by the repeated 12 minutes later by the sweep of the opposite radiance tube.

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The most favorable period for making the measurements would be two hours each side of noon (sun time) since this is the period of minimum change in sun elevation and consequently of minimum change in ambient light level and structure. Allowing 30 minutes at each station would provide for eight stations between 10 AM and 2 PM. Since there would be some ambient light level change during this four hour period a monitoring run was planned for noon. For this run the radiance tubes were set in a vertical position and measurements were made at descrete depths, including all station depths in about 10 minutes. Thus in the final data reduction all radiance distributions for the day could be adjusted to a single sun altitude.

DISCUSSION AND TREATMENT OF DATA

Data recording on April 28 was begun at 0850 at a depth of 66.1m and continued until $\frac{1}{4}$ and $\frac{1}{4}$ and $\frac{1}{4}$ is given in Table II.

 $T_{\rm eff}$ original data clearly shown the environment. The environment of the environment. The environment σ of the sunlit barge wall is obvious at the shallow stations and the shadow of the barge can be seen at all stations although it is not obvious at the deeper ones (See Fig. 3.) The position of the "bright spot" is always recognizable, and at the shallow stations the edge of the "man hole" can be seen, as can the shadow of the instrument itself. In addition, changes in the sun's azimuth position and in its elevation can be detected in the data.

TABLE III

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In order to remove as much of this distortion from the data as possible the following proceedure was adopted:

1. Calibration Correction

The original data, which is very nearly linearly proportional to the log of the radiance, was read at ten degree intervals of tilt with a special "rule" which converted the data to radiance units and at the same time removed any small departures from linearity that might be present. Each information channel, consisting of the multiplier phototube plus the chassis plus the recorder, *A* has its own calibration "rule." The data for these rules *fa*. obtained at the site *»**-i~*~* of the experiment just before and just after the measurements.

2. Changes in Ambient Light Level

Inspection of the data for the nadir direction indicated the extent of the ambient light level change due to altitude changes of the sun. When ω *''*". necessary this has been removed by normalizing all tilt sweeps for the station to the average nadir reading.

3. Azimuth Motion of the Sun

The data thus obtained *tfi£ re plotted on semi-log paper, as shown in **"***"*"* **A.** Fig. 3. In this plot, tilt angle and azimuth angle are both plotted along the $\mathcal A$ abscissa. The known azimuth angle between the sun and the instrument heading was used to locate each tilt sweep on the plot. Thus the azimuth motion of the sun relative to the instrument is not superimposed on the data. m

4. Barge Image and Shadows

The data points resulting from step three were joined by smooth curvos to give azimuth sweeps at constant tilt as well as tilt sweeps at constant azimuth \mathcal{N} (See Fig. 3.) The image of the barge, its shadow \mathcal{A} atc., can be positively \varnothing \varnothing

Figure 3

Typical working plot of the data showing the image of the barge as seen through 16.6 m of water, and $a \pm b$ the shadow of the barge. In this plot the angles at the top marked "instrument heading" are relative to a fixed compass point. They are plotted progressively short of the azimuth scale of the graph in order to remove motion of the sun relative to the instrument heading. The ϕ scale at the top applies to the azimuth sweep at $\theta = 0^{\circ}$ and must be moved $\sim 10^{\circ}$ to the right for each succeeding azimuth Ą sweep. Azimuth sweeps in the upper hemisphere are marked with their appropriate tilt angles (0) at the right. The 0 scale at the bottom applies to each individual tilt sweep by relocating it to the right or left in 20° steps. It is shown lined up with the ϕ = 0° curve.

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identified in these plots. Sections of the data which were distorted by these spurious signals were not used in the data reduction.. The physical location of the instrument was such that only about 54° of the horizontal sweep was distorted by the presence of the barge. The position of the sun was located by the maximums of the azimuth sweeps and checked with the known position of the sun in each case.

5. Read Out

The resulting azimuth curves were read out to give the appropriate data with the sun located at 0° azimuth.

