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Mapping the global distribution of C₄ vegetation using observations and optimality theory

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Plants with the C₄ photosynthesis pathway typically respond to climate change differently from more common C₃-type plants, due to their distinct anatomical and biochemical characteristics. These different responses are expected to drive changes in global C₄ and C₃ vegetation distributions. However, current C₄ vegetation distribution models may not predict this response as they do not capture multiple interacting factors and often lack observational constraints. Here, we used global observations of plant photosynthetic pathways, satellite remote sensing, and photosynthetic optimality theory to produce an observation-constrained global map of C₄ vegetation. We find that global C₄ vegetation coverage decreased from 17.7% to 17.1% of the land surface during 2001 to 2019. This was the net result of a reduction in C₄ natural grass cover due to elevated CO₂ favoring C₃-type photosynthesis, and an increase in C₄ crop cover, mainly from corn (maize) expansion. Using an emergent constraint approach, we estimated that C₄ vegetation contributed 19.5% of global photosynthetic carbon assimilation, a value within the range of previous estimates (18–23%) but higher than the ensemble mean of dynamic global vegetation models (14 ± 13%; mean ± one standard deviation). Our study sheds insight on the critical and underappreciated role of C₄ plants in the contemporary global carbon cycle.

C₄ is one of the three photosynthetic pathways for terrestrial plants¹ and is reported to account for 18–23%^{2–4} of global photosynthesis. C₄ plants also drive wildfire dynamics in tropical and subtropical ecosystems⁵. C₄ plants first evolved in the low atmospheric CO₂ environment of the Oligocene Epoch, roughly 24–35 million years ago⁶. They developed distinct biochemical and anatomical characteristics to enrich CO₂ concentration at the site

of Rubisco carboxylation in leaves, thereby reducing photorespiration and enhancing carbon-fixation rates⁷. These characteristics produce different climate sensitivities in C₄ plants compared to more prevalent C₃ plants^{4,8}, and thus are expected to cause a shift in C₄ plant distributions and their contribution to global photosynthesis under contemporary and future climate change^{8–10}.

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Many previous studies have examined C_4 plant responses to multiple environmental factors. A consensus is that since most C_4 species originated in lower atmospheric CO_2 concentrations^{11,12}, they are expected to benefit less from rising CO_2 concentrations compared to C_3 plants. Meanwhile, higher temperatures are expected⁴ and reported^{13–15} to favor C_4 over C_3 photosynthesis, because the affinity of O_2 to Rubisco relative to CO_2 becomes stronger with increasing temperature and also due to differing solubilities of CO_2 and O_2 with increasing temperature. This should produce an advantage for the carbon concentrating mechanism of C_4 species, especially under high temperatures¹⁶. Hence C_4 species are characteristic of tropical and subtropical ecosystems. Correspondingly, since C_4 photosynthesis is less limited by CO_2 than C_3 photosynthesis, it achieves a higher photosynthetic quantum yield and photosynthetic rates under high light, especially under high temperatures¹⁷. C_4 species also should have a carbon assimilation advantage in arid environments^{18,19} due to their higher water use efficiency (i.e., less water loss through stomata for equivalent carbon gain) than C_3 species, though under humid conditions this advantage could be limited⁹. Contemporary climate change, such as elevated CO_2 , rising temperatures, and changing rainfall patterns, can therefore lead to temporal and spatial shifts in the relative advantages of C_4 to C_3 photosynthesis. For instance, the differential response to a changing environment has been linked to observed woody plant encroachment in tropical Africa, where an increase in precipitation and elevated atmospheric CO_2 levels are hypothesized to have caused a net decrease in C_4 grassland distribution^{20,21}. However, we currently lack a consensus on how the relative advantages of C_4 to C_3 photosynthesis change at the global scale, as regional studies have reported contrasting results and different driving factors—such as increased C_4 grass distribution due to increased temperature²², decreased distribution due to elevated CO_2 ¹⁴ or no overall trend²³. Understanding of how climate change has impacted C_4 vegetation constitutes a major challenge due to its role in global photosynthesis and the terrestrial carbon cycle.

C_4 vegetation overwhelmingly consists of natural grasses and crops using the C_4 pathway. One prominent approach to estimate the distribution of C_4 natural grasses is based on the crossover-temperature model, which predicts that a particular month is determined to favor C_4 grasses over co-occurring C_3 grasses when the mean daytime air temperature is $>22^\circ C$ and precipitation in that same month is $\geq 25\text{ mm}$ ^{2,24,25}. This approach is based on each pathway's relative carbon assimilation as a function of temperature, and thus the crossover temperature is dependent on atmospheric CO_2 concentration with higher crossovers at higher CO_2 levels. A few efforts to model C_4 vegetation distribution have further incorporated the seasonality of precipitation^{26–29}, or mean annual temperature and precipitation^{8,22,26}, but so far they are only validated and applied at the regional scale. Some dynamic global vegetation models (DGVMs) allow adjustment of C_3 and C_4 grass distribution based on the difference between simulated C_3 and C_4 photosynthesis or the difference between their net primary productivity³⁰, or based on the simulations from bioclimate distribution models in each time step, with the baseline C_4 map acquired from remote sensing land cover classifications^{31,32}. Some cohort-based DGVMs further consider competition for resources³³ and disturbances³⁴ when simulating C_4 distributions. In general, current estimates of the distribution of C_4 vegetation adopt a wide range of assumptions and generate rather different results¹⁰.

Uncertainty in global C_4 grass distribution is further exacerbated by the lack of ground observations for validation and then for model extrapolation, since previous models often relied on either local datasets^{26,27} or literature reviews of C_4 grass presence and absence² for validation. This issue has become less prominent recently as some studies have used continental scale (i.e., North America) C_4 plots²⁵ and ^{13}C isotopic records^{33,35} to validate C_4 grass distribution models. Meanwhile, the distribution of C_4 crops has been collated and

estimated in some open datasets^{36–39}. These datasets are based on Food and Agriculture Organization (FAO) census and national reporting of the harvested area for major C_4 crops (i.e., maize, sorghum, millet, sugarcane), which comprised 24% of the global harvested area³⁷, and are supplemented by total cropland area change from FAO and remote sensing⁴⁰. In particular, the Land-Use Harmonization dataset version 2 (LUHv2) is the principal gridded land use dataset for the assessment of global carbon budgets³⁶ and future climate change in CMIP6⁴¹, in which C_4 crop area over time is explicitly reported. Changes in global C_4 crop distribution and the related contribution to global photosynthesis have yet to be evaluated, except for a few studies that have examined the C_4 crop distribution for specific years^{2,42,43}.

Here we quantify the global C_4 vegetation distribution (including natural grasses and crops) and its contribution to global photosynthesis, as well as examine changes in C_4 vegetation distribution over the past two decades. To do so, we use photosynthetic optimality theory to estimate the relative advantage of C_4 to C_3 photosynthesis over the global land surface, and then use the estimated difference in combination with observations to infer global C_4 grass distribution. The optimality model includes a wide array of selective drivers for C_4 grass distribution - CO_2 , temperature, light, aridity, nitrogen, and their interactions⁴⁴ (see Methods), which are advances over previous crossover-temperature approaches which include CO_2 , temperature, and a precipitation threshold. The optimality model estimates the optimal leaf photosynthetic rate for C_3 and C_4 plants, along with optimal stomatal conductance and root/shoot carbon allocation based on growing season climate, with a target to maximize carbon gain with minimized water loss⁴⁴. We further take advantage of multiple open-access databases of C_4 species richness and coverage (i.e., the global TRY database⁴⁵, a dataset for the contiguous United States (the DG dataset)²³ and a subset of the Nutrient Network (NutNet)⁴⁶), global grassland fraction maps from remote sensing (i.e., as the majority of C_4 plant cover is non-woody⁴⁷), in combination with the optimality model simulations to acquire data-constrained estimates of C_4 grass distribution for the past 20 years. Meanwhile, we obtain and examine C_4 crop distribution using multiple open datasets^{36,39}. We further use an emergent constraint technique—a method to infer an unobservable variable from an observable variable based on the large spread of estimates of both variables from DGVMs (see Methods) —to estimate the contribution of C_4 plants to global photosynthesis. By quantifying how C_4 vegetation distribution and photosynthesis have changed over recent decades, our study improves understanding of historical changes in terrestrial photosynthesis and the global carbon cycle.

Results

C_4 photosynthetic advantage and C_4 grass coverage

We found a strong positive relationship between the observed C_4 grass coverage (i.e., the % of grassland area covered by C_4 grass species) and the relative advantage of C_4 photosynthesis (AC_4) over C_3 photosynthesis (AC_3) estimated by the optimality model (denoted as the AC_4/AC_3 - C_4 coverage relationship hereafter; see Methods; Fig. 1b). With the increase in modeled AC_4/AC_3 , the observed C_4 coverage increased and then gradually plateaued. When $AC_4/AC_3 = 1$, C_4 accounts for only 5.2% of grassland cover; when $AC_4/AC_3 = 2.5$, C_4 coverage approaches 100% (Fig. 1b). Across global non-woody regions, AC_4/AC_3 ranged from 0.5 to 2.5, with a mean of 1.9 (Fig. 1a). Importantly, we obtained C_4 coverage observations from multiple sources (i.e., TRY and DG; see Methods), which have different geographic representations and spatial resolutions. However, the AC_4/AC_3 - C_4 coverage relationships are similar when using different observations (Fig. 1b), affirming the robustness of the relationship for C_4 coverage estimation. Using the relationship between AC_4/AC_3 and C_4 coverage (Fig. 1b), and the global AC_4/AC_3 estimated from the optimality model (Fig. 1a), we estimated the global C_4 grass coverage (% of grassland

covered by C_4 ; Fig. 1c). We found C_4 grass coverage followed a clear climatic gradient, and tended to be greater under warmer conditions (Fig. 1d).

