

Lawrence Berkeley National Laboratory

LBL Publications

Title

CO2 flow modeling in a coupled wellbore and aquifer system: Details of pressure, temperature, and dry-out

Permalink

<https://escholarship.org/uc/item/9ph2d6xs>

Authors

Zamani, Nematollah

Oldenburg, Curtis M

Solbakken, Jonas

et al.

Publication Date

2024-02-01

DOI

10.1016/j.ijggc.2024.104067

Copyright Information

This work is made available under the terms of a Creative Commons Attribution-NoDerivatives License, available at <https://creativecommons.org/licenses/by-nd/4.0/>

Peer reviewed

1 CO₂ Flow Modeling in a Coupled Wellbore and Aquifer System: Details of Pressure, Temperature, and
2 Dry-Out

3 Nematollah Zamani ^{a*}, Curtis M. Oldenburg^b, Jonas Solbakken ^a, Morten G. Aarra ^a, Patrick Kowollik^c,
4 Hakan Alkan^d, Mohd Amro^d, Taofik Nassan^d, Jose K.P. Estrada^e, Oleksandr Burachok^c

5 ^a NORCE Norwegian Research Centre., Bergen, Norway

6 ^b Energy Geosciences Division, Lawrence Berkeley National Laboratory, Berkeley, CA, United States

7 ^c Wintershall Dea AG, Germany

8 ^d Technical University Bergakademie Freiberg (TU-BAF), Germany

9 ^e Montanuniversität Leoben, Austria

10 * Corresponding author:

11 Nematollah Zamani

12 Nematollah.zamani@norceresearch.no

13 Nygårdsgaten 112, 5008 Bergen, Norway

14 Telephone: (+47) 56 10 71 50

15

16 Note: This is the manuscript form of the published paper which should be cited as:

17 Zamani, N., Oldenburg, C.M., Solbakken, J., Aarra, M.G., Kowollik, P., Alkan, H., Amro, M.M., Nassan,
18 T.H., Estrada, J.K. and Burachok, O., 2024. CO₂ flow modeling in a coupled wellbore and aquifer
19 system: Details of pressure, temperature, and dry-out. *International Journal of Greenhouse Gas*
20 *Control*, 132, p.104067.

21 <https://doi.org/10.1016/j.ijggc.2024.104067>

22

1 Abstract:

2 In order to understand the details of thermal and hydrologic processes attending CO₂ injection into a
3 deep aquifer in the context of Carbon Capture and Storage (CCS), we have carried out coupled well-
4 reservoir simulations of CO₂ injection using the simulator T2WELL-ECO2M. We focus on the injection
5 of cold, dry CO₂ into a warm aquifer and analyze in detail the thermal and hydraulic processes of the
6 coupled well-reservoir system. The results demonstrate the effectiveness of T2WELL in accurately
7 modeling non-isothermal, multiphase flow, phase changes, and identifying dry-out regions in porous
8 media.

9 We simulated heat exchange with the ambient environment, friction effects, convection, exothermic
10 dissolution in brine, and cooling due to both Joule-Thomson effect and water vaporization. The
11 temperature profile within the wellbore deviated from the geothermal profile, impacting CO₂
12 properties at the bottomhole. The simulation revealed the presence of three fronts in the formation:
13 the CO₂ saturation, thermal, and evaporation fronts. The thermal and evaporation fronts were located
14 farther behind the saturation front, indicating limited dry-out and thermal effects near the wellbore.

15 This simulation capability and insights gained in this study form a foundation for ongoing work such as
16 sensitivity analyses, injection optimization, performance assessment, and operational decision
17 support.

18

19 Keywords: Coupled well-reservoir system, T2WELL-ECO2M, Thermal processes, Dry-out region,
20 Injectivity

1 **1. Introduction**

2 Carbon Capture and Storage (CCS) is a practical approach to decreasing the concentration of CO₂ in the
3 atmosphere and mitigating the greenhouse effect, as well as a transitional step towards renewable
4 energy sources. (e.g., Vilarrasa & Rutqvist, 2017). CCS includes capturing CO₂ from industrial sources,
5 transport to a storage site, and injection through the wellbore into specific geologic underground
6 formations. The injection of CO₂ into subsurface formations is a well-established technology in the oil
7 and gas industry, serving as an enhanced oil recovery method for several decades (Hill et al., 2013;
8 Kumar et al., 2022). Despite relatively slow development of CCS worldwide, several pilot and
9 commercial scales CO₂ storage projects have been or are being performed to prove the reliability and
10 applicability of CO₂ injection and to reduce knowledge gaps in providing safe CCS technology (Alkan et
11 al., 2023; Finley et al., 2013; Hovorka et al., 2006; Litynski et al., 2013; Mathieson et al., 2011;
12 Movahedzadeh et al., 2021; Page et al., 2020; Preston et al., 2005; Rangriz Shokri et al., 2021; Sato et
13 al., 2011; Torp & Gale, 2004; Würdemann et al., 2010). Nevertheless, significant CO₂ injection over an
14 extended period for storage purposes may pose multiple challenges, including impairments to
15 injectivity and potential leakage from the storage complex. These challenges may hinder the fast
16 implementation of CCS projects (Kelemen et al., 2019).

17 Proper assessment of a CCS project requires accurately evaluating reservoir behavior in response to
18 physical, chemical, and thermal perturbation induced by large-scale CO₂ injection over a long injection
19 period (André et al., 2010). Thermophysical properties of pure CO₂, such as density, viscosity, enthalpy,
20 and phase behavior, depend on pressure and temperature, and play a crucial role in CO₂ injectivity,
21 storage capacity, and storage safety (Sokama-Neuyam et al., 2020, 2022, Buursink, 2014; Al-
22 Khdheewi et al., 2018). Due to significant variations in temperature and pressure profiles along the
23 wellbore, thermophysical properties of CO₂ at the bottomhole are different from its properties at the
24 wellhead (Henninges et al., 2011; Lindeberg, 2011; Lu & Connell, 2014b; Vilarrasa et al., 2013). Multiple
25 processes can influence the temperature profile along the wellbore. Heat exchange through the casing

1 walls during CO₂ flow in the wellbore through different geological formations is the primary source of
2 temperature change. This heat exchange depends on flow rate, well completion and isolation status,
3 geothermal gradient, and thermal properties of the surrounding formations (Vilarrasa et al., 2013). In
4 addition, frictional loss, thermal conduction, convection, and the Joule-Thomson effect can alter
5 temperature profiles (André et al., 2010; Vilarrasa & Rutqvist, 2017).

6 In addition, mutual interaction between the wellbore and storage formation, strong transient flow
7 during CO₂ injection operation start-up and shut-in, phase transition along the wellbore, dominance of
8 inertia force within the wellbore, and possible leakage of fluid through casing and cement near the
9 wellbore emphasize the importance of having a reliable simulation tool that can adequately describe
10 CO₂ flow in the wellbore. Despite its importance for evaluating CO₂ storage, only a few papers have
11 addressed CO₂ flow in the well (Liu et al., 2016; Lu & Connell, 2014b; Pan et al., 2011; Piao et al., 2018;
12 Strpic et al., 2021). Most CO₂ injection simulation studies have focused on CO₂ flow in porous media,
13 while the flow in the wellbore has either been disregarded (Dalkhaa et al., 2022; Kim et al., 2014) or
14 coupling separate software for wellbore and reservoir simulation (Aakre et al., 2018), or treated as a
15 part of porous media by assigning equivalent parameters (André et al., 2010). This latter approach is
16 called the equivalent porous media approach (EPM). However, this approach overlooks the distinct
17 nature of flow in porous media and the wellbore, which we will elaborate on further below.

18 As CCS gains attention as a viable solution for mitigating global warming, there is a growing need for
19 reliable simulation tools to design and optimize CCS projects, particularly in the near-wellbore region.
20 The success of CCS projects relies on the ability to predict the behavior of CO₂ in the near wellbore
21 region, which is influenced by complex thermodynamic and fluid flow processes. One of the critical
22 challenges in CCS projects is accurately modeling the behavior of CO₂ as it flows through the wellbore
23 and storage formation. To address this challenge, we investigate the strength and limitations of
24 T2WELL-ECO2M, a simulation code from the TOUGH family, for modeling non-isothermal flow in the
25 wellbore and storage formation, focusing on answering critical industrial questions related to the

1 injection of CO₂ into saline aquifers. Specifically, we aim to understand how T2WELL-ECO2M can be
2 used to model the behavior of CO₂ during injection, including thermal effects, phase changes,
3 multiphase flow, and the occurrence of the dry-out region in the near wellbore region. By doing so, we
4 hope to provide insights into the capabilities and limitations of this simulation tool and the potential
5 challenges and opportunities for designing and optimizing CO₂ storage projects. The finding of this
6 study serves as the basis for ongoing research on parameter sensitivity analysis for CO₂ flow in the
7 wellbore and reservoir to address some of the fundamental industrial challenges summarized as
8 follows:

- 9 • Near-wellbore effect: Under what conditions does a dry-out region emerge, and how far does
10 it extend into the reservoir? What is the impact of cold CO₂ injection on the integrity of the
11 caprock/wellbore and surrounding reservoir, and how can this impact be minimized?
- 12 • How do thermodynamic processes, such as heat transfer and phase change, affect CO₂ flow
13 behavior during injection into the reservoir and subsequent propagation within the storage
14 formation?