6. Duplicate Runs

Both information channels functioned perfectly during the entire operation and consequently duplicate runs were available at all depth stations. These were treated independently through step five above and the read outs were then \leq averaged. A double run was made at the 66.1m station giving a total of four complete determinations at this depth. All determinations are averaged together in the final read out.

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7. Depth-Difference Correction

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The ends of the brightness tubes are about a half a meter from the center of rotation of the measuring head. As a result the depth does not remain constant but changes continuously with tilt angle according to the equation: *4*

Z_+ = Z $+$ r cos Θ

Where Z_t is the true depth, Z is the reported station depth to the center of rotation of the instrument and r is the distance from the center of rotation to the end of the radiance tube.

The averaged data $_{\rm warg}$ corrected to give the radiance distribution at a point $~$ $\rm \omega$ A by determining the slope of the curve of path radiance vs depth for every pair of

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values of tilt angle and azimuth angle and making the proper correction along this slope.

8. Sun Altitude Changes

In addition to changes in ambient light level on the surface, large changes in the altitude of the sun result in a change in the ratio of the zenith to nadir path radiance and a reorientation of the whole radiance distribution solid in the direction of the sun. The depth stations were taken in the order shown in Table III, in order that the shallow stations would be at noon and the others would be clustered around nocn in such a way that a complete set of stations cculd be selected from those nearest noon. Changes in the shape of the distribution solid due to changes in the sun's altitude would in this way be $\omega \rightarrow \gamma$ minimized.-

9. Normalization to the Vertical Run

Before the seven morning runs could be normalized to the single vertical run at noon it was necessary to demonstrate that the radiance distribution solid for the 66.1m station was substantially the same in shape at 0908 as it would have been at 1200, that the radiance distribution solid for the 53.7m station was substantially the same in shape at 0937 as it would have been at 1200, etc. To do this the ratio of the average zenith to average nadir reading for each depth statiun was compared with the ratio of the zenith to nadir reading found from the noon vertical run. For all depth stations duplicate ratios were found \sim indicating that no significant error is introduced by adjusting these runs to a single sun altitude. It was also found that the complete data for the 66.1m and 53.7m morning stations duplicated the data for the 66.1m and 53,7m afternoon stations. For the remaining morning stations the change in sun altitude between station time and 1200 is very small indeed. After refraction the change in sun angle from 1000 to 1200 is only four tenths of the resolving power of the \sim

radiance tube. During this period the greatest incremental change in sun angle takes place at the depth where it has the smallest effect on the shape of the distribution solid. It is, in fact, progressively true throughout the morning data that the large incremental changes in the sun's altitude occur at the depths where they have the least effect and the near zero changes in sun's altitude occure near the surface where large changes would have had a very large effect.

} On the strength of the above evidence, the data for the seven stations $\mathcal{O}_{\mathcal{A}}$ on the strength of the above evidence, the data for the seven stations the seven stations of the seven stations given in Tables Iv, v, vi, vii, viii, ii, and X $\frac{1}{2}$ presented as data for one sun $\frac{1}{2}$ altitude.

DATA FOR THE CLEAR, SUNNY OASE

Data representing the clear_x sunny case was obtained on 28 April 1957. The voice recorded notes for the day read as follows:

> "28 April 1957. It would not be possible to have a more perfect day for the sunny-sky case than today. Between seven in the morning and three-thirty in the afternoon there were no overhead clouds what so ever. The few small clouds that did appear just over the mountain peaks rapidly evaporated. At no time was it possible to see evidence of a high altitude cirrus layer. In addition to this the lake was practically flat calm all day."

Later computations place the clouds mentioned at an altitude of 13^o in the southern sky. They appeared one at a time and evaporated within 10 minutes; Their angular subtense was never more than 1.5° . The optical state of the lake surface is shown in the photograph (Fig. 4) which was taken on April 28 at about 1500.

