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1 Downwind Dispersion of CO₂ from a Major Subsea Blowout in Shallow Offshore Waters

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9 Keywords: CO₂ well blowout, offshore well blowout, atmospheric dispersion, dense gas
10 dispersion, CO₂ leakage risk assessment, reduced order models

11 **Abstract**

12 Growing interest in offshore geologic carbon sequestration (GCS) motivates risk assessment of
13 large-scale subsea CO₂ well blowouts or pipeline ruptures. For major leaks of CO₂ from wells or
14 pipelines, significant fluxes of CO₂ may occur from the sea surface depending on water depth. In
15 the context of risk assessment of human health and safety, we have used previously simulated
16 coupled well-reservoir and water column model results as a source term for dense gas dispersion
17 of CO₂ above the sea surface. The models are linked together by one-way coupling, i.e., output
18 of one model is used as input to the next model. These first-of-their-kind coupled flow results are
19 applicable to assessing the hazard of CO₂ to people at and downwind of the sea surface location
20 of emission. Hazard is quantified by plotting the downwind dispersion length (DDL), which we
21 define in the study as the distances from the emission source to the point at which the emitted
22 CO₂ has been diluted to 5% and 1.5% in air by volume. Results suggest that large-scale blowouts
23 in shallow water (10 m) may cause hazardous CO₂ plumes extending on the order of several
24 hundred meters downwind. Details of the modeling show DDL has a maximum for windspeed
25 (at an elevation of 10 m) of approximately 5 m/s, with smaller DDL for both weaker and stronger
26 winds. This is explained by the fact that wind favors transport but also causes dispersion;
27 therefore there is a certain wind speed that maximizes DDL.

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34

35 **Introduction**

36 Offshore geologic carbon sequestration (GCS) largely avoids a variety of concerns related to
37 onshore GCS in the U.S., e.g., potential for impacting groundwater quality in the case of leakage
38 of CO₂ into underground sources of drinking water [1, 2], inducing seismicity that is felt and/or
39 damaging to houses and buildings [3, 4], and securing subsurface storage rights from a potential
40 multitude of property owners [5, 6]. In addition, a great deal of capacity for GCS has been
41 identified in the offshore waters of the U.S., e.g., in the near offshore waters of the Texas Gulf
42 Coast where there are numerous existing industrial CO₂ sources and networks of transportation
43 infrastructure [7, 8].

44 In addition to the decrease in concerns as mentioned above, health, safety, and environmental
45 (HSE) risks of offshore GCS are generally lower than for onshore GCS if for no other reason
46 than there are fewer people offshore. Nevertheless, there are people present periodically and
47 typically clustered in small areas in the offshore environment, e.g., workers on offshore
48 platforms, fishermen and other people on ships and boats, and these populations will be
49 vulnerable to potential impacts arising from leakage of CO₂ from wells (injection wells along
50 with other wells that may leak unexpectedly), and from leakage from pipelines or ships that are
51 transporting CO₂. As such, it is important to assess health and safety risk in the offshore
52 environment arising from large-scale CO₂ leakage incidents, e.g., those arising from sub-sea well
53 blowouts or pipeline ruptures.

54 Carbon dioxide is a physiologically active chemical when inhaled by humans, and causes
55 increasingly dangerous effects to humans as its concentration increases in inhaled air [9]. The
56 hazard considered in this paper is the size of the region/extent of elevated CO₂ concentrations in
57 air due to CO₂ emissions out of the sea surface related to a subsea well blowout. We do not
58 consider the likelihood of occurrence of the blowout, or the likelihood of given wind speeds and
59 directions, that would be a part of a complete risk assessment. Instead we focus only on the
60 hazard of high concentrations of CO₂ in air and related potential impact (people breathing air
61 with elevated CO₂ concentrations). We further do not consider the hydrodynamic hazards
62 associated with large-scale upwellings of gas from the sea surface such as turbulence,
63 fountaining, or radial outflow of seawater, all of which can destabilize ships and boats in the
64 vicinity of large gas emissions [10].

65 With a focus on scoping-type HSE risk assessment that uses modeling-based information on the
66 concentration of CO₂ in air above the sea surface that could arise from a large-scale offshore CO₂
67 blowout, we present here modeling results of the CO₂ sea-surface leakage flux and resulting
68 dense gas dispersion in the air above the sea surface. The approach used for the atmospheric
69 dispersion modeling was described previously [11]. The purpose of this paper is to summarize
70 the results of modeling the flow, absorption, dilution, and dispersion of CO₂ originating from a
71 major CO₂ blowout in shallow offshore waters of the Texas Gulf Coast. This present study,
72 along with a recently published study of the coupled reservoir-well-water column modeling of
73 the CO₂ blowout [12], combine to form a first-of-its-kind reservoir-to-atmosphere study of CO₂

74 transport and dispersion following a large-scale offshore CO₂ incident. The context is HSE risk
75 assessment, and the approach of the atmospheric dispersion part of the study is aimed at
76 generating fast approximate estimates of the extent of hazardous CO₂ plumes that could develop
77 on the sea surface from major leakage incidents in shallow water. More detailed and site-specific
78 studies will be required for quantitative risk assessments of specific scenarios at actual offshore
79 GCS project sites.