15

16 **2. Method**

17 **2.1. Numerical simulator**

18 The T2WELL-ECO2M simulator (Pan & Oldenburg, 2014; Strpic et al., 2021) has been used for this study
19 because it provides coupled well-reservoir simulation capabilities for CO₂-brine systems over a wide
20 range of *P-T* conditions. T2WELL is an extension to TOUGH2, a general-purpose numerical simulation
21 code for non-isothermal flows of multi-component, multiphase fluids in multi-dimensional porous and
22 fractured media (Pruess et al., 2012). T2WELL-ECO2M uses an explicit tabulated file to describe the
23 density, viscosity, and enthalpy of pure CO₂. It allows modifying this file and including a new data set
24 based on a different equation of state (EOS). Compared with the earlier version of T2WELL, where the
25 ECO2N module is applied, ECO2M covers the thermophysical properties of CO₂ in a broader range of

1 temperature (between -18 and 360 °C) and pressure (between 0.001 MPa and 200 MPa). The
2 simulation tool uses Altunin's correlation (1975) to describe the thermophysical properties of CO₂ in
3 its original state. One of the main advantages of ECO2M is its capability to describe all possible phase
4 changes of CO₂ between gaseous, supercritical, and liquid states, which is crucial for studying different
5 operational pressure and temperature conditions.

6 The other necessary functionality of T2WELL is coupling the wellbore and reservoir by considering two
7 different sub-domains controlled by different underlying flow physics and the ability to solve strongly
8 non-linear and non-isothermal multiphase CO₂ flow equations within the wellbore and storage
9 formation. The multiphase Darcy's law describes three-dimensional flow through the porous media
10 (i.e., reservoir). Viscous flow within the wellbore is modeled by a one-dimensional momentum
11 equation known as the drift flux model (Shi et al., 2005). In addition, brine evaporation and CO₂
12 solubility are included as analytical functions. The partitioning of H₂O and CO₂ among co-existing
13 aqueous and CO₂-rich phases are derived from the equality of the chemical potential of all components
14 in various phases as a function of pressure, temperature, and salinity. T2WELL-ECO2M utilizes a semi-
15 analytical heat exchange calculation approach to model the heat exchange between wellbore fluid and
16 confining beds and regions outside the flow domain. This approach significantly enhances
17 computational efficiency. Further details of the formulations used in the T2WELL-ECO2M simulator can
18 be found in Pan and Oldenburg 2014. Below we focus on demonstrating the applicability of the code
19 by focusing on the details of the interesting behavior of a model well-reservoir system undergoing CO₂
20 injection and dry out.

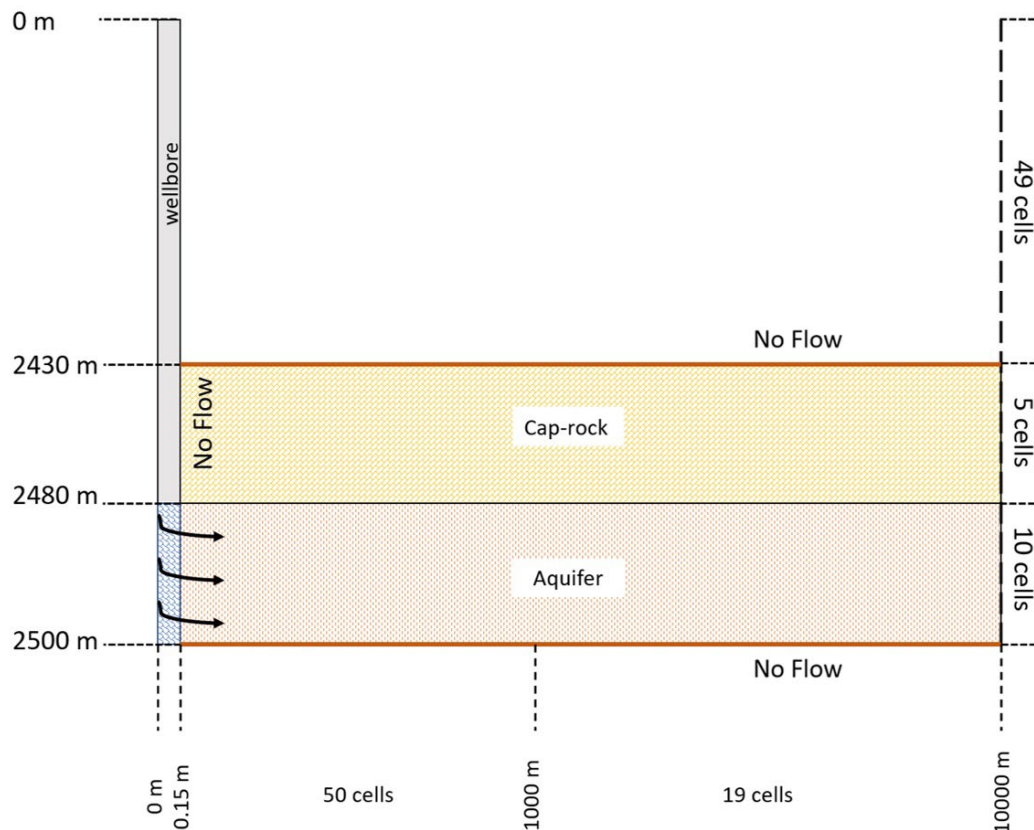
21

22

23 **2.2 Conceptual model description**

24 A synthetic two-dimensional radial model of 10 km in radius is considered, schematically shown in
25 Figure 1. The storage reservoir consists of a 20 m thick permeable formation ($\phi = 0.2$ and $K = 250$ mD)

1 with a depth range of 2480 to 2500 meters, overlaid by a 50 m thick cap rock ($\phi = 0.1$ and $K = 0.01$ mD)
2 with the depth range 2430 to 2480 meters. Homogeneous domains are assumed for both the storage
3 formation and cap rock. The ratio of horizontal to vertical permeability is set to 1. The wellbore and
4 porous media are initially filled with water. The water salinity was set to zero to study only the dry-out
5 phenomenon due to evaporation in the well-near reservoir region and to decouple salt precipitation
6 and its potential effects on pore blockage and pressure build-up. The assumed surface conditions of
7 20 °C and 1 atm, hydrostatic pressure gradient of 10.4 kPa/m, and geothermal gradient of 0.016 °C/m
8 resulted in storage formation pressure and temperature equal to 26 MPa and 60°C, respectively. In the
9 model, the wellbore is situated on the far-left side and extends from the surface down to the reservoir,
10 positioned at a depth of 2500 meters. The internal diameter of the wellbore is 15 cm. The wellbore is
11 solely connected to the storage section located between 2480 and 2500 meters, with no connection
12 assumed along other parts. In these regions, a semi-analytical approach for heat exchange is utilized
13 instead.



14

1 *Figure 1: Schematic illustration of the model with the wellbore, storage formation, and cap rock*

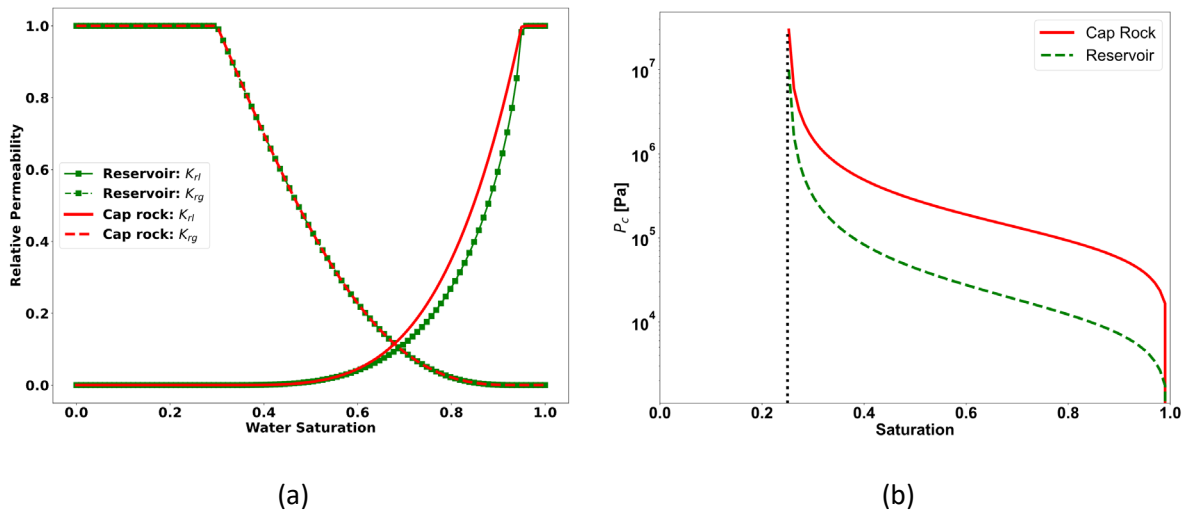
2 The wellbore is discretized into 49 grids, each with a uniform length of 50 meters, up to the top of the
 3 cap rock at a depth of 2430 meters. Beyond that point, the cap rock section is divided into five layers
 4 with thicknesses that vary from 30 meters at the top to 2 meters at the bottom. The storage formation
 5 is divided into ten layers, each with a constant thickness of 2 meters. The porous medium is also
 6 discretized radially with variable mesh size so that finer mesh is used near the wellbore to capture
 7 different phenomena more accurately and gradually becomes coarser towards the outer boundary.
 8 The simulation model comprises a total of 1100 cells, including both the reservoir and wellbore. Of
 9 these cells, 65 are located within the wellbore, while 690 and 345 are within the reservoir and cap rock
 10 sections, respectively.

11 An infinite volume grid is utilized at the rightmost side of the reservoir and caprock (i.e., the Dirichlet
 12 boundary condition) to maintain constant pressure and temperature at the initial reservoir conditions.
 13 No-flow boundaries are assumed for the caprock's upper portions and the reservoir's lower portions.
 14 The leftmost boundary condition is governed by the injection rate assigned to the topmost cell in the
 15 wellbore. Dry CO₂ is injected from the topmost wellbore grid, entering the reservoir along the ten
 16 bottommost cells (Figure 1). An extra high-volume cell above the wellhead is considered to maintain
 17 the operational condition. The Van Genuchten-Mualem model (Genuchten, 1980; Mualem, 1976) is
 18 used for capillary pressure and relative permeability inside the reservoir, and Corey's model for relative
 19 permeability inside caprock. Typically, the capillary pressure is more significant in the caprock than in
 20 the sandstone reservoir because the cap rock has smaller pore sizes. The parameters of the Van
 21 Genuchten model are selected to imitate this phenomenon. Additional information about the
 22 conceptual model can be found in Table 1, while capillary pressure and relative permeability curves
 23 are depicted in Figure 2.

