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SOURCE PROVENANCE OF OBSIDIAN ARTIFACTS FROM A NUMBER OF SITES AND ISOLATED FINDS IN SOUTHERN NEW MEXICO

by

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INTRODUCTION

 The analysis here of 67 artifacts and geological samples from a number of temporal contexts in southeastern New Mexico exhibits a very diverse obsidian source provenance including sources from northern Chihuahua, western New Mexico, and the Rio Grande Quaternary alluvium, as well as one sample that had to be originally procured from Valles Caldera in northern New Mexico. The mix of sources in the assemblage is quite similar to that reported for a Late Pithouse period site in the area (Shackley 2004a).

ANALYSIS AND INSTRUMENTATION

All samples were analyzed whole with little or no formal preparation. The results presented here are quantitative in that they are derived from "filtered" intensity values ratioed to the appropriate x-ray continuum regions through a least squares fitting formula rather than plotting the proportions of the net intensities in a ternary system (McCarthy and Schamber 1981; Schamber 1977). Or more essentially, these data through the analysis of international rock standards, allow for inter-instrument comparison with a predictable degree of certainty (Hampel 1984).

The EDXRF trace element analyses were performed in the Archaeological XRF Laboratory, Department of Earth and Planetary Sciences, University of California, Berkeley, using a Spectrace/ThermoNoranTM QuanX energy dispersive x-ray fluorescence spectrometer. All samples were analyzed whole with little or no formal preparation. The results presented here are quantitative in that they are derived from "filtered" intensity values ratioed to the appropriate x-ray continuum regions through a least squares fitting formula rather than plotting the proportions of the net intensities in a ternary system (McCarthy and Schamber 1981; Schamber 1977). Or more essentially, these data through the analysis of international rock standards, allow for inter-instrument comparison with a predictable degree of certainty (Hampel 1984). The spectrometer is equipped with an air cooled Cu x-ray target with a 125 micron Be window, an xray generator that operates from 4-50 kV/0.02-2.0 mA at 0.02 increments, using an IBM PC

based microprocessor and WinTraceTM software. The x-ray tube is operated at 30 kV, 0.14 mA. using a 0.05 mm (medium) Pd primary beam filter in an air path at 200 seconds livetime to generate x-ray intensity K α -line data for elements titanium (Ti), manganese (Mn), iron (as Fe^T), rubidium zinc (Zn), (Rb), strontium (Sr), yttrium (Y), zirconium (Zr), niobium (Nb), and thorium (Th). Weight percent iron (Fe₂O₃^T) can be derived by multiplying ppm estimates by 1.4297(10-4). Trace element intensities were converted to concentration estimates by employing a leastsquares calibration line established for each element from the analysis of international rock standards certified by the National Institute of Standards and Technology (NIST), the US. Geological Survey (USGS), Canadian Centre for Mineral and Energy Technology, and the Centre de Recherches Pétrographiques et Géochimiques in France (Govindaraju 1994). Further details concerning the petrological choice of these elements in Southwest obsidian is available in Shackley (1992, 1995, 2004; also Mahood and Stimac 1991; and Hughes and Smith 1993). Specific standards used for the best fit regression calibration for elements Ti through Nb include G-2 (basalt), AGV-1 (andesite), GSP-1, SY-2 (syenite), BHVO-1 (hawaiite), STM-1 (syenite), QLO-1 (quartz latite), RGM-1 (obsidian), W-2 (diabase), BIR-1 (basalt), SDC-1 (mica schist), TLM-1 (tonalite), SCO-1 (shale), all US Geological Survey standards, and BR-N (basalt) from the Centre de Recherches Pétrographiques et Géochimiques in France (Govindaraju 1994). In addition to the reported values here, Ni, Cu, and Ga were measured, but these are rarely useful in discriminating glass sources and are not generally reported.

The data were translated directly into Excel™ for Windows software for manipulation and on into SPSS™ for Windows for statistical analyses. In order to evaluate these quantitative determinations, machine data were compared to measurements of known standards during each run. An analysis of the repeated runs of source standard RGM-1 is included in Table 1. Source nomenclature follows Baugh and Nelson (1987), Glascock et al. (1999), and Shackley (1988, 1995, 1998a, 1998b, 2004b). Further information on the laboratory instrumentation and source nomenclature can be found at: http://www.swxrflab.net/ and Shackley (1998a). Trace element data exhibited in Table 1 and Figures 1 and 2 are reported in parts per million (ppm), a quantitative measure by weight.

SUMMARY AND CONCLUSION

 As with the earlier study, the vast majority of obsidian sources present in the assemblage suggests considerable contact or procurement to the south in northwestern Chihuahua (Sierra Fresnal and Los Jaguëyes), and western New Mexico/eastern Arizona (Mule Creek and Antelope Wells, Cow Canyon; Tables 1 and 2 and Figures 1 through 3 here). The Chihuahuan sources, particularly Sierra Fresnal have been found in alluvium considerably north of the primary domes almost to the international border, so the obsidian used to produce these artifacts could actually be nearly "local" in origin. Similarly, the artifacts produced from Mount Taylor and Cerro Toledo Rhyolite glass, could have been procured in the Rio Grande alluvium just to the east of Florida Mountains toward Las Cruces (see Church 2000). The Antelope Wells obsidian is not distributed in secondary deposits, so had to be originally procured from the area near the source at El Berrendo, Chihuahua or immediately north of the border. More interestingly, the one artifact produced from Valle Grande obsidian must have been originally procured from the Valles Caldera proper since it does not erode outside the rim of the caldera (Shackley 2004b).

 The Florida Mountain vitrophyre (perlitic glass) submitted for analysis exhibits an elemental composition similar to Chihuahuan basin and range obsidian, but displays that quite variable elemental composition typical of crystalline rocks and particularly perlite (Zielinski et al. 1977; Table 3 here). Perlite or vitrophyre is commonly attributed to the Little Florida Mountains by local rockhounds, and this study does suggest that there is no artifact quality obsidian derived from that rhyolite. Again, there is a possibility that the Sierra Fresnal obsidian has been transported into alluvial context not far south of these sites. Further work on the Chihuahuan basin and range sources is certainly necessary.

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Sample	Ti	Mn	Fe	Rb	Sr	Y	Zr	Nb	Source
63L10	1111	345	7762	427	137	52	158	29	vitrophyr
									e
63L4	1207	427	8574	302	63	51	157	39	vitrophyr
									e
63L6	1329	468	9404	311	153	57	160	38	vitrophyr
									e
63L7	1016	281	7317	269	419	56	138	35	vitrophyr
									e
63L8	987	380	7257	272	57	50	145	30	vitrophyr
									e
63L9	1250	433	8377	326	81	54	155	39	vitrophyr
									e
63	1249	428	8692	294	75	47	142	32	vitrophyr
									e
64	1229	465	7933	328	35	46	135	28	vitrophyr
									e

Table 3. Elemental concentrations for three vitrophyric glass samples from Florida Mountains.

Figure 1. Y versus Nb biplot of archaeological data.

Figure 2. Rb versus Y biplot of archaeological data with outlier sources Los Jaguëyes and Mount Taylor removed for clarity.

Figure 3. Distribution of obsidian source provenance in the assemblage.