Lawrence Berkeley National Laboratory

LBL Publications

Title

Coupled THM Modeling of Hydroshearing Stimulation in Tight Fractured Volcanic Rock

Permalink https://escholarship.org/uc/item/73f4g872

Journal Transport in Porous Media, 108(1)

ISSN 0169-3913

Authors

Rinaldi, AP Rutqvist, J Sonnenthal, EL <u>et al.</u>

Publication Date 2015-05-01

DOI

10.1007/s11242-014-0296-5

Peer reviewed

1 COUPLED THM MODELING OF HYDROSHEARING STIMULATION IN						
2	TIGHT FRACTURED VOLCANIC ROCK					
3						
4	A. P. Rinaldi ^(a) , J. Rutqvist ^(a) , E. L. Sonnenthal ^(a) , T. T. Cladouhos ^(b)					
5						
6	(a) Lawrence Berkeley National Laboratory					
7	1 Cyclotron Road					
8	Berkeley, CA, 94720, USA					
9	e-mail: aprinaldi@lbl.gov, jrutqvist@lbl.gov, elsonnenthal@lbl.gov					
10						
11	(b) AltaRock Energy					
12	7900 E. Green Lake Drive N					
13	Seattle, WA, 98115, USA					
14	e-mail: tcladouhos@altarockenergy.com					
15						
16						
17						
18						
19						
20						
21	submitted to Transport in Porous Media					
22						
23						

24ABSTRACT

25In this study, we use the TOUGH-FLAC simulator for coupled thermo-hydro-mechanical (THM) 26modeling of well stimulation for an Enhanced Geothermal System (EGS) project. We analyze the 27potential for injection-induced fracturing and reactivation of natural fractures in a porous 28medium with associated permeability enhancement. Our analysis aims to understand how far the 29EGS reservoir may grow and how the *hydroshearing* effect depends upon system conditions. We 30analyze the enhanced reservoir, or hydrosheared zone, by studying the extent of the failure zone 31using an elasto-plastic model, and accounting for permeability changes as a function of the 32induced stresses. Results, for both fully saturated and unsaturated medium cases, demonstrate 33how the EGS reservoir growth depends on the initial fluid phase, and how the reservoir extent 34changes as a function of two critical parameters: the coefficient of friction and the permeability-35enhancing factor. Moreover, whereas the stimulation is driven by the pressure exceeding a 36hydroshearing threshold, the modeling also demonstrates how the injection-induced cooling 37boosts the stimulation a bit farther.

38INTRODUCTION

39Thermo-hydro-mechanical (THM) analysis is essential when dealing with energy extraction from 40hot dry rock. The main concept of the Enhanced Geothermal System (EGS) is the exploitation of 41high thermal gradient regions through the creation (or reactivation) of a fracture network by 42increasing rock permeability, hence enhancing the circulation of water. EGS sites generally 43feature very low permeability formation (such as igneous rock in volcanic regions), where cold 44water is injected at moderate to high pressure to achieve *hydroshearing* (e.g., Cladouhos et al., 452009). Contrary to the more common hydrofracturing (or simply *fracking*) process, during 46hydroshearing reactivation, the injection pressure is kept below the minimum principal stress

1

47magnitude, causing existing fractures to dilate, slip and shear. Hydroshearing can permanently 48enhance permeability of natural fractures that conceptually should remain open because of self-49propping of fractures due to surface roughness when the stimulation period ends and fluid 50pressure is reduced. As pointed out by Riahi and Damjanac (2013), the enhancement of the 51reservoir permeability during hydroshearing will depend on many *in situ* parameters, such as 52regional stress, geological, hydromechanical, and thermal parameters (e.g., frictional coefficient, 53intact rock permeability, heat capacity), as well as chemical processes associated with the 54injection of cold water.

55

56The concept of hydroshearing is not new, but was first recognized and developed in the early 571980s, when Pine and Batchelor (1984) confirmed that the creation of new fractures was not the 58 dominant process during the injection of water into the rock mass at great depth. Far more 59 important was the shearing of natural joints and, in particular, those aligned with the principal 60stresses of the local stress field. In this concept, the joints fail in shear because the fluid injection 61reduces the normal stress across them and allows frictional slippage to occur before jacking, or 62creating of new hydraulic fractures. This was first demonstrated for the Cornwall hot dry rock 63project in the Carnmenellis granite where injection was conducted at depths greater than 2 km 64below ground level. Microseismic events detected during the high-flow rate stimulations 65 indicated strike-slip shear consistent with the orientation of the natural joints and in-situ stress 66conditions. Meanwhile, in the field rock mechanics, the effects of fracture shear on permeability 67was studied through laboratory and in situ block experiments, with the first comprehensive study 68conducted at the Norwegian Getechnical Institude by Makurat et al., (1990). The concept of 69hydroshearing have since been employed at a number of EGS sites worldwide, including Hijoiro 70and Ogachi Japan as well as Soultz-sous-Forêts and Cooper Basin in Australia (Tester, 2006;

2

71Ziagos et al., 2013). Experience gained in the last thirty years at EGS field projects has shown 72the critical importance of understanding and mapping the natural fracture system and the in situ 73stress field (Evans et al., 1999, Tester 2006). Trans-tensional environments (e.g., grabens) may 74be more amenable to successful manipulation than compressive stress regimes in EGS reservoir 75creation (Baria et al., 1999). Moreover, in the recent few years a number of EGS demonstration 76projects have been launched in the U.S., in which different variants of hydroshearing is 77employed in stimulation and permeability enhancement of the reservoir (Ziagos et al., 2013). 78These are funded by the U.S. Department of Energy's Geothermal Technology Program and 79include EGS demonstrations at Desert Peak and Brady's Hot Springs, in Nevada, The Geysers in 80California, and finally the Newberry Volcano, Oregon (Ziagos et al. 2013). Three of these EGS 81demonstrations projects (Dessert Peak, Brady's, and The Geysers) are located within or on the 82margins of existing hydrothermal fields, whereas one (Newberry Volcano), is located at an 83unexplored and undeveloped site (Ziagos et al. 2013).

84

85In 2009, AltaRock Energy company was awarded a grant by the U.S. Geothermal Technology 86Program to plan and demonstrate an EGS at the Newberry Volcano, Oregon. During Phase I of 87the project, completed in April 2012, pre-stimulation field investigations were performed to 88understand the tectonic and volcanic setting, characterize the volume around the proposed EGS 89demonstration area, and plan the stimulation parameters. The stimulation plan was developed 90along the lines of Alta Rock Energy's approach to hydroshearing in which the rock is stimulated 91in stages by injection in multiple isolated sections of the well bore using a injection pressure that 92is slightly lower than the pressure required for hydrofracturing, i.e. slightly lower than the 93minimum principal stress (Cladouhos et al., 2009).

94

3

95A preliminary 3-D model of stress and fracture patterns was presented by Davatzes and Hickman 96(2011). Faulting is mainly evident along the caldera rim about 3 km from the designated 97injection well. There is no evidence of ring fractures or faults in the injection well (NWG 55-29) 98from drilling logs. Furthermore, Newberry Volcano has a very low seismicity rate (Cladouhos et 99al., 2013). An analysis of the natural fractures shows that there are two dominant sets that strike 100N-S and dip approximately 50° to the east and west (Davatzes and Hickman, 2011).

