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Accounting for canopy structure improves hyperspectral radiative transfer and sun-induced chlorophyll fluorescence representations in a new generation Earth System model

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### Authors

Braghiere, Renato K  
Wang, Yujie  
Doughty, Russell  
et al.

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1 **Accounting for canopy structure improves hyperspectral radiative transfer and sun-**  
2 **induced chlorophyll fluorescence representations in a new generation Earth System**  
3 **model**

4 Renato K. Braghieri<sup>1,2\*</sup>, Yujie Wang<sup>3</sup>, Russell Doughty<sup>3</sup>, Daniel Sousa<sup>1</sup>, Troy Magney<sup>4</sup>, Jean-Luc  
5 Widlowski<sup>5</sup>, Marcos Longo<sup>1</sup>, A. Anthony Bloom<sup>1</sup>, John Worden<sup>1</sup>, Pierre Gentine<sup>6</sup>, Christian  
6 Frankenberg<sup>1,3</sup>

7 <sup>1</sup> Jet Propulsion Laboratory, California Institute of Technology, 4800 Oak Grove Drive,  
8 Pasadena, CA, 91109 USA

9 <sup>2</sup> Joint Institute for Regional Earth System Science and Engineering, University of California at  
10 Los Angeles, Los Angeles, CA, 90095 USA

11 <sup>3</sup> Division of Geological and Planetary Sciences, California Institute of Technology, Pasadena,  
12 CA 91125

13 <sup>4</sup> Department of Plant Sciences, University of California, Davis, CA, USA

14 <sup>5</sup> European Commission-DG Joint Research Centre, Institute for Environment and Sustainability,  
15 Ispra, VA I-21027, Italy

16 <sup>6</sup> Department of Earth and Environmental Engineering, Columbia University, New York, New  
17 York, 10027 USA

18 \*Corresponding author: Dr. Renato K. Braghieri ([renato.k.braghiere@jpl.nasa.gov](mailto:renato.k.braghiere@jpl.nasa.gov))

19 Current address: NASA Jet Propulsion Laboratory, M/S 233–305F, 4800 Oak Grove Drive,  
20 Pasadena, CA 91109, USA.

21

22 **Highlights**

- 23 • Horizontal canopy structure is included in a hyperspectral radiative transfer scheme.
- 24 • Clumping index improves the hyperspectral shortwave radiation partitioning.
- 25 • Accounting for horizontal structure improves calculated SIF against NASA's OCO-3 data.
- 26 • SIF correlation to  $\text{NIR}_v$  improves when canopy clumping is considered.
- 27 • SIF canopy escape fraction better correlates with fAPAR when clumping is considered.

28 **Abstract**

29 Three-dimensional (3D) vegetation canopy structure plays an important role in the way radiation  
30 interacts with the land surface. Accurately representing this process in Earth System Models  
31 (ESMs) is crucial for the modeling of the global carbon, energy, and water cycles and hence future  
32 climate projections. Despite the importance of accounting for 3D canopy structure, the inability to  
33 represent such complexity at regional and global scales has impeded a successful implementation  
34 into ESMs. An alternative approach is to use an implicit clumping index to account for the  
35 horizontal heterogeneity in vegetation canopy representations in ESMs at global scale. This paper  
36 evaluates how modeled hyperspectral shortwave radiation partitioning of the terrestrial biosphere,  
37 as well as Sun-Induced Chlorophyll Fluorescence (SIF) are impacted when a clumping index  
38 parameterization is incorporated in the radiative transfer scheme of a new generation ESM, the  
39 Climate Machine (CliMA). An accurate hyperspectral radiative transfer representation within  
40 ESMs is critical for accurately using of satellite data to confront, constrain, and improve land model  
41 processes. The newly implemented scheme is compared to Monte Carlo calculations for idealized  
42 scenes from the Radiation transfer Model Intercomparison for the Project for Intercomparison of  
43 Land-Surface Parameterizations (RAMI4PILPS), for open forest canopies both with and without  
44 snow on the ground. Results indicate that it is critical to account for canopy structural heterogeneity

45 when calculating hyperspectral radiation transfer. The RMSE in shortwave radiation is reduced for  
46 reflectance (25%), absorptance (66%) and transmittance (75%) compared to the scenario without  
47 considering clumping. Calculated SIF is validated against tower and satellite remote sensing data  
48 with the recently launched NASA Orbiting Carbon Observatory (OCO) 3, showing that including  
49 vertical and horizontal canopy structure when deriving fluorescence can improve model predictions  
50 in up to 51% in comparison to the scenario without clumping. By adding a clumping index into the  
51 CliMA model, the relationship between canopy structure and SIF, Gross Primary Productivity  
52 (GPP), hyperspectral radiative transfer and viewing geometry at the canopy scale can be explored  
53 in detail.

54 **Keywords:** canopy structure, Sun-Induced Chlorophyll Fluorescence, hyperspectral radiative  
55 transfer scheme, Earth System models, energy balance, carbon cycle, NASA Orbiting Carbon  
56 Observatory 3

## 57 **1.0 Introduction**

58 Terrestrial vegetation is the largest carbon sink globally, consistently absorbing almost a third  
59 of all anthropogenic carbon emissions (Friedlingstein et al., 2020). However, the fate of the  
60 terrestrial carbon sink in the future is unclear (Friedlingstein et al., 2014; Schimel et al., 2015;  
61 Wieder et al., 2015; Arora et al., 2020) and addressing this important uncertainty lies in improving  
62 Earth System Models (ESMs) (Sellers, 1997; Prentice et al., 2015; Bonan and Doney, 2018).

63 Most state-of-the-art land surface models (LSMs) within ESMs are confined to one-  
64 dimensional (vertical) radiation transfer, often following a plane-parallel turbid media assumption  
65 based on pioneering work from Sellers (1985) and Verhoef (1984). The radiative transfer within  
66 vegetation canopies is rather complex because it involves multiple scattering and mutual

67 shadowing of leaves, which are non-infinitesimal elements arranging themselves in hundreds of  
68 thousands of different angular configurations.

69 A number of studies have shown that neglecting 3D vegetation canopy structural features may  
70 result in significant biases in estimating the land surface energy and carbon balances. For example,  
71 Sprintsin et al. (2012) showed that differences between sunlit and shaded leaves can lead to a  
72 significant underestimation of the canopy gross primary productivity (GPP), similar to other studies  
73 (Chen et al., 2012; Loew et al., 2014; Braghieri et al., 2019, 2020). In alignment with these previous  
74 results, Loew et al. (2014) found that in extreme cases GPP might be underestimated by as much  
75 as 25% and surface albedo might be overestimated by up to 36%, leading to a radiative forcing of  
76 the order of  $-1.25 \text{ W.m}^{-2}$ .

77 Although highly accurate 3D canopy radiative transfer models have been developed and  
78 validated against observations (Wang and Jarvis, 1990; Gastellu-Etchegorry, 2008; Duursma and  
79 Medlyn, 2012), they often demand extreme computational power and cannot be employed at large  
80 scales over long periods of time (Song et al., 2009). Therefore, these highly parameterized 3D  
81 radiative transfer models are unsuitable for direct implementation into ESMs. To account for the  
82 structural effects of vegetation on radiation partitioning, different parameterizations were  
83 developed and applied in radiative transfer models within LSMs, which often work by modulating  
84 the optical depth, or the leaf area index (LAI), of the vegetation canopy through the addition of an  
85 effective variable, the so-called clumping index (Nilson, 1971; Baldocchi and Harley, 1995;  
86 Kucharik et al., 1999; Pinty et al., 2006; Ni-Meister et al., 2010; Braghieri et al., 2019, 2020).

87 The clumping index characterizes the horizontal spatial distribution of trees and leaves, from  
88 small to whole-canopy scales (Nilson, 1971; Norman and Jarvis, 1974), and it can be derived from  
89 gap size distribution measured *in-situ* with ceptometers or digital hemispherical photography  
90 (DHP) (Chen and Cihlar, 1995; Leblanc et al., 2002; Leblanc et al., 2005; Ryu et al., 2010b; Fang

91 et al., 2018; Yan et al., 2019), as well as from space with multi-angular remote sensing data (Pisek  
92 et al., 2015; He et al., 2016) and, more recently, from LiDAR data (Wang and Kumar, 2019).

93 Although the clumping index has been commonly used to account for the impacts of vegetation  
94 structure on radiative transfer modeling and further impacts on land surface processes (Baldocchi  
95 et al., 2002; Ryu et al., 2010a; Chen et al., 2012; Braghieri et al., 2019, 2020), previous studies are  
96 often limited to broadband spectral analysis in the photosynthetically active radiation (PAR, 400-  
97 700 nm) and Near Infrared (NIR, 700-2500 nm), mainly due to the direct applicability of these two  
98 broadbands in current ESMs, as well as the limited information about hyperspectral canopy optical  
99 properties. However, new generation ESMs should be able to include hyperspectral canopy  
100 radiative transfer schemes because high resolution spectral data is now available from aircrafts and  
101 will soon be available from space, on the International Space Station (ISS) and later, globally, via  
102 the US Surface Biology and Geology (SBG) concept (Schimel and Schneider, 2019).

103 Hyperspectral data can provide a wide range of unique constraints on plant functional traits  
104 (Butler et al., 2017). For instance, imaging spectroscopy can map terrestrial vegetation properties,  
105 such as canopy water content, leaf nitrogen and phosphorus compositions, as well as a wide range  
106 of traits related to photosynthesis, respiration, and decomposition of leaf material (Singh et al.,  
107 2015). However, current state-of-the-art ESMs are not able to make use of all the extra information  
108 provided by hyperspectral measurements of vegetation, nor are they able to calculate radiative  
109 transfer in such high spectral resolution.

110 The benefits of using a hyperspectral radiative transfer scheme versus the general broadband  
111 spectral analysis used in current LSMs are linked to: (i) the direct inversion of ecosystem related  
112 parameters from remotely-sensed data (Dutta et al., 2019; Cheng et al., 2020), that has been broadly  
113 used as predictors of ecology related variables, e.g., maximum photosynthetic capacity (Meacham-  
114 Hensold et al., 2019), GPP (Dechant et al., 2019), leaf pigments (Féret et al., 2017), plant traits

115 (Féret et al., 2019), and other morphological and physiological properties (Serbin et al., 2014); and,  
116 (ii) the reduction of uncertainty in surface albedo (Majasalmi and Bright, 2019), and therefore  
117 radiative partitioning and forcing, by moving away from the time-invariant look-up tables of  
118 broadband (PAR and NIR) canopy optical properties originally based on a study published more  
119 than 30 years ago (Dorman and Sellers, 1989). In addition, biases associated with surface  
120 reflectance derivation from remotely-sensed data products are often found when converting  
121 hyperspectral radiation to multispectral radiation through convolution across multiple sensors  
122 (Burggraaff, 2020).

123 Previous studies have developed coupled LSMs to simulate Sun-Induced Chlorophyll  
124 Fluorescence (SIF) (e.g., the Community Land Model (CLM) 4 (Lee et al., 2015), the Biosphere  
125 Energy Transfer Hydrology (BETHY) model (Norton et al., 2019), and the Boreal Ecosystem  
126 Productivity Simulator (BEPS) (Qiu et al., 2019)). In studies with CLM and BETHY, the authors  
127 coupled the original LSMs, capable of simulating carbon assimilation, ecosystem respiration, as  
128 well as the energy and water balances, with the SCOPE (Soil Canopy Observation, Photosynthesis  
129 and Energy fluxes) model (van der Tol et al., 2009; Van Der Tol et al., 2014). The SCOPE model  
130 is a 1D (vertical) radiative transfer and energy balance model that calculates photosynthesis and  
131 chlorophyll fluorescence. SCOPE is based on the 4-stream radiative transfer theory from the SAIL  
132 (Scattering by Arbitrarily Inclined Leaves) model (Verhoef, 1984) and the leaf radiative transfer  
133 model of Fluspect (Vilfan et al., 2016), which is based upon leaf optical properties from the  
134 PROSPECT model (Jacquemoud and Baret, 1990). Apart from recent developments of the SCOPE  
135 model to include some representation of canopy vertical heterogeneity (mSCOPE; Yang et al.,  
136 2017), a limitation of mSCOPE is that it only accounts for vertical variation in canopy properties,  
137 and it has no information about horizontal canopy structure.

138 While the study with BEPS-SIF (Qiu et al., 2019) has explored the impacts of canopy clumping  
139 on SIF emission, the ‘two-leaf’ radiation regime in BEPS (i.e., one vertical vegetation layer with  
140 sunlit and shaded leaves) is different from a vertical multi-layered radiative transfer scheme (e.g.,  
141 two-stream scheme (Sellers, 1985) and 4-stream (Verhoef, 1984)), which had led to divergent  
142 impacts of clumping on GPP (Braghiere et al., 2019) and other aspects of the land surface (Bonan  
143 et al., 2021).

144 The main goal of this study is to introduce and evaluate a clumping index parameterization  
145 scheme used to represent horizontal vegetation canopy structure within a vertically resolved 1D  
146 canopy model, the Climate Model Alliance (CliMA)-Land, within a new generation ESM, the  
147 CliMA model. Here, we aim to investigate the impacts of horizontal vegetation canopy structure  
148 on hyperspectral shortwave radiation partitioning, as well as to determine if by using a  
149 parameterization scheme of vegetation canopy structure through the clumping index, it is possible  
150 to make the commonly used SAIL 4-stream theory (Verhoef, 1984) match the shortwave radiation  
151 partitioning of a more complex 3D radiative transfer model, raytran (Govaerts and Verstraete,  
152 1995, 1998; Widlowski et al., 2011; Hogan et al., 2018).