The global distribution of C_4 vegetation

After predicting the C_4 grass coverage (% of grassland covered by C_4 grasses; Fig. 1c), we overlaid a global grassland fraction map from remote sensing (see Methods and Fig. S1) to estimate the actual C_4 natural grass area abundance (% of the land surface covered by C_4 grasses; Fig. 2a). From 2001 to 2019, C_4 natural grass accounted for $14.8 \pm 1.3\%$ (mean \pm one standard deviation) of the non-frozen land surface area (Fig. 2a), while C_4 crops accounted for $2.8 \pm 0.3\%$ (Fig. 2c). The total estimated C_4 area abundance was $17.5 \pm 1.4\%$ (Fig. 2e). There were several C_4 natural grass hotspots (i.e., $>30\%$ C_4 area abundance) across continents (Fig. 2a): the Great Plains in North America, the savannas in Southern Brazil, the savannas in Africa, the grasslands in Central Asia, and Northern Australia. Meanwhile, we found the main C_4 crop zones were in central North America, the Sahel region, and the west coast of India (Fig. 2c). The disagreement between remote sensing-based grassland fraction maps (Fig. S4), along with the uncertainty in the AC_4/AC_3 - C_4 coverage relationship (Fig. 1b), incurred uncertainties in the C_4 natural grass distribution (Fig. 2b)—the uncertainty typically ranges between 1 and 3% of the land surface area, though in regions like Australia the uncertainty could be as high as 6–7% (Fig. 2b).

The changes in C_4 vegetation distribution

Based on our simulation of C_4 natural grass distribution for the past two decades and the C_4 cropland distribution from the LUHv2 dataset (see Methods), we found the overall area of C_4 vegetation decreased from $17.7 \pm 1.4\%$ (mean \pm one standard deviation) in 2001–2005 to $17.1 \pm 1.4\%$ in 2015–2019, as a net effect of a decrease in C_4 natural grasses from $15.0 \pm 1.3\%$ to $14.2 \pm 1.3\%$, and an increase in C_4 crops from $2.6 \pm 0.3\%$ to $3.0 \pm 0.3\%$ (Fig. 3b). The change in C_4 shows large spatial heterogeneity (Fig. 3a). In particular, C_4 natural grass area decreased all over the globe, except for the central Europe and parts of the western U.S. (Fig. 3c). C_4 crop area increased in most parts of the world, except for central North America where there was the largest decrease, and Europe where there were slight decreases (Fig. 3e).

The increase in C_4 crop area was often at the expense of decreasing C_4 natural grasses, as we found that in regions where there were both C_4 crops and C_4 natural grasses, more than 50% of the regions showed C_4 natural grasses decreased but C_4 crops increased. Meanwhile, only 14% showed that C_4 crops and C_4 natural grasses increased simultaneously, 28% showed they both decreased, and only 7% of the region showed that C_4 natural grasses increased and C_4 crops decreased (Fig. 3d). Our attribution analysis suggested that elevated CO_2 was the dominant reason for the decrease in C_4 natural grass distribution, while the impacts of temperature and water stress (i.e., soil moisture and vapor pressure deficit) were positive (i.e., increase C_4 coverage) over the study period (Fig. 3f). Most of the increase in C_4 crops, as we analyzed from another independent dataset

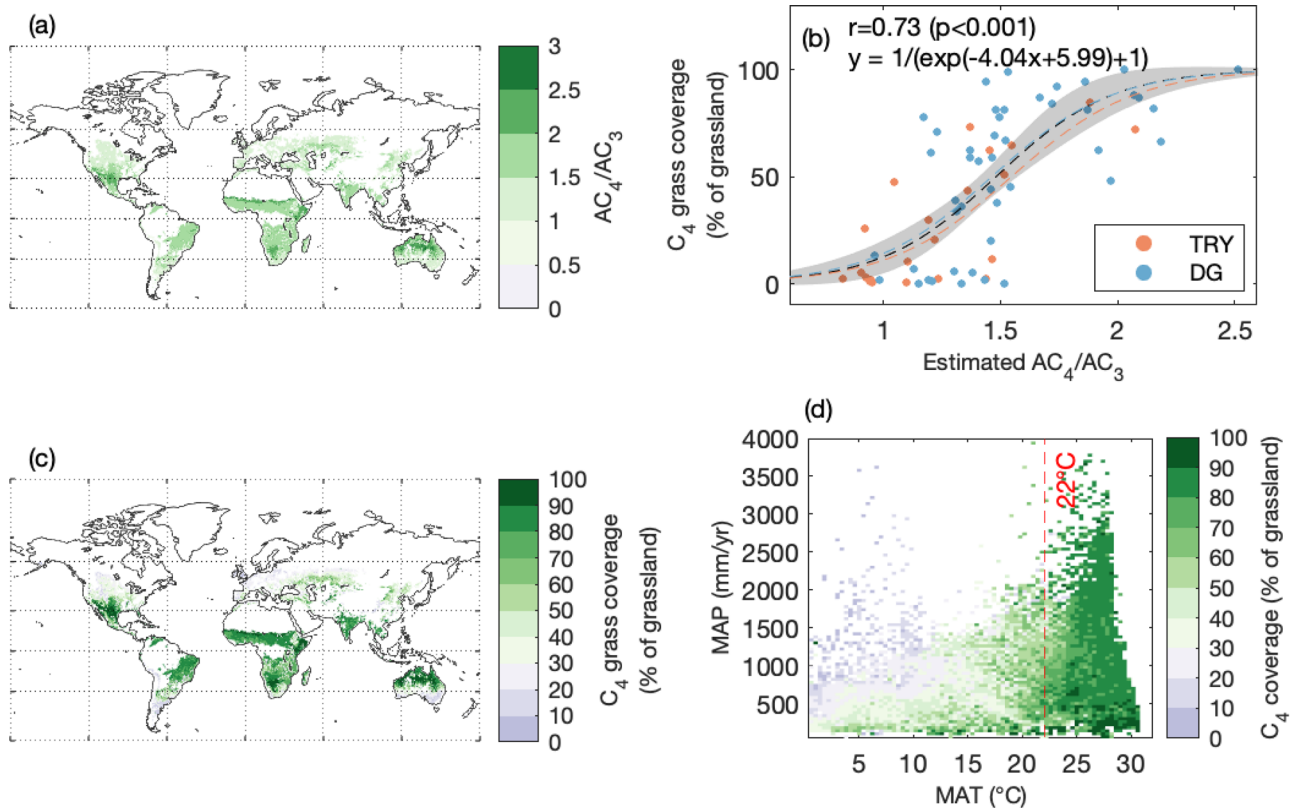


Fig. 1 | C_4 natural grass coverage estimated by the optimality model. a the ratio of C_4 to C_3 photosynthesis estimated by the optimality model (AC_4/AC_3) over global non-woody regions; **(b)** the relationship between observed C_4 coverage (% of grassland) and estimated C_4/C_3 photosynthetic ratio by the optimality model. C_4 coverage observation obtained from difference sources (i.e., TRY, DG datasets; please see methods); gray shaded area indicates the uncertainty range for the relationship between AC_4/AC_3 and C_4 coverage (i.e., 95% confidence interval). The

black line represents the regression using both the TRY and DG datasets, while the red and blue dash lines represent the regression using either the TRY or the DG dataset. **c** C_4 grass coverage (% of grassland) over the globe, which can be regarded as the potential C_4 area abundance when grassland covers 100% of the land surface; **(d)** C_4 coverage in a climate space of mean annual temperature (MAT: °C) and mean annual precipitation (MAP: mm/yr).