101The stimulation took place between October and December 2012, during which three zones were 102created by injecting thermal-degrading zonal isolation materials (chemical diverters) to isolate 103already stimulated zones (Petty et al., 2013). Over 40,000 m³ of water were injected in about 7 104weeks of stimulation, reaching a maximum well-head pressure of about 16 MPa with over 200 105induced microseismic events registered (Petty et al., 2013; Cladouhos et al., 2013). The 106maximum magnitude event (Mw = 2.39) was recorded during the period of highest well head 107pressure (P = 16.7 MPa). However, results show a strong correlation between the cumulative 108injected volume and cumulative logarithmic seismic moment, pressure does not correlate as 109strongly. The cloud of microearthquakes extends about 500–800 m from the injection well 110(Cladouhos et al., 2013).

111A simultaneous analysis of flow rates, pressure, and seismicity that occurred at the Newberry 112EGS Demonstration site after the stimulation shows that the injectivity increased after 113reactivation, indicating an increase in permeability. The maximum well-head pressure (controlled 114at pump) resulted only after the injection of the chemical diverters, which reduced the flow rates. 115Moreover, no changes in pressure or flow rate seem to indicate the occurrence of hydrofracking 116or tensile failure, meaning that the pressure was below the minimum principal stress, but well 117within the range to achieve hydroshearing and its associated permeability enhancement.

118

119Starting from the results obtained at Newberry Volcano, here we aim to study the hydroshearing 120effects during stimulation at an EGS site. We use as input the same stress field and rock 121properties (such as low permeability and porosity) measured at Newberry, and simulate the 122stimulation using a transient well-head pressure similar to the one recorded on the field. We do 123not aim to reproduce neither the seismicity nor the flow rate in exact detail, but aimed at a model 124that broadly represent the conditions and injection data at the Newberry. For example, we used a 125simplified but representative well pressure and a resulting flow rate that was similar to the 126measured one, though not matched in great detail.

127

128The main goal of this work is to understand how hydroshearing may occur during stimulation, 129using the Alta Rock's hydroshearing approach (Claduohus et al., 2009). Through the use of the 130simulator TOUGH-FLAC (Rutqvist et al., 2002), we simulate a porous medium, which deforms 131when subjected to stress change. Fracture effect is simulated by assuming an anisotropic field of 132permeability, with the most of the fluid flow occurring then along the assumed fracture direction. 133Generally fracture aperture may change as subject to stress, and we simulate such a process by a 134permeability-stress relation. Furthermore, if the stresses reach a critical value, defined by the 135Mohr-Coulomb criterion, shear failure will occur, and the permeability of the porous medium 136will be further enhanced by a certain factor (generally between 2 and 3 orders of magnitude – 137Rutqvist and Stephansson, 2003). We study the extent of the EGS reservoir after the stimulation 138(hydrosheared zone) both in a single-phase (liquid), single-component (water) system and in a 139two-phase (gas, liquid), two-component (CO₂, water) system. Effects of some key parameters, 140such as the frictional coefficient and a permeability-enhancing factor, are studied as well.

141MODEL SETUP

142The coupled THM analysis was conducted using the simulator TOUGH-FLAC (Rutqvist et al., 1432002; Rutqvist, 2011) based on the geothermal reservoir simulator TOUGH2 (Pruess et al., 1442011), which allows the modeling of multiphase and multicomponent fluids in a porous medium, 145and the geomechanical code FLAC3D (Itasca, 2009), for the stress changes induced by pressure 146and temperature. The TOUGH-FLAC simulator has been recently applied and tested over a wide 147 range of research fields, such as carbon sequestration (e.g., Cappa and Rutqvist, 2012; Rinaldi 148and Rutqvist, 2013), nuclear waste disposal (e.g., Rutqvist and Tsang, 2012, and references 149therein), hydrothermal systems and volcanology (Todesco et al., 2004), as well as studies related 150to water injection in geothermal fields (e.g., Rutqvist et al., 2013a; Vasco et al., 2013). 151Following the approach of Rutqvist et al. (2013a) for modeling of the stimulation injection at the 152Northwest Geysers EGS Demonstration Project, we studied the stimulation at a generic EGS 153 reservoir with low initial permeability, suitable for observing hydroshearing, such as at Newberry 154Volcano. Although it is beyond the aim of this paper to reproduce the data and observation 155recorded at the Newberry EGS Demonstration, we preferred to perform our simulation study of 156the hydroshearing and EGS reservoir extent starting from some reliable parameters such as those 157 obtained from Newberry Volcano site, where data were collected for more than one year before 158 starting the stimulation. In this study, we extended Rutqvist et al.'s (2013a) approach to calculate 159the actual permeability enhancement during the injection.

160We considered a one-quarter symmetric model, with the injection well (corresponding to the 161Newberry well NWG 55-29) located in one corner (Fig. 1). The model domain is a parallelepiped 162of dimensions 1.5 × 1.5 × 3.5 km and consists of four layers, representing the main geological 163formations of the Newberry area (Sonnenthal et al., 2012). Hydrological properties are listed in 164Table 1.

165During hydroshearing and hydrofracturing, the medium permeability changes as a result of 166injection-induced fluid pressure and effective stress changes, and is strongly dependent on *in situ* 167stress magnitude and orientation as well as fracture orientation. In this study, we consider a 168stress-dependent permeability (hence also an anisotropic initial permeability) with maximum 169permeability in the NS-direction, in order to simulate a highly fractured low permeability 170formation (such as the intruded John Day formation at the Newberry volcano).

171The injection well was simulated as a porous medium with high vertical permeability and very 172high porosity). It is divided into two sections. The first section represents a cased well (high 173vertical permeability and very low horizontal permeability), which allow heat exchange only and 174prevent fluid escaping from the well to the host rock. The second section represents an open well 175completion (very high vertical permeability and same horizontal permeability as the host rock) 176that allows the injection of cold water into the highly fractured low permeability formation 177(between 2000 m and 3000 m depth).

178Initial temperature and pressure distribution were extracted from former analyses of the pre-179stimulation steady-state conditions at Newberry Volcano (Sonnenthal et al., 2012). The 180temperature follows a high gradient of about 100 °C/km, with a maximum temperature of about 181360°C at the bottom of the domain. The pressure is slightly lower than hydrostatic, with a linear 182gradient of about 8.3 MPa/km. Constant pressure was set at the top and bottom boundaries, 183whereas side boundaries were assumed to be closed to fluid flow. No-flow side boundaries are 184needed to simulate only a quarter of a symmetric domain.

185Mechanical properties follow the results of Li et al. (2012). We chose to simulate a model with 186homogeneous mechanical properties and use a Young's modulus of E = 15 GPa and Poisson's 187ratio [] = 0.3. Homogeneous mechanical properties should be adequate in this case, since we 188simulate a short-term stimulation (slightly less than two months) that should affect only the

7

189injection zone (EGS reservoir). Indeed mechanical properties (such as shear and bulk modulus, 190and friction angle) may depend on medium heterogeneities, and then affect the stress and 191deformation distribution at macro scale. However, we are already accounting for an anisotropic 192permeability field, which affect the pore pressure distribution, and then the stress evolution as 193well. Random heterogeneities in mechanical properties may be present in a real field, but only at 194microscopic scale (i.e. too small to be analyzed in this study).

195Initial geomechanical conditions follow those used by Cladouhos et al. (2011) for the pre-196stimulation analysis performed with the AltaStim simulator. We considered a vertical stress 197gradient of 24.1 MPa/km (\Box_v , maximum principal stress in *z*-direction). The intermediate 198principal stress is oriented in the NS-direction (*y*-axis, \Box_H) with a gradient of 23.5 MPa/km. 199Finally, the minimum principal stress is oriented in the EW-direction with a gradient of 14.9 200MPa/km (\Box_h , *x*-axis). Thermal effects on stress were taken into account as well, choosing a 201coefficient of linear thermal expansion $\Box_t = 10^{-5} \, {}^{\circ}C^{-1}$.