153 Part of the SCOPE model has been incorporated into BETHY but without the inclusion of  
154 horizontal canopy heterogeneity. Whereas for the clumping index, several LSMs have used this  
155 parameterization scheme in the past (Ni-Meister et al., 2010; Yang et al., 2010; Chen et al., 2012),  
156 but without the fully resolved hyperspectral shortwave radiation. Therefore, the main advantage of  
157 the clumping index implementation in CliMA-Land is bridging the hyperspectral radiative transfer  
158 with explicit consideration of the horizontal canopy heterogeneity. First, the shortwave radiation  
159 partitioning calculated with CliMA-Land is compared with reference values generated in the  
160 Radiation transfer Model Intercomparison for the Project for Intercomparison of Land-Surface  
161 Parameterizations (RAMI4PILPS) experiment (Widlowski et al., 2011), a radiative transfer model



162 intercomparison exercise. Within the RAMI4PILPS framework, models can be evaluated under  
163 perfectly controlled experimental conditions, i.e., all structural, spectral, illumination, and  
164 observation related characteristics are known without ambiguity. Therefore, possible deviations  
165 between model simulations can thus be directly attributed to the assumptions and shortcuts entering  
166 model-specific implementations of the radiative transfer equations. The parameters of a structural  
167 parameterization scheme of clumping index (Pinty et al., 2006) are tested in the CliMA-Land  
168 hyperspectral radiative transfer scheme under different scenarios with and without snow.

169         Second, we use the updated hyperspectral radiative transfer scheme with clumping index  
170 to explore the impact of vegetation structure on the estimation of SIF emission (He et al., 2017;  
171 Magney et al., 2017; Yang et al., 2019; Zeng et al., 2019; Dechant et al., 2020) and related  
172 vegetation indices, commonly used as GPP predictors, such as the fraction of absorbed PAR  
173 (fAPAR), absorbed PAR (APAR), and the near-infrared reflectance of vegetation (NIRv) (Badgley  
174 et al., 2017; Zeng et al., 2019). We validate the estimation of SIF emission using SIF retrievals  
175 from the NASA Orbiting Carbon Observatory 3 (OCO-3) (Eldering et al., 2019) over a subalpine  
176 evergreen needle-leaf forest in Niwot Ridge, Colorado, and a deciduous broadleaf forest at the  
177 University of Michigan Biological (UMB) Station, Michigan, USA. OCO-3's new “snapshot  
178 mode” feature enabled by the instrument’s ability to swivel and point rapidly, produces  
179 measurements over an area of about 80 by 80 kilometers, which allows scanning across a range of  
180 view zenith angles over a single overpass within about 2 minutes. OCO-3 is also unique as far as  
181 spaceborne SIF instruments because it samples over the day following the ISS orbit, which also  
182 allows a broad coverage of different sun zenith angles.

183         The rationale behind the SIF evaluation with and without clumping index lies in a number  
184 of recent studies suggesting that APAR is among the dominant factors explaining the variability of  
185 SIF, and the strong relationship between SIF and GPP (Miao et al., 2018; Wieneke et al., 2018;

186 Yang and van der Tol, 2018; Li et al., 2020; Magney et al., 2020). More recently, a growing number  
187 of studies have suggested that APAR alone cannot explain observed SIF variability, and that other  
188 factors, such as the physiological SIF emission yield ( $\Phi_F$ ) and the fluorescence escape ratio ( $f_{esc}$ )  
189 would also play a significant role in determining SIF (Du et al., 2017; Migliavacca et al., 2017;  
190 Yang et al., 2018; Zeng et al., 2019; Dechant et al., 2020).  $f_{esc}$  has been linked to canopy structure,  
191 commonly described in terms of LAI and leaf angular distribution, and more recently to the  
192 clumping index (Zeng et al., 2019). In this study we also explore some of the impacts of clumping  
193 index on the variability of SIF and its linkage to canopy structural heterogeneity.

## 194 **2.0 Materials and Methods**

195 In this section, firstly, a description of the CliMA-Land radiative transfer model is presented,  
196 followed by a description of independent methods of derivation of SIF relationship with other  
197 vegetation indices, as well as how canopy structure can impact these relationships. Secondly, a  
198 description of the experimental setup and its elements are presented following: (i) a 1D – 3D model  
199 validation exercise against the RAMI4PILPS dataset (Widlowski et al., 2011), as well as the  
200 methodology used to allow a direct intercomparison between broadband and hyperspectral  
201 radiative transfer; and (ii) an independent validation against SIF estimates via satellite remote  
202 sensed observations over an area of evergreen needleleaf forest canopy with heterogeneous canopy  
203 architecture.

### 204 **2.1 CliMA-Land Radiative Transfer Scheme**

205 In this study, we present and evaluate a new important feature of the canopy radiative transfer  
206 model in the land component of a new generation of Earth System Model developed by the Climate  
207 Modeling Alliance (CliMA). The CliMA-Land model addresses soil water movement, plant water  
208 transport, stomatal regulation, canopy radiation, and the fluxes of water, carbon, and energy in a

209 highly modular manner. Code and documentation of the in-progress CliMA-Land model are freely  
210 and publicly available at <https://github.com/CliMA/Land>.

211 The CliMA-Land Radiative Transfer model is based on the vertically heterogeneous mSCOPE  
212 (Yang et al., 2017), which uses Fluspect (Vilfan et al., 2016) to simulate leaf reflectance,  
213 transmittance, and fluorescence at the leaf level, and SAIL based models to compute spectrally  
214 resolved radiative transfer, as well as emitted fluorescence (van der Tol et al., 2016).

215 The CliMA-Land Radiative transfer model was adapted to overcome the assumption of  
216 horizontal vegetation homogeneity following a parameterization scheme proposed by Pinty et al.  
217 (2006), which accounts for horizontal structural heterogeneity with the addition of an extra  
218 parameter, referred to as the clumping index (Nilson, 1971). Nilson (1971) first introduced the  
219 clumping index ( $\Omega$ ) into the Beer-Lambert's law, to describe plant canopy direct transmittance, or  
220 the gap fraction probability ( $P_{gap}(\theta)$ ) as:

$$221 \quad P_{gap}(\theta) = \exp\left(\frac{-G(\theta) \cdot LAI \cdot \Omega}{\cos\theta}\right) \quad (1.0)$$

222 where  $\theta$  is the sun zenith angle, LAI is the leaf area index, and  $G(\theta)$  is the projection coefficient of  
223 unit foliage area on a plane perpendicular to the view direction (Ross, 1981).

224 Analogously to the clumping index, Pinty et al. (2004) developed a parameterization scheme  
225 that modulates the canopy optical depth in order to replicate the behavior of more complex 3D  
226 radiative transfer schemes but accounting for zenith angular variations of canopy structure. The  
227 hypothesis behind this scheme suggests that throughout the day and year, solar radiation crosses  
228 different pathways associated with different structures. Therefore, the clumping index also varies  
229 with sun zenith angle following:

$$230 \quad \Omega(\theta) = \zeta(\theta) = -\ln(1 - F_c) \frac{2}{LAI} + b \cdot (1 - \cos\theta) \quad (2.0)$$

231 where  $\theta$  is the sun zenith angle, LAI is the leaf area index and  $F_c$  is the vegetation cover  
232 corresponding to the ground fractional cover by all vegetation elements including canopy gaps.  
233 The parameter 'b' has no empirical formulation but it can be derived from observations (Braghiere  
234 et al., 2020). Here 'b' is set to zero throughout all the experiments because of its lack of an empirical  
235 formulation that would further limit the applicability of CliMA-Land to other sites on Earth where  
236 information about clumping zenithal variation is not directly available from remotely-sensed  
237 datasets. Therefore, the zenith variation of clumping index is not considered. The clumping index  
238 varies with the radiation pathway, which is linked to the viewing zenith angle, but also to the sun  
239 zenith angle. The clumping index varying with sun zenith angle can be interpreted as the radiation  
240 pathlength varying with sun zenith angle (Kucharik et al., 1999; Pinty et al., 2006; Ryu et al.,  
241 2010b). This parameterization scheme was previously implemented, validated, and tested with the  
242 land surface model of the UKESM, JULES following Braghiere et al. (2018, 2019, 2020).

243 The parameterization scheme can be directly implemented into the classical SAIL 4-stream  
244 model by assuming that the canopy optical depth is equal to an 'effective LAI' ( $LAI \cdot \Omega$ ) instead of  
245 the 'true LAI' (LAI). Hence, the SAIL 4-stream theory can be recast as:

$$246 \quad \frac{dE_s}{\Omega \cdot LAI dx} = kE_s \quad (3.a)$$

$$247 \quad \frac{dE^-}{\Omega \cdot LAI dx} = -sE_s + aE^- - \sigma E^+ \quad (3.b)$$

$$248 \quad \frac{dE^+}{\Omega \cdot LAI dx} = s'E_s + \sigma E^- - aE^+ \quad (3.c)$$

$$249 \quad \frac{dE_o}{\Omega \cdot LAI dx} = wE_s + \nu E^- + \nu' E^+ - KE_o \quad (3.d)$$

250 where  $E_s$  is the direct solar flux,  $E^-$  is the downward diffuse flux,  $E^+$  is the upward diffuse flux, and  
251  $E_o$  is the flux in the viewing direction.  $x$  is the so-called relative optical height,  
252 which runs from  $-1$  at the bottom to zero at the canopy top, and LAI is the leaf area index.  $k$  and  $K$

253 are the extinction coefficients dependent on canopy geometrical characteristics, such as the leaf  
254 angular distribution, the angular positioning of the sun for  $K$ , and the sun-observer geometry for  $k$ .  
255 The remaining scattering coefficients ( $s$ ,  $a$ ,  $\sigma$ ,  $s'$ ,  $w$ ,  $\nu$ ,  $\nu'$ ) depend on canopy and sun-observer  
256 geometry, as well as the canopy optical properties (i.e., leaf reflectance and transmittance). These  
257 coefficients were first described in Verhoef (1984) and revisited in Yang et al. (2017).

## 258 **2.2 Determining SIF, $f_{esc}$ , and NIR<sub>v</sub>**

259 CliMA-Land calculates SIF emission following the mSCOPE model approach (Yang et al.,  
260 2017), where the incident radiation is converted into emitted chlorophyll fluorescence on each side  
261 of the leaf across all canopy layers and leaf angular orientations. The mSCOPE model framework  
262 was used to simulate light scattering within the canopy but using the ‘effective LAI’ ( $LAI \cdot \Omega(\theta)$ )  
263 as the canopy optical depth, instead of ‘true LAI’ (LAI), in order to consider the effects of  
264 horizontal canopy heterogeneity on SIF determination via the addition of a clumping index ( $\Omega(\theta)$ ).

265 The emitted SIF at the top of the canopy in the viewing direction, as well as the hemispherical  
266 integration are calculated following the same radiative transfer equations, but also accounting for  
267 the emitted radiation. Therefore, SIF estimates depend on the radiative transfer throughout the  
268 canopy, the conversion of incident radiation into chlorophyll emission, and finally, the propagation  
269 of re-emitted chlorophyll fluorescence through the canopy (van der Tol et al., 2009; Yang et al.,  
270 2017).

271 The far-red part of SIF ( $>740$  nm) is an optical signal in the NIR spectrum in which radiation  
272 is highly scattered by leaves allowing only a part of it to escape the vegetation canopy (Knyazikhin  
273 et al., 2013; Yang and van der Tol, 2018; Zeng et al., 2019; Dechant et al., 2020). Studies found  
274 that reflectance can be used to explain part of the SIF scattering signal (Liu et al., 2016; van der  
275 Tol et al., 2016; Badgley et al., 2017; Yang and van der Tol, 2018), but the observed SIF from a

276 tower or from space cannot be totally explained by the cumulative signal of SIF emitted by leaves  
277 due to variabilities in canopy structure (Guanter et al., 2014; Zeng et al., 2019; Dechant et al.,  
278 2020). Therefore, observed SIF ( $SIF_{obs}$ ) can be described as:

$$279 \quad SIF_{obs} = APAR \times \Phi_F \times f_{esc} \quad (4.0)$$

280 where  $\Phi_F$  is the physiological SIF emission quantum yield of the whole canopy and  $f_{esc}$  is the  
281 fluorescence escape ratio, which is a fraction of SIF emitted from leaves that actually escape from  
282 the vegetation canopy.

283 Determining  $f_{esc}$  is rather a difficult task because it requires information about: i) canopy  
284 structural properties, such as LAI (Fournier et al., 2012; Yang and van der Tol, 2018), leaf angular  
285 distribution (Du et al., 2017; Migliavacca et al., 2017), and the clumping index (Zeng et al., 2019;  
286 Dechant et al., 2020); ii) leaf spectral properties; and iii) observation-illumination geometry (Zeng  
287 et al., 2019). While a number of studies have explored the influence of  $\Phi_F \times f_{esc}$  together on  $SIF_{obs}$   
288 (Yang et al., 2015; Miao et al., 2018; Wieneke et al., 2018; Li et al., 2020), the potentially strong  
289 impact of leaf angular orientation and canopy clumping on  $f_{esc}$  has often been neglected, or overly  
290 simplified by treating  $f_{esc}$  as a constant (Guanter et al., 2014). Recently, the whole canopy far-red  
291 SIF emission  $f_{esc}$  was approximated by a relationship of  $NIR_v$  and  $fAPAR$  following Zeng et al.  
292 (2019):

$$293 \quad f_{esc} \approx \frac{NIR_v}{fAPAR} \quad (5.0)$$

294 where  $NIR_v$  is the product of NIR reflectance at 792 nm and NDVI  
295 ( $(R_{792nm} - R_{687nm}) / (R_{792nm} + R_{687nm})$ ; Tucker, 1979), a variable that has been shown to be strongly  
296 correlated with  $SIF_{obs}$  at large spatiotemporal scales (Badgley et al., 2017). In order to test the  
297 impact of clumping index on the validity of **Eq. (5)**, an independent study (Yang and van der Tol,  
298 2018) showed that  $f_{esc}$  can be estimated over a black soil condition as:

299 
$$f_{esc} = \frac{R}{i \times \omega_l} \quad (6.0)$$

300 where  $R$  is the NIR reflectance (740 nm),  $i$  is the canopy interceptance, which represents the  
301 probability of a photon interacting with the canopy and it is defined as one minus the directional  
302 gap fraction (Smolander and Stenberg, 2005),  $\omega_l$  is the leaf single scattering albedo and it  
303 corresponds to the fraction of photons at a specific wavelength that escape the canopy (Knyazikhin  
304 et al., 2013).