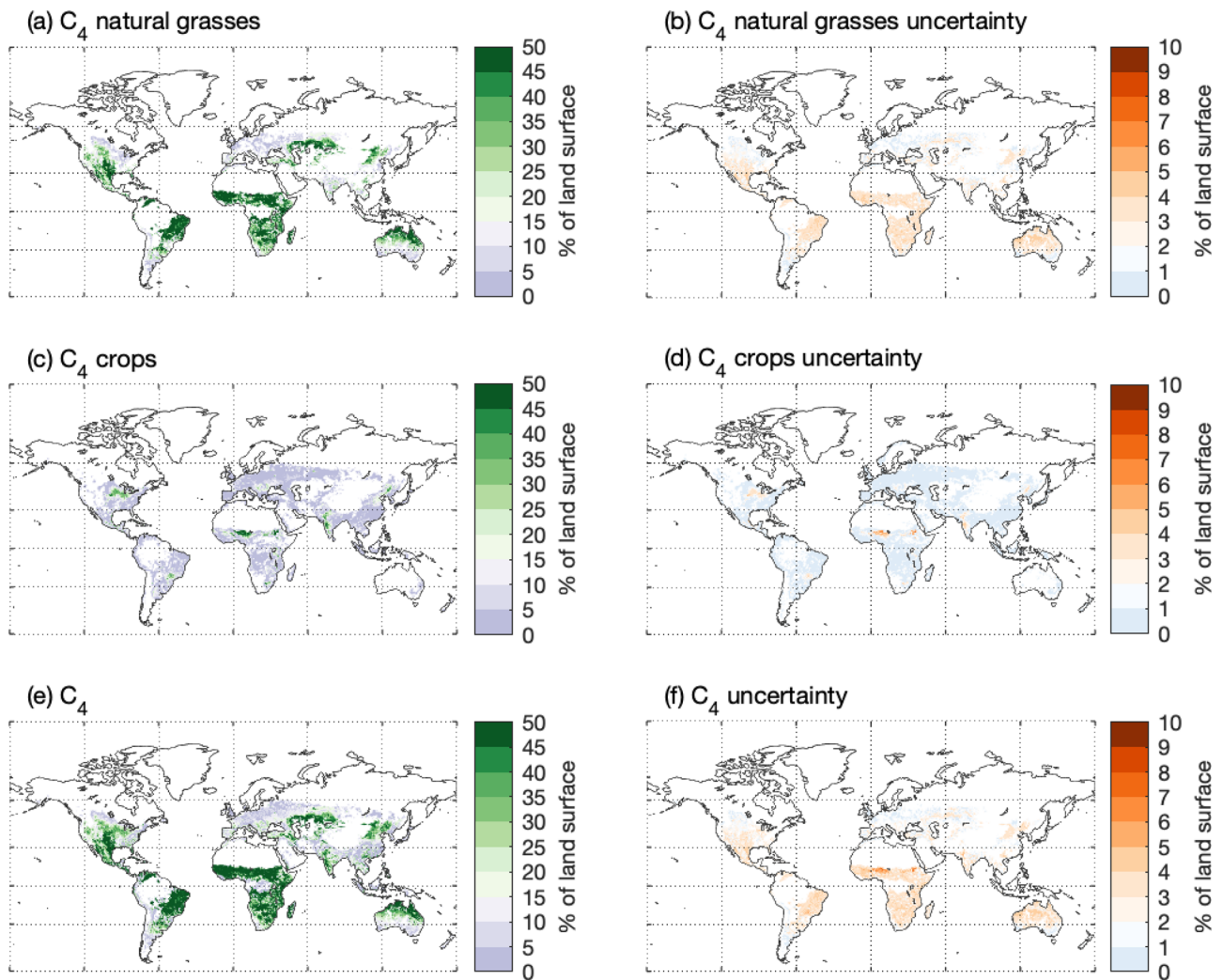


Fig. 2 | The modeled global distribution of C₄ vegetation and associated uncertainties. The area occupied by (a) C₄ natural grasses, (c) C₄ croplands and (e) all C₄ vegetation (unit: % of the land surface). The uncertainties of the area

abundance of (b) C₄ natural grasses, (d) C₄ croplands and (f) all C₄ vegetation (unit: % of the land surface).

on major C₄ crop distributions³², came from the expansion of maize in South America and eastern Europe (Fig. 3f; Fig. S9; see Methods).

The contribution of C₄ vegetation to global photosynthesis

The changes in the C₄ area can cause associated changes in total C₄ photosynthesis, thus impacting global carbon cycle dynamics. Current ensemble of DGVMs predicted that C₄ vegetation contributed from 2% to 40% of global photosynthesis, on 7% to 23% of the global vegetated land surface area (Figs. S6, S7). The large spread of model estimates indicates the various assumptions adopted in C₄ vegetation distribution and potentially the different parameterizations for C₄ photosynthesis in DGVMs. Despite these wide inter-model variations, it is possible to infer emergent constraints on C₄ vegetation contributions to the carbon cycle. To quantify the contribution of C₄ photosynthesis to global photosynthesis, we established an emergent constraint ($p < 0.01$) between the DGVM-simulated occupied area and percentage contribution of C₄ natural grasses and crops to global photosynthesis, respectively. We found that with a 1% increase in area, C₄ natural grass contribution to global photosynthesis increased by 1.10% (Fig. 4a), while the contribution of C₄ crops increased by 1.16% (Fig. 4b). We also conducted a grid cell-level emergent constraint analysis and acquired similar ranges of slopes (Fig. S10). The lower coefficient of emergent constraint for C₄ grass (i.e., 1.10) than the

coefficient for C₄ crop (i.e., 1.16) suggests that croplands tend to have a higher ecosystem photosynthetic rate compared to grasslands over the same area.

We further explored the changes in the coefficients of emergent constraints over the past two decades. We note that for C₄ grasslands and C₄ croplands, the coefficients all slightly decreased in the past two decades. The coefficient of C₄ crops decreased from 1.16 to 1.15, while the coefficient of C₄ grasses decreased from 1.11 to 1.10 from 2001 to 2019 (Fig. 4c). Interestingly, the coefficients are all greater than 1, highlighting that the per unit area photosynthetic rate of C₄ is generally higher than that of the remaining C₃ vegetation. With the likely increase of global photosynthesis in recent decades, the decreasing coefficients of C₄ indicated that C₄ photosynthesis increased at a slower pace than other (mostly C₃) vegetation. By applying the estimated area of C₄ to the annual emergent constraint coefficients (Fig. 4c), we found that the global C₄ natural grass contribution to photosynthesis decreased from $16.5 \pm 1.5\%$ (mean \pm one standard deviation) in 2001–2005 to $15.5 \pm 1.5\%$ in 2015–2019, the C₄ crop contribution to photosynthesis increased from $3.0 \pm 0.3\%$ to $3.4 \pm 0.4\%$, and in total the C₄ contribution to global GPP decreased from $19.7 \pm 1.9\%$ to $19.0 \pm 1.9\%$ (Fig. 4d). The reported value is greater than the ensemble mean reported by the DGVMs ($14 \pm 13\%$), and is within the range of previously modeled values ($18\text{--}23\%^{2-4}$).

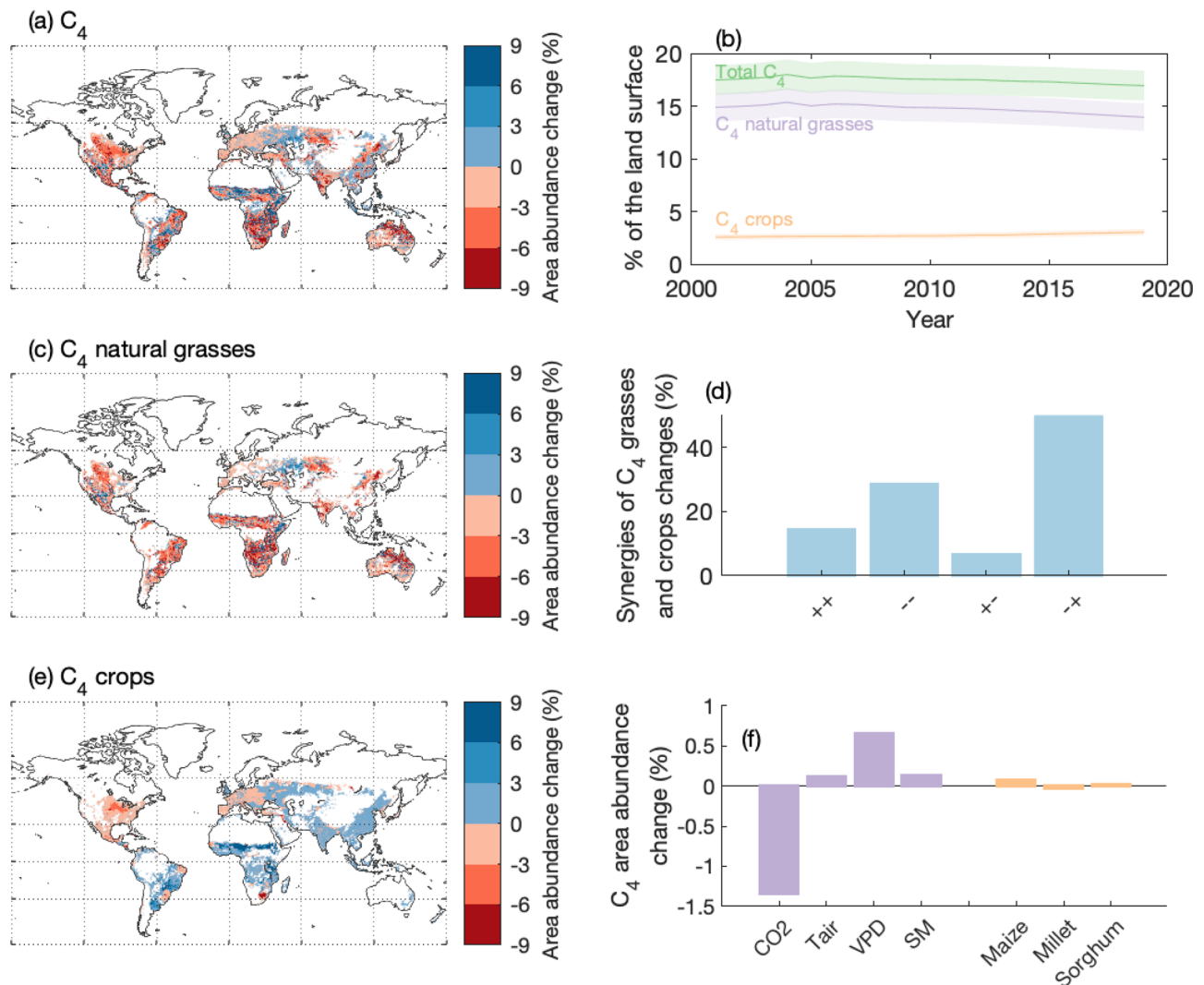


Fig. 3 | Changes in the global distribution of C₄ vegetation between 2001–2019. Spatial distributions of changes in (a) total C₄ vegetation, (c) C₄ natural grasses and (e) C₄ crops from 2001 to 2019; b The changes in the total area of C₄ vegetation, C₄ natural grasses and C₄ crops, in percentages of global vegetated land surface; d the synergies of changes in C₄ natural grasses and C₄ crops, where ++ means both C₄ natural grasses and C₄ crops area abundance increased, -- means both decreased, +- means C₄ natural grasses increased and C₄ crops abundance decreased, -+ means the opposite; f the drivers for the change in C₄ natural grasses

and C₄ crops area abundances. Climate drivers include atmospheric CO₂ concentration, air temperature (T_{air}), vapor pressure deficit (VPD) and soil moisture (SM) during the growing seasons. In (b), the uncertainty for C₄ crop area reported, the uncertainty for C₄ natural grass area is the combination of the uncertainty in remote sensing-based grassland fraction and the uncertainty of the AC₄/AC₃ · C₄ coverage relationship (Fig. 1b), and uncertainty for C₄ vegetation is the combination of uncertainties of C₄ crop and C₄ natural grass areas.