The radiance distribution data for the clear sunny case $\frac{1}{2}$ given for seven depth stations in Tables IV through X inclusive.

DISCUSSION OF DATA AND EVALUATION OF SOURCES OF ERROR.

In the body of the tables, the overall variation in the value of radiance at any one setting is + *5%* of the radiance at that setting. This includes

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RADIANCE DISTRIBUTION CLEAR SUNNY SKY

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Figure 4

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Optical state of the lake surface on 28 April. Arrows indicate points where wires pass thru interface.

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instrument errors, reading and plotting errors, errors made in setting and and holding azimuth positions, and in fact all errors that have entered the $\mathscr A$ measurements before their presentation in the tables.

In the direction of the sun, azimuth steps of 20° move the acceptance cone by a distance which is almost equal to its base diameter. If the air-water boundary were flat the sun's image would therefore be within the cone only once during this group of azimuth settings. However, the direction of the sun is a glitter pattern whose size depends on the optical state of the surface (which in Ù, turn depends on the wind velocity $_{\mathcal{C}}$ During the measurements of 28 April the wind $^{\prime}$ velocity was less than 1 $\frac{m}{\sec A}$ which would indicate a glitter pattern considerably smaller than 10° in angular subtense^{*} but larger than a single sun image. Bue to the optical state of the surface on 28 April this glitter pattern probably consisted of widely separated points of light; Because of the geometry discussed above it is possible that the solid angle of acceptance never covered more than $\mathcal S$ half of the glitter pattern. This would mean that the reading obtained in the sun's direction could be low by a factor of 2 in the worst case.

In certain directions and especially at the shallow stations, the presence $\mathcal A$ noise signal which varies in both frequency and amplitude as a function of depth. Typical high and low values of this noise together with its frequency are given in Table XI. In Tables IV thru X average experimental values for radiance are reported in these directions.

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 \ast Cox and Munk, (9) have found that the glitter pattern subtends an angle of $10^{\sf o}$ when the wind velocity is $7.\sf m/sec.$.

Table XI

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Amplitude and frequency of the noise signal in the direction of the sun.

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ACKNOWLEDGEMENTS

Radiance distribution data taken at the Lake Winnepesaukee Field Station at a fixed depth (in 1956) established the need for an instrument with remote azimuth control which could be hung from a single supporting cable. Concepts like this come with ease, but the reduction to practice can be quite another matter. The difficult job of engineering, building, and testing this new underwater photometer was done by Mr. Roswell W. Austin, Research Engineer at the Visibility Laboratory, and a group of co-workers, including Theodore Petzold, Mechanical Engineer and Robert Howarth, Electronic Engineer. Their excellent workmanship is amply demonstrated by the success of this experiment, and is gratefully acknowledged. Procurement of a Sperry Gyrosyn Compass system was a special problem ably solved by Commander Carrol Walsh whose help in this and *'3* other respects is gratefully acknowledged.

Permission to use the N.E.L. Pend Oreille Calibration Station was arranged by Mr. Geoffrey Gould of the Navy Electronics Laboratory.. His assistance and the generous cooperation of the group at the Calibration Station under Mr.. ar e Donald Huckle is also gratefully acknowledged.

During the experiment the author was priviledged to have the help and During the experiment the author was priviledged to have the help and council of Mr. Robert W. Holmes, Assistant Research Biologist at the Scripps Institution of Oceanography, in the examination of the lake water for homo-Institution of Oceanography, in the examination of the lake water for homogeneity.

Mr. George Tate. Their agility and thoroughness in preparing the experimental set-up and in keeping it running ie also gratefully acknowledged.

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A large part of the tedious work of data reduction has been done by Mr. W. H. Richardson_nwhose diligence and perseverance in double_pchecking each *9* step of the worketias been most helpful and gratifying. Perhaps most important hand A. of all has been the holp and council of Dr. S. Q. Duntley, Director of the Visibility Laboratory, whose long-time interest in radiative transfer in scattering-absorbing media has made this work possible.

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