305 Re-writing **Eq. (6)** in terms of the escape probability theory (Huang et al., 2007), the  
306 recollision probability theory (Smolander and Stenberg, 2005), and the fraction of diffuse radiation,  
307  $f_{esc}$  can be written as:

308 
$$f_{esc} = (1 - f_d) \times \frac{\rho_s}{1 - p_s \times \omega_l} + f_d \times \frac{\rho_d}{1 - p_d \times \omega_l} \quad (7.0)$$

309 where  $f_d$  is the fraction of diffuse solar radiation,  $\rho_{s/d}$  is the escape probability of  
310 sunlit/shaded leaves,  $p_{s/d}$  is the recollision probability of sunlit/shaded leaves, and  $\omega_l$  is  
311 the leaf single scattering albedo. More details on the derivation of **Eq. (7)** and the equations  
312 for  $\rho_{s/d}$  and  $p_{s/d}$  can be found in Appendix A. The impact of clumping index on the  
313 relationship described in **Eq. (5)** is independently tested following the derivation of  $f_{esc}$   
314 through **Eq. (7)**, and  $fAPAR$  and  $NIR_v$  directly calculated from CliMA-Land.

315 In order to verify that the version of CliMA-Land radiative transfer with clumping index is  
316 indeed a better approximation of the relationship proposed by Zeng et al. (2019), two popular  
317 measures of model parsimony (Aho et al., 2014) were also calculated to determine: the Akaike  
318 information criterion (AIC; (Akaike, 1973)) and the Bayesian information criterion (BIC;  
319 (Schwarz, 1978)). The AIC and BIC are statistical variables used to represent how accurately a  
320 determined model fits the data. A better model presents smaller values of AIC and BIC.

### 321 **2.3 RAMI4PILPS benchmarking**

322 Evaluating models can be challenging, especially when it focuses on highly accurate details,  
323 such as 3D architectural features of a scene (Kobayashi et al., 2012). There are different ways to  
324 evaluate the performance of a specific radiative transfer model including comparisons against  
325 different sources of observed data, such as bidirectional reflectance (North, 1996; Malenovský et  
326 al., 2008), transmittance (Wang and Jarvis, 1990; Norman and Welles, 1983; Tournebize and  
327 Sinoquet, 1995; Law et al., 2001; Sinoquet et al., 2001), and gap fraction measurements (Cescatti,  
328 1997; Kucharik et al., 1999; Yang et al., 2010). The use of these observed datasets is often limited  
329 by a restricted spatiotemporal coverage, as well as by a restricted number of suitable instruments.  
330 To eliminate uncertainties arising from an incomplete or erroneous knowledge of the structural,  
331 spectral, and illumination conditions related to canopy characteristics, typical of model validations  
332 with *in-situ* observations, the RAdiative transfer Model Intercomparison (RAMI) (Pinty et al.,  
333 2001, 2004; Widlowski et al., 2007, 2011, 2013, 2015) have been used to evaluate models against  
334 the extensively verified 3D reference Monte Carlo model, raytran (Govaerts and Verstraete, 1995,  
335 1998) under perfectly controlled conditions. In particular, the RAMI4PILPS (Project for  
336 Intercomparison of Land-Surface Parameterizations) suite of experiments (Widlowski et al., 2011)  
337 was designed to evaluate the accuracy and consistency of shortwave radiative transfer formulations  
338 as commonly used in ESMs. Here we use the RAMI4PILPS heterogeneous canopy scenario where  
339 tree crowns were approximated by woodless spheres in an open forest canopy scene. Details of the  
340 RAMI4PILPS experiments used in here are summarized in **Table 1**. For each scenario, simulations  
341 for different LAI values and varying soil albedos are performed, assuming direct radiation for three  
342 different sun zenith angles.

343 We simulate all three components of the radiative partitioning: (i) canopy reflectance,  
344 which is defined as the ratio of reflected to incident radiation at the top-of-canopy, (ii) canopy



345 absorption, which is defined as the fraction of radiation entering the canopy through a reference  
 346 plane at the top-of-canopy, and absorbed by the elements in the scene, and (iii) canopy  
 347 transmittance, which is defined as the amount of spectral energy transmitted through the vegetation.

348  
 349 **Table 1.** Summary of variables defining structurally heterogeneous scenes (see Widlowski et al.  
 350 (2011) for details). Different soil albedos are defined as BLK = black, MED = medium, SNW =  
 351 snow.

<b>Variable Identification</b>	<b>Values (Units)</b>
<b>Leaf Area Index/whole canopy</b>	0.50 <sup>S</sup> , 1.50 <sup>M</sup> and 2.50 <sup>D</sup> (m <sup>2</sup> .m <sup>-2</sup> )
<b>Leaf Area Index/each tree</b>	5.0 <sup>S</sup> , 5.0 <sup>M</sup> and 5.0 <sup>D</sup> (m <sup>2</sup> .m <sup>-2</sup> )
<b>1 – P<sub>gap</sub> (θ = 0°)</b>	0.09 <sup>S</sup> , 0.26 <sup>M</sup> and 0.43 <sup>D</sup>
<b>Tree density</b>	12.80 <sup>S</sup> , 38.24 <sup>M</sup> and 63.68 <sup>D</sup> (trees/hectare)
<b>Maximum canopy height</b>	16 m
<b>Minimum sphere center height</b>	7 m
<b>Maximum sphere center height</b>	11 m
<b>α<sub>soil, PAR</sub> / α<sub>soil, NIR</sub></b>	BLK: 0.00/0.00; MED: 0.12/0.21; SNW: 0.96/0.56
<b>Soil scattering law</b>	Lambertian

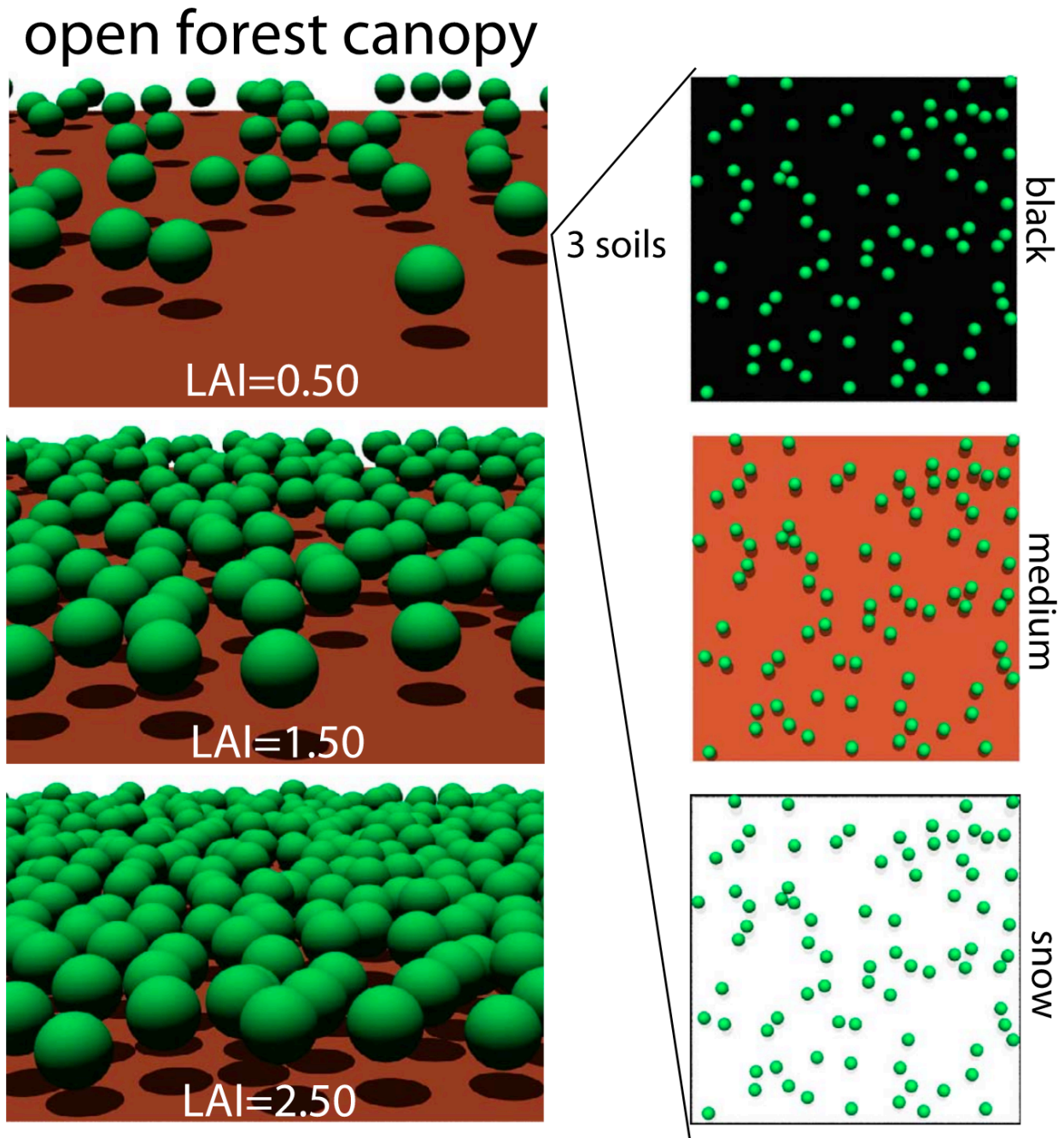
$\rho_{\text{leaf, PAR}} / \rho_{\text{leaf, NIR}}$	0.0735/0.3912
$\tau_{\text{leaf, PAR}} / \tau_{\text{leaf, NIR}}$	0.0566/0.4146
<b>Leaf scattering law</b>	Bi-Lambertian
<b>Sun zenith angle</b>	27.0°/60.0°/83.0°
<b>Scatterer Normal Distribution</b>	spherical
<b>Woody area index</b>	0.0 (m <sup>2</sup> .m <sup>-2</sup> )

352 <sup>s</sup> Sparse vegetation condition.

353 <sup>m</sup> Medium vegetation condition.

354 <sup>d</sup> Dense vegetation condition.

355



356

357 **Figure 1.** Graphical representation of the open forest canopy environments used in the  
358 RAMI4PILPS experiment. Three different leaf area index (LAI) values and three different  
359 background soil albedos (adapted from Widlowski et al. (2011)).

360

## 361 **2.4 Moving the reference values from two broadbands to hyperspectral resolution**

362 The RAMI4PILPS experiment focused on two separate broadbands (PAR and NIR) to be  
363 directly comparable to ESMs, which often make use of the two-stream radiative transfer scheme  
364 in these only two broadbands, separately. Therefore, the canopy spectral properties, i.e., leaf  
365 reflectance and leaf transmittance, are given as an average value representing the entire broadbands  
366 PAR and NIR. In order to move from a broadband radiative transfer scheme to a hyperspectral one,  
367 the reference spectral properties were fitted using the Fluspect model (**Table 2**).

368 The average broadband values of leaf reflectance PAR ( $\rho_{\text{leaf,PAR}}$ ), leaf reflectance NIR  
369 ( $\rho_{\text{leaf,NIR}}$ ), leaf transmittance PAR ( $\tau_{\text{leaf,PAR}}$ ), and leaf transmittance NIR ( $\tau_{\text{leaf,NIR}}$ ) were prescribed  
370 as  $\rho_{\text{leaf,PAR}} = 0.0735$ ,  $\rho_{\text{leaf,NIR}} = 0.3912$ ,  $\tau_{\text{leaf,PAR}} = 0.0566$ , and  $\tau_{\text{leaf,NIR}} = 0.4146$  (**Table 1**), as  
371 previously defined in the RAMI4PILPS experiment. To find the optimal combination of leaf  
372 parameters described in **Table 2** that approximate the prescribed values of leaf optical properties,  
373 each one of the 9 parameters (N,  $C_{\text{AB}}$ ,  $C_{\text{AR}}$ ,  $A_{\text{NT}}$ ,  $C_{\text{S}}$ ,  $C_{\text{W}}$ ,  $C_{\text{M}}$ ,  $C_{\text{X}}$ , and FQE) in its range of plausible  
374 values were minimized independently, following the sum of squared difference between modeled  
375 and prescribed average  $\rho_{\text{leaf,PAR}}$ ,  $\rho_{\text{leaf,NIR}}$ ,  $\tau_{\text{leaf,PAR}}$ , and  $\tau_{\text{leaf,NIR}}$ .

376 A publicly available customized multiple dimensional optimization algorithm was used to  
377 fit leaf spectral parameters (see **Data availability**). In this method: (i) each parameter in **Table 2**  
378 is initialized with an initial guess value; (ii) The first parameter (i.e., N) is calculated to minimize  
379 the sum of squared error, while holding all the other parameters constant; (iii) this method is  
380 repeated for the other variables; (iv) when the set of leaf spectral parameters reaches equilibrium,  
381 the increment step decreases in 10%; and (v) steps ii-iv are repeated until all steps were below their  
382 solution tolerances ( $10^{-9}$ ).