Discussion

In this study, we estimated the global C₄ vegetation distribution and quantified the changes in C₄ vegetation distribution and photosynthesis over the past two decades, using an optimality photosynthesis model, photosynthetic pathway records from global/regional databases and remote sensing observations. On average, from 2001 to 2019, C₄ plants occupied $17.5 \pm 1.4\%$ of the global vegetated surface and contributed $19.5 \pm 1.9\%$ of global photosynthesis, within the range of previous estimates (i.e., 18–23% for photosynthesis) but are greater than the estimates from the ensemble mean of DGVMs ($13 \pm 8\%$ for area and $14 \pm 13\%$ for photosynthesis). C₄ total area and C₄ contribution to global photosynthesis both decreased over this period (i.e., 0.6% of the land surface and 0.7% of global GPP), which resulted from the increases in C₄ crop area and its contribution to global photosynthesis, and the decreases in C₄ natural grass area and its contribution to global photosynthesis.

Our study suggests that the decrease in C₄ natural grass distribution was primarily driven by elevated CO₂, in accordance with

previous theoretical and experimental works that showed C₄ advantage in carbon assimilation over C₃ decreased with rising CO₂^{48–50}. There is also evidence showing many grassland and savanna areas have been invaded by C₃ woody species, with increased atmospheric CO₂ proposed as a major driver for the encroachment^{51,52}. The decrease in the emergent constraint coefficients demonstrated that the impact of elevated CO₂ on C₄ and C₃ were included in most DGVMs (Fig. 4c). Evidence for the historical expansion of C₄ over geological time scales seems to support our conclusion, in particular for Africa where CO₂ dominates the C₄ grassland expansion or decline^{50,53,54}, while in Central Asia⁵⁵, Australia⁵⁶, central China⁵⁷ and central US grasslands⁵⁸ reports show hydroclimatic change impacted C₄ grass distribution. Our results also show that the changes in soil moisture and VPD over the past two decades largely induced positive impacts on C₄ grass distribution, though globally their impacts were unable to offset the negative changes driven by elevated CO₂ (Fig. 3f; Fig. S8).

To assess the distribution of C₄ vegetation, we noted that there were three types of abundance used in previous literatures: for

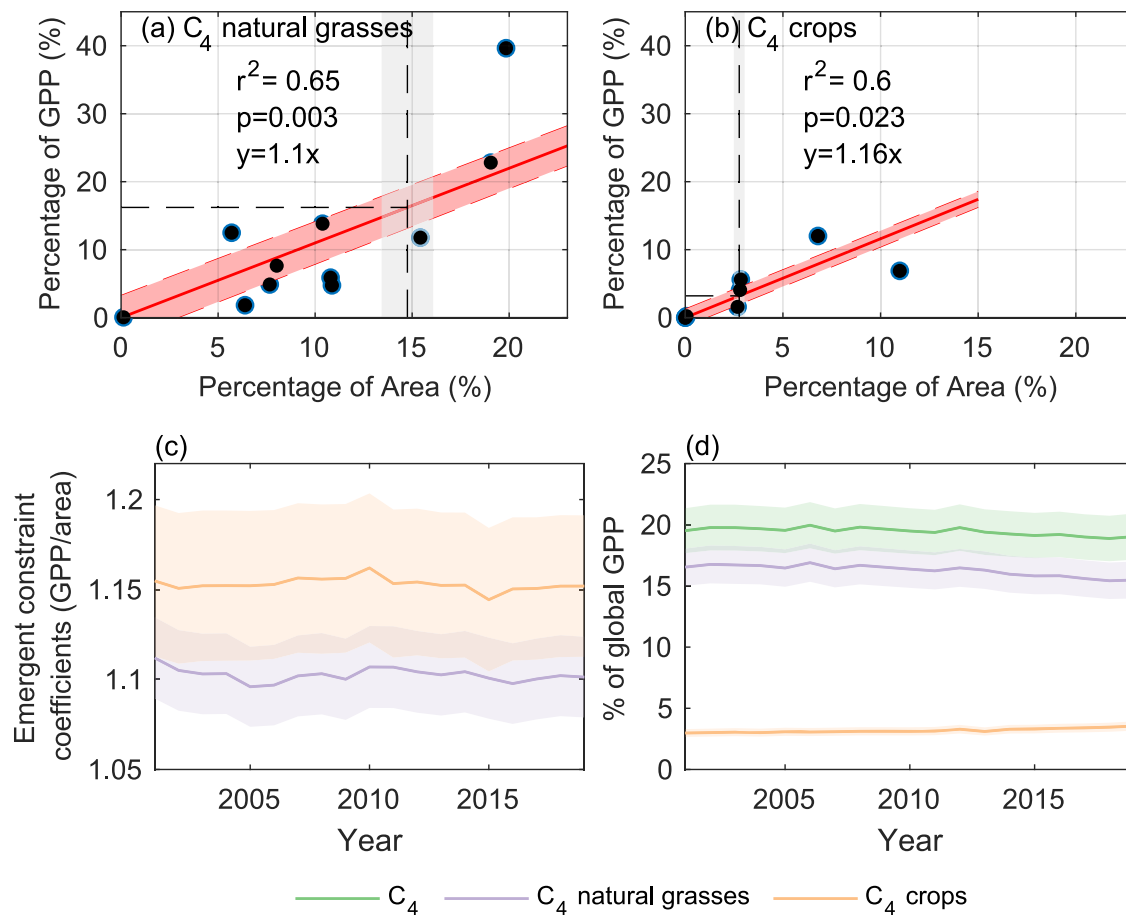


Fig. 4 | The contribution of C₄ vegetation to global photosynthesis. The emergent constraints between the percentage of area occupied and the percentage of global photosynthesis contributed by (a) C₄ natural grasses and (b) C₄ crops, based on the estimates from an ensemble of the DGVMs; c changes in emergent constraint coefficients (i.e., the slopes of the linear regressions in (a, b) from 2001 to 2019 for C₄ natural grasses and C₄ crops). The uncertainties in (a, b, c) were

quantified as one standard error (i.e., SE) by bootstrapping models when getting emergent constraint; d The contributions of C₄, C₄ natural grasses and C₄ crops to global GPP from 2001 to 2019, whereas the uncertainty was quantified as one SE by bootstrapping the uncertainty range of C₄ areas and the uncertainty range of the emergent constraint coefficients.

modeling practice, we often used the area abundance, that is, the percentage of the land surface occupied by C₄ plants. However, the area abundance was challenging to observe over large scales, and it was rare to have direct observations of large-scale area abundance other than approximations from remote sensing³². Most observations of C₄ presence and cover are at the species level and, therefore, observation-based studies often report the relative species richness of C₄ plants^{13,26}. A few studies reported biomass abundance, i.e., the percentage of local biomass contributed by C₄ plants, and biomass abundance is often much higher than the species abundance^{27,29}, echoing some studies suggesting that species abundance should be used with caution to infer biomass abundance and productivity⁵⁶. In this study, we developed a conversion factor to translate C₄ species richness into C₄ area abundance, using plot-level concurrent measurements of both from NutNet. We found 1% increase in C₄ species richness led to $1.51 \pm 0.15\%$ in C₄ area abundance (Fig. S3). The conversion factor enabled us to translate global observations of C₄ species richness into C₄ coverage to develop the AC₄/AC₃ - C₄ coverage relationship in our study (Fig. 1b, S1).

Compared to the previously estimated C₄ vegetation distribution from a crossover-temperature model² (Fig. S5), our study provided a similar area occupied by C₄ vegetation (20.5 million km² versus 21.1 million km²). We find a lower estimate for Africa, where our estimate of C₄ area abundance showed a more nuanced gradient compared to the estimate from the crossover-temperature

model (Fig. S5). We suspect this is partly due to their difference in relating C₄ photosynthetic advantage to C₄ grass coverage (% of grassland covered by C₄ grasses) - while our study used the AC₄/AC₃ - C₄ grass coverage relationship to gradually adjust C₄ grass coverage, the crossover-temperature model assumes all grasslands in the pixel are C₄ grassland as long as the monthly climate satisfies the crossover criteria (e.g., mean daytime air temperature is >22 °C and precipitation in that same month is ≥25 mm). Therefore, for regions where monthly climate meets the crossover criteria (e.g., sub-Saharan Africa), the crossover-temperature model tends to estimate 100% C₄ grass, however, it was not the case for the optimality approach in those regions (Fig. 1c). Meanwhile, we estimated considerably higher C₄ grass cover in Central Asia, which is consistent with the reported prevalence of C₄ species in the region^{26,59} - the high abundance in these inland regions were potentially due to harsh environments characterized by high maximum temperature and aridity levels, which favor the growth of C₄ plants over C₃. We also note that our approach may underestimate C₄ grass distribution in some mesic savanna ecosystems - such as the longleaf pine savannas in the Southeastern US - where a C₄ understory exists beneath a C₃ canopy^{60,61}. The underestimation is likely because the optimality model predicted no photosynthetic advantage for C₄ plants over C₃ under low light conditions in understory (Fig. 1a), and the remote sensing products reported low grassland fraction in the region (Fig. S4).