80 **Background and Prior work**

81 *Dense Gas Dispersion Modeling*

82 The modeling of atmospheric dispersion of leaking CO₂ from hypothetical CO₂ pipelines, wells,
83 and surface seeps has been the subject of numerous papers in the area of HSE risk assessment
84 related to GCS. Because dense gas dispersion experiments and modeling pre-date research on
85 GCS, we can call the early work on the subject the zeroth generation of research related to CO₂
86 atmospheric dispersion. In fact, this early work was not for CO₂ at all but rather was motivated
87 by the need to understand risks associated with cryogenic gas transport such as liquefied natural ,
88 gas (LNG) [13, 14]. A particularly good review of the field of dense gas dispersion in terms of its
89 context, physics, and modeling is given by Britter [15]. We mention this early work first because
90 the method that we will apply in the current study was built from the observations and synthesis
91 of results of these early experiments as explained previously[11]. But before describing our
92 methods, it is important to mention the considerable advances in dense gas dispersion modeling
93 applicable to GCS risk assessment that have been made in the last 30-plus years.

94 Some of the work on dense-gas dispersion modeling in the GCS context has considered fixed
95 source terms specified in terms of flow rate of single-phase (gaseous) CO₂ , e.g., from ruptured
96 pipelines [16] or from the ground surface [17]. The effects of topography are profound for risk
97 assessment of dense gases for well blowouts or pipeline leaks [18]. For surface seeps/emissions,
98 considerable modeling work has been carried out to understand topographic effects with much of
99 the work motivated by the Lake Nyos event [19, 20, 21] and other natural volcanic sources [22].
100 Other modeling work considered coupling of the source of CO₂ to the flux at the ground surface
101 or out of the pipe rupture [23, 24]. Advances in the level of detail of the physics occurring upon
102 leakage from high-pressure pipelines or tanks have been made by the consideration of
103 multiphase aspects of leaking CO₂ that arise when CO₂ decompresses, e.g., causing formation of
104 solid (dry ice) particles [25, 26, 27]. An industry-led collaborative project with both experimental
105 and modeling components demonstrated that complex phase-change-related processes in CO₂
106 leakage can be evaluated and modeled [28, 29]. The determination and modeling of the nature of
107 the source terms for atmospheric dispersion following pipeline leakage are improving up to the
108 present through combined experimental and model development work that includes phase
109 change [30, 31].

110 ***Offshore Atmospheric Dispersion***

111 In the specific area of dispersion of leaking gas in the offshore environment, i.e., gas being
112 emitted from the sea surface, there is an enormous literature from the worldwide offshore oil and
113 gas industry. One limitation to this knowledge base provided by the long experience of spill and
114 leakage incidents, along with preventive risk assessment work on offshore oil and gas production
115 and transport, is that the gas of interest is mostly natural gas (nominally methane) which is a light
116 gas relative to air rather than a dense gas like CO₂. In short, natural gas emanating from the sea
117 surface tends to rise and disperse, whereas CO₂ will generally be a dense gas and tend to spread
118 out on the sea surface. Nevertheless, in the area of risk assessment of offshore GCS well and
119 pipeline blowouts, there is an extensive knowledge base and established technical expertise that
120 the GCS community can build upon. For example, the physical controls, basic physics, and
121 simple mathematical models of subsea well blowouts and pipeline leaks were developed and
122 tested experimentally several decades ago[32, 33, 34]. Jumping ahead to more recent times with
123 powerful computers and software available, computational fluid dynamic (CFD) models have
124 been demonstrated capable of simulating the dynamics of gas flow and dispersion at the meter-
125 scale and smaller representing every desired detail of an offshore platform or ship [35, 36, 37,
126 38, 39]. As for the differences between light and dense gases with differing solubility in
127 seawater, even the coupled subsea-atmosphere systems during blowout scenarios of natural gas
128 and CO₂ have been simulated using CFD and other approaches [40].