202 Permeability changes

203Laboratory tests have shown how the state of stress may affect the hydraulic properties in 204samples (e.g. Liu et al., 2004). Specifically, the medium permeability is related to the fracture 205mechanical aperture *b* and to the effective stress normal to the fracture \Box_n according to the 206following exponential function (Liu et al., 2004):

207 $b=b_r+b_{max}\exp(\alpha \sigma_n)(1)$

208where b_r is the initial mechanical aperture, b_{max} is the mechanical aperture corresponding to zero 209normal stress, \Box is a parameter related to the curvature of the fitting function. The mechanical 210aperture change can also be simply related to the initial state of stress (Rutqvist et al., 2008): 211 $b = b_i + b_{max} (\exp(\alpha \sigma_n) - \exp(\alpha \sigma_i))(2)$ 212, and \Box_{ni} is the initial stress normal to the fractures. In our formulation, compressive stresses are 213considered negative.

214

215Generally most of the EGS sites feature a fracture system striking a certain direction. The 216Newberry Volcano, for example, features a NS-striking fracture system. Such a direction would 217be the *y*-axis in our formulation, and by assuming an initial anisotropic permeability field in a 218porous medium (higher permeability on *y*-axis), we can simulate the effects of the fracture 219network. The permeability would change mostly in the fracture direction (i.e. *y*-axis) and 220changes will be negligible in the other directions. We can calculate the changes in permeability 221along the *y*-direction (\Box_y) as a function of the normal stress (\Box_x) using the cubic law of parallel-222plate flow (Witherspoon at al., 1980):

223
$$\kappa_y = f \frac{b_y^3}{12}(3)$$

224where \Box_y is the permeability in the fracture direction (*y*-axis in our case) and *f* is the fracture 225spacing. b_y is the fracture aperture from Eq. 1 or 2, and it is a function of the stress normal to the 226fracture (\Box_x).

227Using the approach described by Eqs. 1-3 would require a high number of unknown parameters 228that need to be calibrated (b_r or b_i , b_{max} , f, and []). However, we can reduce the number of 229parameters by following an approach for scaling the fracture properties with the initial 230permeability (Liu et al., 2004). Following Eq. 3, we can relate the ratio between initial, 231unstressed permeability and the final permeability to the ratio between the aperture at initial state 232and aperture at final stage:

233
$$\frac{\kappa_y}{\kappa_{yi}} = \left(\frac{b_y}{b_{yi}}\right)^3 (4)$$

234then using a dimensionless parameter $R_b=b_r/b_{max}$, and combining Eq. 4 with Eq. 1, we can write 235for permeability changes (Liu et al., 2004):

236
$$\frac{\kappa_y}{\kappa_{yi}} = \left[\frac{R_b + \exp(\alpha \sigma_x)}{R_b + \exp(\alpha \sigma_{xi})}\right] (5)$$

237where the stress aperture function is related to the dimensionless parameter $R_b = b_r/b_{max}$. Assuming 238the fractures to be identical, R_b will be a constant through the model domain. Using R_b , the 239permeability change factor is independent of initial permeability. We implemented Eq. 5 into 240TOUGH-FLAC and calibrate our model for two parameters only (R_b and []) using data recorded 241during an injection test (see following section).

242

243The variation of stress is not the only process that may affect the permeability. Most of the 244changes will occur after shear reactivation, which is the main mechanism for creating permanent 245permeability enhancement within the EGS reservoir. However, while stress-induced permeability 246changes occur everywhere in the domain along the direction of fractures subjected to aperture 247changes, the shear-induced permeability changes only occur in the portion of the domain 248subjected to shear reactivation. In this work we assumed that the permeability would change by a 249fixed factor if a gridblock were subjected to shear reactivation:

250 $\kappa_i = K_{HS} \cdot \kappa_i^{bHS}(6)$

251for the *i*-direction. K_{HS} is a constant value (set to 500 for the base case analyses, i.e., between 2 252and 3 orders of magnitude), and the index *bHS* refers to the permeability before the 253hydroshearing. Reactivation may occur with random orientation if a threshold pressure is 254reached, then we cannot attribute the changes in a single direction, but we assume the 255permeability changes isotropically if shear reactivation occur. **256**In case of multi-stage shearing, the permeability changes accordingly, i.e. the permeability \Box^{bHS} **257**represents the one before the actual shear stage.

258A similar approach was also recently used by Kelkar et al. (2012) during the study of shear259stimulation at Desert Peak Geothermal Field (Nevada), though limiting the permeability change260to a factor of 15 upon shear failure.

261

262Alternative approaches for modeling hydroshear may involve discrete fracture network models 263or combinations of fracture network and continuum models. For example when using the distinct 264element codes 3DEC or UDEC in which each fracture is explicitly represented, the permeability 265changes in individual rock fractures may be calculated as a result of aperture changes due to 266shear induced dilation based on some constitutive law for single fractures (e.g. Min et al. 2004). 267Other examples involve discrete fracture network models originally developed for groundwater 268 flow and transport extended through a simplified geomechanics approach in which shear failure 269on each fracture is evaluated in an assumed and constant and homogenous external stress field 270(e.g. Willis-Richards 1996; Bruel, 2007). In such an approach, the fracture responses upon shear 271failure may be calculated based on a local elastic solution for an assumed circular shaped 272 fracture of certain radius. For example, based on the radius of the fracture and shear modulus of 273the surrounding rocks, the shear stress drop and maximum shear displacement and associated 274 fracture dilation and permeability change can be calculated. Such an approach has the potential 275of handling a large number of fractures explicitly, but the mechanics is simplified as described 276and does not consider shear induced stress changes or relaxation of the stress field in stimulated 277 areas. A combination approach may involve considering a back-ground fracture network used for 278calculating equivalent continuum properties that for a very fine continuum mesh can be used to 279 represent fractures explicitly by changing properties of elements that are intersected by

11

280individual fractures surfaces (Tezuka et al., 2005; Rutqvist et al., 2013b). The approach for 281permeability change adopted in this study can be considered a rational approach applied to a 282continuum model. However, regardless of the model adopted it usually involves some level of 283calibration against field data as will be discussed in the next section.

284

285 Model calibration

286Model calibration is necessary for understanding whether the system is correctly responding to 287the injection of fluids and whether boundary and initial conditions are properly set. 288Pore compressibility (c_p) and thermal conductivity ([]) within the injection well were calibrated 289to match field data. Moreover, a calibration is needed to assign appropriate values to the 290parameters [] and R_b for the stress-dependent permeability function (Eq. 5).

291The calibration was made simulating a low-pressure injection test, and comparing the resulting 292pressure and temperature profiles along the well with data collected at the NWG 55-29 well 293during a field injection test (September-October 2010).

294According to Davatzes and Hickman (2011) the injection test was performed in two steps. 295During the first period, lasting three days, the injection rate was 0.6 L/s (10 gpm) with an 296injection temperature of 10°C and a well-head pressure of about 5 MPa (750 psi). This period 297was then followed by two weeks of shut-in, before restarting the injection for nine days at a rate 298of 1.4 L/s (22 gpm), with an injection temperature of 10°C and a wellhead pressure of about 8 299MPa (1153 psi).