Formation Properties	Values	
	Storage	Cap rock

Thickness [m]	20	50
Porosity [-]	0.2	0.1
Permeability [mD]	250	1e-2
Thermal conductivity [W/m.°C]	2.5	1.72
Rock grain Specific Heat [J/Kg °C]	1000	1000
Compressibility [1/Pa]	8.5e-10	8.5e-10
Transport Parameters	Values	
	Storage	Cap rock
Relative permeability	Van Genuchten model: $\lambda = 0.7, S_{lr} = 0.3,$ $S_{ls} = 0.95, S_{gr} = 0.05$	Corey's model: $S_{lr} = 0.3, S_{gr} = 0.05$
Capillary pressure	Van Genuchten model: $\lambda = 0.457, S_{lr} = 0.25, 1/P_0 =$ $8 \cdot 10^{-5}, P_{max} = 10^7, S_{ls} =$ 0.999	Van Genuchten model: $\lambda = 0.5, S_{lr} = 0.25,$ $1/P_0 = 10^{-5}, P_{max} = 10^8, S_{ls} = 1$
Wellbore Properties	Values	
Internal diameter (I.D.) [m]	0.15	
Length [m]	2500	
Injection temperature [°C]	5	
Wellbore roughness [m]	5e-5	
Injection rate [kg/s] @ wellhead	10	
Thermal Conductivity [W/m. °C]	2.5	
General Properties	Values	
Temperature gradient [°C/m]	0.016	
Pressure gradient [Pa/m]	1.04e4	
Surface pressure [Pa]	1.01e5	
Surface temperature [°C]	20	

1 Table 1: Model properties



1 *Figure 2: (a) relative permeability curves and (b) capillary pressure curves used within cap rock and reservoir.*

2 Case description

3 A synthetic case is utilized to investigate the CO₂ flow inside the wellbore, comprising phase transitions,
 4 pressure/temperature profiles, and the resulting CO₂ front, thermal front, and dry-out region within
 5 the reservoir. The injection of dry CO₂ at a constant mass flow rate of 10 kg/sec as a boundary
 6 condition, a temperature of 5°C, and a pressure of 10.1 MPa right above the wellhead is considered for
 7 this purpose. This injection condition results in a liquid CO₂ state above the wellhead. Liquid CO₂
 8 injection is selected for several reasons, including testing the minimum temperature that T2WELL-
 9 ECO2M can handle, accounting for phase changes in the wellbore, and potentially offering an energy-
 10 efficient injection concept (Vilarrasa et al., 2013). When injection starts, CO₂ flows 2500 meters to
 11 reach the bottomhole while experiencing thermal processes such as heat convection, conduction,
 12 Joule-Thomson, and frictional flow. As a combined result of these processes, the bottomhole
 13 temperature may significantly differ from the initial injecting temperature (i.e., 5 °C) and initial
 14 reservoir temperature (i.e., 60 °C), which affect CO₂ propagation within the storage formation.

15 This study investigates the flow and heat propagation within the reservoir along three lines at distinct
 16 depths. The top layer is the first row of cells beneath the interface between the cap rock and storage
 17 formation at 2480 meters, while the center layer is in the middle of the reservoir at a depth of 2490

1 meters. The bottom layer is positioned at the bottommost row of the reservoir at a depth of 2500
2 meters (Figure 1).

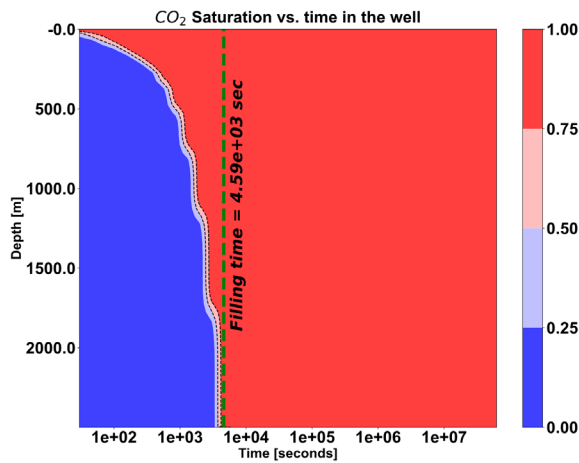
3 **3. Results**

4 ***3.1. Characterization of CO₂ flow within the wellbore***

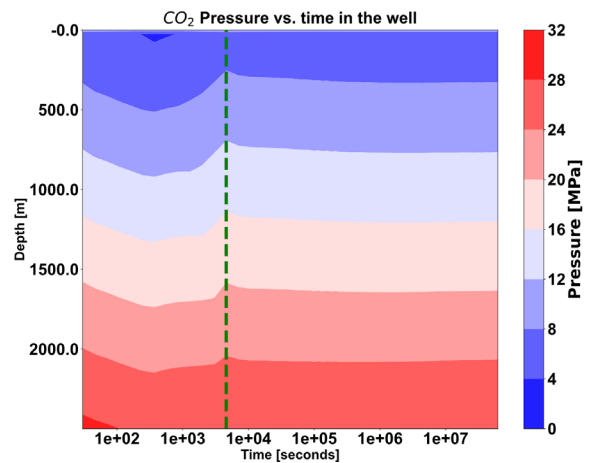
5 Figure 3 shows the evolution of CO₂ a) saturation, b) pressure, and c) temperature within the wellbore
6 over the whole simulation time (i.e., two years). The double dashed line represents the front between
7 water that initially filled the wellbore and injected CO₂, and the vertical dashed line reflects the filling
8 time. Filling time refers to the required time for CO₂ to push the water from the wellbore into the
9 reservoir. In these figures, three regions are recognized, summarized as follows: (I) at the initial state
10 (t = 0 sec), the well is filled by water and pressure, and temperature follows the hydrostatic pressure
11 and geothermal gradient along the wellbore. (II) after injection starts, a narrow transition zone can be
12 recognized where the wellbore flow changes into a multiphase where both CO₂ and brine are flowing
13 together. As injection starts, CO₂ saturation within the wellbore increases, displacing the water within
14 the wellbore into the reservoir (Figure 3-a). In this period, according to Figure 3-b, for a short time, ca.
15 500 seconds (ca. 8.5 mins), the overall pressure profile decreases, and then until the filling time,
16 pressure build-up is observed. It can be attributed to the higher flow resistance within the porous
17 medium compared to the wellbore. The pressure build-up reaches its peak at the time of filling. The
18 whole transient time lasts for 77 minutes. From the practical point of view, before starting the CO₂
19 injection, water within the wellbore will be replaced by injecting cushion gas to prevent contact of
20 water with CO₂ and possible corrosion within the tubing. However, T2WELL-ECO2M only models CO₂,
21 water, and salt and therefore cannot model this procedure. (III) after filling time, only CO₂ is flowing
22 within the wellbore. Following the filling time, the pressure profile along the wellbore approaches a
23 steady state and remains relatively stable.

24 Figure 3-c shows the temperature profile evolution over time in the wellbore. Initially, the water
25 temperature in the wellbore follows the assigned geothermal gradient. According to the temperature

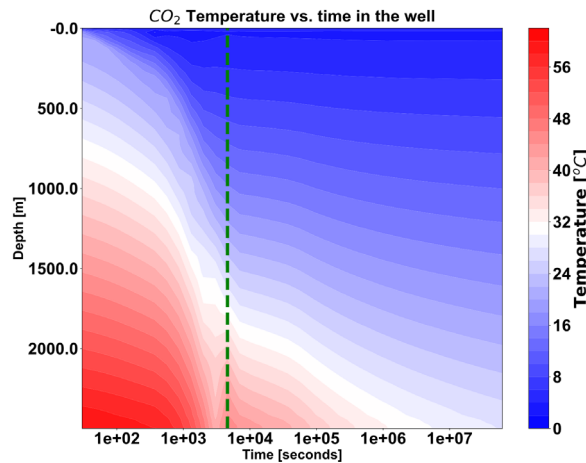
1 gradient, colder water (20 °C) is located at a shallower depth and gradually becomes warmer and
2 reaches 60 °C at the bottomhole. As the injection starts, cold CO₂ at the wellhead pushes the water
3 column within the wellbore into the storage formation. First, the warmer water enters the reservoir,
4 and gradually, colder water from further up in the column enters the reservoir. Due to the high
5 injection rate in the wellbore, the heat transfer mechanism is mainly controlled by convection
6 compared to other mechanisms, such as conduction and frictional effect. The decrease in temperature
7 before the filling time can be captured in Figure 3-c. At the filling time, the CO₂ temperature at the
8 bottomhole is around 40 °C which is significantly higher than the temperature of the injected CO₂ at
9 the surface (5 °C) but still lower than the temperature of the reservoir (60 °C). As illustrated in Figure
10 3-c, after filling time, even though the temperature at the wellhead reaches the steady state, the
11 temperature at the bottomhole is still changing due to convection heat transfer. As injection continues,
12 colder CO₂ will cool down the surrounding formation, reducing the transfer of heat from the formation
13 to the CO₂ flowing in the wellbore. This leads to the continuous arrival of colder CO₂ to the bottomhole.