383 Nine parameters (**Table 2**) were fitted to minimize the sum of square difference between  
384 modeled and prescribed average  $\rho_{\text{leaf,PAR}}$ ,  $\rho_{\text{leaf,NIR}}$ ,  $\tau_{\text{leaf,PAR}}$ , and  $\tau_{\text{leaf,NIR}}$ . To best represent leaf biological

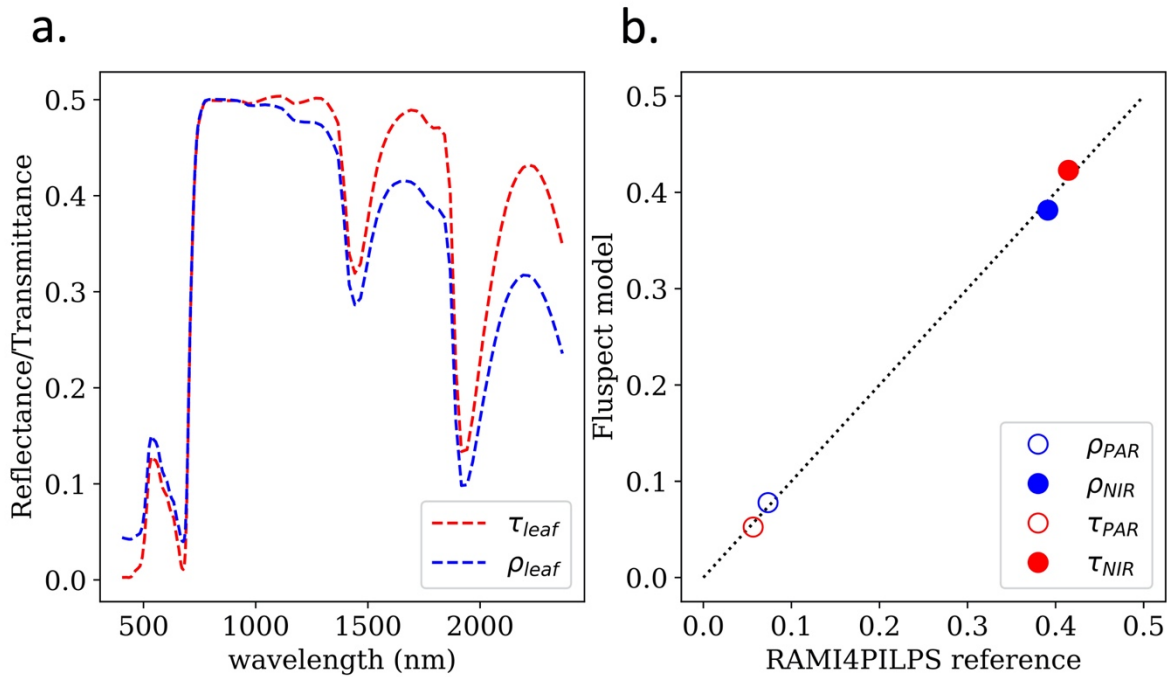
385 properties, we constrained the parameters to their physiological ranges: N in [1,3], C<sub>AB</sub> in [0,100],  
 386 C<sub>AR</sub> in [0,30], A<sub>NT</sub> in [0,40], C<sub>S</sub> in [0,1], C<sub>w</sub> in [0,0.05], C<sub>M</sub> in [0,0.5], C<sub>X</sub> in [0,1], and FQE in [0,1].

387 **Figure 2a** shows the hyperspectral canopy reflectance and transmittance minimized against  
 388 the RAMI4PILPS reference values using Fluspect. The average values for two broadbands  
 389 separately are shown as circles in **Figure2b**.

390  
 391 **Table 2.** Leaf spectral variables and parameters in leaf biochemical model. See (J.-B. Féret et al.,  
 392 2017; Jacquemoud et al., 2009; Jacquemoud and Baret, 1990) for further details.

VARIABLE	DESCRIPTION	UNITS	VALUE
N	Leaf structure parameter	-	1.6
C <sub>AB</sub>	Chlorophyll a + b content	μg cm <sup>-2</sup>	30.0
C <sub>AR</sub>	Carotenoid content	μg cm <sup>-2</sup>	5.0
A <sub>NT</sub>	Anthocyanin content	μg cm <sup>-2</sup>	2.75
C <sub>S</sub>	Senescent material (brown pigments)	fraction	0.0
C <sub>w</sub>	Equivalent water thickness	cm	5.0E-03
C <sub>M</sub>	Dry matter content	μg cm <sup>-2</sup>	0.0
C <sub>X</sub>	Fraction between Zeaxanthin and Violaxanthin in Car (1=all Zeaxanthin)	fraction	0.0
FQE	Leaf fluorescence efficiency	-	0.01

393



394  
395 **Figure 2a.** Hyperspectral leaf reflectance (blue) and leaf transmittance (red) obtained from  
396 Fluspect using values given in **Table 2**; **b.** The average values of these curves are represented by  
397 circles for two broadbands and single scattering albedo term, separately, i.e., PAR (400-700 nm)  
398 and NIR (700-2500 nm); reflectance ( $\rho$ ) and transmittance ( $\tau$ ).

399

## 400 2.5 Study sites

### 401 2.5.1. Niwot Ridge, Colorado, USA

402 The validation study for CliMA-Land radiative transfer simulated SIF was conducted at the  
403 subalpine forest of the Niwot Ridge AmeriFlux Core site (US-NR1) in the Rocky Mountains in  
404 Colorado, USA (40.03°N, 105.55°W, 3050 m elevation). The forest is composed of three dominant  
405 evergreen needleleaf species: lodgepole pine (*P. contorta Douglas ex Loudon*), Engelmann spruce  
406 (*Picea engelmannii Parry ex Engelm.*), and subalpine fir (*Abies lasiocarpa (Hook.) Nutt*). The  
407 vegetation canopy structure consists of an average stem density of 4000 stems.ha<sup>-1</sup>, average tree

408 height of 12.5 m, and LAI of  $3.8 \text{ m}^2 \cdot \text{m}^{-2}$  (Bowling et al., 2018; Magney et al., 2019). Due to its  
409 high elevation, this forest is exposed to cold winters with persistent snowpacks from October to  
410 May (Blanken et al., 2009; Burns et al., 2015).

411 The clumping index at Niwot Ridge was reported as  $0.740 \pm 0.057$  by Sprintsin et al. (2012)  
412 after the remote sensing work of Chen et al. (2005) using POLDER (POLarization and  
413 Directionality of the Earth's Reflectances; 6 km). However, a more recent algorithm based on  
414 MODIS BRDFs (He et al., 2012) reports a clumping index of 0.48 for the 500 m pixel that includes  
415 the US-NR1 flux tower. The main difference from the MODIS clumping index and the one from  
416 POLDER is the spatial resolution.

### 417 **2.5.2. UMB Station, Michigan, USA**

418 The validation study for CliMA-Land radiative transfer simulated SIF was conducted at a  
419 maturing aspen-dominated forest AmeriFlux Core site (US-UMB) in the upper Great Lakes region  
420 in Michigan, USA ( $45.58^\circ\text{N}$ ,  $84.72^\circ\text{W}$ , 234 m elevation). The forest is composed of dominant  
421 deciduous broadleaf species: bigtooth aspen (*Populus grandidentata*) and trembling aspen  
422 (*Populus tremuloides*), but with significant presence of maple (*Acer rubra*, *A. saccharum*), red oak  
423 (*Quercus rubra*), birch (*Betula papyrifera*), and beech (*Fagus gran-difolia*) as well. The vegetation  
424 canopy structure consists of an average stem density of  $700\text{-}800 \text{ stems} \cdot \text{ha}^{-1}$ , average tree height of  
425  $\sim 22 \text{ m}$ , and LAI of  $3.5 \text{ m}^2 \cdot \text{m}^{-2}$  (Schmid, 2003; Gough et al., 2013). The clumping index at UMB  
426 was reported as  $0.700 \pm 0.047$  by Sprintsin et al. (2012) after the remote sensing work of Chen et  
427 al. (2005) and 0.52 from MODIS BRDFs for the 500 m pixel that includes the US-UMB flux tower.

### 428 **2.6 OCO-3 SIF Retrievals**

429 To assess the effect of the clumping index on CliMA Radiative Transfer model estimates  
430 of SIF, we compared simulated SIF computed with and without the clumping index to spaceborne

431 SIF retrievals from the Orbiting Carbon Observatory 3 (OCO-3). We ran the model for each OCO-  
432 3 sounding in three snapshot area maps (SAMs) taken by OCO-3 at Niwot Ridge, Colorado, USA,  
433 two of which were obtained on June 12<sup>th</sup> and June 16<sup>th</sup>, 2020, and two SAMs at UMB Station,  
434 Michigan, USA, taken on August 6<sup>th</sup> and August 11<sup>th</sup>, 2020.

435 OCO-3 is a spectrometer that is similar to OCO-2 and is on the ISS. OCO-3 has the unique  
436 ability to obtain SAMs by scanning a target several times in a single overpass with scans being  
437 offset to obtain a wider sampling of the Earth's surface, which yield large contiguous scans of ~  
438 100 km by 100 km (Eldering et al., 2019). The spatial resolution of each OCO-3 sounding footprint  
439 is  $\leq 4$  km, with the size varying due to viewing geometry. The ISS orbit is precessing rather than  
440 sun-synchronous and it orbits the Earth about 16 times a day, thus overpasses do not occur at the  
441 same local time for any latitude and the amount of time between overpasses for any given target  
442 location is highly variable and unpredictable in the long term.

443 For each sounding footprint, the OCO-3 data provides, among other variables, solar and  
444 viewing zenith and azimuth angles, instantaneous SIF retrieved at 757 nm, landcover classification,  
445 cloud flags, and quality control flags (Frankenberg et al., 2014; Taylor et al., 2020). From these  
446 sun and sensor geometries, we calculated relative azimuth and phase angles for each sounding.  
447 Prior to analysis, we removed soundings classified as barren or urban and also those soundings not  
448 classified as 'best' by the quality control flag and 'clear' by the cloud flag.

449 We also calculated area weighted mean LAI, Cab, and clumping index for each sounding.  
450 We have illustrated SIF<sub>757</sub> and the clumping index for one of the June 12<sup>th</sup>, 2020 overpasses in  
451 **Figure 3**. The LAI map, PROBA-V LAI V2, was produced by Copernicus at 1 km resolution  
452 (Fuster et al., 2020) without consideration of any canopy, understory, or foliage clumping effects,  
453 as stated in their Algorithm Theoretical Basis Documents (ATBD) (Verger et al., 2019). The  
454 temporal resolution is variable, but the file we used had a start date of January 3<sup>rd</sup>, 2020 and an end



455 date of June 30<sup>th</sup>, 2020. The Cab map had a spatial resolution of 0.5 degrees and a weekly temporal  
456 resolution for the years 2003-2011 (Croft et al., 2020). To approximate differences in Cab between  
457 pixels during the OCO-3 overpass, we computed weekly means using all years and used Cab  
458 concentrations from the week in which the overpasses occurred (weeks 24 and 25).

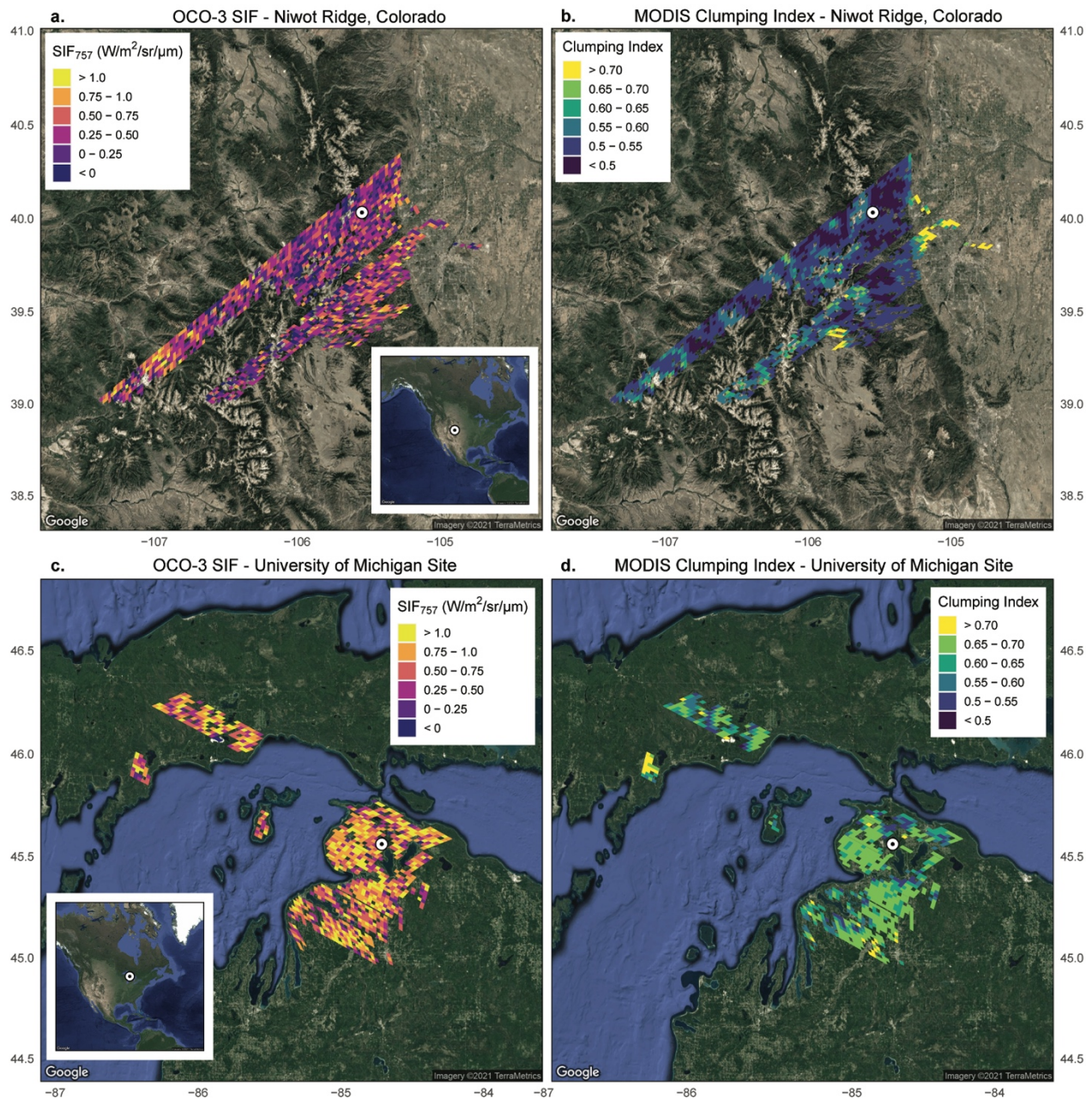
459         The global MODIS-derived clumping map produced by He et al. (2012) was used to  
460 provide a clumping index estimate for the CliMA-Land radiative transfer model. The global  
461 clumping index map has a spatial resolution of 500 m and was produced for the year of 2006. We  
462 assume that the global clumping index map derived for 2006 data is reliable for usage in 2020 since  
463 the interannual variability of clumping index is generally small (He et al., 2016). The data were  
464 derived from the NASA-MODIS BRDF/albedo product (MCD43) by considering the difference in  
465 forward and backward scattering from the surface, which is primarily controlled by the structure  
466 of the vegetation (Braghiere et al., 2019). The MODIS clumping index (He et al. 2012) is an  
467 average for all view zenith angles, not specific to nadir or other angles. It can be derived from  
468 different combinations of hotspot and dark spot values, but the authors used nadir for hotspot and  
469 47.7° for dark spot in order to produce a map that correlates well with observed *in-situ*  
470 measurements.