300Here we performed a simulation with the same, transient injection and we reproduced the same 301observed profiles along the well for pressure and temperature after 3 days at 10 gpm (Fig. 2a and 302b for pressure and temperature, respectively) and after 9 days at 22 gpm (Fig. 2c and d, for 303pressure and temperature, respectively), considering the permeability changes that may arise

304with the evolving effective stresses. Parameters for permeability changes were set constant 305during the two stages as $R_b = 0.2$ and $\Box = 0.13$ MPa⁻¹ after calibration (see Eq. 5). The calibration 306for R_b and \Box is not unique, but the used of the two parameters only reduce the degree of freedom 307of the system.

308The pore compressibility and thermal conductivity were calibrated as well, and the values 309allowing a good match between simulated and measured profiles are listed in Table 1. As stated 310by Davatzes and Hickman (2011), after nine days of 22 gpm injection the well-head pressure was 311lowered to allow the well to be logged; hence, the pressure field data in Figure 2d needs to be 312recalibrated to match a wellhead pressure of 8 MPa (1153 psi) during the majority of the inject-313to-cool operation.

314 STIMULATION AND HYDROSHEARING MODELING

315The stimulation of an EGS reservoir requires that an elevated amount of water be injected into 316the system. For example, at the Basel geothermal system (Switzerland) more than 11500 m³ of 317water were injected in about 5 days (about 30 L/s average flow rate) before peaking at a well-318head pressure of almost 30 MPa and inducing a M_L = 3.4 event (Bachmann et al., 2011). In 319contrast, at the Northwest Geysers EGS Demonstration at The Geysers Geothermal Field 320(California), flow rates reached more than 50 L/s, but bottom-hole pressures were relatively low 321(typically less than 8 MPa), resulting in a large number of small-magnitude seismic events and a 322maximum magnitude event of 2.87 (Garcia et al., 2012; Vasco et al., 2013; Rutqvist et al. 2013a). 323Flow rates during the stimulation at the Newberry EGS Demonstration ranged from about 5 L/s 324up to 20 L/s, with a well-head pressure that peaked at about 16 MPa (Fig. 3, Petty et al., 2013). 325The stimulation was conducted in three different stages, and thermally degradable zonal isolation 326materials (TZIM) were injected between each stage to partially seal stimulated permeable 327fractures and activate stimulation in a new zone. The use of chemical diverters (TZIM) helped to 328stimulate multiple zones in the well bore, resulting then in a more injection or flow capacity (last 329ten days in Fig. 3). Detailed description of the three stages performed during the stimulation can 330be found in Petty et al. (2013).

331Here we are interested in the effects of the stimulation on hydroshearing and how far EGS 332reservoirs can extend under different system conditions. In most of the EGS applications, 333operators try to maintain a fixed injection rate, monitoring the injection pressure as outcome. 334However, we preferred to keep the pressure constant rather the injection rate, since the pressure 335is the main variable for reactivation. There was no fundamental reason, and it was done for better 336understanding of the process of hydroshearing, which is mainly based on pressure rather than on 337the flow rate. The stimulation injection is simulated by fixing the pressure at the top of the well, 338following the average values recorded at the Newberry Volcano EGS Demonstration. Again, it is 339not our goal to reproduce the observed injectivity and flow rates in detail, but it is essential to 340keep our model as close as possible to a real case, in order to achieve reasonable simulation 341results (Fig. 3).

342We simulated the injection in two stages. The first stage lasted for about 28 days with a fixed 343well-head pressure of 7.8 MPa. This is followed by a 10 day shut-in period, before the injection 344restarts for the second stage with 3 days at 7 MPa and 10 days at 14 MPa (Fig. 3, green line). 345Note that we did not consider the effect of diverters (TZIM), which means that an increase in 346pressure results in an increase in flow rate, whereas, as observed at Newberry, the diverters 347sealed the permeable zones, permitting an higher well-head pressure without a substantial 348increase in flow rate prior to the start of subsequent stimulation steps.

349The increase in pressure is necessary to allow the system to reduce the shear strength of fractures 350to less than the shear stress across the fracture, and enable the hydroshearing of fractures with a

14

351wider variety of fracture orientations to occur (Cladouhos et al, 2011). Rather than keeping a 352fixed, high value for the well-head pressure, we preferred to study a case of transient evolution, 353in which the injection starts at a relatively low pressure (7 MPa, first stage) and then is doubled 354after a shut-in (no-injection) period (14 MPa, second stage). The values we chose are within the 355range needed for hydroshearing, but never exceeded the minimum principal stress, so that only 356shear failure can occur, rather than tensile failure or fracturing.

357To estimate the extent of the EGS reservoir, we looked at the zone where the system is subjected 358to hydroshearing. This can be done with a Mohr-Coulomb model: considering a cohesionless 359solid, shear reactivation will occur when the following criterion is satisfied:

360
$$\sigma'_{1c} = N_{\phi}\sigma'_{3}, N_{\phi} = \frac{1 + \sin\phi}{1 - \sin\phi}(7)$$

361where \Box_{1c} is the critical maximum principal effective stress (\Box_{V} or \Box_{zz} in our case), and \Box_{3} is the 362minimum principal effective stress (\Box_{h} or \Box_{xx}). \Box is the frictional angle (frictional coefficient 363 \Box =tan \Box), which is set to 30° for the base-case simulation. In our model we considered as Mohr-364Coulomb solid the Intruded John Day formation only, and the upper John Day formation, which 365is not a highly fractured formation, acts as a barrier for fracture propagation (Fig. 1). This may 366not be true in the field, and the fractures may propagate to shallow depth. 367Equation 7 corresponds to the case in which the media contains fractures with a uniform 368distribution of orientations and equally spaced throughout the reservoir. Using this approach, 369shear reactivation would be induced whenever the maximum principal stress is $N\Box$ times higher 370than the minimum principal stress.

371Stimulation of a single phase, single component system

372The first model we analyze does not take into account the presence of gas within the system. 373Basically, we simulated the stimulation of a geothermal system injecting water into a fully water-

374saturated system. We are fully aware that this is a limiting case in this high temperature 375environment, but would certainly be applicable in lower temperature systems. For this base case 376the friction angle \Box was set to 30°, i.e., a standard value corresponding to a frictional coefficient 377of about 0.6. The constant K_{HS} for shear-enhanced permeability changes was set to 500, i.e. 378corresponding to a 2.7-order magnitude change in permeability to all directions when a shear 379reactivation occurs. This constant was set to a sensible value (Rutqvist and Stephansson, 2003). 380For example, Lee and Cho (2002) have shown after laboratory tests that two orders magnitude 381increase in fracture permeability may arise upon shear displacement.