In our study, we estimated the annual distribution of C_4 grasses using the growing season mean climate; however, many locations have seasonal shifts between C_4 and C_3 grass dominance depending on seasonal climate variations⁴. For example, in the grasslands of south-east Australia, a recent study suggests C_4 dominance is the highest in summer when there is high temperature and low precipitation, while other seasons have more C_3 vegetation²⁸. Therefore, for modeling the seasonal variation of carbon fluxes from seasonal changes in C_4 grass distribution, the crossover-temperature approach based on monthly climate and weighted by a vegetation index like NDVI could be more useful². We acknowledge that the distribution of our observations was not uniform across the globe, with North America being better represented compared to other regions (Fig. S2). As an additional test to validate our estimation of C_4 vegetation distribution, we compared the C_4 grass coverage estimated in our study (Fig. 1c) with the C_4 coverage estimated from isotopic measurements and remote sensing in Australia⁶² (Fig. S11). This validation demonstrates a strong agreement ($r=0.69$, $p<0.01$) between the two independent estimates, affirming the robustness of our estimates in under-sampled regions.

We also need to highlight that the optimality approach was based on the assumption that C_4 grass distribution is determined by the photosynthetic advantage of C_4 compared to C_3 , and the photosynthetic advantage is largely dependent on local climate⁴⁴. While this assumption is also adopted by the crossover-temperature model, it neglects the role of grass phylogeny in determining C_4 grass distributions. Some studies have suggested that grass clades (i.e., Pooideae for C_3 , and PACCMAD for C_3 and C_4) are perhaps more critical than photosynthetic pathway for determining C_4 and C_3 grass distributions, at least along temperature gradients^{19,63,64}. This lack of consideration on phylogeny in C_4 grass distribution models may impact predictions of future distributions, as C_3 grasses from certain clades have less competitive disadvantage compared to C_4 grasses in a warmer world.

Other environmental changes that can impact C_4 grass distribution include fire and nitrogen deposition. Fire can influence C_4 coverage and productivity either through creating open canopies for light capture by C_4 plants or as a characteristic of semi-arid environments that provide a photosynthetic advantage for C_4 ⁶⁵. Recent woody plant encroachment suggests fire had a central role^{51,66} in the formation of grasslands and the rise of C_4 dominant grasslands in late Neogene⁶⁷ and late Miocene⁶⁸. For recent decades, since globally the trend of fire occurrence is still very uncertain with strong regional variations in the trend⁶⁹, we were unable to quantify its impact on C_4 grass distribution. Neither the crossover-temperature C_4 model, nor the optimality model we used, incorporates the role of fire in C_4 dynamics at present. However, since our approach used annual grassland fractional maps based on satellite remote sensing, the impacts of fire at the annual scale were implicitly considered. Some DGVMs have incorporated fire-relevant processes, however, we found they estimated lower C_4 grass abundance (i.e., $9.6 \pm 6.7\%$) than those models that do not include fires (i.e., $11.8 \pm 3.2\%$). Additionally, since C_4 plants have higher photosynthetic nitrogen use efficiency than C_3 ⁷⁰, anthropogenic nitrogen deposition⁷¹ might have impacted the relative advantage of C_4 to C_3 photosynthesis. Previous studies have suggested C_4 plants tend to have a higher photosynthetic rate than C_3 across a spectrum of nitrogen supply - meaning C_4 plants have photosynthetic advantage on infertile soils and the advantage will be enhanced by increased nitrogen availability, which can be used to increase C_4 leaf area⁷². In this study, we used a data-driven product of leaf nitrogen content and remote sensing leaf area index to simulate C_4 grass distribution (see Methods), which may have implicitly accounted for the effect of nitrogen deposition or limitation on C_4 photosynthesis.

In this study, we used LUHV2—the primary dataset used in current global carbon cycle modeling and climate forecasting^{36,41}—to examine the changes in C_4 cropland and reported an increase in C_4 cropland

area. However, we note a key source of uncertainty in this dataset: the historical simulation of C_4 crop distribution used a constant fraction of C_4 crop cover for each global grid cell, based only on observations circa 2000⁴¹. Therefore, the increase in C_4 crop area in LUHV2 could just reflect an increase in all croplands rather than real C_4 expansion. To reduce the uncertainty caused by this issue, we used another cropland dataset³⁹, which dynamically simulated the area of 17 major crop types (including main C_4 crops maize, millet and sorghum) based on annual FAO census of crop harvested areas. This analysis confirmed our results regarding the increase in C_4 cropland mainly due to the expansion of maize (Fig. 3f), with a similar spatial pattern reported (Fig. S9). However, the sum of the three main C_4 species only increased by 0.1%, relatively lower than what we see from the LUHV2 dataset (i.e., 0.4%). We can conclude there was an increase in C_4 cropland, though the magnitude of increase should be subject to further examinations.

In conclusion, we used a combination of plant photosynthetic pathway records, remote sensing, and an optimality-based photosynthesis model to estimate the global C_4 coverage and the magnitudes of C_4 photosynthesis and their variations over the past two decades. We infer that C_4 vegetation covered on average 17.5% of the global land surface over the period from 2001 to 2019, while C_4 grass cover decreased due to elevated CO_2 and C_4 crop cover increased because of corn (maize) expansion. We predict that C_4 photosynthesis accounted for 19.5% of the global total photosynthesis, with an increased contribution from C_4 crops and a decrease from C_4 natural grasses during this period. Our study offers an updated and more observationally constrained estimate of C_4 vegetation distribution and photosynthesis, thereby improving our understanding of potential future C_4 changes and enhancing the quantification of the global carbon budget.

Methods

The overarching framework

The distribution of C_4 vegetation overwhelmingly consists of C_4 natural grasses and C_4 crops. To estimate the C_4 natural grass distribution, we first used an optimality photosynthesis model⁴⁴ to simulate the optimal photosynthetic assimilation rates of C_4 and C_3 plants (noted as AC_3 and AC_4 , respectively) using 0.5×0.5 degree gridded historical climate (i.e., CRU-JRA2020), soil⁷³ and leaf nitrogen content⁷⁴. We calculated the ratio of AC_4 to AC_3 , and established a statistically significant ($p<0.01$) relationship between AC_4/AC_3 and the observed C_4 coverage from multiple databases—the TRY database⁴⁵ and the DG dataset²⁵ based on an assumption that larger AC_4/AC_3 indicates higher C_4 grass coverage (% of grassland covered by C_4 grasses). Using the AC_4/AC_3 - C_4 coverage relationship, we estimated the C_4 grass coverage (a.k.a. potential C_4 abundance when grasses cover 100% of the land surface) from estimated AC_4/AC_3 for the globe. We lastly overlaid the C_4 coverage map to a global map of grassland fraction from remote sensing to acquire actual C_4 grass abundance (% of the land surface covered by C_4). The workflow is presented in Fig. S1. Meanwhile, for C_4 crop distribution, we directly used the estimates from the LUHV2-2019 dataset, in which C_4 crop distribution is estimated from FAO survey and satellite remote sensing.

Processing observational C_4 records

We acquired 61,588 georeferenced records of photosynthetic pathways from the TRY database (last accessed 2022 June), among them, there were 2269 records of C_4 . The time range of the record covers roughly the past 50 years. We first removed the woody species from the records, based on species names and an index table from the TRY database (<https://www.try-db.org/TryWeb/Data.php#3>), as our study aimed to examine C_4 grass distribution and the global cover of C_4 forests is not extensive⁴⁷. After the step, we kept 13,919 records for non-woody species, among which 1963 were C_4 . We further removed 82 records that belong to major C_4 crops (i.e., maize, sugarcane, millet,

and sorghum) and kept 1881 records. We then aggregate these records to 10×10 degree cells, in each cell we calculate the species richness of C_4 (i.e., number of C_4 species/total number of herbaceous species, the numbers were derived from the available records in the TRY database). The gridded values of C_4 species richness would be further used to constrain the optimality model to estimate global C_4 grassland coverage. Here we use the large-size grid cell to make sure there were enough samples in each cell to acquire a meaningful estimate of C_4 abundance – in this analysis, each cell should have at least 50 species (i.e., C_3 and C_4 in total). We used 1619 (out of 1881) C_4 species records in this aggregation step, and obtained 23 10×10 degree cells for the analysis (Fig. S2).

Note here the C_4 abundance from the TRY database was species richness, not equal to the area abundance that is often used in DGVMs. To acquire C_4 grass coverage (% of grassland covered by C_4) from C_4 species richness (% of grass species that is C_4), we used an open dataset from the global nutrient network (NutNet) that has paired C_4 species richness and C_4 area abundance (% of the land surface covered by C_4) to infer their relationship⁴⁶ (Fig. S3). The dataset includes species-specific coverage records as well as the grass species richness data collected in 25 m² plots across 34 sites. Each site has between 1 and 6 control plots. We only used the data from the control plots, excluding plots that underwent nutrient addition treatments. To avoid the uneven distribution of data samples, we grouped the paired observations by their C_4 species richness, and for each species richness we get a mean C_4 area abundance and the standard deviation of the C_4 grass coverage. We then conducted 1000 linear fittings (i.e., with an intercept of 0, since C_4 grass coverage should be 0 when C_4 species abundance is 0), and for each fitting we used randomly sampled C_4 grass coverage values (i.e., based on mean and the standard deviation) value against C_4 species richness values. The slopes of the linear regressions represented a conversion factor between C_4 species richness and C_4 grass coverage (Fig. S3).

In addition to the TRY database that has a global representation, we also used a gridded C_4 grass coverage data compiled for the contiguous United States (denoted as the DG dataset)²³. The DG dataset provides C_4 grass coverage (% of grassland) aggregated at a 100 km resolution grid, which was sampled from roughly 40,000 plots over the past 40 years. Please note that the DG dataset only surveyed C_4 grass species. We used the DG dataset and the TRY database to establish the relationship between AC_4/AC_3 and C_4 grass coverage (Fig. 1b).