(a)



(b)



(c)

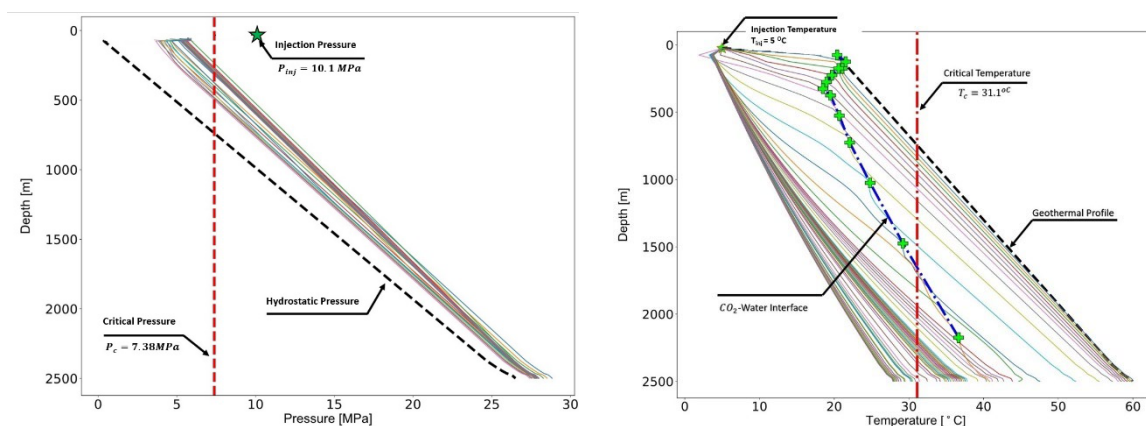
1 *Figure 3: (a) saturation, (b) pressure, and (c) temperature profiles vs. time within the wellbore*

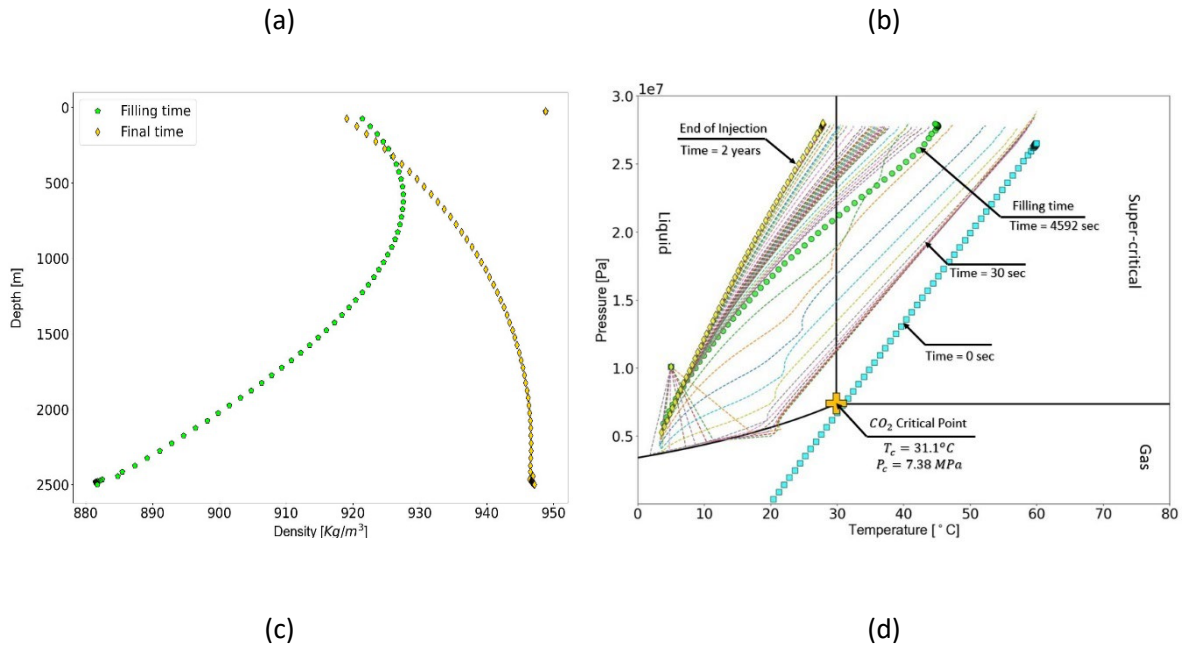
2 As a complement to Figure 3, we present Figure 4 to illustrate the pressure and temperature profiles
3 along the wellbore at different times. The thick dashed lines correspond to the initially assigned
4 geothermal temperature and hydrostatic pressure. According to Figure 4a-b, it is shown that the
5 pressure profile quickly deviates from the initial hydrostatic pressure, and during the injection, the
6 pressure profile remains constant. It is due to the injection conditions that the CO₂ density becomes
7 large at 940 kg/m³ which is very close to the original brine density (999 kg/m³), which was used to
8 calculate the hydrostatic pressure profile (Figure 4-c). Conversely, the temperature profile deviates
9 very slowly from the geothermal gradient, and this is due to an interplay between different heat
10 processes, including convective flow in the wellbore and heat transfer from the ambient formation.
11 The temperature profile clearly illustrates that CO₂ is heating up as it goes down the wellbore but with
12 a slower gradient than the geothermal gradient. Therefore, the temperature at the bottomhole is
13 significantly lower than the reservoir temperature, consistent with field observations reported in
14 numerous articles (Bissell et al., 2011; Lu & Connell, 2014a, 2014b; Paterson et al., 2008, 2010).

15 In addition, the data in Figure 4a, b can be used for two purposes: (a) To indicate at which depth phase
16 transition can occur and track its development, and (b) to show how fast phase changes arise/diminish,
17 and the pressure and temperature reached at steady-state conditions. In this case, phase transitions
18 can be captured twice: the first is at the early stages, when phase change involves the transition from

1 liquid (at the wellhead) to gaseous state. This phase change happened right after the injection started
2 and gradually expanded within the wellbore. After 244 seconds, the phase-change front expanded up
3 to 300 meters, and after that, the expansion length reduced and completely diminished after 858
4 seconds. The second transition occurred at a depth of 1750 meters at $t = 3019$ s, marking the initial
5 point at which CO_2 achieved the conditions for a supercritical state (Figure 4-b). Over approximately
6 40 days, CO_2 flowed in a supercritical state below a depth of 1750 meters, gradually diminishing until
7 it flowed solely in the liquid state.

8 Phase transition and CO_2 properties continuously change over the injection period, initially from liquid
9 to gas and later from liquid to supercritical and vice versa. To gain a more comprehensive
10 understanding of the phase transitions and behavior during the injection period, we plotted
11 temperature and pressure profiles on the CO_2 phase transition diagram, as shown in Figure 4-d. It
12 highlights three different times, (I) first, the initial condition as injection starts (Time = 0 s), (II) second,
13 when all of the water in the wellbore is replaced by CO_2 , which is stated by the filling time (Time = 4592
14 s) and (III) third, at the end of the injection after two years. However, the temperature and pressure
15 variation rate after the filling time is significantly reduced compared to the initial transient flow period.
16 Meanwhile, the Pressure-Temperature profile follows the iso-enthalpy lines. This behavior aligns with
17 previous findings in various studies, indicating that pressure and temperature profiles within the
18 wellbore can transiently change during CO_2 injection. (Li et al., 2015; Liu et al., 2015; Lu & Connell,
19 2008, 2014a, 2014b).



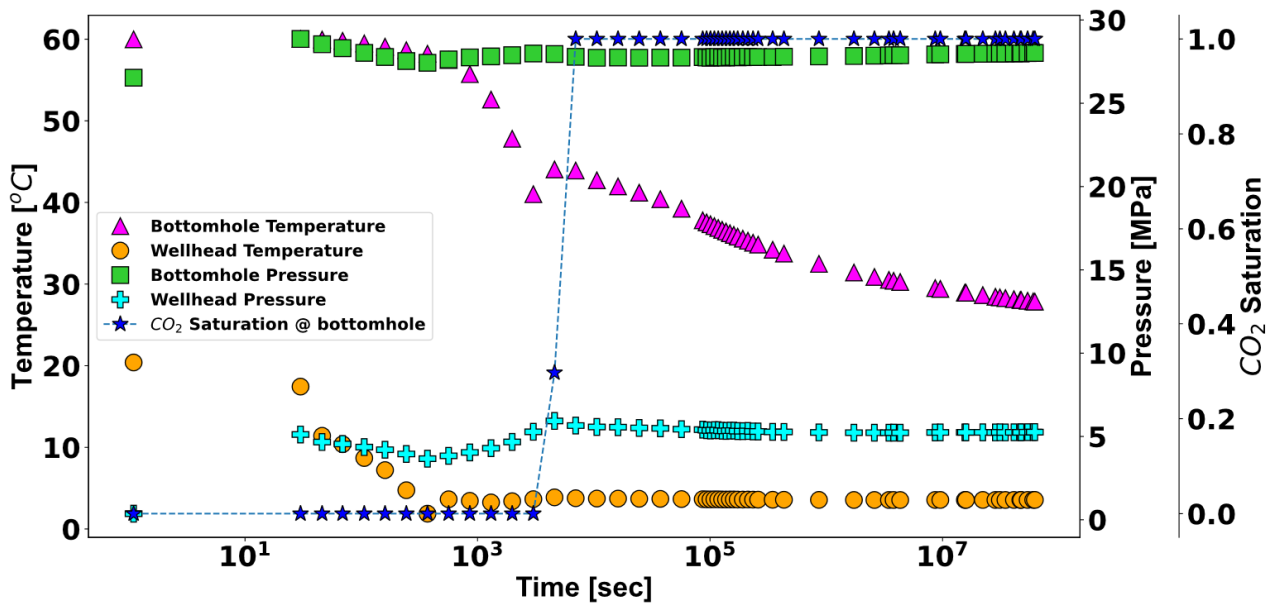


1 *Figure 4: Pressure and temperature profile evolution along the wellbore: (a) temperature profile vs. depth, (b) pressure*
 2 *profile vs. depth and (c) density profile, and (d) pressure-temperature plot vs. CO₂ phase behavior*

3 Figure 5 illustrates the CO₂ saturation, pressure, and temperature evolution at the wellhead and
 4 bottomhole. The CO₂ saturation profile is used to track the location of CO₂ within the wellbore. The
 5 temperature profile at the bottomhole is mainly influenced by the colder water in the wellbore before
 6 the filling time. As the CO₂ is injected into the reservoir, the colder water is displaced toward the
 7 reservoir resulting in a downward trend in the temperature profile. However, just before the filling
 8 time, there is a slight uptick in the bottomhole temperature, with a rise of 3 °C from 41 to 44 °C. This
 9 increase in temperature can be attributed to the compression of the CO₂ as it enters the storage
 10 formation. Following the filling time, a sustained cooling effect is observed, which persists throughout
 11 the injection period. This cooling effect is significant, with the bottomhole temperature gradually
 12 decreasing to around 30 °C after two years of injection, which is half the actual bottomhole
 13 temperature of 60 °C and six times higher than the injection temperature. This continuous cooling can
 14 be attributed to the decreasing heat transfer capabilities of the surrounding environment, which
 15 reduces its ability to heat the fluid in the wellbore.