382Figure 4a shows the pressure transient evolution applied at the top of the injection well (red line) 383and the resulting flow rate associated with the water-saturated system (blue line). Results show a 384 first period during which the flow rate increases up to about 30 kg/s, then hydroshearing begins 385occurring and the flow rate stabilizes to a constant value of about 20 kg/s (or L/s) for the rest of 386the first stage (0 - 28 days with well-head pressure at 7.8 MPa). During the 10 day shut-in 387period, the flow rate is almost nil. Then, during the second stage of stimulation, the well-head 388pressure reaches 14 MPa, and the flow rate peaks at more than 60 kg/s, only to decrease to 40 389kg/s at the end of the stimulation. The simulated flow rates are similar to values observed at EGS 390demonstration sites such as The Geysers (Rutqvist et al., 2013a; Vasco et al., 2013) and 391Newberry Volcano (Petty et al., 2013). Again, note that we are not simulating any injection of 392diverters, and then, during our second stage of the stimulation for a higher well-head pressure, 393we achieve a higher flow rate. The use of diverters proved to be effective at Newberry, where the 394sealing properties of the injected isolating materials sealed the existing permeable fracture 395network, thus resulting in higher pressures but with the same flow rate (Petty et al., 2013). 396Results for pressure changes within the system, and the resulting zone affected by shear 397 reactivation after 28 days of stimulation, are shown in Figure 4b and c, respectively. Pressure

398changes and the hydrosheared zone both extend up to about 400 m from the injection well in the 399NS direction (y-axis), i.e., along the direction we set as the primary fracture strike, which has 400higher permeability (Table 1). Growth of the hydrosheared zone is much smaller in the EW 401direction (*x*-axis), since both the initial permeability and the stress-induced permeability changes 402are smaller in the EW direction, which means that pressure changes do not propagate much in 403that direction. Although the pressurization within the well reaches about 8 MPa (same as the 404wellhead pressure during this first stage), within the system the pressure changes are a few MPa 405smaller, reaching a maximum of 6 MPa at the bottom of the well (Fig. 4b). However, these few 406MPa changes are enough to satisfy the failure criterion and activate the shearing process (Fig. 4074c).

408After the second stage, the increased well-head pressure results in higher-pressure changes 409 within the system. In fact, Figure 4d shows the changes to be around 6 MPa, with maximum 410value of 8 MPa in the region close to the bottom of the well, although these values are still a little 411smaller than the pressure change within the well (about 12 MPa). The resulting hydrosheared 412 region expands somewhat during the second stage, reaching a maximum value of about 500 m 413along the NS-direction (Fig. 4e). We can estimate the extent of the EGS reservoir and calculate 414the volume of the region affected by pressure change and where reactivation occurred: such a 415stimulated volume corresponds to about $9 \cdot 10^7$ m³, i.e., about 0.1 km³.

416Stimulation of a two phase, two component system

417In this section, we consider a system that before the stimulation is completely dry, saturated with 418gaseous CO₂, subjected to cold-water injection. Hydrological and mechanical properties were 419kept the same as the fully water-saturated case. We used an equation of state for a two-420components system (CO_2 and water – EOS2). Carbon dioxide can dissolve in water according to 421the Henry's law and the equation of state is applicable up to the water critical temperature (350 18

422°C), although it does not account for chemical reaction. More details can be found elsewhere 423(Pruess et al., 2011).

424This conceptual model may be somewhat unrealistic, but it represents a good case study as 425compared to the previous water-saturated system. Moreover, CO₂ is a good approximation for a 426volcanic gas: there are a few examples in literature of CO₂ degassing in volcanic regions, such as 427Campi Flegrei caldera (Italy) (e.g. Todesco et al., 2004) or Furnas (Azores) (e.g. Rinaldi et al., 4282012). As previously implemented, we set the friction angle to 30° and the shear-enhanced 429permeability-changes factor to 500.

430Results of the stimulation for this system involving two fluid phases and two fluid components 431are shown in Figure 5. The pressure transient evolution imposed at the well is the same as 432previously (red line, Fig. 5a), and only a few changes are evident in the resulting flow rate 433compared to the case of a water-saturated system (blue line, Fig. 4a and 5a). The small variations 434are at the beginning of the simulation, during which a system with gas requires a higher flow rate 435to displace the gas from the region close to the injection well.

436The resulting pressure increase after the first stage (28 days) is still close to the injection well, 437with an average value of about 5 MPa (Fig. 5c). Although the pressure changes do not propagate 438far from the injection well, the average variation is still similar to the previous water-saturated 439case (Fig. 4b). As a consequence of the poorly distributed pressure changes, at this stage the 440region affected by hydroshearing is limited to a small region around the injection well and 441extends only for about 60 m along NS direction and 13 m along EW direction (Fig. 5d). 442After the second stage (51 days), the pressure changes are still averaging around 5 MPa, 443although within the well the pressure is 14 MPa (Fig. 5d). In any case, at this stage the pressure 444perturbation propagated more, resulting in a larger region where the shear reactivation occurred.

445The hydrosheared region extends about 100 m along the NS direction and about 25.5 m along the 446EW direction (i.e., almost twice that after the first stage, Fig. 5e).

447The results suggest that fractures will propagate much less in a medium initially saturated with a 448compressible gas. In fact, the stimulated volume resulting for an unsaturated medium is about 44910⁷ m³, i.e., about 1 order magnitude smaller than the volume that can be stimulated in a medium 450fully saturated with water. This effect can be explained because of the compressibility of the gas 451phase. In a saturated medium, the water within the system is pushed away from the volume that 452is injected, allowing for the pressure perturbation to move faster, thus reactivating a larger 453region. In an unsaturated medium, the gas phase will be compressed by the injected water. The 454injected water will propagate only to a region close to the injection well, with the pressure 455perturbation following the water front, resulting in the stimulation of a much smaller region.

456 Thermal effects on hydroshearing

457The injection of cold water produces changes in temperature distribution. The changes are 458mostly confined around the injection well, and extend only a few tens of meters, with changes up 459to more than 30 °C. Resulting temperature distribution for both the water-saturated and 460unsaturated cases are shown in Figure 6a and c, respectively. Although we have seen how the 461hydrosheared zone can differ between a water-saturated and an unsaturated system, changes in 462temperature are very similar, as already showed for the resulting flow rates (Fig. 4a and Fig. 5a). 463However, these small and confined changes in temperature may have an effect on the resulting 464stress. In essence, the cooling caused by the injection along the permeable (stimulated) zone 465causes cooling shrinkage that in turn tends to cause an additional reduction in effective stress and 466shear strength. Such shear strength reduction will tend to promote shear failure and propagation 467of the stimulation zone. In fact, if we compared a case that considered a hydro-mechanical (HM, 468i.e. considering $\Box_t = 0$) coupling only with a full THM modeling, we found that the temperature

469changes may help the EGS reservoir growth. In fact, the when conspiring thermal effects, EGS 470reservoir grow about 100 m farther along the NS direction for the case of water-saturated system 471(Fig6b, THM brown, HM yellow). Some small differences are observed for the case of gas-472saturated system, with about 20 m difference between THM and HM modeling, with a slightly 473larger reservoir when thermal effects are taken into account (Fig6d).

474<u>Sensitivity analysis</u>

475We have seen how two systems can respond differently to stimulation if we account for the 476presence of gas. However, the presence of gas within a system is only one of the parameters that 477should be taken into account when simulating a complex system such as a geothermal reservoir. 478Some of the parameters can be taken from field studies: for example, the stress field, which plays 479a huge role in shear reactivation, can be evaluated by *in situ* tests, and by looking at local 480seismicity. The same can be done for the permeability, although most of the parameters studied 481in the laboratory analysis can produce results quite different from those observed in the field. 482Here, we aim to focus only on two main parameters that affect the resulting stimulated volume: 483(1) the constant K_{HS} for the shear-enhanced permeability (Eq. 6) and (2) the frictional angle [] for 484the Mohr-Coulomb criterion (Eq. 7). All the analyses presented in this section were done at the 485end of the second stage, i.e., after 51 days of simulation.