Processing cropland and land use data

We used a gridded C_4 crop distribution from the LUHv2 dataset (version: LUHv2-GCB2019)³⁶. It was estimated based on the FAO census and national reporting of >170 major crop types (including the main C_4 types), supplemented by the total cropland area collated by HYDE3.2⁴⁰ which came from FAO census and remote sensing products. The C_4 crop fraction of each grid cell was only acquired based on observations circa 2000³⁷ and the fraction was kept constant over the study period. The most recent versions of the LUHv2 dataset have been used in CMIP6 for the IPCC AR6 report and the global carbon budget. Since the LUHv2 dataset did not contain an estimate of uncertainty, we relied on an independent study that compared four different land use products (including LUHv2) and reported the uncertainty of cropland estimation between products was about 10%⁷⁵. We thus used 10% to represent the uncertainty range of the C_4 cropland area.

In addition to LUHv2, we used another open dataset reporting the area of 17 main crop types from 1961 to 2014³⁹. Unlike the LUHv2 dataset which almost exclusively relied on observations circa 2000 to quantify the C_4 crop fraction, this other dataset used annual FAO census records for crop area fraction estimation, including those of the three primary C_4 crops: maize, millet, and sorghum. We used this

dataset to examine the changes in global C_4 croplands and compare to the values obtained from the LUHv2 dataset.

The optimality model for C_4 and C_3 photosynthesis

We used optimal C_3 and C_4 photosynthesis models to simulate optimal C_3 and C_4 photosynthesis⁴⁴. The soil-plant-air water continuum was incorporated in C_3 photosynthesis models⁷⁶ and C_4 photosynthesis models⁷⁷ to examine interactions of CO₂, water availability, light and temperature. The model considered optimal stomatal resistance and leaf/fine-root allocation to maximize the carbon gain regarding water loss, and successfully predicted the ancient distribution of C_4 species in Oligocene and Miocene⁴⁴.

In the current study, we improved the modeling processes through the following aspects. (1) We used different parameters for C_3 and C_4 species specifically to better represent the diversity of C_3 and C_4 species variability (Table S2 in the Supplementary Note). (2) We considered the effects of nitrogen availability and optimal nitrogen allocation between C_3 and C_4 species. Specifically, we adjusted the maximum carboxylation rate (V_{cmax}) and maximum electron transport rate (J_{max}) values using optimal J_{max}/V_{cmax} ratio (i.e., 2.1 for C_3 and 5.0 for C_4 , which were supported by both measurements and theoretical modeling)^{78,79} and available leaf nitrogen content for C_3 and C_4 respectively. (3) Since a large majority of C_4 species are herbaceous, when we modeled closed canopy biomes (e.g., those pixels dominated by tree and shrubs), we used estimates of understory photosynthetic active radiation (PAR) to model the relative advantage of the herbaceous species. A full model description and the parameterization is in the Supplementary Note.

Using the models, we were able to calculate the optimal assimilation rates for C_3 and C_4 (i.e., AC_3 and AC_4) over the globe at the 0.5-degree resolution (i.e., dependent on the spatial resolution of climate input), where the relative advantage of C_4 to C_3 is defined as AC_4/AC_3 . The simulation was conducted at an annual time step and there was no need for model initialization. When establishing the relationship between AC_4/AC_3 and C_4 grass coverages (Fig. 1b) from the DG and TRY datasets, we aggregated the simulations from 0.5-degree to 1-degree (approximately 100 km at the equator) and 10-degree. As the relationships derived from both 10-degree (i.e., TRY) and 1-degree (i.e., DG) data were similar, we assumed that the relationship is scale-independent. Consequently, we applied it to 0.5-degree estimates of AC_4/AC_3 to infer global C_4 grass coverage. We also assumed that the relationship between AC_4/AC_3 and C_4 grass coverage was time-invariant.

Running the optimality model

We used annual growing season average soil water potential, vapor pressure deficit (VPD), 2 m daytime air temperature (T_{air}), photosynthetic active radiation (PAR) and leaf nitrogen content as inputs for the optimality photosynthesis model. The growing season was defined using the MODIS phenology product (MCD12Q2)⁸⁰.

T_{air} was acquired from the CRU-JRA2020 dataset. VPD was estimated using the specific humidity and air temperature from CRU-JRA2020. Soil water potential was estimated from soil texture properties from soil grids and global soil water content datasets, using the Clapp & Hornberger equation⁸¹. The global soil water content datasets came from GLEAM v3⁸². To avoid extreme low soil water potential that does not allow plant growth in the optimality model, we set the minimal soil water potential to -3 MPa. The leaf nitrogen content was acquired from a machine learning upscaled leaf traits product⁷⁴. PAR was also acquired from the CRU-JRA2020 dataset, which is a reanalysis from CRU⁸³ and JRA⁸⁴. We directly used PAR for 'open' ecosystems (i.e., grasslands, savannas), however, for dense forests and shrublands we used understory PAR (i.e., as C_4 grasses often exist in understories), which was derived from PAR and multi-year average

MODIS LAI⁸⁵ (i.e., assume they are overstory LAI) following a radiation gradient mandated by the Beer's Law.

We ran the photosynthetic optimality models multiple times in the process. We first modeled the growing season AC_4/AC_3 in the study period using the climatology of the variables mentioned from 2001 to 2019. To model the growing season AC_4/AC_3 from 2001 to 2019, for each year we used the 20-year climatology (i.e., 20 years before the target year) of the driving variables. In addition, we also conducted simulation for four scenarios, in which we replaced the climate input for 2001 simulation with the CO_2 , T_{air} , VPD and soil moisture from 2019, respectively. Then we used AC_4/AC_3 to estimate the C_4 grass distribution for each year or for each scenario. By calculating the difference between the C_4 grass distribution of four scenarios and the C_4 grass distribution in 2001, we quantified the contribution of CO_2 , T_{air} , VPD and soil moisture to the changes in C_4 grass distribution.

Remote sensing estimates of global grassland fraction

Multiple remote sensing products provide information on grassland distributions. Some directly provide continuous fraction values (i.e., GLC⁸⁶ at 100 meter and Dynamic World⁸⁷ at 10 meter) and some provide categorical information on grassland and savannas (i.e., MODIS⁸⁸ at 500 meter and ESA-CCI⁸⁹ and 300 meter). For the former, we can directly calculate the grassland fraction value at 0.5-degree resolution; for the latter, we assign the grassland/savanna type pixel to 100% and others to 0% grassland, and then obtain the mean value for each 0.5 grid cell. We found that those four estimates of grassland fraction vary considerably (Fig. S4). Based on a visual comparison of the four estimates (i.e., GLC, Dynamic World, MODIS and ESA-CCI) against vegetation map estimates⁹⁰, we found Dynamic World and ESA-CCI substantially underestimate grassland fraction. We therefore used only MODIS and GLC estimates of grassland fractions in our study.

The MODIS grassland fraction product is available from 2001 to 2019. The GLC product is only available from 2015 to 2019. To extend the GLC product back to 2001, we employed a random forest approach to estimate GLC estimates based on surface reflectance, climate, and soil type and extrapolate it to 2001 (i.e., the training accuracy is 99% and the validation accuracy is 95%). We used the average of MODIS and GLC estimates to represent the grassland fraction. To quantify the uncertainty of the approach, for each pixel we bootstrapped 1000 times between the MODIS estimate and the GLC estimate, and use the one standard deviation of these 1000 values to represent the uncertainty in grassland fraction.

Dynamic Global Vegetation Models (DGVMs)

We used 11 DGVMs participating in the global carbon project⁹¹(Table S1) in our study. Though all of the 11 DGVMs provided simulations for C_4 natural grasses, only 7 of them have simulations for C_4 crops (Table S1). We established an emergent constraint between C_4 area and C_4 photosynthesis contribution using the estimates from the model ensemble. We used the S3 scenario (i.e., considering elevated CO_2 , climate change and land use change) of model simulations in our analysis.

Emergent constraint approach

The emergent constraint technique is widely used in climate and modeling communities to infer unobserved quantities of interest in land surface processes^{92,93}. The underlying assumption is that although there is a large spread in the model estimates of an observed variable X and an unobserved variable Y across models, the relationship linking the two is tightly constrained across models. Based on the strong and robust relationship across models between X and Y, observations of X can be used to generate a constraint on unobserved Y. This approach has been termed 'emergent' because the functional relationship cannot be diagnosed from a single model, but rather emerges from the spread of the model estimates. The emergent constraint identified in

this study links the contribution of C_4 grasses/crops to total GPP to the percentages of area covered by C_4 grasses/crops as estimated from the ensemble of DGVM simulations.

Reporting summary

Further information on research design is available in the Nature Portfolio Reporting Summary linked to this article.

Data availability

The global C_4 vegetation distribution map is available at <https://zenodo.org/records/10516423>. The CS C_4 map was acquired from https://daac.ornl.gov/cgi-bin/dsvviewer.pl?ds_id=932. The CRU TS4.02 climate data is available at <https://crudata.uea.ac.uk/cru/data/hrg/>, the soil moisture data can be downloaded from <https://www.gleam.eu/#datasets>. The global dataset of leaf photosynthetic pathway was acquired from the TRY database <https://www.try-db.org/TryWeb/Home.php>, by selecting those records with the field "photosynthesis pathway (traitID: 22)". The DG dataset was obtained from the supporting information of <https://onlinelibrary.wiley.com/doi/10.1111/jbi.13061>. The subset of the observations from the nutrient network (NutNet) are accessible at <https://portal.edirepository.org/nis/mapbrowse?packageid=edi.1037.2>.