129 ***Fast Modeling for Risk Assessment***

130 Fast modeling is important in the context of uncertainty analysis when carrying out risk
131 assessment for systems with unknown or variable properties typical of GCS projects. The U.S.
132 National Risk Assessment Partnership (NRAP) has been putting the collective understanding and
133 knowledge of GCS into the development of a set of reduced-order models (ROMs) and an
134 associated integrated assessment model (IAM) [41]. The purpose of using ROMs rather than full-
135 physics models is that many model runs can be carried out quickly for many different parameter
136 values and scenarios to capture ranges of uncertainty inherent in scoping-type risk assessments.
137 One of the ROMs developed in the NRAP project, the Multi-Source Leakage ROM (MSLR), is a
138 simple model for atmospheric dispersion of the CO₂ [11]. In this paper, we will apply the MSLR
139 to the offshore blowout scenario reported in a prior paper [12], the results of which are
140 summarized below.

141 The reader will wonder why, with all of the advanced modeling capabilities in existence as
142 described in the prior subsection, we use the very simple MSLR. Aside from the need for fast
143 simulations for uncertainty analysis, in the context of GCS where work to date is mostly on
144 hypothetical or yet-to-be-built GCS projects, there is a need for scoping-type risk assessments in
145 which the goal is a general idea of risk rather than tightly constrained quantitative risk
146 assessment. Simply put, for scoping studies there is no single site to characterize in detail so use
147 of a detailed and highly sophisticated model does not match the level of detail of the input data.
148 This is the case for the Gulf of Mexico Partnership for Offshore Carbon Storage (GoMCarb)

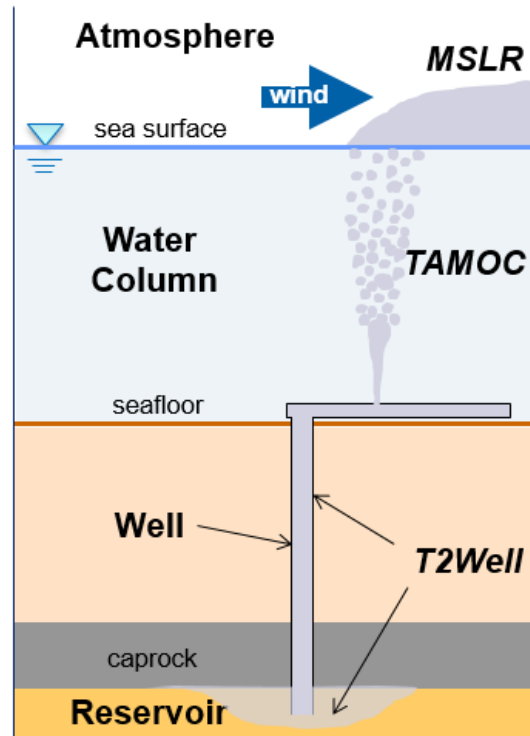
149 where multiple sites are being evaluated generally for safe, long-term, and economically-viable
150 offshore CO₂ storage.

151 As to the applicability of the MSLR, the offshore environment and the relatively long-term (> 10
152 days) and steady-state blowout scenario we model matches the assumptions made in the simple
153 nomograph approach, e.g., no topography, steady-state winds, constant source strength, and
154 uniform pressure and temperature. Finally, as described in our previous paper [11], the MSLR
155 estimates of downwind concentration match closely with an established CFD code [42] for an
156 applicable test problem, so the use of a simple model like the MSLR is sufficient for the system
157 of interest and for our present purposes.

158 **Modeling to Generate the Source Term**

159 ***Coupled Reservoir-Well and Water Column Subsea CO₂ Blowout Model***

160 Full details of previous modeling of a scenario of an offshore CO₂ well blowout into shallow
161 seawater are given in the prior paper [12]. To summarize this prior work briefly, the scenario
162 involved a well blowout at the seafloor as shown in Figure 1. The objective of the work was to
163 understand roughly the potential for such an incident to cause hazardous CO₂ emissions (high
164 flow rates of CO₂) at the sea surface. The system comprises a reservoir under injection of CO₂, a
165 long well (3 km), the water column, and the atmosphere (Figure 1). Each of these flow domains
166 is important in the transport and/or dispersion/absorption of CO₂ that could lead to high CO₂
167 concentrations in air at the sea surface. The three different models used in the four domains are
168 shown in Figure 1, specifically T2Well for the coupled reservoir and well, TAMOC for the water
169 column, and MSLR for the atmosphere, each of which will be described below.