486In the base-case simulations we set a shear-enhanced permeability change factor to achieve 487between a 2 and 3 orders magnitude change (K_{HS} = 500). However, in the field this factor may be 488greater (or smaller), leading to a different system response. Figure 7a shows how the EGS 489reservoir extent varies along the NS direction as a function of this constant for both an 490unsaturated (red line) and saturated system (blue line). Figure 7b shows the stimulated volume as 491a function of the shear-enhanced permeability factor. The extent of the region subject to 492hydroshearing may reach up to 800 m in a water-saturated system, when we consider a factor 10⁴

21

493of permeability changes, or may extend for few tens of meters when we do not consider changes 494in permeability due to hydroshearing ($K_{HS} = 0$). These variations are much smaller for an 495unsaturated system. Even if we use a factor 10⁴, in a system that is initially gas-dominated, the 496reservoir may extend up to about 200 m only. From both Figures 7a and b, we found that the 497hydrosheared region (extent or volume) varies as a logarithmic function of the shear-enhanced 498permeability factor greater than 10 (i.e. $\log_{10}(K_{HS})$ greater than one):

499 $V_{HS}(L_{EGS}) \log_{10}(K_{HS})(K_{HS})(5)$

500where V_{HS} represents the volume subjected to hydroshearing, which is proportional to the EGS 501reservoir length (L_{EGS}). Note the logarithmic scale that Figure 7b.

⁵⁰²The second parameter for which we perform a sensitivity analysis is the friction angle. For the

503 base-case simulations, we used an angle of 30°, which means a friction coefficient of about 0.6.

⁵⁰⁴This factor can play a big role, depending upon the initial stress condition. Considering a fixed,

- ⁵⁰⁵linear stress distribution, with no heterogeneities and variation in depth, the friction angle simply
- ⁵⁰⁶ regulates how much overpressure is needed to satisfy the failure criterion (Eq. 7). Results of the

⁵⁰⁷ sensitivity analysis for this parameter are shown in Figures 6c and d, for the EGS extent along

⁵⁰⁸ the NS direction and for the stimulated volume, respectively. Both these variables seem to

⁵⁰⁹ change linearly with the friction angle, with values ranging between 600 m (volume $2 \cdot 10^8 \text{ m}^3$) 510and 250 m (volume $5 \cdot 10^7 \text{ m}^3$) for a water-saturated medium with an increase in the friction angle 511 resulting in a smaller stimulated volume and linear extent, and hydrosheared region is almost 512costant for an unsaturated medium: only a 40 m variation in EGS extent over the considered 513range of values.

514<u>CONCLUSIONS</u>

515During stimulation of an Enhanced Geothermal System (EGS) it is always difficult to predict 516how far the reservoir and fracture network can grow. Moreover, creating a new fracture network 517requires elevated pumping pressure and flow rates in order to fracture the rock. One mechanism 518that has been proposed to reduce the cost is so-called *hydroshearing*, which involves reactivating 519an existing fracture network by a shear process, taking advantage of the fracture surface 520roughness, which should naturally maintain the enhanced permeability. The pressure needed for 521hydroshearing has to be below the minimum principal stress, but without exceeding it, thus 522avoiding then the creation of new tensile fractures. Once the fractures are reactivated, some 523isolating, thermally degrading material may be injected to plug the fracture network, and this will 524permit stimulating multiple fracture zone without drill rig or setting multiple packers. Moreover, 525the injection of chemical diverters will permit injecting at a higher pressure (hence reactivating 526some other, deeper zones) without changing the flow rate.

527However, some questions need to be answered: will the presence of gas within the system help or 528reduce the hydroshearing reactivation? How much can the permeability change after reactivation, 529and how is that change related to the extent of the EGS reservoir and to the stimulated volume? 530The aim of this paper was to answer to these questions. Through the use of the TOUGH-FLAC 531simulator, we carried out a coupled thermo-hydro-mechanical modeling of EGS stimulations. We 532accounted for a Mohr-Coulomb solid that can fail when a criterion is satisfied. Upon fracture 533reactivation, we assumed a change in medium permeability, which allows for a better 534propagation of the pressure perturbation. Taking into account previous simulations performed for

22

535The Geysers Geothermal Field and starting with data collected at the Newberry Volcano EGS 536Demonstration site, we simulated the stimulation by fixing the overpressure at the top of an 537injection well. Pressure transient evolution was taken as the average of field values measured at 538Newberry Volcano during stimulation. Although our aim was not to reproduce any observed 539variations, we used field data to keep our model as realistic as possible.

540We first presented the results for two limiting cases: (1) a water-saturated medium and (2) a 541medium initially saturated with CO₂ in the gas phase. Results suggest that an EGS reservoir will 542extend much further in a medium initially fully saturated with water than in a gas-phase 543dominated system. We explained this effect as owing to the compressibility of the gas phase. In a 544water-saturated medium, the native water within the system is pressurized by the water injected 545into the well, allowing the pressure perturbation to propagate faster and reactivating fractures 546over a larger region. In an unsaturated medium, the much more compressible gas phase will be 547compressed by the injected water, which will propagate only to a region close to the injection 548well, following the water front, finally stimulating a much smaller region. Thermal effects on 549stress may help to reach shear failure. Although temperature changes are small and confined 550within tens of meter from the injection well, thermal effects on stress are evident at earlier time 551and helped the EGS reservoir to grow. A hydro-mechanical modeling resulted in smaller 552hydrosheared region after the stimulation.

553The presence of gas is not the only parameter affecting the growth of an EGS reservoir. Many 554parameters are involved that may play a significant role, such as the medium initial permeability, 555the natural fracture-network orientation, and the stress distribution. We performed a sensitivity 556analysis on two key parameters that are generally hard to measure in the field: (1) the factor for 557the shear-enhanced permeability changes and (2) the friction angle for the failure criterion. 558Results showed that the extent of the EGS reservoir and the volume subjected to shear

23

559reactivation strongly depend upon these two parameters. We found that in our system, the EGS 560extent (or reactivated region volume) will depend logarithmically upon the constant used to 561relate permeability change associated with hydroshearing, and linearly upon the frictional angle. 562These variations are more accentuated in a water-saturated system.

563The volume subjected to hydroshearing (stimulated volume) ideally should also represent the 564region where the microseismicity cloud should be. However, natural-system heterogeneities in 565the stress field and permeability may play a significant role, and the seismicity cloud may not 566exclusively represent only the region that has been stimulated. For example, a brittle material at 567shallow depth may be affected by deformation and stress transfer coming from a deeper 568overpressure, and seismicity may then be induced at a shallower depth than we would expect. 569Technical issues may be involved as well: for example, a leakage from the cased well may allow 570the water to move at shallow depths, where it is much easier to cause hydroshearing, or 571hydrofracturing, since the *in situ* would be stress smaller, generally depending on the depth.

572ACKNOWLEDGMENTS

573This work was supported by the Department of Energy under Award Number DE-EE0002777 574and by the American Recovery and Reinvestment Act (ARRA), through the Assistant Secretary 575for Energy Efficiency and Renewable Energy (EERE), Office of Technology Development, 576Geothermal Technologies Program, of the U.S. Department of Energy under contract no. DE-577AC02-05CH11231. Technical review comments by Matt Uddenberg of the AltaRock company, 578and Patrick F. Dobson and Pierre Jeanne at the Berkeley Lab, as well as editorial review by Dan 579Hawkes at the Berkeley Lab are all greatly appreciated.