471         After simulating instantaneous SIF<sub>757</sub> for each OCO-3 sounding using the CliMA radiative  
472 transfer model and input data from OCO-3 (sun-sensor geometries) and area weighted mean LAI,  
473 Cab, and clumping index, we grouped soundings by phase angle and computed the mean for each  
474 group. Individual SIF retrievals are noisy and differences in sun-sensor geometry between  
475 soundings can contribute to differences in the retrieved SIF values. Thus, it is advised not to use  
476 individual soundings for analysis, but retrievals can be averaged across space and/or time to reduce  
477 their standard errors and offset potential differences in viewing geometry ( Frankenberg et al.,

478 2014; Köhler et al., 2018; Doughty et al., 2019). Thus, the points in **Figure 3** are mean SIF<sub>757</sub>  
479 values of soundings from a single orbit with nearly identical viewing geometries and the error bars  
480 represent the standard error of the mean for that group of soundings. Groups with fewer than 10  
481 soundings ( $n < 10$ ) were excluded from the analysis. We ran the model for each OCO-3 sounding  
482 footprint, not only for the sounding including the flux tower (represented by a white circle with a  
483 black dot in the middle in **Figure 3** for reference). To reduce the error, we take their means where  
484 sun-sensor geometry is nearly identical.

485 Topographic effects can be observed on OCO-3 CO<sub>2</sub> retrievals due to air mass  
486 dependencies, but no effect on retrieved SIF. It appears the main effect is physiological in a direct  
487 comparison of OCO-2 targets and CFIS (airborne) overpasses to tower SIF at Niwot Ridge  
488 (Parazoo et al., 2019).

489



490  
491 **Figure 3.** OCO-3 retrieved SIF at 757 nm over **a.** Niwot Ridge, Colorado, USA on June, 12<sup>th</sup>, 2020,  
492 and **c.** UMB Station, Michigan, USA on August 11<sup>th</sup>, 2020. MODIS derived clumping index map  
493 from He et al. (2012) over **b.** Niwot Ridge, Colorado, USA and **d.** UMB Station, Michigan, USA,  
494 for the year of 2006 matching the OCO-3 scan. The white circle with a black dot in the middle  
495 represents the position of the flux towers for reference.

496 **3.0 Results**

497 **3.1 Validating canopy radiative partitioning: broadbands PAR and NIR**

498 **Figure 4a** shows the three components of the radiation partitioning (lines) using the default  
499 case (no clumping) and the respective RAMI4PILPS reference values (circles) for the sparse  
500 canopy case with LAI = 0.5 m<sup>2</sup>.m<sup>-2</sup> and 10% vegetation cover over a black soil ( $\alpha_{\text{soil}} = 0.0$ ). **Figure**  
501 **4b** shows the same example but including clumping derived from **Eq.(2)**, with  $\Omega = 0.37$  and  $b =$   
502 0.0. For similar figures for all the other canopy structures and soil albedos, see **Supplemental**  
503 **material**.

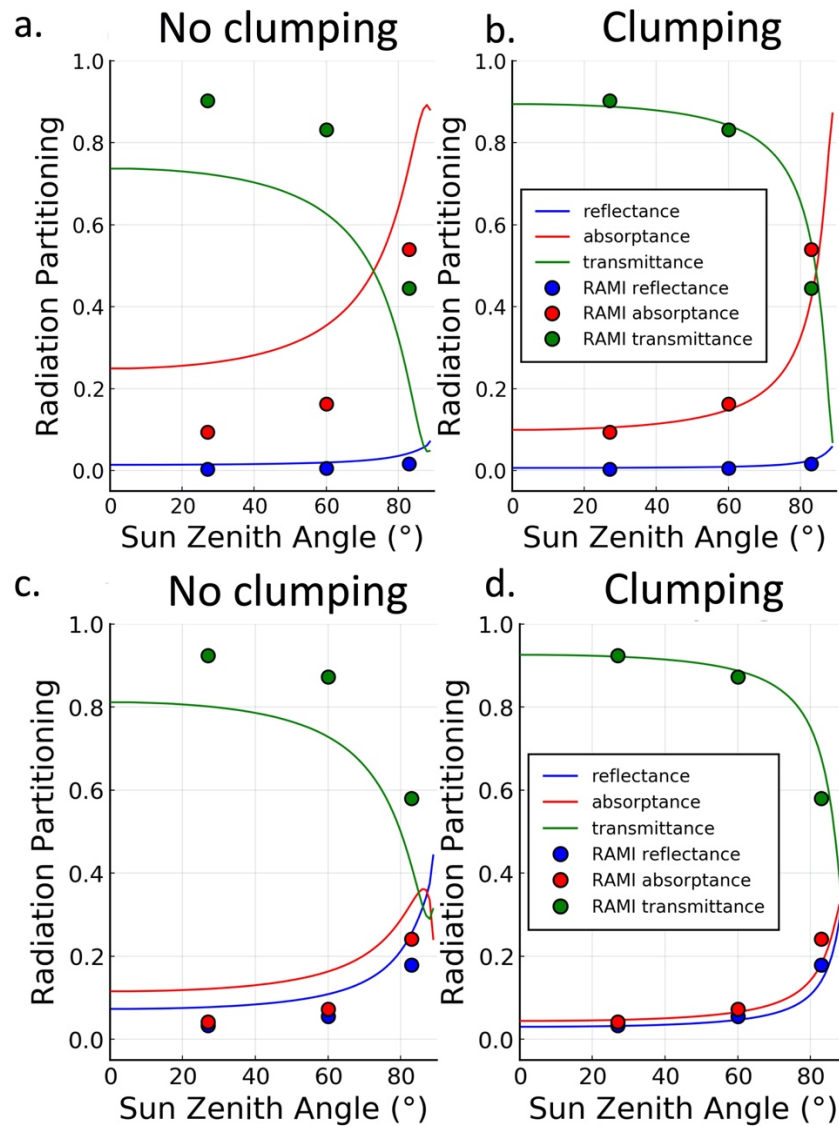
504 **Figure 5** shows a total of 27 cases (3 canopy densities, 3 soil albedos, and 3 sun zenith  
505 angles) for two separate wavebands (PAR and NIR) evaluated separately for reflectance,  
506 absorptance, and transmittance. For the PAR and NIR wavebands, the addition of canopy clumping  
507 improved the agreement between CLIMA-Land and the RAMI4PILPS reference values for all terms  
508 of the radiation partitioning.

509 In the PAR waveband, accounting for clumping index significantly improves the model  
510 predictive skill, as RMSE dropped from 0.12 to 0.03 for reflectance, from 0.21 to 0.06 for  
511 absorptance, and from 0.22 to 0.06 for transmittance. The addition of clumping improved the  $r^2$  for  
512 all terms of the radiative partitioning to  $r^2 > 0.97$ . The 1D case underestimates reflectance and  
513 transmittance, while overestimates absorptance over all the evaluated cases.

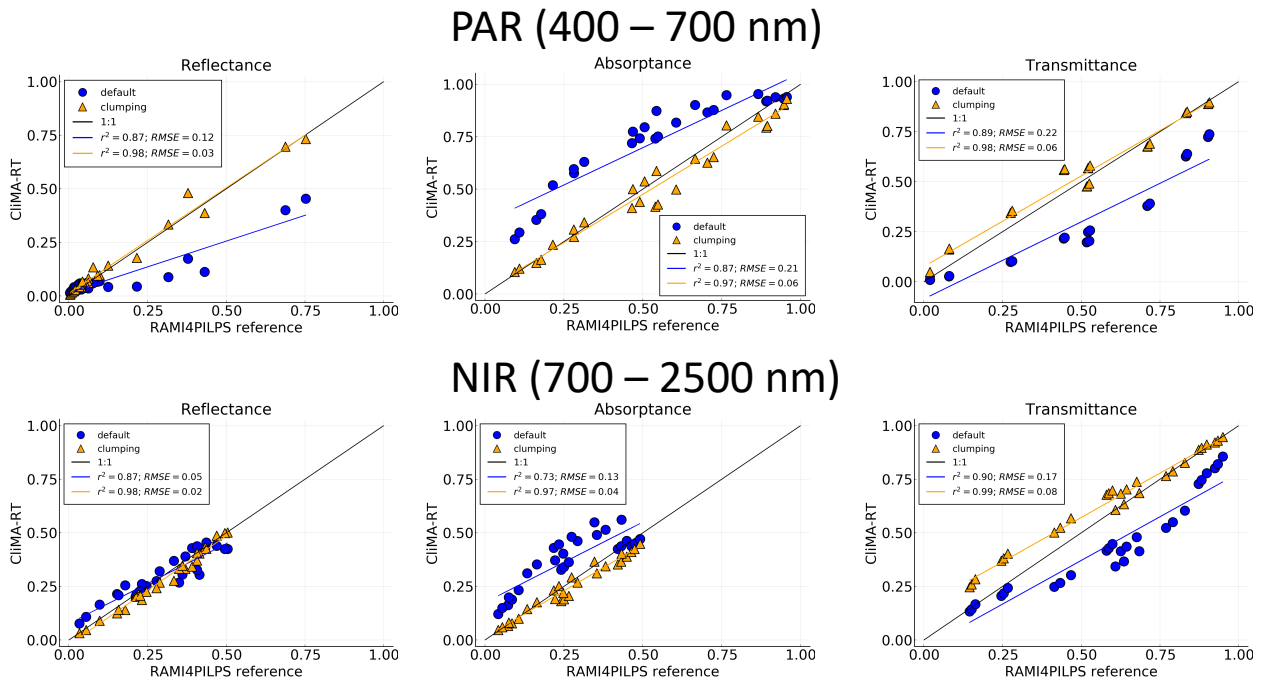
514 In the NIR spectral region, the addition of clumping significantly improves the  $r^2$  for all  
515 terms of the radiative partitioning: from  $r^2 = 0.87$  to  $r^2 = 0.98$  for reflectance; from  $r^2 = 0.73$  to  $r^2 =$   
516 0.97 for absorptance, and for transmittance from  $r^2 = 0.90$  to  $r^2 = 0.99$ . The clumping index  
517 parameterization scheme has decreased the RMSE for reflectance (from RMSE = 0.05 to RMSE =

518 0.02), for absorptance (from RMSE = 0.13 to RMSE = 0.04), and for transmittance (from RMSE  
519 = 0.17 to RMSE = 0.08).

520 These results indicate that the addition of clumping improves the agreement between the  
521 1D and the 3D cases for all terms of the radiation partitioning for both spectral regions.



522  
523 **Figure 4.** Intercomparison of zenith profile of the fraction of direct absorbed (red), reflected (blue),  
524 and transmitted (green) (a-b) PAR (400-700 nm) and (c-d) NIR (700-2500 nm) calculated with 2  
525 different model setups with (clumping) and without clumping (no clumping), and the  
526 RAMI4PILPS reference values obtained with a 3D Monte Carlo ray-tracing model, raytran.



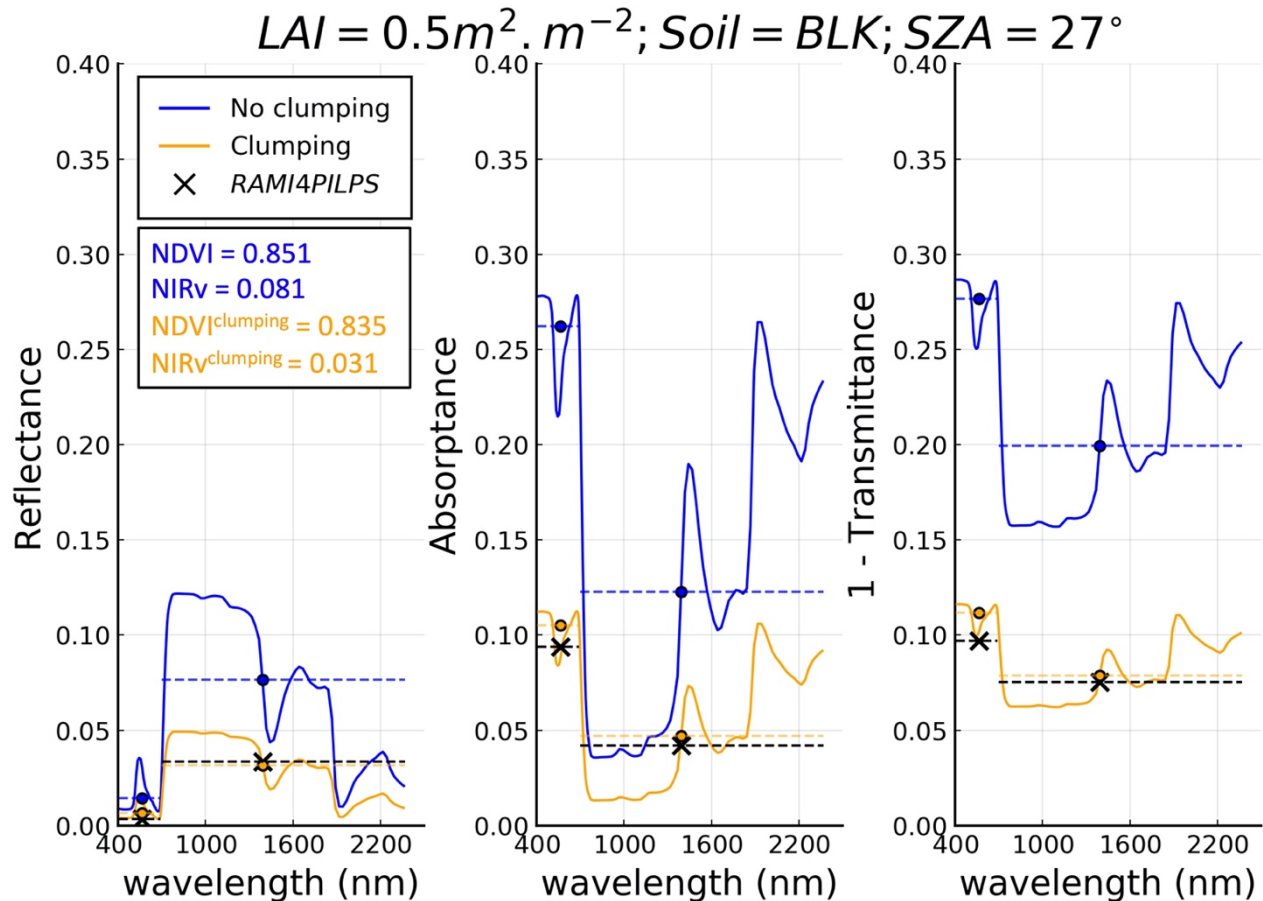
527  
 528 **Figure 5.** Intercomparison of reflected, absorbed, and transmitted PAR (400-700 nm) and NIR  
 529 (700-2500 nm) for 3 canopy densities, 3 soil albedos, and 3 sun zenith angles calculated with 2  
 530 different model setups with clumping (orange) and without clumping (blue) (1D) and the  
 531 RAMI4PILPS reference values (3D) obtained with a 3D Monte Carlo ray-tracing model, raytran.