170

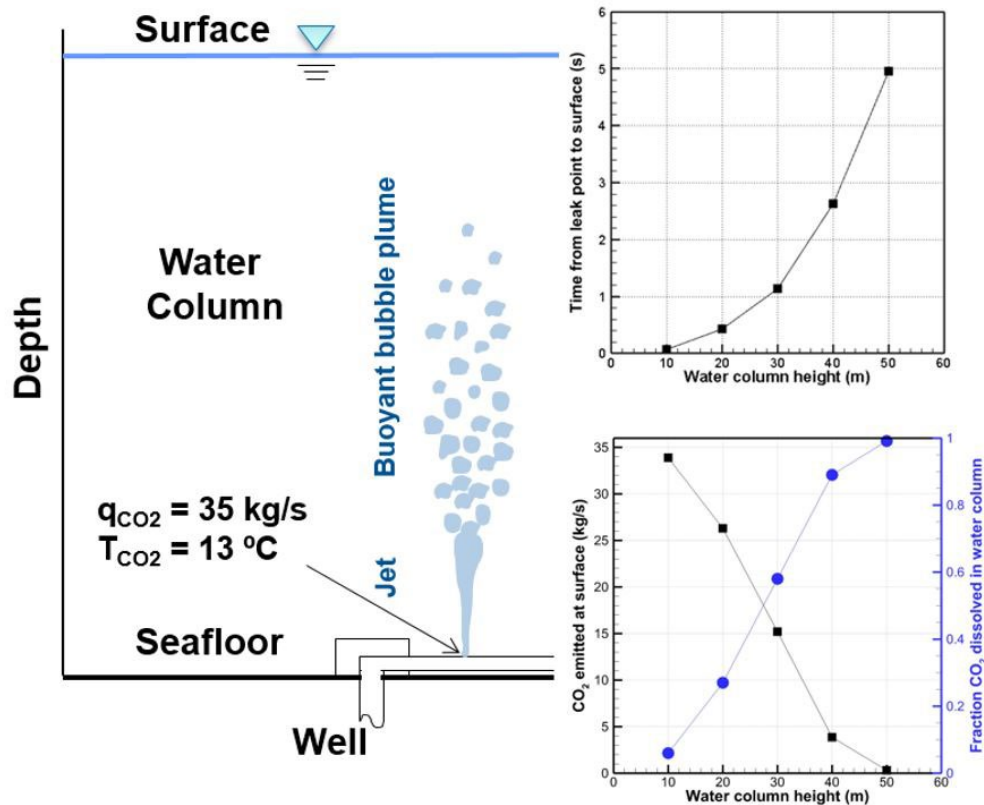
171 Figure 1. Schematic of offshore CO₂ well blowout with four different domains (reservoir, well,
172 water column, and atmosphere) shown along with the three one-way coupled models (T2Well,
173 TAMOC, and MSLR) used to model CO₂ transport and attenuation/dispersion in each domain.
174 While not to scale, the scenario sketched is for a case with small water column height (shallow
175 depth) where there is sufficient CO₂ flux to migrate through the entire water column without
176 significant absorption of CO₂ before being emitted at the sea surface where the CO₂ undergoes
177 dense gas dispersion.

178

179 Referred to here as a well blowout, the incident scenario considered was actually the breaching
180 of the supply line near (10 m away from) the wellhead in the form of a small hole 5 cm (2
181 inches) in diameter. The scenario definition included the assumption that the supply pipeline
182 would be automatically shutdown quickly after pressure-drop detection, but that the blowout
183 preventer on the injection well somehow fails with the result that CO₂ flows up the well
184 unrestricted and leaks into the water column. Because this was an injection well, the CO₂ around
185 the perforations in the well in the injection interval would be highly mobile. The fully coupled
186 reservoir-well system was modeled using T2Well [43] which captured the physics of two-phase
187 non-isothermal CO₂-water flow in the reservoir and well system up to the discharge point at the
188 seafloor. Briefly, T2Well models two-phase flow in a well by the drift-flux model, and is fully
189 coupled to flow in the porous media reservoir through perforations in the well [43].

190 To summarize the results, the CO₂ was emitted into the water column from the hole in the pipe at
191 the seafloor as a jet rapidly entraining seawater and breaking up into small bubbles [12]. All of
192 the water-column bubble plume flow and transport processes were modeling using TAMOC [44,
193 45, 46] with the T2Well leakage flow output as the source term. Briefly, TAMOC estimates a
194 distribution of bubble sizes based on the source term strength (here a jet of CO₂ emanating from
195 the hole in the pipe) and models the rise of the bubble-plume through the water column in an
196 integral model sense including a variety of processes such as seawater entrainment and
197 absorption (dissolution) of CO₂ into seawater. The height of the water column (depth of seafloor
198 and source of back-pressure on CO₂ blowout into water column) was varied in the simulations
199 from 10 m to 50 m representing a range of shallow near-offshore Texas Gulf Coast sites being
200 examined by the GoMCarb project. The TAMOC results showed strong absorption (dissolution)
201 of CO₂ from the small bubbles (mean size approximately 0.5 mm) generated by the vigorous
202 outflow at the hole in the pipe.