580**REFERENCES**

581

24

582Bachmann, C. E., S. Wiemer, J. Woessner, S. Hainzl, Statistical analysis of the induced Basel 5832006 earthquake sequence: Introducing a probability-based monitoring approach for Enhanced 584Geothermal System, *Geophys. J. Int.*, 186(2) (2011). doi:10.1111/j.1365-246X.2011.05068.x. 585

586Baria R, Baumgärtner J, Rummel F, Pine RJ, Sato Y (1999a) HDR/HWR reservoirs: concepts, 587understanding and creation. Geothermics 28:533–552.

588

589Bruel, D., 2007, Using the migration of the induced seismicity as a constraint for fractured hot 590dry rock reservoir modeling: International Journal of Rock Mechanics and Mining Sciences & 591Geomechanics Abstracts, 44, 1106–1117.

592

593Cappa, F., J. Rutqvist, Seismic rupture and ground accelerations induced by CO₂ injection in the 594shallow crust, *Geophys. J. Int.*, 190, 1784-1789 (2012).

595

596Cladouhos, T. T., M. Clyne, M. Nichols, S. Petty, W. L. Osborn, L. Nofziger, Newberry Volcano 597EGS Demonstration Stimulation Modeling, *GRC Transactions*, 35 (2011).

598

599Cladouhos, T. T., W. L. Osborn, S. Petty, D. Bour, J. Iovenitti, O. Callahan, Y. Nordin, D. Perry,
600P. L. Stern, Newberry Volcano EGS Demonstration – Phase I results, *Proceedings of 37th*601*Workshop on Geothermal Reservoir Engineering*, Stanford, California, January 30 – February 1,
6022012.

603

604Cladouhos, T. T., S. Petty, B. Larson, J. Iovenitti, B. Livesay, R. Baria, Toward More Efficient 605Heat Mining: A Planned Enhanced Geothermal System Demonstration Project, *GRC* 606*Transactions*, 33, 165-170 (2009).

607

608Cladouhos, T. T., S., Petty, Y. Nordin, M. Moore, K. Grasso, M. Uddenberg, M. Swyer, B. Julian, 609G. Foulger, Microseismic Monitoring of Newberry Volcano EGS Demonstration, *Proceedings of* 610*the 38th Workshop on Geothermal Reservoir Engineering*, Stanford, California, February 11 - 13, 6112013.

612

613Davatzes N. C. and S. H. Hickman, Preliminary Analysis of Stress in the Newberry EGS Well 614NWG 55-29, *GRG Transactions*, 35, 323-332 (2011).

615

616Evans KF, Cornet FH, Hashida T, Hayashi K, Ito T, Matsuki K, Wallroth T (1999) Stress and 617rock mechanics issues of relevance to HDR/WDR engineered geothermal systems: review of 618developments during the past 15 years. Geothermics 28:455–474

619

620Garcia, J., Walters, M., Beall, J., Hartline, C., Pingol, A., Pistone, S., and Wright, M., Overview 621of the Northwest Geysers EGS demonstration project, *Proceedings of the 37th Workshop on* 622*Geothermal Reservoir Engineering*, Stanford, California, January 30 – February 1, 2012 623

624Kelkar, S., K. Lewis, S. Hickman, N. C. Davatzes, D. Moos, G. Zyvoloski, Modeling coupled
625thermal-hydrological-mechanical processes druing shear stimulation of an EGS well,
626*Proceedings of the 37th Workshop on Geothermal Reservoir Engineering*, Stanford, California,
627January 30 - February 1, 2012.

629Lee, H. S., and T. F. Cho, Hydraulic Characteristic of Rough Fractures in Linear Flow under 630Nomrla and Shear Load, *Rock Mech. Rock Eng.*, 35 (4), 299-318, 2002.

631

632Li, Y., J. Wang, W. Jung, A. Ghassemi, Mechanical Properties of Intact Rock and Fractures in
633Welded Tuff from Newberry Volcano, *Proceedings of 37th Workshop on Geothermal Reservoir*634*Engineering*, Stanford, California, January 30 – February 1, 2012.

635

636Liu, H. H., J. Rutqvist, G. Zhou, G. S. Bodvarsson, Upscaling of normal stress-permeability
637relationship for fracture network obeying the fractional levy motion. In: Stephansson O., Hudson
638J. A., Jing L., editors. *Coupled THMC processes in geo-system: fundamentals, modelling,*639experiments and applications. Oxford: Elsevier; p.263-268 (2004).

640

641Itasca, FLAC3D, Fast Lagrangian Analysis of Continua in 3 Dimensions, Version 4.0, 642Minneapolis, Minnesota, Itasca Consulting Group (2009).

643

644Makurat A, Barton N, Rad NS (1990) Joint conductivity variation due to normal and shear 645deformation. In: Barton N, Stephansson O (eds) Rock joints. Balkema, Rotterdam, pp 535–540 646

647Min KB, Rutqvist J., Tsang C.-F., and Jing L. Stress-dependent permeability of fracture rock 648masses: a numerical study. *Int. J. Rock Mech. & Min. Sci*, 41, 1191-1210 (2004).

649

650Petty, S., Y. Nordin, W. Glassley, T. T. Cladouhos, M. Swyer, Improving Geothermal Project 651Economics with Multi-Zone Stimulation: Results from the Newberry Volcano EGS

652Demonstration, *Proceedings of the 38th Workshop on Geothermal Reservoir Engineering*,
653Stanford, California, February 11 - 13, 2013.

654

655Pine RJ, Batchelor AS (1984) Downward migration of shearing in jointed rock during hydraulic 656injections. Int J Rock Mech Mining Sci 21:249–263

657

658Pruess, K., C. M. Oldenburg, and G. Moridis, *TOUGH2 User's Guide, Version 2.1*, Report
659LBNL-43134 (revised), Lawrence Berkeley National Laboratory, Berkeley, Calif. (2011).
660

661Riahi, A., B. Damjanac, Numerical Study of Hydro-Shearing in Geothermal Reservoirs with a 662Pre-Existing Discrete Fracture Network, *Proceedings of the 38th Workshop on Geothermal* 663*Reservoir Engineering*, Stanford, California, February 11 - 13, 2013.

664

665Rinaldi, A. P., J. Rutqvist, Modeling of deep fracture zone opening and transient ground surface 666uplift at KB-502 CO₂ injection well, In Salah, Algeria, *Int. J. Greenh. Gas Contr.*, 12, 155-167 667(2013). doi:10.1016/j.ijggc.2012.10.017.

668

669Rinaldi, A. P., J. Vandemeulebrouck, M. Todesco, F. Viveiros, Effects of atmospheric conditions 670on surface diffuse degassing, *J. Geophys. Res. - Solid Earth*, 117, B11201 (2012). 671doi:10.1029/2012JB009490.

672

673 Rutqvist, J., Status of TOUGH-FLAC simulator and recent applications related to coupled fluid 674 flow and crustal deformations. *Comput. Geosci.*, 37, 739-750 (2011).

675

676Rutqvist J., Dobson P.F., Garcia J., Hartline C., Jeanne P., Oldenburg C.M., Vasco D.W., Walters 677M. The Northwest Geysers EGS Demonstration Project, California: Pre-stimulation Modeling 678and Interpretation of the Stimulation. Mathematical Geosciences, In press, (2013a)

679

680Rutqvist J., Leung C., Hoch A., Wang Y., and Wang Z. Linked multicontinuum and crack tensor 681approach for modeling of coupled geomechanics, fluid flow and transport in fractured rock. 682International Journal of Rock Mechanics and Geotechnical Engineering, 5, 18–31 (2013b).. 683

684Rutqvist, J., B. Freifeld, K.-B. Min, D. Elsworth, Y. Tsang, Analysis of thermally induced 685changes in fractured rock permeability during 8 years of heating and cooling at the Yucca 686Mountain Drift Scale Test, *Int. J. Rock Mech. Min. Sc.*, 45, 1373-1389 (2008).