Code availability

The code for analysis is available at <https://github.com/lxzswr/C4distribution/> and <https://zenodo.org/records/10516423>. The code of optimality photosynthesis model is available at <https://github.com/zhouhaoran06/C3C4OptPhotosynthesis->. All maps in the study were generated with the assistance of the 'M_Map' package in Matlab, the instruction of which is available at <https://www.eoas.ubc.ca/rich/map.html>.

References

- Lambers, H., Chapin, F. S. & Pons, T. L. *Plant Physiological Ecology*. (Springer New York, 2008). <https://doi.org/10.1007/978-0-387-78341-3>.
- Still, C. J., Berry, J. A., Collatz, G. J. & DeFries, R. S. Global distribution of C_3 and C_4 vegetation: Carbon cycle implications. *Glob. Biogeochem. Cycles* **17**, 6-1-6-14 (2003).
- Lloyd, J. & Farquhar, G. D. ^{13}C discrimination during CO_2 assimilation by the terrestrial biosphere. *Oecologia* **99**, 201-215 (1994).
- Ehleringer, J. R., Cerling, T. E. & Helliker, B. R. C_4 photosynthesis, atmospheric CO_2 , and climate. *Oecologia* **112**, 285-299 (1997).
- Randerson, J. T. et al. Fire emissions from C_3 and C_4 vegetation and their influence on interannual variability of atmospheric CO_2 and $\delta^{13}C_{CO_2}$: C_4 FIRE EMISSIONS. *Glob. Biogeochem. Cycles* **19**, n/a-n/a (2005).
- Sage, R. F. The evolution of C_4 photosynthesis. *N. Phytol.* **161**, 341-370 (2004).
- Christin, P. & Osborne, C. P. Tansley review: the evolutionary ecology of C_4 plants. *N. Phytol.* **204**, 765-781 (2014).
- Collatz, G. J., Berry, J. A. & Clark, J. S. Effects of climate and atmospheric CO_2 partial pressure on the global distribution of C_4 grasses: present, past, and future. *Oecologia* **114**, 441-454 (1998).
- Baldocchi, D. The grass response. *Nature* **476**, 160-161 (2011).
- Still, C. J., Cotton, J. M. & Griffith, D. M. Assessing earth system model predictions of C_4 grass cover in North America: from the glacial era to the end of this century. *Glob. Ecol. Biogeogr.* **28**, 145-157 (2019).
- Pagani, M., Zachos, J. C., Freeman, K. H., Tipple, B. & Bohaty, S. Marked decline in atmospheric carbon dioxide concentrations during the paleogene. *Science* **309**, 600-603 (2005).
- Edwards, E. J. et al. The origins of C_4 grasslands: integrating evolutionary and ecosystem science. *Science* **328**, 587-591 (2010).

13. Bremond, L., Boom, A. & Favier, C. Neotropical C₃/C₄ grass distributions - present, past and future. *Glob. Change Biol.* **18**, 2324–2334 (2012).
14. Morgan, J. A. et al. C₄ grasses prosper as carbon dioxide eliminates desiccation in warmed semi-arid grassland. *Nature* **476**, 202–205 (2011).
15. Teeri, J. A. & Stowe, L. G. Climatic patterns and the distribution of C₄ grasses in North America. *Oecologia* **23**, 1–12 (1976).
16. Ehleringer, J. & Björkman, O. Quantum yields for CO₂ Uptake in C₃ and C₄ plants: dependence on temperature, CO₂, and O₂ concentration. *Plant Physiol.* **59**, 86–90 (1977).
17. Ehleringer, J. R. Implications of quantum yield differences on the distributions of C₃ and C₄ grasses. *Oecologia* **31**, 255–267 (1978).
18. Young, H. J. & Young, T. P. Local distribution of C₃ and C₄ grasses in sites of overlap on Mount Kenya. *Oecologia* **58**, 373–377 (1983).
19. Edwards, E. J. & Still, C. J. Climate, phylogeny and the ecological distribution of C₄ grasses. *Ecol. Lett.* **11**, 266–276 (2008).
20. Sage, R. F. A portrait of the C₄ photosynthetic family on the 50th anniversary of its discovery: species number, evolutionary lineages, and Hall of Fame. *J. Exp. Bot.* **67**, 4039–4056 (2016).
21. Stevens, N., Lehmann, C. E. R., Murphy, B. P. & Durigan, G. Savanna woody encroachment is widespread across three continents. *Glob. Change Biol.* **23**, 235–244 (2017).
22. Epstein, H. E. et al. The relative abundance of three plant functional types in temperate grasslands and shrublands of North and South America: effects of projected climate change. *J. Biogeogr.* **29**, 875–888 (2002).
23. Griffith, D. M., Cotton, J. M., Powell, R. L., Sheldon, N. D. & Still, C. J. Multi-century stasis in C₃ and C₄ grass distributions across the contiguous United States since the industrial revolution. *J. Biogeogr.* **44**, 2564–2574 (2017).
24. Wei, Y. et al. The north american carbon program multi-scale synthesis and terrestrial model intercomparison project - Part 2: environmental driver data. *Geosci. Model Dev.* **7**, 2875–2893 (2014).
25. Griffith, D. M. et al. Biogeographically distinct controls on C₃ and C₄ grass distributions: merging community and physiological ecology: climate disequilibrium in C₄ grass distributions. *Glob. Ecol. Biogeogr.* **24**, 304–313 (2015).
26. Wang, R. & Ma, L. Climate-driven C₄ plant distributions in China: divergence in C₄ taxa. *Sci. Rep.* **6**, 27977 (2016).
27. Winslow, J. C., Hunt, E. R. & Piper, S. C. The influence of seasonal water availability on global C₃ versus C₄ grassland biomass and its implications for climate change research. *Ecol. Model.* **163**, 153–173 (2003).
28. Xie, Q. et al. Satellite-observed shifts in C₃/C₄ abundance in Australian grasslands are associated with rainfall patterns. *Remote Sens. Environ.* **273**, 112983 (2022).
29. Paruelo, J. M. & Lauenroth, W. K. Relative abundance of plant functional types in grasslands and shrublands of North America. *Ecol. Appl.* **6**, 1212–1224 (1996).
30. Fox, D. L. et al. Climatic controls on C₄ grassland distributions during the neogene: a model-data comparison. *Front. Ecol. Evol.* **6**, 147 (2018).
31. Krinner, G. et al. A dynamic global vegetation model for studies of the coupled atmosphere-biosphere system. *Glob. Biogeochem. Cycles* **19**, 1–33 (2005).
32. Bonan, G. B., Levis, S., Kergoat, L. & Oleson, K. W. Landscapes as patches of plant functional types: an integrating concept for climate and ecosystem models: plant functional types and climate modELS. *Glob. Biogeochem. Cycles* **16**, 5–15–23 (2002).
33. Smith, B. et al. Implications of incorporating N cycling and N limitations on primary production in an individual-based dynamic vegetation model. *Biogeosciences* **11**, 2027–2054 (2014).
34. Longo, M. et al. The biophysics, ecology, and biogeochemistry of functionally diverse, vertically and horizontally heterogeneous ecosystems: the ecosystem demography model, version 2.2 – Part 1: model description. *Geosci. Model Dev.* **12**, 4309–4346 (2019).
35. Cotton, J. M., Cerling, T. E., Hoppe, K. A., Mosier, T. M. & Still, C. J. Climate, CO₂, and the history of North American grasses since the Last Glacial Maximum. *Sci. Adv.* **2**, e1501346 (2016).
36. Chini, L. et al. Land-use harmonization datasets for annual global carbon budgets. *Earth Syst. Sci. Data* **13**, 4175–4189 (2021).
37. Monfreda, C., Ramankutty, N. & Foley, J. A. Farming the planet: 2. Geographic distribution of crop areas, yields, physiological types, and net primary production in the year 2000: global crop areas and yields in 2000. *Glob. Biogeochem. Cycles* **22**, n/a–n/a (2008).
38. Yu, Q. et al. A cultivated planet in 2010 – Part 2: the global gridded agricultural-production maps. *Earth Syst. Sci. Data* **12**, 3545–3572 (2020).
39. Jackson, N. D., Konar, M., Debaere, P. & Estes, L. Probabilistic global maps of crop-specific areas from 1961 to 2014. *Environ. Res. Lett.* **14**, 094023 (2019).
40. Goldewijk, K. K., Beusen, A., Doelman, J. & Stehfest, E. Anthropogenic land use estimates for the Holocene - HYDE 3.2. *Earth Syst. Sci. Data* **9**, 927–953 (2017).
41. Hurtt, G. C. et al. Harmonization of global land use change and management for the period 850–2100 (LUH2) for CMIP6. *Geosci. Model Dev.* **13**, 5425–5464 (2020).
42. Powell, R. & Still, C. Biogeography of C₃ and C₄ vegetation in South America. *An. XIV Simpósio Brasileiro Sensoriamento Remoto* **14**, 2935–2942 (2009).
43. Powell, R. L., Yoo, E.-H. & Still, C. J. Vegetation and soil carbon-13 isoscapes for South America: integrating remote sensing and ecosystem isotope measurements. *Ecosphere* **3**, art109 (2012).
44. Zhou, H., Helliker, B. R., Huber, M., Dicks, A. & Akçay, E. C₄ photosynthesis and climate through the lens of optimality. *Proc. Natl. Acad. Sci. USA* **115**, 12057–12062 (2018).
45. Katge, J. et al. TRY plant trait database – enhanced coverage and open access. *Glob. Change Biol.* **26**, 119–188 (2020).
46. Carroll, O. et al. Nutrient identity modifies the destabilising effects of eutrophication in grasslands. *Ecol. Lett.* **25**, 754–765 (2022).
47. Sage, R. F. & Sultmanis, S. Why are there no C₄ forests? *J. Plant Physiol.* **203**, 55–68 (2016).
48. Sage, R. F. & Kubien, D. S. The temperature response of C₃ and C₄ photosynthesis. *Plant, Cell Environ.* **30**, 1086–1106 (2007).
49. Raubenheimer, S. L. & Ripley, B. S. CO₂-stimulation of savanna tree seedling growth depends on interactions with local drivers. *J. Ecol.* **110**, 1090–1101 (2022).
50. Kgope, B. S., Bond, W. J. & Midgley, G. F. Growth responses of African savanna trees implicate atmospheric [CO₂] as a driver of past and current changes in savanna tree cover: response of african savanna trees to CO₂. *Austral Ecol.* **35**, 451–463 (2009).
51. Bond, W. J., Midgley, G. F. & Woodward, F. I. The importance of low atmospheric CO₂ and fire in promoting the spread of grasslands and savannas: FIRE, LOW CO₂ and trees. *Glob. Change Biol.* **9**, 973–982 (2003).
52. Buitenwerf, R., Bond, W. J., Stevens, N. & Trollope, W. S. W. Increased tree densities in South African savannas: >50 years of data suggests CO₂ as a driver. *Glob. Change Biol.* **18**, 675–684 (2012).
53. Polissar, P. J., Rose, C., Uno, K. T., Phelps, S. R. & deMenocal, P. Synchronous rise of African C₄ ecosystems 10 million years ago in the absence of aridification. *Nat. Geosci.* **12**, 657–660 (2019).
54. O'Mara, N. A. et al. Pleistocene drivers of Northwest African hydroclimate and vegetation. *Nat. Commun.* **13**, 3552 (2022).
55. Shen, X. et al. Increased seasonality and aridity drove the C₄ plant expansion in Central Asia since the Miocene–Pliocene boundary. *Earth Planet. Sci. Lett.* **502**, 74–83 (2018).