203 The main results of the Oldenburg and Pan (2020) study [12] are compactly shown in Figure 2
204 by the sketch on the left-hand side showing a cartoon of one variation, and the two insets on the
205 right-hand side showing quantitative results. Specifically, the upper frame on the right-hand side
206 of Figure 2 shows the travel time of CO₂ from the breach in the pipe to the sea surface as a
207 function of water depth. Travel time for the 10 m-depth case was less than one second, while it
208 took approximately 5 s for CO₂ exiting the pipe at the seafloor to reach the sea surface for the 50
209 m-depth system. In the lower inset on the right-hand side is shown the amount of CO₂ that leaves
210 the sea surface and enters the atmosphere both as a mass flow rate and fraction of the amount
211 emitted from the pipe breach. As shown, the modeling determined that nearly all of the CO₂ in
212 the deepest system (50 m) was absorbed by seawater, while very little CO₂ was absorbed in the
213 shallowest system (10 m), and there were intermediate results for water column heights in
214 between. For reference, one million tonnes per year of CO₂ is 31.7 kg/s; the shallow depth
215 system is predicted to emit CO₂ at more than this rate for this scenario). These results suggests
216 that from the point of view of human health risk assessment related to sea-surface CO₂ emissions
217 from sea floor blowouts, deeper sites are preferred. In the case where shallow-water sites are
218 chosen for CO₂ injection wells and/or CO₂ pipelines exist, and if a large-scale containment
219 failure scenario occurs, CO₂ can be expected to be emitted out of the sea surface and therefore it
220 is important to understand CO₂ plume dispersion in air above the sea surface. Full details of the
221 reservoir-well-water-column modeling study are presented in [12].



222

223 Figure 2. Offshore leakage scenario and summary of CO₂ well blowout simulations of [12]
 224 showing travel time in the water column (upper right-hand figure) and amount of CO₂ emitted
 225 and dissolved during transit through the water column (lower right-hand figure) as a function of
 226 water column height (water depth).

227 **Atmospheric Dispersion Modeling Method**

228 The method for estimating atmospheric dispersion of CO₂ (a dense gas) above the sea surface
 229 used here was described fully in a prior paper [11]. To briefly summarize, the approach is based
 230 on a nomograph that was developed from empirical data on dense gas dispersion [47]. In this
 231 earlier work [47], correlations were determined between measured downstream concentrations of
 232 various dense gases and wind speed and source strengths in terms of release rate and
 233 concentration. The multiple variables of the flow and dispersion field experiments were
 234 condensed into dimensionless quantities and correlated using two-dimensional nomographs. The
 235 nomographs allow one to estimate the distance (downstream of the source) to the locations of
 236 arbitrary fractions of initial concentration [47]. Zhang et al. (2016) [11] extended the nomograph
 237 approach to include multiple leakage source locations, where CO₂ leakage is quantified in the
 238 GCS context by leakage mass flow rate (kg CO₂/s). This enhanced model was converted by into
 239 a computational tool and named the Multi-Source Leakage ROM, or MSLR [11]. The National
 240 Risk Assessment Partnership (NRAP) project sponsored by the U.S. Department of Energy

241 developed a graphical user interface for the MSLR with ability to take user input on leakage
242 locations, emission source strengths (kg/s), windspeed (m/s) at a height of 10 m (approximate
243 thickness of the (atmospheric) surface boundary layer), and critical concentration. The main
244 output of the MSLR is the downwind dispersion length (DDL), i.e., the downwind distance at
245 which the dispersing CO₂ reaches the critical concentration (i.e., a given level of dilution relative
246 to the concentration at the source). The MSLR is one-dimensional and provides DDL for user-
247 specified wind speed. Downwind transport makes the DDL larger than a corresponding
248 transverse dispersion length and therefore considering downwind extent is a conservative
249 approach for hazard analysis. Because the wind direction could be random and variable, a
250 conservative method is to define the DDL as a circular exclusion zone within which CO₂
251 concentration can be expected to be unhealthy to breathe for any wind direction.

252 Although the original work [47] to develop the nomograph underlying the MSLR pre-dates
253 interest in GCS and was focused not on CO₂ risk assessment but rather on accidents that could
254 happen during transport of cryogenic liquid hydrocarbons (e.g., liquefied natural gas, propane,
255 etc.), the ranges of leakage rates and density contrasts used to develop the original nomographs
256 overlap those relevant to CO₂. For example, we note the molecular weight of propane and CO₂
257 are the same at 44 g mole⁻¹ making CO₂ and propane densities in ambient air roughly the same.
258 In addition, incidents involving leakage of liquefied gas or propane may result in similar flow
259 rates and processes because high-pressure pipeline and ship transport and related infrastructure
260 are common to both liquefied gas and CO₂ transport.