1 In contrast to the bottomhole, the temperature profile at the wellhead displays a distinctly different
2 pattern. Initially, the CO₂ cools down after injection begins, but the temperature remains higher than
3 the injected temperature (5°C) due to heat exchange with the surrounding environment (20°C). The
4 cooling observed downstream is a result of decompression or expansion. The pressure at surface (one
5 cell above wellhead) is maintained at 10.1 MPa to keep the CO₂ in a liquid state, whereas the pressure
6 at the wellhead drops to approximately 5 MPa. This sudden pressure drop causes expansion, leading
7 to a phase transition from liquid to gas and a decrease in temperature due to the Joule-Thomson effect.
8 The cooling effect at the wellhead stabilizes within a few hundred seconds, reaching a steady
9 temperature of around 3.5°C. The combined effect of these heat processes results in a temperature at
10 the wellhead significantly lower than the ambient temperature (20°C) and slightly lower than the
11 injection temperature.

12 The pressure profile at the bottomhole is mainly governed by the injection rate. In contrast, the
13 pressure value at the wellhead includes the pressure profile within the wellbore and overpressure from
14 the reservoir. As illustrated in Figure 5, the pressure profile at the bottomhole is relatively constant
15 and around 2 MPa higher than reservoir pressure. The pressure profile at the wellhead shows little
16 pressure build-up due to higher resistance toward flow within the reservoir. This variation is limited to
17 the filling time; after that, no pressure change is observed in the wellhead.



1

2 *Figure 5: Tracking the pressure, temperature, and saturation profile at the wellhead and bottomhole as a function of time.*

3

4 **3.2. Characterization of CO_2 flow within the storage formation**

5 Monitoring pressure, temperature, and saturation propagation within the reservoir is crucial for
 6 practical purposes such as tracking CO_2 plume migration and identifying potential issues related to
 7 wellbore and storage integrity. The properties of CO_2 in the wellbore can impact the flow
 8 characteristics and features of the reservoir, including heat propagation, the dry-out region near the
 9 wellbore, and the movement of CO_2 away from the wellbore. The dynamics of CO_2 plume are primarily
 10 governed by the density of CO_2 , which in turn is controlled by temperature. Figure 6 depicts the
 11 propagation of CO_2 saturation, temperature, and pressures over four different time frames, 100 days,
 12 one year, 1.5 years, and two years. The gravity override caused by the difference between CO_2 and
 13 water densities is evident in Figure 6 a-d. (Alkan et al., 2023).

14 As cold CO_2 is injected into the reservoir, it creates a cool zone near the wellbore. However, as it travels
 15 deeper into the reservoir, its temperature gradually increases, causing a change in density. The CO_2
 16 density is higher near the wellbore (916 kg/m^3) and lower at the plume front (790 kg/m^3). This change
 17 in density causes the CO_2 to migrate preferentially through the upper portion of the reservoir and

1 exhibit a more stable front near the wellbore. Another notable observation is that even though the
2 caprock has considerably lower permeability than the storage formation, the CO₂ still advances slightly
3 into the caprock near the wellbore. This region is of practical importance, as it could potentially result
4 in leakage, and further investigation is necessary.

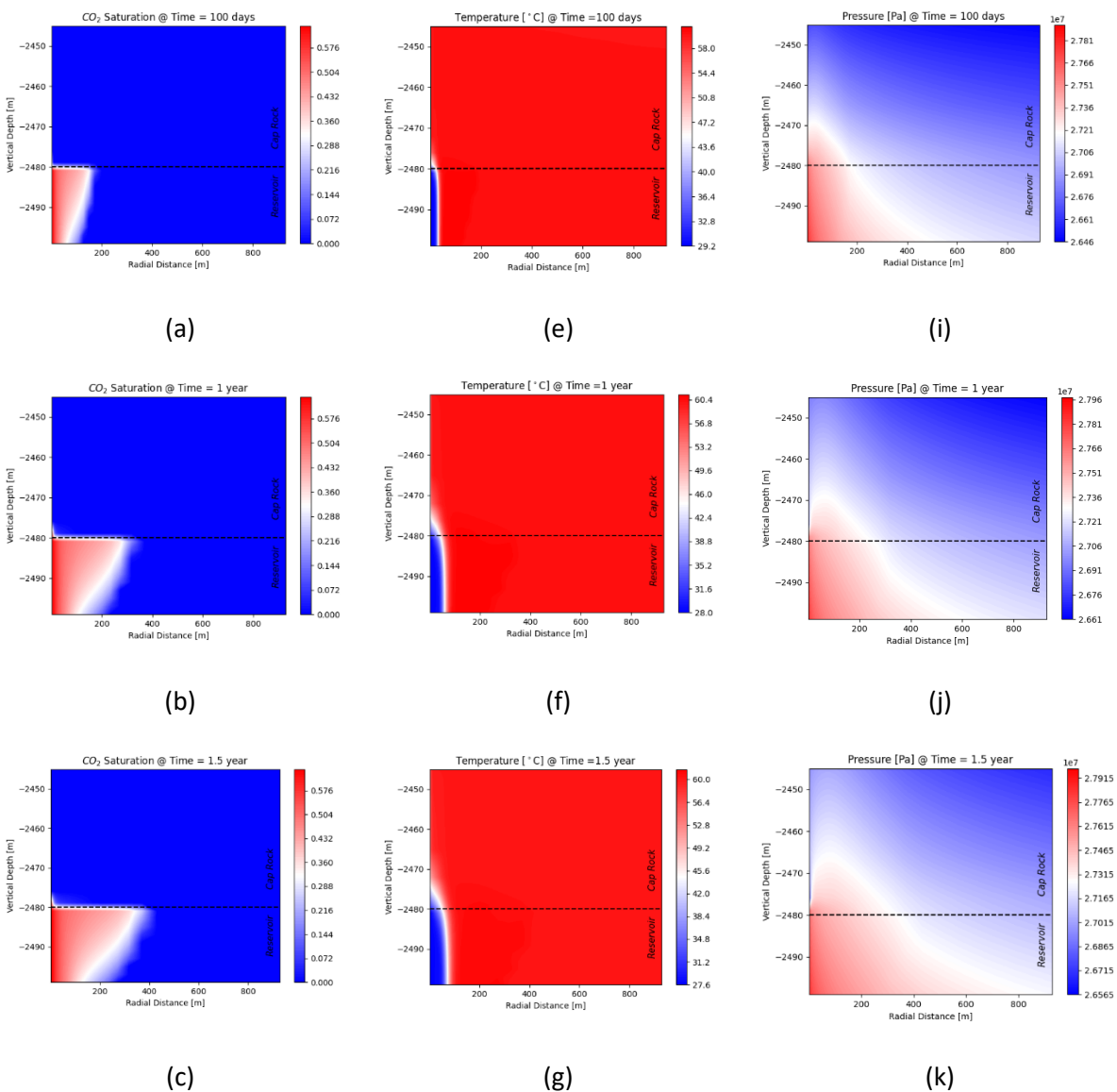
5 Temperature propagation is illustrated in Figure 6 e-h, which shows a significantly lower rate of
6 propagation in comparison with CO₂ migration. To better understand the low propagation of heat in
7 comparison with CO₂ saturation, thermal and pressure diffusivity can be compared (Pruess et al.,
8 2012). Thermal diffusivity can be defined by Equation 1, in which λ is thermal conductivity, ρ_r is rock
9 density, and C_r is rock-specific heat. Pressure diffusivity is calculated by Equation 2, in which K is
10 permeability, ϕ is rock porosity, c is fluid compressibility, and μ is the fluid density. The corresponding
11 parameter values are summarized in Table 1, and based on the values, pressure diffusivity is around
12 one [m²/s] while thermal diffusivity is 10⁻⁶ [m²/s], several orders smaller than pressure diffusivity.

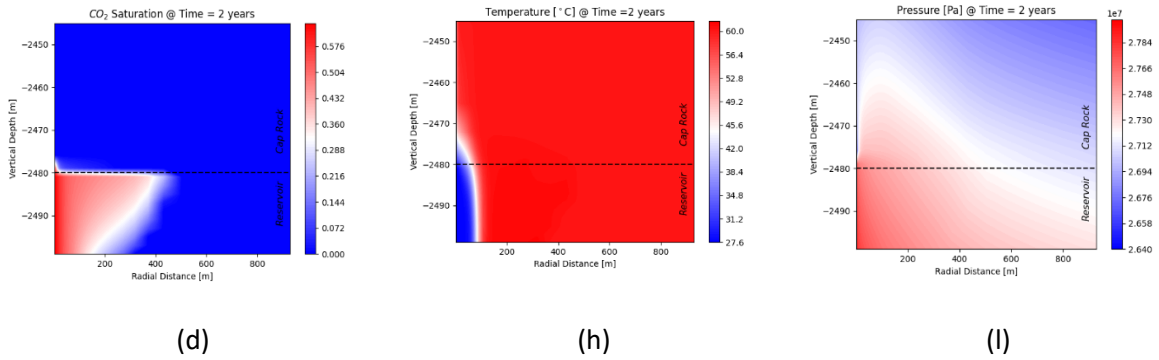
$$\Theta = \frac{\lambda}{\rho_r C_r} \quad \text{Equation 1}$$

$$D_p = \frac{K}{\phi c \mu} \quad \text{Equation 2}$$

13 The cooling of the storage formation is limited to a relatively small area near the wellbore and back of
14 the casing, as Rutqvist (2012) noted. This cooling can cause a significant decrease in temperature,
15 which may alter the stress state and activate new fractures closer to failure conditions. Furthermore,
16 cooling the lower part of the cap rock that is in contact with the storage formation can also activate or
17 destabilize fractures within the cap rock, leading to safety concerns and the risk of leakage (Sagu &
18 Pao, 2013; Vilarrasa & Laloui, 2016). In light of these observations, it is crucial to consider coupled
19 thermal-hydrologic-mechanical (THM) geomechanical simulations when planning CO₂ storage, as this
20 can help identify potential issues with cap rock stability, cap rock leakage, and induced seismicity
21 (Rutqvist, 2012).