56. Murphy, B. P. & Bowman, D. M. J. S. Seasonal water availability predicts the relative abundance of C₃ and C₄ grasses in Australia. *Glob. Ecol. Biogeogr.* **16**, 160–169 (2007).
57. Wang, H. et al. Asian monsoon rainfall variation during the Pliocene forced by global temperature change. *Nat. Commun.* **10**, 5272 (2019).
58. Knapp, A. K. et al. Resolving the Dust Bowl paradox of grassland responses to extreme drought. *Proc. Natl. Acad. Sci. USA* **117**, 22249–22255 (2020).
59. Toderich, K. et al. C₃/C₄ plants in the vegetation of Central Asia, geographical distribution and environmental adaptation in relation to climate. in *Climate Change and Terrestrial Carbon Sequestration in Central Asia* (eds. Lal, R., Suleimenov, M., Stewart, B., Hansen, D. & Doraiswamy, P.) 33–63 (Taylor & Francis, 2007). <https://doi.org/10.1201/9780203932698.ch3>.
60. Noss, R. F. *Forgotten Grasslands of the South: Natural History and Conservation*. (Island Press, 2013).
61. Pau, S., Griffith, D. M., Zampieri, N. & Costanza, J. Longleaf pine savannas reveal biases in current understanding of savanna biogeography. *Glob. Ecol. Biogeogr.* **32**, 2047–2052 (2023).
62. Munroe, S. E. M. et al. A vegetation carbon isoscape for Australia built by combining continental-scale field surveys with remote sensing. *Landsc. Ecol.* **37**, 1987–2006 (2022).
63. Edwards, E. J. & Smith, S. A. Phylogenetic analyses reveal the shady history of C₄ grasses. *Proc. Natl. Acad. Sci. USA* **107**, 2532–2537 (2010).
64. Pau, S., Edwards, E. J. & Still, C. J. Improving our understanding of environmental controls on the distribution of C₃ and C₄ grasses. *Glob. Change Biol.* **19**, 184–196 (2013).
65. Lehmann, C. E. R., Archibald, S. A., Hoffmann, W. A. & Bond, W. J. Deciphering the distribution of the savanna biome. *N. Phytol.* **191**, 197–209 (2011).
66. Venter, Z. S., Cramer, M. D. & Hawkins, H. J. Drivers of woody plant encroachment over Africa. *Nat. Commun.* **9**, 1–7 (2018).
67. Scheiter, S. et al. Fire and fire-adapted vegetation promoted C₄ expansion in the late Miocene. *N. Phytol.* **195**, 653–666 (2012).
68. Karp, A. T., Behrensmeyer, A. K. & Freeman, K. H. Grassland fire ecology has roots in the late Miocene. *Proc. Natl. Acad. Sci. USA* **115**, 12130–12135 (2018).
69. Bowman, D. M. J. S. et al. Vegetation fires in the Anthropocene. *Nat. Rev. Earth Environ.* **1**, 500–515 (2020).
70. Long, S. P. Environmental Responses. in *C₄ Plant Biology* 215–249 (Elsevier, 1999). <https://doi.org/10.1016/B978-012614440-6/50008-2>.
71. Vitousek, P. M. et al. Human alteration of the global nitrogen cycle: sources and consequences. *Ecol. Appl.* **7**, 737–750 (1997).
72. Ripley, B. S., Abraham, T. I. & Osborne, C. P. Consequences of C₄ photosynthesis for the partitioning of growth: a test using C₃ and C₄ subspecies of *Alloteropsis semialata* under nitrogen-limitation. *J. Exp. Bot.* **59**, 1705–1714 (2007).
73. Hengl, T. et al. SoilGrids250m: Global gridded soil information based on machine learning. *PLoS ONE*. **12**, e0169748 (2017).
74. Moreno-Martínez, Á. et al. A methodology to derive global maps of leaf traits using remote sensing and climate data. *Remote Sens. Environ.* **218**, 69–88 (2018).
75. Bayer, A. D. et al. Uncertainties in the land-use flux resulting from land-use change reconstructions and gross land transitions. *Earth Syst. Dynam.* **8**, 91–111 (2017).
76. Farquhar, G. D., von Caemmerer, S. & Berry, J. A. A biochemical model of photosynthetic CO₂ assimilation in leaves of C₃ species. *Planta* **149**, 78–90 (1980).
77. Caemmerer, S. V. *Biochemical models of leaf photosynthesis*. (CSIRO Publishing, 2000).
78. Zhou, H., Akçay, E. & Helliker, B. Optimal coordination and reorganization of photosynthetic properties in C₄ grasses. *Plant Cell Environ.* **46**, 796–811 (2023).
79. Pignion, C. P. & Long, S. P. Retrospective analysis of biochemical limitations to photosynthesis in 49 species: C₄ crops appear still adapted to pre-industrial atmospheric CO₂. *Plant Cell Environ.* **43**, 2606–2622 (2020).
80. Friedl, M., Gray, J. & Sulla-Menashe, D. *MCD12Q2 MODIS/Terra +Aqua Land Cover Dynamics Yearly L3 Global 500m SIN Grid V006*. (2019) <https://doi.org/10.5067/MODIS/MCD12Q2.006>.
81. Clapp, R. B. & Hornberger, G. M. Empirical equations for some soil hydraulic properties. *Water Resour. Res.* **14**, 601–604 (1978).
82. Martens, B. et al. GLEAM v3: satellite-based land evaporation and root-zone soil moisture. *Geosci. Model Dev.* **10**, 1903–1925 (2017).
83. Harris, I., Osborn, T. J., Jones, P. & Lister, D. Version 4 of the CRU TS monthly high-resolution gridded multivariate climate dataset. *Sci. Data* **7**, 109 (2020).
84. Kobayashi, S. et al. The JRA-55 reanalysis: general specifications and basic characteristics. *J. Meteorol. Soc. Jpn.* **93**, 5–48 (2015).
85. Myneni, R. B. et al. Global products of vegetation leaf area and fraction absorbed PAR from year one of MODIS data. *Remote Sens. Environ.* **83**, 214–231 (2002).
86. Tsendbazar, N. et al. Towards operational validation of annual global land cover maps. *Remote Sens. Environ.* **266**, 112686 (2021).
87. Brown, C. F. et al. Dynamic world, near real-time global 10 m land use land cover mapping. *Sci. Data* **9**, 251 (2022).
88. Sulla-Menashe, D. & Friedl, M. A. *User Guide to Collection 6 MODIS Land Cover (MCD12Q1 and MCD12C1) Product*.
89. ESA. *Land Cover CCI Product User Guide Version 2. Tech. Rep.* https://www.esa-landcover-cci.org/?q=webfm_send/84 (2014).
90. Lehmann, C. E. R. et al. *Functional diversification enabled grassy biomes to fill global climate space.* <http://biorxiv.org/lookup/doi/10.1101/583625> (2019). <https://doi.org/10.1101/583625>.
91. Friedlingstein, P. et al. Global Carbon Budget 2019. *Earth Syst. Sci. Data* **11**, 1783–1838 (2019).
92. Cox, P. M. Emergent constraints on climate-carbon cycle feedbacks. *Curr. Clim. Change Rep.* **5**, 275–281 (2019).
93. Eyring, V. et al. Taking climate model evaluation to the next level. *Nat. Clim. Change* **9**, 102–110 (2019).

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Author contributions

XL conceived the idea. XL and HZ designed the study and carried out the analysis. XL, HZ, DG, TK, NS, SS and CS participated in discussions at various stages. SS provided the TRENDY v9 DGVMs simulations. JT, TS, RZ and NS aided in data collection and analysis. All authors contributed to writing.

Competing interests

The authors declare no competing interests.

Additional information

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