261 For the offshore CO₂ blowout case considered here, the resulting plume of interest is a direct
262 emission of CO₂ from a source area on the sea surface as opposed to evaporation from a puddle,
263 high-energy pipe or tank emission, or smoke-stack or other elevated source. The characteristic
264 length scale of the emission area on the sea surface modeled by TAMOC was between 3.2-16 m
265 [12]. This length scale is used by the MSLR to determine if dense gas dispersion is applicable for
266 the given mass flow rate of the source. We further assume the transport above the sea is
267 isothermal at 25 °C and 1 atm (0.101325 MPa) and that CO₂ does not absorb (dissolve) into
268 seawater as it flows over the sea surface (i.e., the sea surface is a closed boundary). By
269 neglecting absorption of CO₂ into the seawater as the CO₂ flows over the sea surface, and by
270 using the single-temperature results of the prior study [12] as the emission source term, the
271 present study does not address the sensitivity of DDL to seawater temperature. Knowing that
272 colder water absorbs more CO₂, we can safely speculate that colder seawater would result in a
273 smaller emission source term and shorter DDL, but we leave quantification of this effect for
274 future study. Although vigorous boil areas with fountain heights up to 70 m have been observed
275 for offshore blowouts of natural gas and these can create significant marine hazards to boats and
276 ships [48], here we assume the sea surface is horizontal and model only the downwind dispersion
277 and CO₂ concentration in air as the hazard of interest. Atmospheric stability was reported to play
278 a minor role in the empirical data upon which the nomograph was constructed [47], and therefore
279 atmospheric stability is not a factor in the MSLR.

280 **Results**

281 The output of T2Well for five different water-column heights was fed to TAMOC, and the
282 output of TAMOC for the five different water column heights in terms of flow rate (kg/s) out of
283 the sea surface were used as inputs to the MSLR. Seven different windspeeds (at a height of 10
284 m) above the sea surface were modeled for three different water-column heights (10, 30, and 50
285 m--the intervening 20 m- and 40 m-depth cases being intermediate results not shown here for
286 brevity) for a single leakage source. The three different water column heights produce different
287 source strengths for many reasons, e.g., the longer time and distance traveled upward through the
288 water column increases absorption of CO₂ into the seawater, the spreading of the bubble plume,
289 and the amount of entrained water. In addition, there is a small effect of the hydrostatic back
290 pressure on the hole in the pipe that slightly affects leakage flow at the sea floor. Inputs to the
291 MSLR are shown in Table 1. The MSLR also needs input on release duration which was chosen
292 to be 1×10^6 s (~11 days) to represent a long period relative to the travel time to reach the DDL
293 (i.e., represents a steady-state condition for the leak relative to wind transport times) even for the
294 lowest wind speed. Specifically, consider that the wind speeds vary from 0.5 m/s (~1.1 mph) to
295 40 m/s (~88 mph) which corresponds to downwind travel times for 500 m radius from the source
296 varying from 1000 s (17 mins) and 12.5 s, respectively. In this sense, winds that persist for 10
297 mins or more are easily able to advect leaking CO₂ to the farthest extent of the DDL, and this
298 time is short relative to the duration of the leakage (~11 days).

299 The results of the MSLR model are shown in Figure 3 as the downwind dispersion length (DDL)
300 at which the concentration of CO₂ in the air has been diluted from pure CO₂ to a mixture of air
301 with CO₂ at 5% (dilution factor of 20) and 1.5% (dilution factor of 67) by volume, respectively.
302 Note that these distances are referred to as downwind dispersion lengths (DDL) rather than a
303 downwind safety lengths because inhalation of air with 5% CO₂ by volume may present a health
304 and safety hazard; 4% CO₂ by volume is the U.S. national standard concentration considered
305 immediately dangerous to life or health [49]. The lower concentration value, 1.5% CO₂ by
306 volume, is a concentration at which some people will experience mild respiratory stimulation
307 [50], i.e., a concentration with non-zero but low impact to human health over short exposure
308 periods.