532  
 533 **3.2 Validating canopy radiative partitioning: hyperspectral shortwave radiation**

534 The three hyperspectral components of the radiation partitioning were compared to the  
 535 RAMI4PILPS reference values. **Figure 6** shows one example of the three components of the  
 536 hyperspectral radiation partitioning (lines) using the default case (no clumping) and the modified  
 537 version with clumping. The average values for PAR and NIR are shown as circles and the  
 538 respective RAMI4PILPS reference values are shown as crosses. **Figure 6** shows the sparse canopy  
 539 case with LAI = 0.5 m<sup>2</sup>.m<sup>-2</sup> and 9% vegetation cover over a black soil ( $\alpha_{\text{soil}} = 0.0$ ) for a sun zenith  
 540 angle of 27°. For similar figures for all the other canopy structures and zenith angles, see

541 **Supplemental material.** The hyperspectral cases were only evaluated over a black soil albedo due  
542 to complexities involved in scaling up soil albedos in the presence of snow. Polar plots showing  
543 the difference in Far-Red SIF, NDVI, and NIRv between the clumped and non-clumped cases can  
544 be found in **Supplemental material**.

545



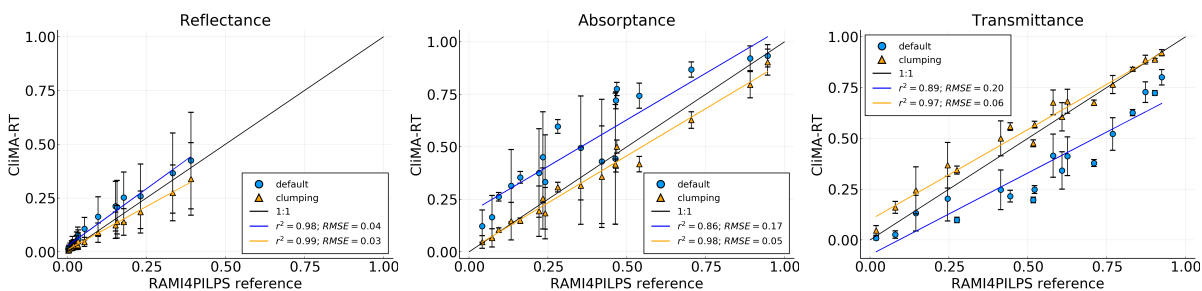
546  
547 **Figure 6.** Intercomparison of reflected, absorbed, and transmitted hyperspectral shortwave  
548 radiation (400-2500 nm) for a sparse case ( $LAI = 0.50 m^2 \cdot m^{-2}$  and 9% vegetation cover), over black  
549 soil, with sun zenith angle =  $27^\circ$  calculated with 2 different model setups with clumping (orange)  
550 and without clumping (blue) (1D). The RAMI4PILPS reference values (3D) obtained with a 3D  
551 Monte Carlo ray-tracing model, raytran (black crosses represent the average PAR and NIR,  
552 separately). The average values for PAR and NIR are shown as points and horizontal dashed lines

553 for clumping (orange) and no clumping (blue). The values of NDVI and NIRv, with and without  
554 clumping, are also indicated.

555  
556 **Figure 7** shows a total of 18 cases (3 canopy densities, 3 sun zenith angles, and two spectral  
557 regions) for reflectance, absorptance, and transmittance. The addition of canopy clumping  
558 improved the agreement between CliMA-Land and the RAMI4PILPS reference values for all terms  
559 of the radiation partitioning.

560 For reflectance, the RMSE between CliMA-Land and the RAMI4PILPS reference values  
561 dropped from 0.04 to 0.03 when clumping was considered. For absorptance, the RMSE between  
562 CliMA-Land and the RAMI4PILPS reference values dropped from 0.17 to 0.05 when clumping  
563 was considered. For transmittance, the RMSE between CliMA-Land and the RAMI4PILPS  
564 reference values dropped from 0.20 to 0.06 when clumping was considered. The 1D case  
565 overestimates reflectance and absorptance, while underestimates transmittance over all the  
566 evaluated cases. The addition of clumping has also improved the  $r^2$  for all terms of the radiative  
567 partitioning (from  $r^2 = 0.98$  to  $r^2 = 0.99$  for reflectance; from  $r^2 = 0.86$  to  $r^2 = 0.98$  for absorptance;  
568 and from  $r^2 = 0.89$  to  $r^2 = 0.97$  for transmittance). These results indicate that clumping has improved  
569 the agreement between the 1D and the 3D cases throughout all wavelengths in the shortwave  
570 radiation spectrum from 400 to 2500 nm.

571



572



573 **Figure 7.** Intercomparison of reflected, absorbed, and transmitted averaged in the PAR (400-700  
574 nm) and NIR (700-2500 nm) wavebands for 3 canopy densities, 3 sun zenith angles, and a black  
575 soil albedo calculated with 2 different model setups with clumping (orange) and without clumping  
576 (blue) (1D). The RAMI4PILPS reference values (3D) were obtained with a 3D Monte Carlo ray-  
577 tracing model, raytran. The vertical black bars indicate the standard deviation of the mean values  
578 for each waveband considered in 10 nm spectral resolution.

579

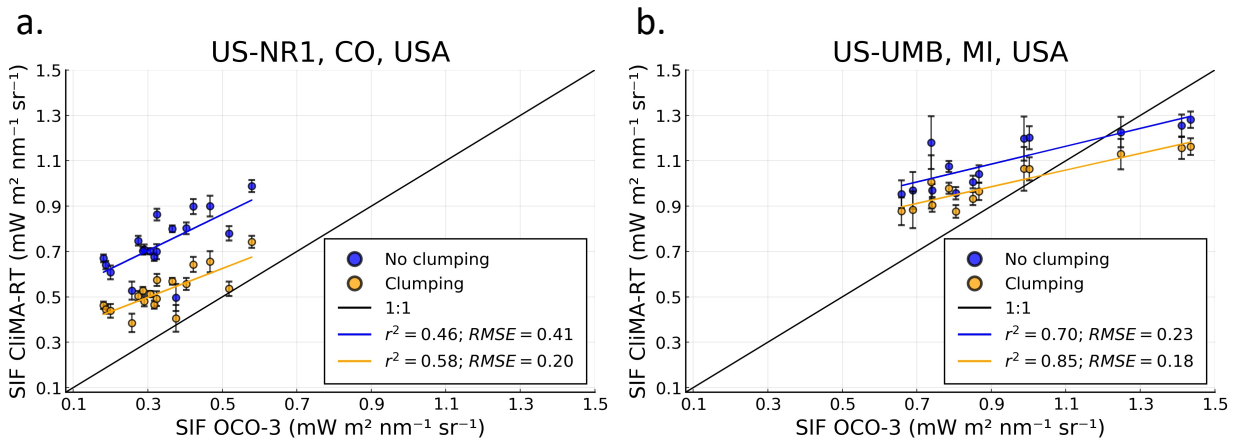
### 580 **3.3 Validating SIF emission with OCO-3 observations**

581 In order to estimate the effect of the clumping index on model estimates of SIF from  
582 CliMA-Land radiative transfer, we also compared simulated SIF computed with and without the  
583 clumping index to canopy-scale remote sensing SIF retrievals from OCO-3 on board of the ISS, at  
584 Niwot Ridge, Colorado and UMB Station, Michigan, USA.

585 **Figure 8** shows a scatter plot of Far-Red SIF (at 757 nm) from CliMA-Land radiative  
586 transfer (with clumping in yellow and without clumping in blue) versus Far-Red SIF derived from  
587 OCO-3 for both sites in 2020. The individual points in the linear fit represent the whole scan area  
588 shown in **Figure 3**. For each OCO-3 overpass, there are several scans for the SAMs. Basically,  
589 each scan has very similar sun-sensor geometry and the soundings can be grouped based on phase  
590 angle. Each point in **Figure 8** represents the mean of all the soundings with approximately the same  
591 phase angle in order to reduce the error associated with sensor geometry.

592 The estimates of Far-Red SIF from CliMA-Land radiative transfer with clumping index  
593 indicate an improvement with observations. The linear fit between model and observations shows  
594 a higher  $r^2$  (0.58 for Niwot Ridge and 0.85 for UMB Station) and a lower RMSE (0.20 for Niwot  
595 Ridge and 0.18 for UMB Station) when considering canopy structure with a clumping index, versus

596 the original version of the model (default) without clumping index. The reduction of 51.2% in  
597 RMSE over Niwot Ridge and 21.7% over UMB Station when considering canopy structure through  
598 clumping highlights the importance of considering canopy structure when deriving SIF products  
599 from remote sensing.  
600



601  
602 **Figure 8.** Intercomparison of SIF (757 nm) between CiMA-Land radiative transfer (with clumping  
603 in yellow and without clumping in blue) and two SAMs that were taken by OCO-3 at **a.** Niwot  
604 Ridge, Colorado, USA obtained on June 12<sup>th</sup> and June 16<sup>th</sup>, 2020, and **b.** UMB Station, Michigan,  
605 USA obtained on August 06<sup>th</sup> and August 11<sup>th</sup>, 2020. The  $r^2$  and RMSE of the linear fits are also  
606 shown. Each point represents the mean of all the soundings with approximately the same phase  
607 angle in order to reduce the error associated with sensor geometry, represented by the error bars.  
608

### 609 3.4 The impact of canopy clumping on vertical APAR, fAPAR, and NIRv

610 The radiation partitioning from the CiMA-Land radiative transfer model has been validated  
611 against a detailed model benchmarking, as well as the SIF estimates from the model have been  
612 tested against SIF observation from satellite remote sensing data. In both cases, results indicate that  
613 whenever the clumping index parameterization scheme is considered when modeling the transfer

614 of radiation, the agreement between both model and highly accurate 3D radiative transfer models  
615 and model and satellite observations is higher (RMSE ~50% smaller).

616 To further evaluate the impacts of canopy structure on the carbon and water cycles, the  
617 impacts of clumping on vertical fAPAR and APAR should be tested because these variables drive  
618 the light limiting regime of photosynthesis in ESMs. **Figure 9** shows the vertical zenith profile of  
619 the difference in APAR between the modified CliMA-Land radiative transfer with clumping index  
620 minus the default version (without clumping) for 3 canopy densities (0.5, 1.5, and 2.5 m<sup>2</sup>.m<sup>-2</sup>) over  
621 3 soil albedos (BLK, MED, SNW). The CliMA-Land version without clumping is equivalent to  
622 the mSCOPE, and so, the validation with the mSCOPE model is indirectly present in all  
623 evaluations.

624 Throughout all the evaluated scenarios, APAR increases when clumping is considered, with  
625 a stronger difference towards the bottom of the evaluated canopy. This result is not straightforward,  
626 because the vertical fAPAR does not follow the same behavior as the vertical APAR (see  
627 **Supplemental material**). While the clumping index acts to decrease the total optical depth of the  
628 vegetation canopy, fAPAR decreases at the top of the canopy and increases at the bottom. The  
629 effect of soil albedo is mostly noted when the value of soil albedo is high (i.e., over SNW with  
630  $\alpha_{\text{soil,PAR}} = 0.96$ ), and the zenith angle of incident radiation is small ( $\text{SZA} = 27^\circ$ ), because at nadir  
631 the optical path length is the shortest. For the sparse canopy, the clumping index reduces the total  
632 fAPAR in approximately half of the one obtained by the default CliMA-Land radiative transfer,  
633 and the distribution of fAPAR throughout the vertical canopy is homogenous. Over a bright soil,  
634 the fAPAR at the bottom of the canopy is relatively larger than at the top because of the scattering  
635 effects from the background soil underneath. This effect has also been shown by Pinty et al. (2006)  
636 and Braghieri (2018), whose work reaffirms that for low vegetation densities, fAPAR is rather  
637 small and so the differences between the 1D canopy and the 3D canopy remain limited over a

638 darker soil. For the medium and dense canopies, the clumping index affects the vertical profile of  
639 fAPAR in two primary ways: i) it reduces the total amount of PAR absorption at the top layers,  
640 and; ii) it increases fAPAR at the bottom of the canopy, especially over brighter soils. Over a bright  
641 soil, fAPAR at the bottom of the canopy is more than twice as large as the one calculated by the  
642 default version of the model for the dense canopy, and about one and a half times larger than for  
643 the medium canopy. This effect is observed throughout all sun zenith angles, with an increase  
644 towards larger zenith angles.