687

688Rutqvist, J., O. Stephansson, The role of hydromechanical coupling in fractured rock 689engineering, *Hydrogeology Journal*, 11 (1), 7-40 (2003).

690

691Rutqvist, J., C. F. Tsang, Multiphysics processes in partially saturated fractured rock: 692Experiments and models from Yucca mountain, *Rev. Geophys.*, 50(3) (2012). 693doi:10.1029/2012RG000391.

694

695Rutqvist, J., Y.-S. Wu, C.-F. Tsang, G. Bodvarsson, A modeling approach for analysis of coupled
696multiphase fluid flow, heat transfer, and deformation in fractured porous rock, *Int. J. Rock Mech.*697*Min. Sc.*, 39, 429-442 (2002).

698

699Sonnenthal, E. L., N. Spycher, O. Callahan, T. Cladouhos, and S. Petty, A thermal-hydrological-700chemical model for the Enhanced Geothermal System Demonstration Project at Newberry 701Volcano, Oregon, *Proceedings of the 37th Workshop on Geothermal Reservoir Engineering*, 702Stanford, California, January 30 - February 1, 2012.

703

704Tester, J.W. et al., 2006 The future of geothermal energy. Part 1—Summary and part 2—Full705report. Massachusetts Institute of Technology, Cambridge, MA.

706

707Tezuka K., Tamagawa T., and Watanabe K. 2005. Numerical Simulation of Hydraulic Shearing in 708Fractured Reservoir. Proceedings World Geothermal Congress 2005. Antalya, Turkey, 24-29 709April 2005

710

711Todesco, M., J. Rutqvist, G. Chiodini, K. Pruess, C. M. Oldenburg, Modeling of recent volcanic 712episodes at Phlegrean Fields (Italy): geochemical variations and ground deformation, 713*Geothermics*, 33, 531-547 (2004).

714

715Vasco, D. W., J. Rutqvist, A. Ferretti, A. Rucci, F. Bellotti, P. Dobson, C. M. Oldenburg, J.
716Garcia, M. Walters, C. Hartline, Monitoring deformation at The Geysers geothermal field,
717California using C-band and X-band Interferometric Synthetic Aperture Radar, *Geophys. Res.*718Lett., accepted for publication (2013). doi:10.1002/grl.50314.

719

720Willis-Richards, J., K.Watanabe, and H. Takahashi, 1996, Progress toward a stochastic rock 721mechanics model of engineered geothermal systems: Journal of Geophysical Research, 101, 17, 722481–17, 496. 723

724Witherspoon, P. A., J. S. Y. Wang, K. Iwai, J. E. Gale, Validity of cubic law for fluid flow in a 725deformable rock fracture, *Water Resour. Res.*, 16, 1016-1024 (1980).

726

727Ziagos J., Phillips B.R., Boyd L., Jelacic A., Stillman G. and Eric Hass E. A technology roadmap 728for strategic development of enhanced geothermal systems. Proceedings of the 38th Workshop on 729Geothermal Reservoir Engineering , Stanford, California, February 11 - 13, 2013.

Tables

	5	0	1 1			/1	5	0	– 1	5
733 \Box_{rock} rock density. <i>D</i> rock grain specific heat. \Box thermal conductivity. c_p pore compressibility										
		5	Ū	Ĩ	_		5 1 1		U U	
		New	/herry_D	eschutes	John Dav	Intruded	John Dav	Cased well	Onen w	ell

Table 1. Hydrological properties. \Box_i initial (stress-free) permeability along *i*-direction. \Box porosity.

	NewDerry-Deschutes	Joini Day	Intruded John Day	Cased well	Open wen
$\Box_x (m^2)$	10 ⁻¹⁷	2.6·10 ⁻¹⁶	5·10 ⁻¹⁸	10-20	10-16
$\prod_{y} (\mathbf{m}^2)$	10-17	$2.6 \cdot 10^{-16}$	10 ⁻¹⁷	10 ⁻²⁰	10^{-16}
\Box_{z} (m ²)	10-17	2.6·10 ⁻¹⁶	5·10 ⁻¹⁸	10-8	10-8
	10	5	3	95	100
\Box_{rock} (kg/m ³)	2400	2400	2400	-	-
D (J/kg °C)	1000	1000	1000	800	800
[][(W/m °C)	1.80	2.15	2.20	2.20	1.80
c_{p} (Pa ⁻¹)	3.2·10 ⁻⁹	3.2·10 ⁻⁹	3.2·10 ⁻⁹	-	-

734Figure captions



736*Figure 1*. Mesh and boundary conditions for modeling the stimulation of an EGS reservoir.

737Initial pressure, temperature and stress condition follows pre-stimulation analysis at the

738Newberry Volcano EGS Demonstration, as well as rock properties and distribution.



Figure 2. Model calibration. (a) Pressure well log (blue, dashed line) and simulated pressure (red 741line) after 3 days (09/27/2010) of 0.6 kg/s (10 gpm) injection rate. (b) Temperature well log 742(blue, dashed line) and simulated temperature (red line) after 3 days (09/27/2010) of 0.6 kg/s (10 743gpm) injection rate. (c) Pressure well log (black, dashed line) and simulated pressure (red line) 744after 9 days (10/20/2010) of 1.4 kg/s (22 gpm) injection rate. (d) Temperature well log (blue, 745dashed line) and simulated temperature (red line) after 9 days (10/22/2010) of 1.4 kg/s (22 gpm) 746injection rate.



748*Figure 3*. Example of stimulation. Wellhead pressure (blue) and injection rate (red) observed at 749Newberry Volcano EGS Demonstration. Gap in timeline is when stimulation pumps were offline. 750The green line represents an average of the wellhead pressure, which is used as input for our 751modeling of an EGS system to study the hydroshearing. The shut in period between our first and 752second stage is indicated in figure. Note that at Newberry during that period the injection 753continued at a very low rate.



Figure 4. Simulation results for a system fully saturated with water. (a) Applied wellhead 756pressure (red) and resulting flow rate (blue). (b, c) Resulting pressure and hydrosheared zone at 757the end of the first stage (28 days). (d, e) Resulting pore pressure changes and hydrosheared zone 758at the end of the second stage (51 days).



Figure 5. Simulation results for a system initially saturated with carbon dioxide. (a) Applied 761wellhead pressure (red) and resulting flow rate (blue). (b, c) Resulting pressure and hydrosheared 762zone at the end of the first stage (28 days). (d, e) Resulting pore pressure changes and 763hydrosheared zone at the end of the second stage (51 days).



Figure 6. (a) Temperature changes after the stimulation and (b) resulting hydrosheared zone for
766THM (brown) and HM (yellow) modeling for a water-saturated system. (c) Temperature changes
767after the stimulation and (d) resulting hydrosheared zone for THM (brown) and HM (yellow)
768modeling for a system initially fully saturated with gas.



770*Figure 7*. Sensitivity analysis. Blue and red lines refer to saturated and unsaturated systems, 771respectively. EGS reservoir extent along NS-direction (a) and hydrosheared volume (b) as a 772function of the constant for shear-enhanced permeability (K_{HS} , Eq. 6). EGS reservoir extent along 773NS-direction (c) and hydrosheared volume (d) as a function of the friction angle for the Mohr 774Coulomb criterion (\Box , Eq. 7).