1 Figure 6 i-l illustrates the propagation of pressure waves within storage formation and caprock
 2 resulting from CO₂ injection. The pressure distribution follows the expected pattern of higher pressure
 3 at the bottom and lower pressure at the top, which also causes lower fluid density at the top and
 4 higher density at the bottom. On the other hand, from a geomechanical standpoint, pressure
 5 propagation through the cap rock may activate fractures and increase the risk of CO₂ leakage from the
 6 storage formation. Therefore, it is crucial to consider the impact of pressure propagation on the
 7 caprock stability and potential leakage when planning CO₂ storage.

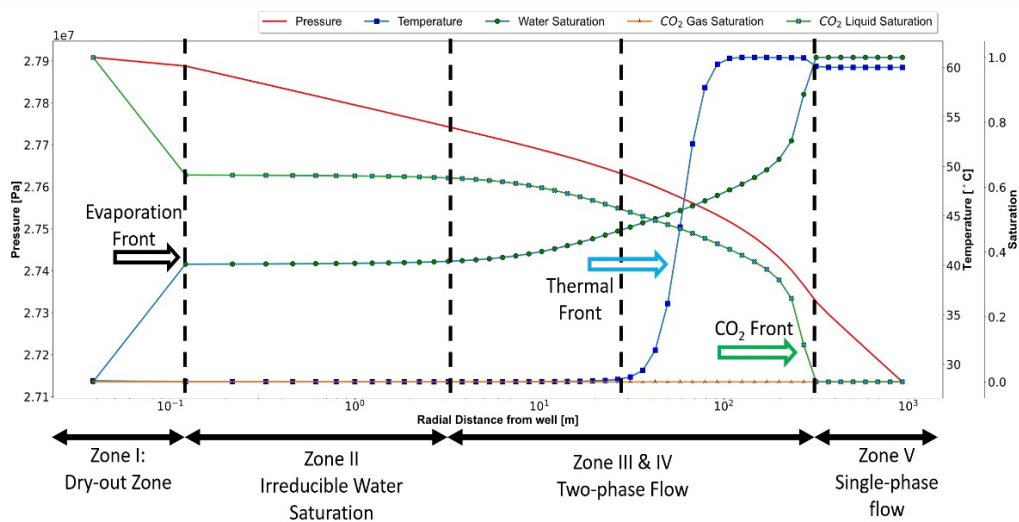




1 *Figure 6: CO₂ Saturation, temperature, and pressure profiles within cap rock and storage formation at four different times:*
 2 *100 days, 1, 1.5, and 2 years.*

3 Figure 7 demonstrates the propagation of pressure, temperature, and saturation of both CO₂ and
 4 water along the center of the reservoir as a function of distance from the wellbore after two years of
 5 injection. To better visualize the profiles near the wellbore, the x-axis is shown on a logarithmic scale.
 6 When CO₂ injection reaches the reservoir, a two-phase region forms where the CO₂-rich phase and
 7 brine flow together. Behind this front, residual brine will remain trapped and be exposed to dry CO₂,
 8 initiating the evaporation regime. As a result, the molar water fraction in the CO₂ stream increases. If
 9 the injection continues for a sufficient period, all the remaining water will evaporate and dissolve into
 10 dry CO₂, forming a dry-out region.
 11 Figure 7 presents three different fronts, namely the CO₂ front, thermal front, and evaporation front.
 12 The CO₂ front is where the CO₂ saturation becomes zero, and water saturation becomes one. The
 13 thermal front represents where the reservoir temperature decreases from its initial value. Just ahead
 14 of this front, there is a slight increase in temperature due to the exothermic reaction of CO₂ dissolution
 15 in water (e.g., up to 5°C). The evaporation front is where all the water in the reservoir is evaporated
 16 by dry CO₂, creating a dry-out region. In this zone, the water saturation gradually decreases below the
 17 irreducible water saturation and eventually reaches zero, whereas the CO₂ saturation reaches one as
 18 no salt is included in this simulation. Along the CO₂ profile, four regions can be identified, which include
 19 the dry-out region, the transition region where CO₂ flows at irreducible water saturation and causes
 20 continuous evaporation of water, the two-phase region where both CO₂ and brine flow simultaneously,
 21 and the single-phase brine flow that occurs far deep in the reservoir where no CO₂ has yet arrived.

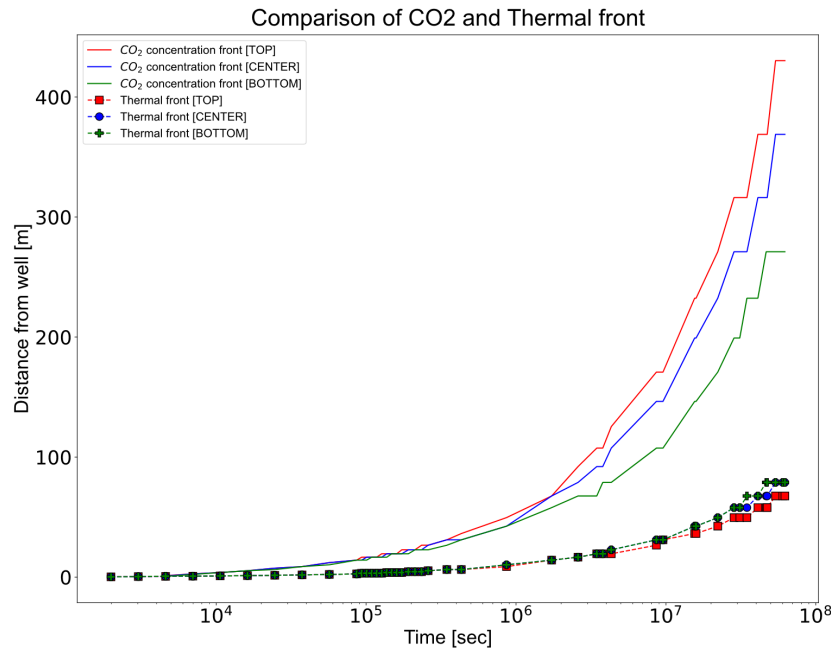
1 The evaporation front is determined by two key parameters: (a) the evaporation onset time and (b)
 2 the extent of the dry-out region. Two main factors govern these parameters: i) the migration speed of
 3 the dry-out zone, which is influenced by viscous forces such as injection rate and horizontal
 4 permeability of the storage formation, and ii) the buoyancy force, which is affected by factors such as
 5 CO₂ density, vertical permeability, and the rate of brine counterflow. (Miri & Hellevang, 2016).
 6 The relative migration rates of the various fronts in our simulations are as follows: CO₂ front >> thermal
 7 front > evaporation front. The thermal front lags behind the CO₂ front due to the rock heat capacity
 8 causing a delay in heat propagation. The evaporation front is much farther behind and may not occur
 9 during injection in some cases. Figure 8 illustrates the location of the CO₂ and thermal fronts at
 10 different layers within the reservoir over time to highlight the discrepancies between them. The speed
 11 of the thermal front is far lower than the CO₂ front, which agrees with the pressure and thermal
 12 diffusivity calculation presented earlier and simulation results (Rayward-Smith & Woods, 2011). The
 13 migration speed of the fronts also varies at different depths. The migration rate of the CO₂ front is
 14 higher in the top layer due to the buoyancy force, while the thermal front propagates slower due to
 15 its interaction with the cap rock.



16

17

Figure 7: Radial propagation of main reservoir parameters after two years of continuous CO₂ injection



1

2

Figure 8: Mapping location of CO₂ front and thermal front as a function of time along layers located at different depth

3

4 4. Discussion

5

Numerical investigations of cold CO₂ injection in a deep aquifer using T2WELL-ECO2M reveal distinct

6

thermal and hydraulic processes. In the wellbore, CO₂ experiences heating due to ambient heat

7

exchange, compression, and frictional loss during its downward flow. However, the heating rate is

8

typically lower than the geothermal gradient, resulting in CO₂ arriving at the bottomhole with a lower

9

temperature (30 °C) than the reservoir temperature (60 °C). Upon entering the reservoir, the colder

10

CO₂ cools down the surrounding rock. Additionally, minor cooling (1-2 °C) can occur due to the Joule-

11

Thomson effect and water vaporization. However, in the aquifer, the Joule-Thomson effect is less

12

pronounced than we expect it to be in depleted hydrocarbon reservoirs. Dissolving CO₂ into brine is

13

exothermic and can raise the temperature by up to 5 °C, consistent with prior simulation findings and

14

field observations. (Bissell et al., 2011; Han et al., 2010; Lu & Connell, 2014b, 2014a; Paterson et al.,

15

2008, 2010).

1 The reservoir can be divided into five regions around the wellbore based on hydraulic and thermal
2 processes, as shown in Figure 7. Initially, CO₂ dissolves in the reservoir brine at the interface between
3 brine and CO₂ front, resulting in a local temperature increase, pH reduction, and improving injectivity
4 by mineral dissolution, shown by region IV. Once the dissolution process is complete, CO₂ displaces the
5 brine, creating a two-phase region (region III) where both fluids co-flow. When the CO₂ ultimately
6 pushed the brine into the reservoir away from the wellbore, immobile water was left as water films.
7 The extent of this region is controlled by relative permeability and capillary pressure curves. In region
8 II, irreducible water is exposed to dry CO₂, and the vaporization process occurs continuously. After a
9 while, when irreducible water is sufficiently exposed to dry CO₂, the whole water will be vaporized,
10 and a dry-out region will be formed (region I).