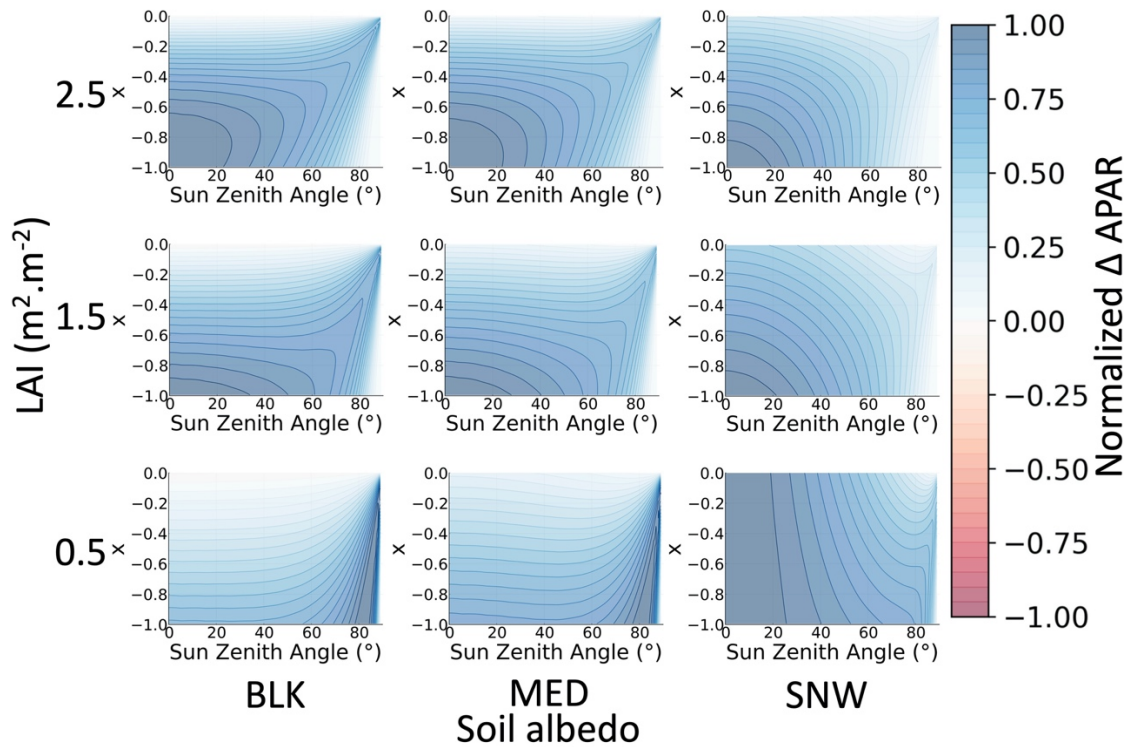
645         However, it is expected that although fAPAR decreases in most cases, APAR increases  
646 throughout all the evaluated scenes and sun zenith angles because more light penetrates the canopy  
647 and, therefore, there is more available energy to be absorbed. For this reason, it is important to  
648 evaluate the impacts on fAPAR together with a change in the incident radiation in the top layers of  
649 the canopy. In order to keep consistency with reality for the evaluations of vertical APAR, the  
650 value of incident PAR at the top of the canopy was modulated following the cosine of the sun  
651 zenith angle.

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656  
657 **Figure 9.** Vertical zenith profile of normalized APAR difference between the modified CliMA-  
658 Land radiative transfer with clumping index minus the non-clumping version for 3 canopy densities  
659 ( $0.5$ ,  $1.5$ , and  $2.5 \text{ m}^2 \cdot \text{m}^{-2}$ ) over 3 soil albedos (BLK, MED, SNW).  $x$  is the relative optical height,  
660 which runs from  $-1$  at the bottom to zero at the top of the canopy.

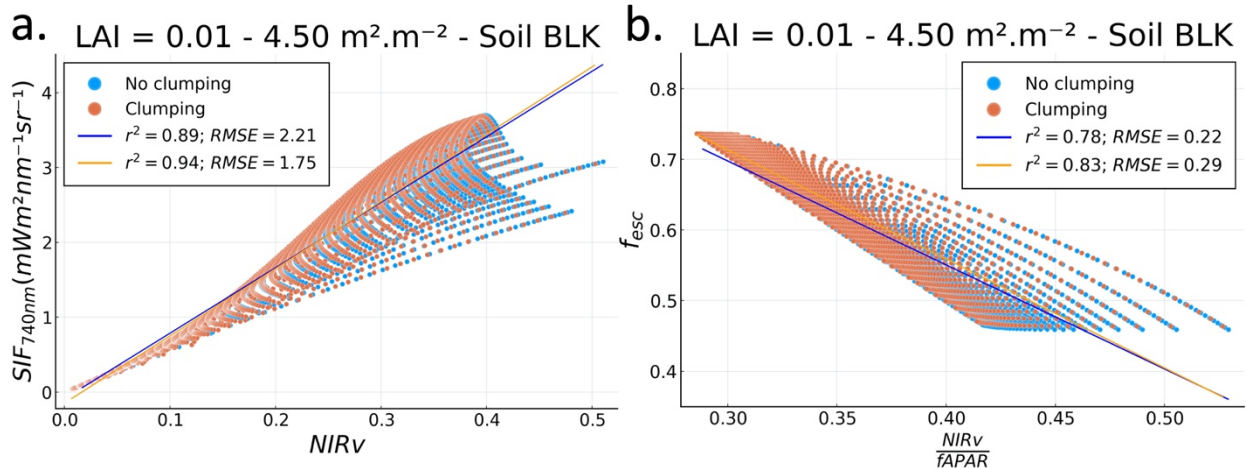
661  
662 To evaluate the impacts of canopy clumping on the relationships between  $\text{NIR}_v$  and  $\text{SIF}_{740\text{nm}}$   
663 described in Badgley et al. (2017), as well as on the relationship between  $f_{\text{esc}}$  and  $\text{NIR}_v \cdot f\text{APAR}^{-1}$   
664 as described in Zeng et al. (2019), **Eq.(2)** was used to recreate multiple canopy densities with  
665 different cover fractions, representing a structurally diverse vegetation canopy with LAI varying  
666 from  $0.01 \text{ m}^2 \cdot \text{m}^{-2}$  to  $\text{LAI} = 4.50 \text{ m}^2 \cdot \text{m}^{-2}$ , and vegetation cover fraction calculated as  $\text{LAI} \text{ over } 5$ . All  
667 scenes were simulated over all possible sun zenith angles with background soil albedo set to black  
668 (BLK;  $0.0$ ).

669 **Figure 10a.** shows the linear fit between calculated  $SIF_{740nm}$  versus  $NIR_v$  for the modified  
670 CliMA-Land radiative transfer with clumping index (in yellow) and the default version (in blue)  
671 for multiple canopy densities. The consideration of canopy clumping improves the relationship  
672 between estimated SIF and  $NIR_v$  from the CliMA-Land radiative transfer model, with an increase  
673 in  $r^2$  from 0.89 to 0.94, and a decrease in RMSE from  $2.21 \text{ mWm}^2\text{nm}^{-1}\text{sr}^{-1}$  to  $1.75 \text{ mWm}^2\text{nm}^{-1}\text{sr}^{-1}$ .  
674 While **Figure 10b.** shows the linear fit between the fluorescence escape ratio ( $f_{esc}$ ) and the  
675  $NIR_v \cdot fAPAR^{-1}$  for the modified CliMA-Land radiative transfer with clumping index and the default  
676 version for multiple canopy densities (from  $LAI = 0.01 \text{ m}^2 \cdot \text{m}^{-2}$  to  $LAI = 4.50 \text{ m}^2 \cdot \text{m}^{-2}$ ) over a black  
677 soil albedo (BLK) with clumping calculated through **Eq.(2)** for sun zenith angles from  $0^\circ$  to  $30^\circ$ .  
678 For similar figures over medium (MED) and snowy (SNW) soil albedos, see **Supplemental**  
679 **material.** The linear fit improves when canopy clumping is considered with an increased  $r^2$  values  
680 from 0.78 to 0.83. While, the RMSE value decreased for the linear relationship when the clumping  
681 index was considered, the relationship described in Zeng et al. (2019) does not refer to an absolute  
682 equal equation, but rather to an approximation of  $f_{esc}$  and  $NIR_v \cdot fAPAR^{-1}$ , and so, the absolute values  
683 should not be strictly considered.

684 In **Figure 10b.**, the linear fit of the CliMA-Land radiative transfer without clumping index  
685 has  $AIC = -4923.44$  and  $BIC = -4907.90$ , while the version with clumping index has  $AIC = -$   
686  $5291.47$  and the  $BIC = -5275.94$ . The AIC and BIC values indicate a stronger relationship between  
687  $f_{esc}$  and  $NIR_v \cdot fAPAR^{-1}$ , as proposed by Zeng et al. (2019), when canopy structure is considered.

688

689



690  
 691 **Figure 10. a.** Linear fit between SIF<sub>740nm</sub> and NIR<sub>v</sub> for the modified CliMA-Land radiative transfer  
 692 with clumping index (yellow) and the default version (blue) for multiple canopy densities (from  
 693 LAI = 0.01 m<sup>2</sup>.m<sup>-2</sup> to LAI = 4.50 m<sup>2</sup>.m<sup>-2</sup>) over a black soil albedo (BLK) with clumping calculated  
 694 through **Eq.(2)** for sun zenith angles from 0° to 89°, and; **b.** linear fit between the fluorescence  
 695 escape ratio ( $f_{esc}$ ) and the NIR<sub>v</sub>/fAPAR for the modified CliMA-Land radiative transfer with  
 696 clumping index and the default version for multiple canopy densities and over a black soil albedo  
 697 (BLK) as in **Fig.10a.** with clumping calculated through **Eq.(2)** for sun zenith angles from 0° to 30°.  
 698 For CliMA-Land radiative transfer without clumping index the AIC = - 4923.44 and the BIC = -  
 699 4907.90, while for CliMA-Land radiative transfer with clumping index the AIC = - 5291.47 and  
 700 the BIC = - 5275.94.

701  
 702 **4.0 Discussion**

703 In this study, we implemented and evaluated a parameterization of horizontal vegetation  
 704 structure on the radiative transfer scheme of a new generation ESM, the CliMA model. We  
 705 benchmarked the radiation partitioning of CliMA-Land radiative transfer with results from a 3D  
 706 Monte-Carlo ray tracer previously presented in Widlowski et al. (2011). In each of the evaluated

707 scenarios, all terms of the radiation partitioning (reflectance, absorptance, transmittance) from the  
708 model version that included the effects of canopy structure showed a better agreement with the  
709 accurate 3D modeling, indicating the importance of considering not only the vertical heterogeneity  
710 of vegetation canopies, but also the horizontal effects of canopy structure. The improvement for  
711 reflectance was smaller than the ones for absorptance and transmittance partly due to the fact that  
712 reflectance values are the smallest terms of the radiation partitioning for the evaluated cases.

713 The main difference between the present study and previous ones is the hyperspectral nature of  
714 the radiative transfer model combined with horizontal canopy structural heterogeneity in CliMA-  
715 Land. By using a single value of clumping index following the work of Pinty et al. (2006), we were  
716 able to account for the effects of vegetation structure on the transfer of radiation across all  
717 wavelengths of the shortwave radiation spectrum with 10 nm spectral resolution. The results  
718 presented here highlight the capability of the new CliMA-Land model to be directly compared with  
719 observed canopy spectroscopy from high resolution spectral data currently available from aircrafts  
720 preparing Earth system modelers for a suite of global hyperspectral measurements that soon will  
721 be available from the US SBG concept (Schimel and Schneider, 2019).

722 We also presented a validation exercise with observations of SIF emission over an evergreen  
723 needleleaf site and a deciduous broadleaf site in the USA from remote sensing with the recently  
724 launched OCO-3 sensor on board of the ISS at spatial resolution of not more than 4 km, including  
725 the footprint of two flux tower sites (US-NR1 and US-UMB), in order to facilitate further  
726 evaluation and comparison to FLUXNET data (Baldocchi et al., 2001). Combining SIF from OCO-  
727 3 with a suite of remote sensing products, including Copernicus LAI (Fuster et al., 2020) at 300 m  
728 spatial resolution, a chlorophyll product from ENVISAT MERIS (Croft et al., 2020), and clumping  
729 index from MODIS (He et al., 2012), we were able to determine a substantial improvement on  
730 modelled SIF when vegetation canopy structure was considered. The importance of directly



731 modeling SIF with an ESM is related to the SIF-GPP relationships required for remote large-scale  
732 estimations of GPP (Ryu et al., 2019; Dechant et al., 2020), as well as the direct assimilation of  
733 SIF data to improve GPP predictions (Norton et al., 2019; Parazoo et al., 2020), which are currently  
734 highly uncertain globally (Braghiere et al., 2019) (see **Supplemental material** for a model  
735 intercomparison with other SIF-enabled LSMs). SIF 740nm estimates from CliMA-Land are  
736 comparable to those of BETHY, while the impact of clumping decreases the total SIF signal. In  
737 the comparison with SCOPE, CliMA-Land slightly underestimates the SIF peak.

738       After thorough validation with accurate 3D modeling and observations, we evaluated the  
739 impact of the clumping index parameterization scheme on proxies of GPP, i.e., vertical APAR, in  
740 order to characterize further impacts on GPP from CliMA-Land when absorbed radiation will be  
741 used to derived photosynthesis through the Farquhar-von Caemmerer-Berry model (Farquhar et  
742 al., 1980). Contrary to expectation, considering horizontal canopy structure through the addition of  
743 clumping on the radiative transfer scheme of CliMA-Land caused fAPAR to vary largely across  
744 different canopy densities, illumination angles, and soil backgrounds albedos, but with one single  
745 impact on the total APAR across the vertical canopy. Throughout all the evaluated scenes, APAR  
746 increased when canopy structure is considered, especially in the bottom layers of the vegetation  
747 canopy. This can be thought of as a reduction on the total optical depth of the canopy and, therefore,  
748 less plant material for the radiation to interact with along its pathway to the ground and back up  
749 after interacting with the surface underneath. These results are in alignment with previous studies  
750 that evaluated the impact of the clumping index on radiative transfer schemes in land surface  
751 models (Braghiere et al., 2020, 2019; Loew et al., 2014).

752       The CliMA-Land model can simulate photosynthesis. However, photosynthesis is a process  
753 that includes many more different sub-models, e.g., the Farquhar ecophysiology model (Farquhar  
754 et al., 1980), model of root development, model of water distribution in soils and plants. Therefore,

755 the current study is limited to the evaluation of the radiative transfer scheme in CliMA-Land, in  
756 order to keep consistency and conciseness without completely leaving photosynthesis behind  
757 through the evaluation of the impact of clumping on vegetation indices. Nevertheless, further  
758 evaluation on CliMA-Land photosynthesis is required.

759 Finally, we tested two relationships that were described in the literature as strongly influenced  
760 by canopy structure and that our new model allowed us to explore. The first one is the relationship  
761 between observed SIF and  $NIR_v$  proposed by Badgley et al. (2017) and further evaluated in a  
762 number of studies (Badgley et al., 2019; Dechant et al., 2020). Here we showed an improved linear  
763 fit between  $NIR_v$  and SIF when considering canopy structure when calculating the transfer  
764 radiation with a reduction of 20% on RMSE. This result reinforces previous evidence relating the  
765 effect of canopy structure, represented by  $f_{esc}$ , on SIF emission, APAR, and GPP using modelling  
766 and observations (Dechant et al., 2020; Du et al., 2017; Migliavacca et al., 2017).