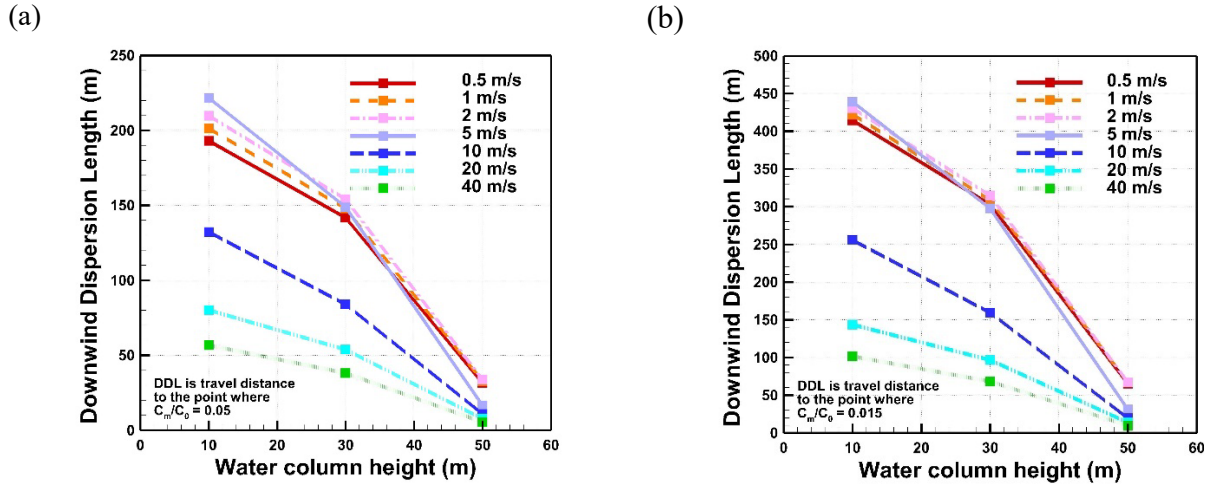
309 Table 1. Properties of the system assumed in the MSLR.

Pressure (atm)	Air, seawater, and source temperature (°C)	Wind speed at 10 m elevation (m/s)	Source strength for 10 m-high water column (kg/s)	Source strength for 30 m-high water column (kg/s)	Source strength for 50 m-high water column (kg/s)	Critical concentrations (fraction by volume)
1.0	25.	0.5	33.4	15.1	0.31	0.05, 0.015
1.0	25.	1.	33.4	15.1	0.31	0.05, 0.015
1.0	25.	2.	33.4	15.1	0.31	0.05, 0.015
1.0	25.	5.	33.4	15.1	0.31	0.05, 0.015
1.0	25.	10.	33.4	15.1	0.31	0.05, 0.015
1.0	25.	20.	33.4	15.1	0.31	0.05, 0.015
1.0	25.	40.	33.4	15.1	0.31	0.05, 0.015

310

311 Note first in Figure 3 that smaller critical concentrations imply larger DDL (note the different y-
 312 axis scales). In other words, the smaller the critical concentration chosen to ensure health and
 313 safety, the larger will be the exclusion zone. The second conclusion from Figure 3 is that higher
 314 wind speeds cause smaller DDL generally because there is more turbulent mixing and air
 315 entrainment to dilute and spread the leaking CO₂ when windspeed is higher. Recall from Figure
 316 2 that surface emissions vary for the different water column heights (depths), and in the deepest
 317 case (50 m water column) the surface leakage flow rate is very low so that for all wind speeds,
 318 the resulting DDL never exceeds approximately 35 m (115 ft) for the 5% critical concentration,
 319 and 65 m (213 ft) for the 1.5% critical concentration. In contrast, for the 10-m water-column
 320 height, the DDL for a CO₂ critical concentration of 1.5% by volume can reach over 400 m.

321 Considering again the shallowest case (10 m water column), an interesting reversal in DDL trend
 322 is observed. First, recall in the 10 m-depth case, nearly all of the CO₂ leaking from the well/pipe
 323 is emitted at the sea surface (i.e., very little absorption of CO₂ occurs in the water column as
 324 shown in Figure 2). For this case of a large CO₂ emission at the sea surface, the maximum DDL
 325 for both critical concentrations (5% and 1.5%) occurs for windspeeds of approximately 5 m/s
 326 (~11 mph), with slightly smaller DDL for weaker winds and much smaller DDL for much higher
 327 winds (10, 20, 40 m/s).