11 In this study, CO₂ phase conditions vary during the injection, leading to a short dry-out region near the
12 wellbore (10 cm) where CO₂ at the bottomhole was in supercritical condition. The extent of this region
13 is highly dependent on various parameters, including reservoir properties, CO₂ injection rate, and
14 phase condition. Preliminary results confirm that the dry-out region only extends a relatively short
15 distance from the wellbore, even over a long injection period.

16 **5. Conclusions**

17 This study demonstrated the capability of T2WELL-ECO2M, a numerical code from the TOUGH family
18 of codes, to accurately model the CO₂ injection process in the wellbore and storage formation. The
19 code successfully captured phase transitions, including gaseous to liquid, liquid to supercritical, and
20 vice versa, which is challenging in numerical simulations. The successful modeling of phase transitions
21 by T2WELL-ECO2M opens up opportunities to study various conditions for CO₂ injection, such as
22 injecting cold CO₂ into the system. The code also simulated thermal processes, such as heat exchange
23 with the ambient, Joule-Thomson effect, water vaporization, CO₂ dissolution in the brine, frictional
24 flow, and convection, providing reliable data.

1 One advantage of T2WELL is its semi-analytical approach to modeling heat exchange, improving
2 computational efficiency. The study observed minor cooling from the Joule-Thomson effect and CO₂
3 heating due to dissolution in brine. Water vaporization caused by dry CO₂ exposure was adequately
4 captured, but in this specific study, the dry-out region was relatively small, depending on injection
5 conditions and reservoir properties. These findings lay the foundation for further research on the
6 impact of parameters like reservoir heterogeneity, injection rate, CO₂ condition, and wellbore
7 properties on CO₂ flow, which is essential for proper CCS project design.

8 However, some limitations of using T2WELL-ECO2M were also encountered. First, assigning proper
9 initial conditions, including pressure and temperature profiles within the wellbore, is crucial for
10 obtaining realistic outcomes. Another perspective is that the code's ability to simulate field operations
11 precisely may be limited from a practical standpoint. For example, injection of a different gas as a
12 cushion gas is typically done before CO₂ injection to prevent corrosion, but T2WELL-ECO2M only allows
13 the simulation of a CO₂-NaCl-water system. Clearly the needs of current applications in the area of
14 detailed design of geologic storage systems for CCS motivate further developments of numerical
15 simulation capabilities.

16 **Acknowledgement**

17 The authors are thankful to Wintershall Dea AG for the possibility to publish this paper.
18 Acknowledgement is given to the CLIMIT DEMO Project, grant no. 621097 "INJECTWELL – Experimental
19 and Numerical Assessments of CO₂ Injectivity and Flow Assurance During Storage in Depleted
20 Hydrocarbon Reservoirs", a joint industry project, sponsored by the CLIMIT Program (collaboration
21 between Gassnova and the Research Council of Norway) and Wintershall Dea AG.

1 **References**

- 2 Aakre, H., Mathiesen, V., & Moldestad, B. (2018). Performance of CO₂ flooding in a heterogeneous
3 oil reservoir using autonomous inflow control. *Journal of Petroleum Science and Engineering*,
4 167, 654–663. <https://doi.org/10.1016/J.PETROL.2018.04.008>
- 5 Alkan, H., Burachok, O., & Kowollik, P. (2023). Geologic carbon storage: key components. *Surface*
6 *Process, Transportation, and Storage*, 325–422. [https://doi.org/10.1016/B978-0-12-823891-](https://doi.org/10.1016/B978-0-12-823891-2.00009-0)
7 2.00009-0
- 8 Al-Khdheewi, E. A., Vialle, S., Barifcani, A., Sarmadivaleh, M., Zhang, Y., & Iglauer, S. (2018). Impact
9 of salinity on CO₂ containment security in highly heterogeneous reservoirs. *Greenhouse Gases:*
10 *Science and Technology*, 8(1), 93–105. <https://doi.org/10.1002/GHG.1723>
- 11 André, L., Azaroual, M., & Menjot, A. (2010). Numerical simulations of the thermal impact of
12 supercritical CO₂ injection on chemical reactivity in a carbonate saline reservoir. *Transport in*
13 *Porous Media*, 82(1), 247–274. <https://doi.org/10.1007/S11242-009-9474-2/METRICS>
- 14 Bissell, R. C., Vasco, D. W., Atbi, M., Hamdani, M., Okwelegbe, M., & Goldwater, M. H. (2011). A full
15 field simulation of the in Salah gas production and CO₂ storage project using a coupled geo-
16 mechanical and thermal fluid flow simulator. *Energy Procedia*, 4, 3290–3297.
17 <https://doi.org/10.1016/J.EGYPRO.2011.02.249>
- 18 Buursink, M. L. (2014). Significance of carbon dioxide density estimates for basin-scale storage
19 resource assessments. *Energy Procedia*, 63, 5130–5140.
20 <https://doi.org/10.1016/J.EGYPRO.2014.11.543>
- 21 Dalkhaa, C., Jiang, T., Burton-Kelly, M. E., Scharenberg, M., Smith, V., Walker, J. L., Duguid, A.,
22 Heinrichs, M. R., Bosshart, N. W., & Sorensen, J. A. (2022). A simulation study of carbon storage
23 with active reservoir management. *Greenhouse Gases: Science and Technology*, 12(1), 4–23.
24 <https://doi.org/10.1002/GHG.2119>
- 25 Finley, R. J., Frailey, S. M., Leetaru, H. E., Senel, O., Couëslan, M. L., & Marsteller, S. (2013). Early
26 Operational Experience at a One-million Tonne CCS Demonstration Project, Decatur, Illinois,
27 USA. *Energy Procedia*, 37, 6149–6155. <https://doi.org/10.1016/J.EGYPRO.2013.06.544>
- 28 Genuchten, M. Th. van. (1980). A Closed-form Equation for Predicting the Hydraulic Conductivity of
29 Unsaturated Soils. *Soil Science Society of America Journal*, 44(5), 892–898.
30 <https://doi.org/10.2136/SSSAJ1980.03615995004400050002X>
- 31 Han, W. S., Stillman, G. A., Lu, M., Lu, C., McPherson, B. J., & Park, E. (2010). Evaluation of potential
32 nonisothermal processes and heat transport during CO₂ sequestration. *Journal of Geophysical*
33 *Research: Solid Earth*, 115(B7). <https://doi.org/10.1029/2009JB006745>
- 34 Henniges, J., Liebscher, A., Bannach, A., Brandt, W., Hurter, S., Köhler, S., & Möller, F. (2011). P-T-p
35 and two-phase fluid conditions with inverted density profile in observation wells at the CO₂
36 storage site at Ketzin (Germany). *Energy Procedia*, 4, 6085–6090.
37 <https://doi.org/10.1016/J.EGYPRO.2011.02.614>
- 38 Hill, B., Hovorka, S., & Melzer, S. (2013). Geologic Carbon Storage Through Enhanced Oil Recovery.
39 *Energy Procedia*, 37, 6808–6830. <https://doi.org/10.1016/J.EGYPRO.2013.06.614>

- 1 Hovorka, S. D., Benson, S. M., Doughty, C., Freifeld, B. M., Sakurai, S., Daley, T. M., Kharaka, Y. K.,
2 Holtz, M. H., Trautz, R. C., Nance, H. S., Myer, L. R., & Knauss, K. G. (2006). Measuring
3 permanence of CO₂ storage in saline formations: the Frio experiment. *Environmental*
4 *Geosciences*, 13(2), 105–121. <https://doi.org/10.1306/EG.11210505011>
- 5 Kim, K.-Y., Shik Han, W., Oh, J., Park, E., Lee, P.-K., Kim, K., Oh, J., Lee, P., Han, W. S., & Park, E. (2014).
6 Flow Dynamics of CO₂ /brine at the Interface Between the Storage Formation and Sealing Units
7 in a Multi-layered Model. *Transp Porous Med*, 105, 611–633. [https://doi.org/10.1007/s11242-](https://doi.org/10.1007/s11242-014-0387-3)
8 014-0387-3
- 9 Kumar, N., Augusto Sampaio, M., Ojha, K., Hoteit, H., & Mandal, A. (2022). Fundamental aspects,
10 mechanisms and emerging possibilities of CO₂ miscible flooding in enhanced oil recovery: A
11 review. *Fuel*, 330, 125633. <https://doi.org/10.1016/J.FUEL.2022.125633>
- 12 Li, X., Xu, R., Wei, L., & Jiang, P. (2015). Modeling of wellbore dynamics of a CO₂ injector during
13 transient well shut-in and start-up operations. *International Journal of Greenhouse Gas Control*,
14 42, 602–614. <https://doi.org/10.1016/J.IJGGC.2015.09.016>
- 15 Lindeberg, E. (2011). Modelling pressure and temperature profile in a CO₂ injection well. *Energy*
16 *Procedia*, 4, 3935–3941. <https://doi.org/10.1016/J.EGYPRO.2011.02.332>
- 17 Litynski, J., Rodosta, T., Vikara, D., & Srivastava, R. (2013). U.S. DOE's R&D Program to Develop
18 Infrastructure for Carbon Storage: Overview of the Regional Carbon Sequestration Partnerships
19 and other R&D Field Projects. *Energy Procedia*, 37, 6527–6543.
20 <https://doi.org/10.1016/J.EGYPRO.2013.06.584>
- 21 Liu, D., Li, Y., Song, S., & Agarwal, R. (2016). Simulation and analysis of lithology heterogeneity on
22 CO₂ geological sequestration in deep saline aquifer: a case study of the Ordos Basin.
23 *Environmental Earth Sciences*, 75(11), 1–13. [https://doi.org/10.1007/S12665-016-5754-](https://doi.org/10.1007/S12665-016-5754-7/FIGURES/15)
24 7/FIGURES/15
- 25 Liu, D., Li, Y., Xu, L., & Yu, Y. (2015). Numerical investigation of the influence of interaction between
26 wellbore flow and lateral reservoir flow on CO₂ geological sequestration. *Environmental Earth*
27 *Sciences*, 74(1), 715–726. <https://doi.org/10.1007/S12665-015-4076-5/FIGURES/11>
- 28 Lu, M., & Connell, L. D. (2008). Non-isothermal flow of carbon dioxide in injection wells during
29 geological storage. *International Journal of Greenhouse Gas Control*, 2(2), 248–258.
30 [https://doi.org/10.1016/S1750-5836\(07\)00114-4](https://doi.org/10.1016/S1750-5836(07)00114-4)
- 31 Lu, M., & Connell, L. D. (2014a). The transient behaviour of CO₂ flow with phase transition in
32 injection wells during geological storage - Application to a case study. *Journal of Petroleum*
33 *Science and Engineering*, 124, 7–18. <https://doi.org/10.1016/J.PETROL.2014.09.024>
- 34 Lu, M., & Connell, L. D. (2014b). Transient, thermal wellbore flow of multispecies carbon dioxide
35 mixtures with phase transition during geological storage. *International Journal of Multiphase*
36 *Flow*, 63, 82–92. <https://doi.org/10.1016/J.IJMULTIPHASEFLOW.2014.04.002>
- 37 Mathieson, A., Midgely, J., Wright, I., Saoula, N., & Ringrose, P. (2011). In Salah CO₂ Storage JIP: CO₂
38 sequestration monitoring and verification technologies applied at Krechba, Algeria. *Energy*
39 *Procedia*, 4, 3596–3603. <https://doi.org/10.1016/J.EGYPRO.2011.02.289>
- 40 Miri, R., & Hellevang, H. (2016). Salt precipitation during CO₂ storage—A review. *International*
41 *Journal of Greenhouse Gas Control*, 51, 136–147. <https://doi.org/10.1016/J.IJGGC.2016.05.015>