767 The impacts of canopy clumping were also evaluated on the relationship demonstrated by Zeng  
768 et al. (2019) and described in **Eq.(5)** where  $f_{esc}$  can be approximated by  $NIR_v \cdot fAPAR^{-1}$ . Zeng et al.  
769 (2019) showed that  $f_{esc}$  can be derived from  $NIR_v$  properly even over sparsely vegetated areas with  
770 minimal effects from background soil albedo. In here, we showed an improved linear fit in **Figure**  
771 **10b** when considering clumping index in CliMA-Land radiative transfer, which highlights the  
772 important effect that horizontal canopy heterogeneity can have on the appropriate usage of **Eq.(5)**.

773

#### 774 **4.1 Data uncertainties and model limitations**

775 The non-linearity of clumping index spatial scaling at the landscape level has been  
776 previously explored using LAI-2000 and digital hemispherical photography datasets (Ryu et al.,  
777 2010a). In our study, the clumping index and LAI values were linearly scaled up as area weighted

778 averages for the OCO-3 SIF validation experiment (<4 km vs. 500 m), which may introduce biases  
779 in our results, mainly due to changes in vegetation heterogeneity with spatial scale. The linear  
780 averaging method in this particular case was preferred due to: (i) the absence of high-resolution  
781 gap fraction and clumping index measurements; and, (ii) the fairly homogeneous clumping index  
782 values in the evaluated area (see **Supplemental material**). In addition, the MODIS clumping index  
783 was retrieved using the Normalized Difference between Hotspot and Darkspot (NDHD) algorithm  
784 (Chen et al., 2005) and validated with *in-situ* measurements over a set of 63 globally distributed  
785 LPV (Land Product Validation) and VALERI (VALidation of Land European Remote sensing  
786 Instruments) sites (Baret et al., 2006; Garrigues et al., 2008; Nightingale et al., 2011; Pisek et al.,  
787 2015b), as well as intercompared with higher resolution (275 m) data from the Multi-angle Imaging  
788 SpectroRadiometer (MISR) satellite (Pisek et al., 2013), showing a particularly good agreement  
789 over needleleaf forests, with MODIS showing a wider range of clumping index values (0.47–0.72  
790 compared to MISR 0.52–0.59) (Pisek et al., 2015b).

791 Further intercomparison between MISR, MODIS, and POLDER clumping index datasets  
792 (Pisek et al., 2010) highlighted the importance of appropriately scaling up the clumping index  
793 values in order to match the scale of the application. For instance, if POLDER clumping index (~6  
794 km resolution) was to be used with our model, an alternative scaling methodology would be  
795 preferred in order to avoid the addition of significant biases due to the usage of coarser resolution  
796 data. Likewise, if an evaluation was to be performed using OCO-3 SIF grouped into larger areas  
797 (e.g., 0.5 degree as current ESMs), a non-linear averaging method would be indicated in order to  
798 limit uncertainty (Ryu et al., 2010a). In future validation studies of CliMA-Land at site level with  
799 scanning spectrometers, e.g., PhotoSpec (Grossmann et al., 2018), clumping index values should  
800 be derived at much finer spatial scales (<1m), taking into account clumping index variations with  
801 canopy height and view zenith/azimuth angles accordingly.

## 802 **5.0 Conclusion**

803           Our work suggests that considering vertical and horizontal vegetation canopy structure  
804 through the addition of a clumping index parameterization scheme may significantly improve the  
805 hyperspectral shortwave radiation partitioning of an ESM without losing efficiency, with a RMSE  
806 reduction on the order of 25% for reflectance, 66% for absorptance, and 75% for transmittance in  
807 comparison to a highly accurate Monte Carlo 3D radiative transfer model. The dominant effect that  
808 introducing clumping has in our study is to allow more shortwave radiation to propagate further  
809 into lower canopy levels increasing APAR levels throughout the vertical canopy and across sun  
810 zenith angles.

811           We also compared SIF emissions against observed data with a satellite spectrometer. The  
812 results presented here strongly support previous evidence that horizontal canopy structural features  
813 are crucial for an accurate estimation of SIF, as do further extrapolations that might come out from  
814 this variable, such as global photosynthesis. The improvement of SIF estimates with a clumping  
815 index indicates that the clumping index can capture the horizontal canopy structural features at  
816 remote sensing scales (<4 km).

817           Finally, we showed how the clumping index parameterization scheme improved the SIF  
818 correlation to  $\text{NIR}_v$ , as well the correlation of  $f_{\text{esc}}$  with  $f\text{APAR}$ , which provides further evidence for  
819 the role of vertical and horizontal canopy structure on SIF emission and the appropriate  
820 determination of other vegetation indices.

## 821 **Appendix A. Appendix**

### 822 **A.1. Calculating escape and recollision probabilities**

823 This Appendix has additional information on the calculation of the escape and recollision  
824 probabilities. For the complete set of equations, see Huang et al. (2007) and Smolander and

825 Stenberg (2005). First, the canopy interceptance ( $i$ ) refers to the probability of an incoming photon  
826 interacting with the vegetation canopy, and it can be approximated by  $1 - P_{\text{gap}}$ , where  $P_{\text{gap}}$  is the  
827 direct transmittance. Second, the recollision probability ( $p$ ) refers to the probability that a photon  
828 recollides with elements of the canopy at an  $n$ -th plus one time, on its  $n$ -th interaction with the  
829 canopy, and it can be obtained by rearranging equation 2 presented in Smolander and Stenberg  
830 (2005) as:

$$831 \quad p_{s/d} = \frac{1 - \left( \frac{1 - \omega_l}{fAPAR_{s/d}} \right) \times (1 - P_{\text{gap}})}{\omega_l} \quad (\text{A1})$$

832 where  $fAPAR$  is the fraction of absorbed PAR,  $P_{\text{gap}}$  is the direct transmittance, and  $\omega_l$  is the single  
833 scattering albedo. Finally, the escape probability ( $\rho$ ) refers to the probability of a photon escaping  
834 the vegetation canopy after interacting with elements of vegetation, and it can be obtained by  
835 rearranging equation 9 presented in Huang et al. (2007) as:

$$836 \quad \rho_{s/d} = \frac{R_{s/d}}{\omega_l \times (1 - P_{\text{gap}}) + \frac{\omega_l^2 \times p_{s/d} \times (1 - P_{\text{gap}})}{1 - p_{s/d} \times \omega_l}} \quad (\text{A2})$$

837 where  $R$  is the canopy albedo,  $P_{\text{gap}}$  is the direct transmittance,  $\omega_l$  is the single scattering albedo, and  
838  $p$  is the recollision probability.

### 839 **Data availability**

840 The CliMA project, code, simulation configurations, model output, and tools to work with the  
841 output are described at <https://github.com/CliMA>. The land model and examples are available at  
842 <https://github.com/CliMA/Land>. The minimization of hyperspectral leaf reflectance and  
843 transmittance was performed using a Julia package available at [https://github.com/Yujie-](https://github.com/Yujie-W/ConstrainedRootSolvers.jl)  
844 [W/ConstrainedRootSolvers.jl](https://github.com/Yujie-W/ConstrainedRootSolvers.jl). The LAI map, PROBA-V LAI V2, was produced by Copernicus at  
845 1 km resolution and it is available at <https://land.copernicus.eu/global/products/lai>. National

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847 mosaic. Provisional data downloaded from <https://data.neonscience.org> on November 30, 2020.  
848 Battelle, Boulder, CO, USA NEON. 2020.

#### 849 **Description of author's responsibilities**

850 RKB: conceptualization, methodology, formal analysis, writing - original draft, writing - review &  
851 editing, implementation of clumping index, research; RKB and YW: spectral properties fitting  
852 package, model coding, and writing; RKB and RD: OCO-3 SIF methodology and writing; DS: soil  
853 spectral properties for Niwot Ridge; TM and JLW: validation datasets and editing; ML, AB, JW,  
854 PG: editing and model conceptualization; CF: model conceptualization, coding, review, and  
855 editing.

#### 856 **Declaration of competing interests**

857 The authors declare no competing interests.

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1369 **List of Figure Captions**

1370  
1371 **Figure 1.** Graphical representation of the open forest canopy environments used in the  
1372 RAMI4PILPS experiment. Three different leaf area index (LAI) values and three different  
1373 background soil albedos (adapted from Widlowski et al. (2011)).

1374  
1375 **Figure 2a.** Hyperspectral leaf reflectance (blue) and leaf transmittance (red) obtained from  
1376 Fluspect using values given in **Table 2**; **b.** The average values of these curves are represented by  
1377 circles for two broadbands and single scattering albedo term, separately, i.e., PAR (400-700 nm)  
1378 and NIR (700-2500 nm); reflectance ( $\rho$ ) and transmittance ( $\tau$ ).

1379  
1380 **Figure 3.** OCO-3 retrieved SIF at 757 nm over **a.** Niwot Ridge, Colorado, USA on June, 12<sup>th</sup>, 2020,  
1381 and **c.** UMB Station, Michigan, USA on August 11<sup>th</sup>, 2020. MODIS derived clumping index map  
1382 from He et al. (2012) over **b.** Niwot Ridge, Colorado, USA and **d.** UMB Station, Michigan, USA,  
1383 for the year of 2006 matching the OCO-3 scan. The white circle with a black dot in the middle  
1384 represents the position of the flux towers for reference.

1385  
1386 **Figure 4.** Intercomparison of zenith profile of the fraction of direct absorbed (red), reflected (blue),  
1387 and transmitted (green) (**a-b**) PAR (400-700 nm) and (**c-d**) NIR (700-2500 nm) calculated with 2  
1388 different model setups with (clumping) and without clumping (no clumping), and the  
1389 RAMI4PILPS reference values obtained with a 3D Monte Carlo ray-tracing model, raytran.

1390  
1391 **Figure 5.** Intercomparison of reflected, absorbed, and transmitted PAR (400-700 nm) and NIR  
1392 (700-2500 nm) for 3 canopy densities, 3 soil albedos, and 3 sun zenith angles calculated with 2

1393 different model setups with clumping (orange) and without clumping (blue) (1D) and the  
1394 RAMI4PILPS reference values (3D) obtained with a 3D Monte Carlo ray-tracing model, raytran.

1395  
1396 **Figure 6.** Intercomparison of reflected, absorbed, and transmitted hyperspectral shortwave  
1397 radiation (400-2500 nm) for a sparse case ( $LAI = 0.50 \text{ m}^2 \cdot \text{m}^{-2}$  and 9% vegetation cover), over black  
1398 soil, with sun zenith angle =  $27^\circ$  calculated with 2 different model setups with clumping (orange)  
1399 and without clumping (blue) (1D). The RAMI4PILPS reference values (3D) obtained with a 3D  
1400 Monte Carlo ray-tracing model, raytran (black crosses represent the average PAR and NIR,  
1401 separately). The average values for PAR and NIR are shown as points and horizontal dashed lines  
1402 for clumping (orange) and no clumping (blue). The values of NDVI and NIR<sub>v</sub>, with and without  
1403 clumping, are also indicated.

1404  
1405 **Figure 7.** Intercomparison of reflected, absorbed, and transmitted averaged in the PAR (400-700  
1406 nm) and NIR (700-2500 nm) wavebands for 3 canopy densities, 3 sun zenith angles, and a black  
1407 soil albedo calculated with 2 different model setups with clumping (orange) and without clumping  
1408 (blue) (1D). The RAMI4PILPS reference values (3D) were obtained with a 3D Monte Carlo ray-  
1409 tracing model, raytran. The vertical black bars indicate the standard deviation of the mean values  
1410 for each waveband considered in 10 nm spectral resolution.

1411  
1412 **Figure 8.** Intercomparison of SIF (757 nm) between CliMA-Land radiative transfer (with clumping  
1413 in yellow and without clumping in blue) and two SAMs that were taken by OCO-3 at **a.** Niwot  
1414 Ridge, Colorado, USA obtained on June 12<sup>th</sup> and June 16<sup>th</sup>, 2020, and **b.** UMB Station, Michigan,  
1415 USA obtained on August 06<sup>th</sup> and August 11<sup>th</sup>, 2020. The  $r^2$  and RMSE of the linear fits are also

1416 shown. Each point represents the mean of all the soundings with approximately the same phase  
1417 angle in order to reduce the error associated with sensor geometry, represented by the error bars.

1418  
1419 **Figure 9.** Vertical zenith profile of normalized APAR difference between the modified CliMA-  
1420 Land radiative transfer with clumping index minus the non-clumping version for 3 canopy densities  
1421 (0.5, 1.5, and 2.5 m<sup>2</sup>.m<sup>-2</sup>) over 3 soil albedos (BLK, MED, SNW). x is the relative optical height,  
1422 which runs from -1 at the bottom to zero at the top of the canopy.

1423  
1424 **Figure 10. a.** Linear fit between SIF<sub>740nm</sub> and NIR<sub>v</sub> for the modified CliMA-Land radiative transfer  
1425 with clumping index (yellow) and the default version (blue) for multiple canopy densities (from  
1426 LAI = 0.01 m<sup>2</sup>.m<sup>-2</sup> to LAI = 4.50 m<sup>2</sup>.m<sup>-2</sup>) over a black soil albedo (BLK) with clumping calculated  
1427 through **Eq.(2)** for sun zenith angles from 0° to 89°, and; **b.** linear fit between the fluorescence  
1428 escape ratio (f<sub>esc</sub>) and the NIR<sub>v</sub>/fAPAR for the modified CliMA-Land radiative transfer with  
1429 clumping index and the default version for multiple canopy densities and over a black soil albedo  
1430 (BLK) as in **Fig.10a.** with clumping calculated through **Eq.(2)** for sun zenith angles from 0° to 30°.  
1431 For CliMA-Land radiative transfer without clumping index the AIC = - 4923.44 and the BIC = -  
1432 4907.90, while for CliMA-Land radiative transfer with clumping index the AIC = - 5291.47 and  
1433 the BIC = - 5275.94.