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SOURCE PROVENANCE OF OBSIDIAN ARTIFACTS FROM PREHISTORIC SITES IN KANSAS

by

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Report Prepared for

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INTRODUCTION

The analysis here of 38 obsidian artifacts from a number of prehistoric sites in Kansas exhibits a mix of source provenance typical of sites from this region. The assemblage is dominated by the two major source groups from the Jemez Mountains in northern New Mexico, and one artifact produced from the Malad, Idaho source.

LABORATORY SAMPLING, ANALYSIS AND INSTRUMENTATION ANALYSIS AND INSTRUMENTATION

 All archaeological samples are analyzed whole. The results presented here are quantitative in that they are derived from "filtered" intensity values ratioed to the appropriate xray continuum regions through a least squares fitting formula rather than plotting the proportions of the net intensities in a ternary system (McCarthy and Schamber 1981; Schamber 1977). Or more essentially, these data through the analysis of international rock standards, allow for interinstrument comparison with a predictable degree of certainty (Hampel 1984).

The trace element analyses were performed in the Archaeological XRF Laboratory, Department of Earth and Planetary Sciences, University of California, Berkeley, using a Spectrace/ThermoNoranTM QuanX energy dispersive x-ray fluorescence spectrometer. The spectrometer is equipped with an air cooled Cu x-ray target with a 125 micron Be window, an xray generator that operates from 4-50 kV/0.02-2.0 mA at 0.02 increments, using an IBM PC based microprocessor and WinTraceTM reduction software. The x-ray tube is operated at 30 kV, 0.14 mA, using a 0.05 mm (medium) Pd primary beam filter in an air path at 200 seconds livetime to generate x-ray intensity Kα-line data for elements titanium (Ti), manganese (Mn), iron (as FeT), zinc (Zn), gallium (Ga), rubidium (Rb), strontium (Sr), yttrium (Y), zirconium (Zr), and niobium (Nb). Zinc and gallium are only reported for the basalt artifacts, since they are generally in low quantities in western North American obsidian. Trace element intensities were converted to concentration estimates by employing a least-squares calibration line established for each element from the analysis of international rock standards certified by the National Institute

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of Standards and Technology (NIST), the US. Geological Survey (USGS), Canadian Centre for Mineral and Energy Technology, and the Centre de Recherches Pétrographiques et Géochimiques in France (Govindaraju 1994). Line fitting is linear (XML) for all elements but Fe where a derivative fitting is used to improve the fit for the high concentrations of iron and thus for all the other elements. Further details concerning the petrological choice of these elements in obsidian is available in Shackley (1995, 1998 and 2005; also Mahood and Stimac 1991; and Hughes and Smith 1993). Specific standards used for the best fit regression calibration for elements Ti through Nb include G-2 (basalt), AGV-1 (andesite), GSP-1, SY-2 (syenite), BHVO-1 (hawaiite), STM-1 (syenite), QLO-1 (quartz latite), RGM-1 (obsidian), W-2 (diabase), BIR-1 (basalt), SDC-1 (mica schist), TLM-1 (tonalite), SCO-1 (shale), all US Geological Survey standards, BR-N (basalt) from the Centre de Recherches Pétrographiques et Géochimiques in France, and JR-1 and JR-2 (obsidian) from the Geological Survey of Japan (Govindaraju 1994). In addition to the reported values here, Ni, and Cu, were measured, but these are rarely useful in discriminating glass sources, are poorly measured with the Cu target, and are not generally reported.

The data from the WinTrace software were translated directly into Excel for Windows software for manipulation and on into SPSS for Windows for statistical analyses. In order to evaluate these quantitative determinations, machine data were compared to measurements of known standards during each run. RGM-1 is analyzed during each sample run for obsidian artifacts to check machine calibration (see Table 1). Compilation and discussion of RGM-1 analyses are available at http://www.swxrflab.net/anlysis.htm. Source assignments were made with reference to the source standard library at Berkeley (Shackley 1995, 1998, 2005).

DISCUSSION

 The provenance of the obsidian artifacts dominated by Jemez Mountains sources is not surprising, and is rather typical of late period sites in the Plains (Baugh and Nelson 1987;

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Hawley and Hughes 1999; Hughes 1988; Hughes and Lees 1991; Hughes and Roper 1999; Tables 1 and 2, and Figure 1 here). The two major sources in the Sierra de los Valles is Valles Rhyolite, often called Cerro del Medio in the vernacular, and Cerro Toledo Rhyolite often called Obsidian Ridge in the vernacular (see Shackley 2005). The two sources have very different eruptive and depositional histories that are important for prehistoric procurement.

 The Cerro Toledo Rhyolite glass was erupted slightly earlier in a caldera collapse creating Plinian events around the east side of the Cerro Toledo caldera (Shackley 2005). Because of the large ash flow tuffs directed to the east and south, a large portion of the ash and rhyolite glass was ultimately eroded into the Rio Grande system, and now can be found as marekanites as far as Chihuahua (Church 2000; Shackley 2005). It is generally impossible to determine whether raw material was procured from the primary domes in the Sierra de los Valles or in secondary contexts except by using a rough nodule size index. The slightly later Valle Grande caldera collapse that produced the Cerro del Medio and other domes producing Valles Rhyolite glass have not yet eroded outside the caldera, and so must have been originally procured in the caldera. For a distance as great as that between northern New Mexico and Kansas, this may not be an issue, but procurement in the late period in the Jemez Mountains likely required some payment, that may not have been necessary in the secondary deposits of the Rio Grande.

 Generally, Idaho and Wyoming sources tend be associated with Woodland and Archaic period sites in the Plains, but they also occur in later period sites, either as scavenged raw material from earlier contexts or through exchange relationships with the northwestern Plains (Hawley and Hughes 1999; Logan et al. 2001).

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Table 1. Elemental concentrations for the obsidian archaeological specimens. All measurements in parts per million (ppm).

¹ These samples were slightly smaller than the optimum sample size for EDXRF and the concentrations may be somewhat outside the source standard data (see Davis et al. 1998). 2 The high Ti concentrations in these samples is likely due to the labeling (white out) that has a high titanium content. The other concentrations appear to be unaffected.

Table 2. Obsidian source provenance by site.

Figure 1. Y, Sr, Nb three-dimensional plot of elemental concentrations for obsidian archaeological samples from all sites.