328

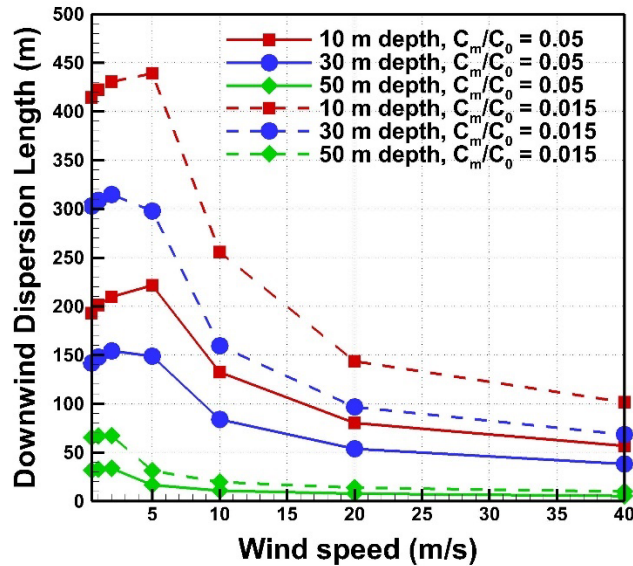
329 Figure 3. Downwind dispersion length (DDL) for the subsea CO₂ blowout scenarios for different
330 water column heights (different surface leakage rates) and wind speeds for two different critical
331 concentrations: (a) $C_m/C_0 = 0.05$; (b) $C_m/C_0 = 0.015$ (note the different y-axis scales).

332

333 The DDL maxima as a function of windspeed are shown explicitly in Figure 4 where DDL is
334 plotted as a function of windspeed for the three water-column-height cases. The reason for a
335 local maximum in DDL is that very slow winds do not transport the gas very effectively leaving
336 advective spreading to occur only by self-generated density-dependent flow effects [15]. On the
337 other hand, faster winds transport the CO₂ downwind effectively but also have more dispersive
338 capacity. So there is a competition between fast-moving air and the dispersive capacity of that air
339 in how far high-CO₂ concentrations can persist downwind. For the scenario modeled here, and
340 depending on the particular case, windspeeds of 2-5 m/s (~4.4-11 mph) are effective at
341 transporting CO₂ without diluting it as much as when the windspeeds are higher and thereby
342 generate a larger DDL, all other things being equal. We emphasize that the MSLR used to
343 generate these results is based on actual field studies of dense gas dispersion, and the nomograph
344 upon which the MSLR is based represents a multitude of lumped physical dispersion processes.

345

346



347

348 Figure 4. DDL as a function of windspeed for three different water column depths and two
349 different critical concentrations. The plot shows that a maximum DDL occurs for windspeeds of
350 2-5 m/s depending on the case.

351 **Conclusions**

352 We have applied a fast model (NRAP MSLR) for atmospheric dispersion of a dense gas (CO₂) to
353 estimate downwind dispersion of CO₂ emitted from the sea surface following a generalized
354 large-scale subsea well blowout. The source term for the fast dispersion model is the mass flow
355 rate out of the sea surface which was simulated previously [12] using TAMOC, which modeled
356 the transport and absorption of CO₂ in a bubble plume rising upward through the water column.
357 The source term for TAMOC was the mass flow rate into the water column from the well which
358 was also simulated previously using the fully coupled T2Well code to couple reservoir and well
359 flow with water depth providing the upper pressure boundary condition on the hole in the pipe
360 [12]. To our knowledge, the work described in the present paper together with that in [12]
361 comprise the first study that has coupled three models of CO₂ flow in these four key domains to
362 estimate CO₂ concentrations downwind of a sea-surface emission from a large-scale blowout.
363 The results show that for very shallow offshore GCS sites (e.g., 10 m water depth) with a large-
364 scale blowout underway, one can expect hazardous CO₂ concentrations to extend a few hundred
365 to several hundred meters from the emissions source, depending of course on CO₂ emission
366 source strength, windspeed, choice of critical concentration, etc. For deeper sites (e.g., 50 m) the
367 large amount of CO₂ absorption in the water column makes a much weaker emission at the sea
368 surface resulting in hazardous CO₂ concentration extending shorter distances from the emissions
369 site, all other things being equal.

370 The results presented here are in the context of providing estimates of DDL for scoping-type risk
371 assessments. For any individual site and for a quantitative risk assessment, a more detailed and
372 advanced CFD-based atmospheric dispersion model should be considered along with site-
373 specific information to characterize the wind speed, direction, and frequency (e.g., a wind rose).
374 Nevertheless, as shown here, even the simple nomograph-based MSLR model captures an
375 interesting effect of the payoff between windspeed in transporting CO₂ downwind while also
376 causing its dispersion. For the well-reservoir system and water column properties considered
377 here, T2Well and TAMOC are appropriate models to capture the details of the physics of CO₂
378 flow and transport. For deeper and/or colder systems for which hydrates (not modeled by T2Well
379 currently) could form at the hole in the pipe due to severe decompression cooling, simulators
380 capable of modeling CO₂ hydrate should be used.

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