- 1 Movahedzadeh, Z., Shokri, A. R., Chalaturnyk, R., Nickel, E., & Sacuta, N. (2021). Measurement,
2 monitoring, verification and modelling at the Aquistore CO₂ storage site. *First Break*, 39(2), 69–
3 75. <https://doi.org/10.3997/1365-2397.FB2021013/CITE/REFWORKS>
- 4 Mualem, Y. (1976). A new model for predicting the hydraulic conductivity of unsaturated porous
5 media. *Water Resources Research*, 12(3), 513–522. <https://doi.org/10.1029/WR012I003P00513>
- 6 Page, B., Turan, G., Zapantis, A., Burrows, J., Consoli, C., Erikson, J., Havercroft, I., Kearns, D., Liu, H.,
7 Rassool, D., Tamme, E., Temple-Smith, L., Townsend, A., & Zhang, T. (2020). *The Global Status of*
8 *CCS 2020: Vital to Achieve Net Zero*.
9 <https://www.h2knowledgecentre.com/content/researchpaper1679>
- 10 Pan, L., & Oldenburg, C. M. (2014). T2Well—An integrated wellbore–reservoir simulator. *Computers*
11 *& Geosciences*, 65, 46–55. <https://doi.org/10.1016/J.CAGEO.2013.06.005>
- 12 Pan, L., Oldenburg, C. M., Pruess, K., & Wu, Y.-S. (2011). Modeling and Analysis Transient CO₂
13 leakage and injection in wellbore-reservoir systems for geologic carbon sequestration †. *Ltd /*
14 *Greenhouse Gas Sci Technol*, 1, 335–350. <https://doi.org/10.1002/ghg>
- 15 Paterson, L., Ennis-King, J., & Sharma, S. (2010). Observations of Thermal and Pressure Transients in
16 Carbon Dioxide Wells. *Proceedings - SPE Annual Technical Conference and Exhibition*, 5, 3449–
17 3460. <https://doi.org/10.2118/134881-MS>
- 18 Paterson, L., Lu, M., Connell, L. D., & Ennis-King, J. (2008). Numerical Modeling of Pressure and
19 Temperature Profiles Including Phase Transitions in Carbon Dioxide Wells. *Proceedings - SPE*
20 *Annual Technical Conference and Exhibition*, 4, 2693–2703. <https://doi.org/10.2118/115946-MS>
- 21 Piao, J., Han, W. S., Choung, S., & Kim, K. Y. (2018). Dynamic Behavior of CO₂ in a Wellbore and
22 Storage Formation: Wellbore-Coupled and Salt-Precipitation Processes during Geologic CO₂
23 Sequestration. *Geofluids*, 2018. <https://doi.org/10.1155/2018/1789278>
- 24 Preston, C., Monea, M., Jazrawi, W., Brown, K., Whittaker, S., White, D., Law, D., Chalaturnyk, R., &
25 Rostron, B. (2005). IEA GHG Weyburn CO₂ monitoring and storage project. *Fuel Processing*
26 *Technology*, 86(14–15), 1547–1568. <https://doi.org/10.1016/J.FUPROC.2005.01.019>
- 27 Pruess, K., Oldenburg, C., & Moridis, G. (2012). *TOUGH2 USER'S GUIDE, VERSION 2*.
- 28 Rangriz Shokri, A., Talman, S., Chalaturnyk, R. J., & Nickel, E. (2021). *On the Temporal Evolution of*
29 *Non-Isothermal Injectivity Behaviour at an Active CO₂ Injection Site*. OnePetro.
30 <https://dx.doi.org/>
- 31 Rayward-Smith, W. J., & Woods, A. W. (2011). Some implications of cold CO₂ injection into deep
32 saline aquifers. *Geophysical Research Letters*, 38(6), 6407.
33 <https://doi.org/10.1029/2010GL046412>
- 34 Rutqvist, J. (2012). The geomechanics of CO₂ storage in deep sedimentary formations. *Geotechnical*
35 *and Geological Engineering*, 30(3), 525–551. [https://doi.org/10.1007/S10706-011-9491-](https://doi.org/10.1007/S10706-011-9491-0/FIGURES/4)
36 [0/FIGURES/4](https://doi.org/10.1007/S10706-011-9491-0/FIGURES/4)
- 37 Sagu, O. L., & Pao, W. K. S. (2013). In-situ stress perturbation due to temperature around borehole
38 during carbon injection. *Asian Journal of Applied Sciences*, 6(1), 40–49.
39 <https://doi.org/10.3923/AJAPS.2013.40.49>

- 1 Sato, K., Mito, S., Horie, T., Ohkuma, H., Saito, H., Watanabe, J., & Yoshimura, T. (2011). Monitoring
2 and simulation studies for assessing macro- and meso-scale migration of CO₂ sequestered in an
3 onshore aquifer: Experiences from the Nagaoka pilot site, Japan. *International Journal of*
4 *Greenhouse Gas Control*, 5(1), 125–137. <https://doi.org/10.1016/J.IJGGC.2010.03.003>
- 5 Shi, H., Holmes, J. A., Durlofsky, L. J., Aziz, K., Diaz Teran Ortegon, L. R., Alkaya, B., & Oddie, G. (2005).
6 Drift-Flux Modeling of Two-Phase Flow in Wellbores. *SPE Journal*, 10(01), 24–33.
7 <https://doi.org/10.2118/84228-PA>
- 8 Sokama-Neuyam, Y. A., Adu-Boahene, F., Boakye, P., Aggrey, W. N., & Ursin, J. R. (2020). Theoretical
9 modeling of the effect of temperature on CO₂ injectivity in deep saline formations. *Greenhouse*
10 *Gases: Science and Technology*, 10(1), 4–14. <https://doi.org/10.1002/GHG.1951>
- 11 Sokama-Neuyam, Y. A., Aggrey, W. N., Boakye, P., Sarkodie, K., Oduro-Kwarteng, S., & Ursin, J. R.
12 (2022). The effect of temperature on CO₂ injectivity in sandstone reservoirs. *Scientific African*,
13 15. <https://doi.org/10.1016/J.SCIAF.2021.E01066>
- 14 Strpic, K., Bonduà, S., Bortolotti, V., Macini, P., Battistelli, A., & Pan, L. (2021). *Evaluation of Time-*
15 *Convolution Approach for Modelling Heat Exchange Between Wellbore and Formation During*
16 *Multiphase CO₂ Injection with T2Well-ECO2M*. OnePetro.
- 17 Torp, T. A., & Gale, J. (2004). Demonstrating storage of CO₂ in geological reservoirs: The Sleipner and
18 SACS projects. *Energy*, 29(9–10), 1361–1369. <https://doi.org/10.1016/J.ENERGY.2004.03.104>
- 19 Vilarrasa, V., & Laloui, L. (2016). Impacts of Thermally Induced Stresses on Fracture Stability During
20 Geological Storage of CO₂. *Energy Procedia*, 86, 411–419.
21 <https://doi.org/10.1016/J.EGYPRO.2016.01.042>
- 22 Vilarrasa, V., & Rutqvist, J. (2017). Thermal effects on geologic carbon storage. *Earth-Science Reviews*,
23 165, 245–256. <https://doi.org/10.1016/J.EARSCIREV.2016.12.011>
- 24 Vilarrasa, V., Silva, O., Carrera, J., & Olivella, S. (2013). Liquid CO₂ injection for geological storage in
25 deep saline aquifers. *International Journal of Greenhouse Gas Control*, 14, 84–96.
26 <https://doi.org/10.1016/J.IJGGC.2013.01.015>
- 27 Würdemann, H., Möller, F., Kühn, M., Heidug, W., Christensen, N. P., Borm, G., & Schilling, F. R.
28 (2010). CO₂SINK—From site characterisation and risk assessment to monitoring and
29 verification: One year of operational experience with the field laboratory for CO₂ storage at
30 Ketzin, Germany. *International Journal of Greenhouse Gas Control*, 4(6), 938–951.
31 <https://doi.org/10.1016/J.IJGGC.2010.08.010>

32